

# Drought assessment in the northern region of Colombia using the Standardized Precipitation Index (SPI): a case study in the department of La Guajira

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## Research Article

**Keywords:** Droughts, Standardized Precipitation Index (SPI), La Guajira, operational characteristics, temporal trends

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2 **(SPI): a case study in the department of La Guajira**

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18  
19 **Abstract**

20  
21 The objective of this study was to assess droughts in the department of La Guajira, Columbia, on the basis  
22 of their operational characteristics. Droughts were assessed using the Standardized Precipitation Index  
23 (SPI) at three- and six-month aggregation periods (SPI-3 and SPI-6), and their operational characteristics  
24 were analyzed by the run theory and via analysis of temporal trends using the modified Mann-Kendall  
25 (MMK) test. The results indicated that droughts were most frequent in La Guajira between 1995 and 2004.  
26 The occurrence of droughts was most accurately identified by SPI-6. It was observed that the central,  
27 southern, and western areas of the department had droughts of greater duration, severity, and intensity. The  
28 MMK test shows negative (decreasing) temporal trends at significance levels  $\alpha$  between 0.1 and 0.01 in  
29 6.12% of the meteorological stations located in the central and southern areas of La Guajira. These results  
30 support the conclusion that droughts are recurrent events in the department of La Guajira as a result of the  
31 arid and semi-arid climate prevalent in significant portions of the department's land area. This elucidates  
32 the vulnerability of agriculture and livestock in such areas that are prone to droughts of greater duration,  
33 severity, and intensity.

34  
35 **Keywords:** Droughts, Standardized Precipitation Index (SPI), La Guajira, operational characteristics,  
36 temporal trends.

## 1. Introduction

Droughts are extreme climatic events that occur worldwide, especially in arid and semi-arid regions (Touchan et al. 2005). They are characterized by a reduction in rainfall over a prolonged period of time and are currently a major concern owing to their negative impacts on ecosystems, agricultural production (Yu et al. 2014), the environment, and people due to climate change, changes in land use, degradation of natural resources, and impact on social and economic systems (Thomas et al. 2015; Touchan et al. 2005).

Drought is primarily caused by deficient rainfall, and the duration, distribution, and intensity of this deficiency in relation to existing water storage, demand, and usage impacts its probability and intensity (Asadi Zarch et al. 2015). Three well-known categories of droughts have been recorded in the literature: (1) meteorological drought, (2) hydrological drought, and (3) agricultural drought. Meteorological drought is defined as the absence of rainfall in a region for a particular period of time (Belal, El-Ramady, Mohamed and Saleh, 2012; Mishra and Singh, 2010; Thomas et al. 2014). Hydrological drought follows a meteorological drought, and is defined as a deficit in the volume of surface and groundwater supply available to a particular water resource management system that is responsible for fulfilling the demands of an area (Thomas et al. 2014). Finally, an agricultural drought occurs when there is insufficient soil moisture and, consequently, leads to crop yield failures (Karabulut, 2015; Luetkemeier, Stein, Drees and Liehr, 2017; Mishra and Singh, 2010).

A drought is generally assessed using drought indices. Many such indices are available globally, and each has its own strengths and weaknesses. The most notable ones are the Palmer Drought Severity Index (Palmer, 1965), Standardized Precipitation Index (Mckee et al. 1993), Standardized Precipitation Evapotranspiration Index (Vicente-Serrano et al., 2010), and Effective Drought Index (Byun and Wilhite, 1999), among others (Mishra and Singh, 2010). The Standardized Precipitation Index (SPI) proposed by Mckee et al. (1993) has become popular worldwide as it can be used for the assessment of different time scales and supports analysis of different drought categories. Therefore, SPI is considered to be among the most robust and effective drought indices. Furthermore, as it only requires rainfall data, it can be calculated with greater ease compared to other, more complex, indices. It also supports the comparison of drought conditions in different regions and for different time scales (Caloiero, 2017).

Several studies that investigated various drought-related issues have demonstrated the wide applicability of SPI in the United States (Logan et al. 2010; Mitra and Srivastava, 2017), Asia (Sönmez et al. 2005; Shahid and Behrawan, 2008; He et al. 2015; Nam et al. 2015; Zuo et al. 2016; Dahal et al. 2016; Ghosh and Srinivasan, 2016; Rahman and Lateh, 2016; Yan et al. 2017), Africa (Guenang and Mkankam Kamga, 2014; Luetkemeier et al. 2017; Manatsa et al. 2010), Europe (Ashraf and Routray, 2015; Bonaccorso et al. 2015; Caloiero, 2017; Ionita et al. 2016; Livada and Assimakopoulos, 2007), and Latin America (Benitez and Domecq, 2014; Guimarães et al. 2017; Mendoza and Puche, 2007; Rivera et al. 2007; Rivera and Penalba, 2014; Toná Juliani and Passos Okawua, 2017).

77 The existing scientific literature on drought as a natural hazard is limited in Colombia. Significant  
78 contributions include an important local-scale investigation by Loaiza et al. (2015) to characterize  
79 meteorological droughts in the Dagua river basin, studies on the spatiotemporal analysis of droughts and  
80 their relationship with the warm phase of the *El Niño/La Niña-Oscilación del Sur* (ENSO) climate  
81 variability phenomenon in the city of Cali (Loaiza-Cerón et al. 2020), comparisons of regional droughts in  
82 the Sumapaz and Lebrija basins based on their Severity-Duration-Frequency curves (Rojas and Díaz-  
83 Granados, 2018), and technical reports from the Institute of Hydrology, Meteorology and Environmental  
84 Studies (IDEAM) to understand droughts in the country (Gómez, 2016; Hurtado, 2012; Mayorga and  
85 Hurtado, 2006) following the recommendation of the World Meteorological Organization (WMO) to use  
86 the SPI for characterizing droughts.

87

88 Geographically, the department of La Guajira, in Colombia, is among the areas that are most sensitive to  
89 the effects of climate change in South America, and is especially sensitive to increased intensity and  
90 frequency of hydrometeorological hazards such as droughts, which pose a threat to human life and  
91 socioeconomic activities. However, there are no studies assessing droughts as a natural hazard in the  
92 department and which contribute to its understanding from an operational perspective.

93

94 Therefore, considering the unique characteristics of the department of La Guajira due to the arid and semi-  
95 arid climate in a large proportion of its territory, and the presence of highly vulnerable areas with water-  
96 dependent human life and socioeconomic activities, this study aimed to assess droughts in the region using  
97 SPI and to define them in terms of their operational characteristics: duration, severity, and intensity.

98

## 99 **2. Materials and methods**

100

### 101 *2.1. Study area*

102

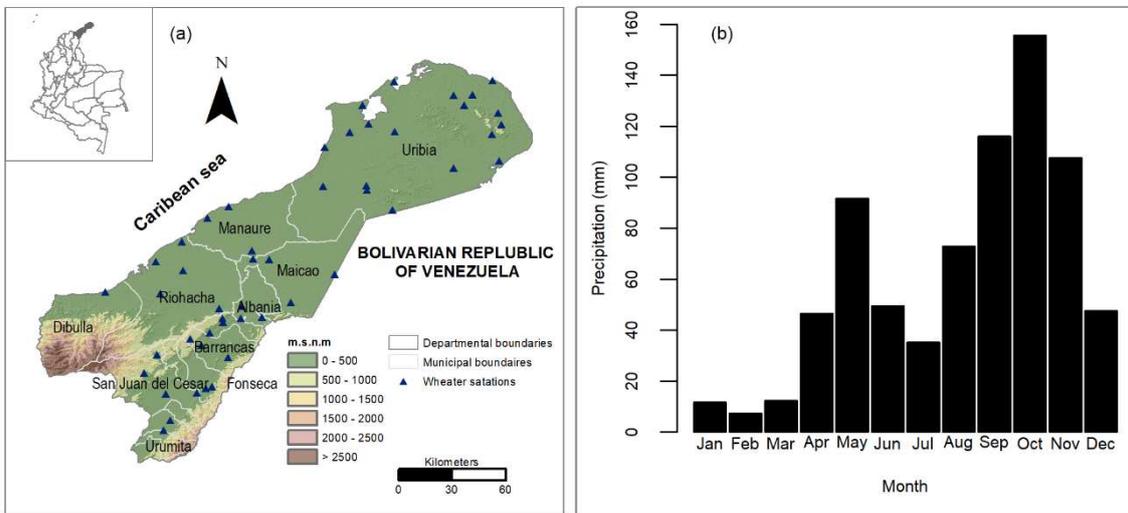
103 The department of La Guajira is located in the northernmost region of the Republic of Colombia and South  
104 America (CORPOGUAJIRA, 2012) on a peninsula of the same name, and is part of the Caribbean region.  
105 It is bordered to the north and west by the Caribbean Sea, to the east by the Bolivarian Republic of  
106 Venezuela and the Gulf of Venezuela, to the south by the department of Cesar, and to the southwest by the  
107 department of Magdalena (Fig. 1a).

108

109 Geographically, La Guajira is located between 10°23'N and 12°28'N, and 71°06'W and 73°39'W, and  
110 occupies an area of approximately 20,670 km<sup>2</sup>. It has an arid, dry, and tropical climate with high  
111 temperature. Rainfall is mainly governed by the movement of the trade winds, which create conditions of  
112 severe dryness and aridity when they pass through the northern mountains and the great plains, and cause  
113 rainfall upon collision with the Sierra Nevada de Santa Marta mountain range. As a result, rainfall increases  
114 from northeast to southwest with mean annual values ranging from 200 mm in the northeast of La Guajira  
115 to approximately 2,000 mm in the southwest regions, near the Sierra Nevada de Santa Marta.

116

117 In general, the rainfall regime in La Guajira is bimodal (Fig. 1b); with scarce rainfall occurring across two  
 118 distinct seasons: the first from April to mid-June, and the second from the end of September to mid-  
 119 December (Government of La Guajira, 2012).  
 120



122 **Fig. 1** Study area: a) Location of the department of La Guajira and spatial distribution of selected  
 123 climatological stations; b) Monthly distribution of historical mean rainfall in the selected climatological  
 124 stations

125  
 126  
 127

## 2.2. Data

128 The historical records of total monthly rainfall for the period from 1985 to 2015 (31 years) in 49  
 129 meteorological stations (Fig. 1a) were provided by the IDEAM (<http://www.ideam.gov.co/>) to analyze  
 130 droughts in the department of La Guajira. One of the limitations of using rainfall data of the department  
 131 was the missing data in the time series of some stations; therefore, a selection criteria was imposed wherein  
 132 only stations with less than 10% of missing data in their records for the selected study period were used  
 133 (Dahal et al. 2016). The imputation of missing data was carried out by using the mean value of rainfall  
 134 totals of the same month from previous and subsequent years (He et al. 2011; Rahman and Lateh, 2016).

135  
 136  
 137

## 2.3. Standardized Precipitation Index (SPI)

138 Drought conditions were assessed using SPI at three- and six-month periods, which reflect soil moisture  
 139 conditions and impact on crop production, i.e., parameters that were associated with the assessment of  
 140 agricultural droughts (Caloiero, 2017). To calculate SPI, rainfall records at each station were fitted to a  
 141 probability distribution function, specifically, the gamma function (Ashraf and Routray, 2015; Caloiero,  
 142 2017; Chang et al. 2016; He et al. 2015; Mansouri et al. 2013). This was then transformed to a standard  
 143 normal distribution, such that the mean SPI was zero and represented the normal moisture conditions of the  
 144 study area, thus indicating the area's mean rainfall. Therefore,  $SPI < 0$  reflects drought conditions in the  
 145 area of interest and  $SPI > 0$  indicates an excess of water (Chang et al. 2016).

146

147 The mathematical procedure for calculating SPI is briefly described below. Initially, the historical records  
 148 of total rainfall for each station at the selected time scale were fitted to a gamma probability distribution  
 149 function:

150

$$151 \quad g(x) = \frac{1}{\beta^\alpha \Gamma(\alpha)} x^{\alpha-1} e^{-x/\beta}, \quad x > 0 \quad (1)$$

152

153 Where,  $\alpha > 0$  is a shape parameter,  $\beta > 0$  is a scale parameter, and  $x > 0$  is the rainfall.  $\Gamma(\alpha)$  is the  
 154 gamma function, which is defined as:

155

$$156 \quad \Gamma(\alpha) = \int_0^{\infty} y^{\alpha-1} e^{-y} dy \quad (2)$$

157

158 The unknown parameters  $\alpha$  and  $\beta$  were estimated using the maximum likelihood approximation of Thom  
 159 (1966) for each season, each time scale (three and six months), and each month of the year as shown below:

160

161

$$162 \quad \hat{\alpha} = \frac{1}{4A} \left( 1 + \sqrt{1 + \frac{4A}{3}} \right) \quad (3)$$

163

$$164 \quad \hat{\beta} = \frac{\bar{x}}{\hat{\alpha}} \quad (4)$$

165

166 Where,

167

$$168 \quad A = \ln(\bar{x}) - \frac{1}{n} \sum_{i=1}^n \ln(x_i) \quad (5)$$

169

170 Here,  $n$  is the number of rainfall observations,  $\bar{x}$  is the mean rainfall for the time scale of interest, and  $A$   
 171 is a measure of the asymmetry of the distribution (Husak et al. 2007).

172

173 Integrating the probability distribution function with respect to  $x$  and inserting the estimated parameters  
 174  $\hat{\alpha}$  and  $\hat{\beta}$  produces the following expression for the cumulative probability  $G(x)$  of an observed rainfall  
 175 event occurring in a given month and time scale:

176

$$177 \quad G(x) = \frac{1}{\hat{\beta}^{\hat{\alpha}} \Gamma(\hat{\alpha})} \int_0^x x^{\hat{\alpha}} e^{-x/\hat{\beta}} dx \quad (6)$$

178

179 Since the gamma distribution is undefined for rainfall values of  $x = 0$ ;  $G(x)$  should be modified as  
180 follows for considering the zero values occurring in the rainfall records:

181

$$182 \quad H(x) = q + (1 - q)G(x) \quad (7)$$

183

184 Where  $q$  is the probability that  $x = 0$  (rainfall equal to zero) occurs, calculated as the quotient of the  
185 number of zeros in the precipitation series ( $m$ ) and the total number of observations ( $n$ ) (Thom, 1966).

186

187 Finally, the cumulative probability function  $H(x)$  is transformed to a variable  $Z$  that follows a standard  
188 normal distribution with zero mean and standard deviation equal to unity, which represents the SPI value,  
189 using the approximate conversion suggested by Abramowitz and Stegun (1965):

190

$$191 \quad Z = SPI = - \left( t - \frac{c_0 + c_1 t + c_2 t^2}{1 + d_1 t + d_2 t^2 + d_3 t^3} \right) \quad 0 < H(x) \leq 0.5 \quad (8)$$

$$192 \quad Z = SPI = + \left( t - \frac{c_0 + c_1 t + c_2 t^2}{1 + d_1 t + d_2 t^2 + d_3 t^3} \right) \quad 0.5 < H(x) < 1.0 \quad (9)$$

193

194 Where,

195

$$196 \quad t = \sqrt{\ln \left( \frac{1}{(\ln H(x))^2} \right)} \quad 0 < H(x) \leq 0.5 \quad (10)$$

$$197 \quad t = \sqrt{\ln \left( \frac{1}{(1 - \ln H(x))^2} \right)} \quad 0.5 < H(x) < 1.0 \quad (11)$$

198

$$199 \quad c_0 = 2.515517, \quad d_1 = 1.432788$$

$$200 \quad c_1 = 0.802853, \quad d_2 = 0.189269$$

$$201 \quad c_2 = 0.010328, \quad d_3 = 0.001308$$

202

203 In order to optimize the calculation of SPI for all climatological stations at the department of La Guajira  
204 and at the selected time scales, the *SPEI* package (Begueria et al. 2017) of R v3.5.1 (R Core Team, 2018)  
205 was used.

206

207 Table 1 shows the classification of droughts according to SPI as proposed by Mckee et al. (1993). Further  
 208 information on SPI can be found in Mckee et al. (1993), Hayes et al. (1999), and Lloyd-Hughes and  
 209 Saunders (2002).

210

211 **Table 1.** Drought categories defined by SPI

SPI	Drought category	Probability of occurrence (%)
$-1.0 \leq SPI < 0.0$	Mild drought	34.1
$-1.5 \leq SPI < -1.0$	Moderate drought	9.2
$-2.0 \leq SPI < -1.5$	Severe drought	4.4
$SPI < -2.0$	Extreme drought	2.3

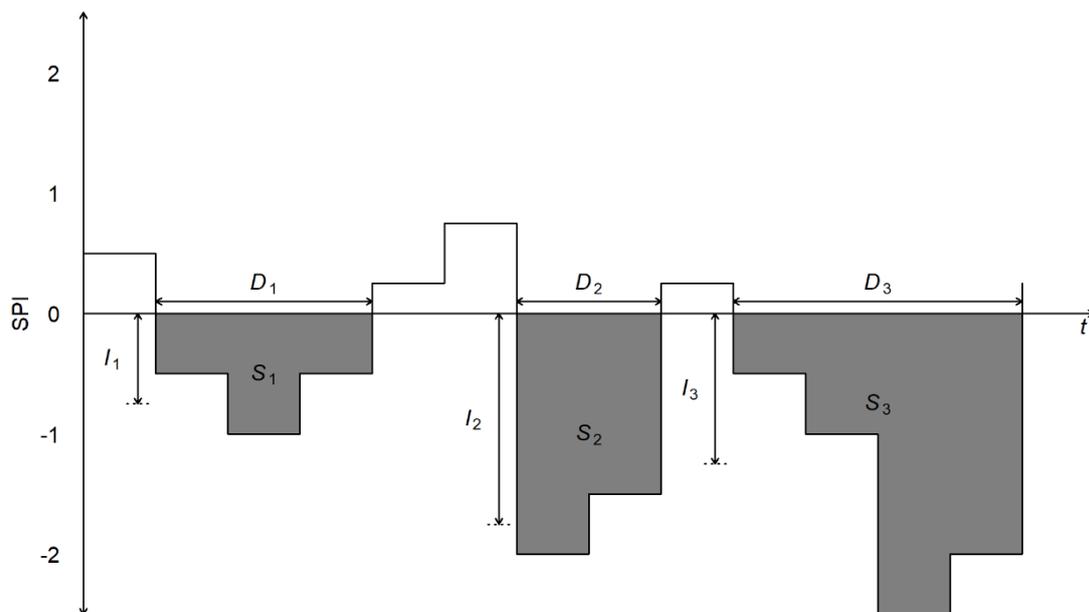
212

213 *2.4. Operational characteristics of droughts*

214

215 The operational characteristics of droughts were determined by the run theory proposed by Yevjevich  
 216 (1967). A run is defined as a portion of a one-variable time series—in this case the SPI time series—in  
 217 which each value is above or below a given threshold (Fig. 2). According to the classification of droughts  
 218 shown in Table 1, the threshold level selected to define droughts is -1. Therefore, based on Fig. 2, the  
 219 drought characteristics were defined as follows: (1) duration of a drought ( $D$ ) is the time elapsed between  
 220 the beginning and the end of a drought, with SPI values continuously below the threshold; (2) drought  
 221 severity (magnitude) ( $S$ ) is the dimensionless sum of SPI values below the threshold during the drought  
 222 event; and (3) drought intensity ( $I$ ) is the mean severity during the drought event and is calculated as the  
 223 severity divided by the duration.

224



226

**Fig. 2** Drought characteristics using the run theory of Yevjevich (1967)

227

228 *2.5. Spatial and temporal drought analysis*

229

230 The spatial distribution of droughts in the department of La Guajira was mapped using the inverse distance  
231 weighted (IDW) interpolation method (Manikandan and Tamilmani, 2015; Parker et al. 2009; Yuan et al.  
232 2015). The IDW interpolation method assumes that the values of the variable to be processed (drought) in  
233 areas with no observed data is a mean of the inverse weighted distance of the observed values of the variable  
234 at sampling points located in the vicinity; therefore, the closer the points with observed data are to the  
235 estimation areas, the more similar the values will be to those of more distant areas. The spatial interpolation  
236 of SPI data in the department of La Guajira using the IDW method was performed with GIS (ArcGis  
237 Desktop 10.5).

238

239 The temporal trends of drought in different meteorological stations of La Guajira were analyzed using the  
240 Mann-Kendall (MK) test, (Mann, 1945; Kendall, 1975); this a non-parametric rank-based test that is widely  
241 used for detecting trends in time series of hydrological variables and drought indicators (Ashraf and  
242 Routray, 2015; Chang et al. 2016; Dahal et al. 2016; Hong et al. 2014). This test considers that the MK  
243 trend test statistic for a series of observations is calculated by

244

$$245 \quad S = \sum_{i < j} \text{sng}(x_j - x_i) \quad (12)$$

246

247 Where,

$$248 \quad \text{sng}(x_j - x_i) = \begin{cases} 1, & x_i < x_j \\ 0, & x_i = x_j \\ -1, & x_i > x_j \end{cases} \quad (13)$$

249

250 Subsequently, under the hypothesis that the observations are independent and randomly ordered or  
251 identically distributed, the statistic  $S$  tends to be normally distributed when  $n$  is large, with mean

252  $E(S) = 0$  and variance given by

253

$$254 \quad \text{Var}(S) = \frac{n(n-1)(2n+5)}{18} \quad (14)$$

255

256 Therefore, the significance of the trends is assessed by comparing the standardized statistical test

257  $Z = S / \sqrt{\text{Var}(S)}$  with the standard normal distribution at the desired level of significance. Hence, the

258 hypothesis of no significant trend is rejected when  $|Z| > Z_{1-\alpha/2}$ . Positive  $Z$  values indicate an increase in

259 the trend and vice versa (Ashraf and Routray, 2015).

260

261 However, the efficacy of the MK test is reduced by the existence of autocorrelation in the time series,  
262 expressed as disturbances in the variance of  $S$ . For this reason, Hamed and Rao (1998) proposed the  
263 modified Mann-Kendall (MMK) test as a procedure that considers the existence of autocorrelation in the  
264 time series and removes its influence on trend detection by using the autocorrelation coefficients at a lag  $i$   
265 ( $\sigma_s(i)$ ) that are significantly different from zero at a level of significance of 0.05 to calculate the modified  
266 variance of  $S$ , which is given by:

267

$$268 \quad V'(S) = \text{Var}(S)\text{Cor} \quad (15)$$

269

270 Where 'Cor' is a correction given to the autocorrelation of the time series calculated as

271

$$272 \quad \text{Cor} = 1 + \frac{2}{n(n-1)(n-2)} \sum_{i=1}^{n-1} (n-i)(n-i-1)(n-i-2)\sigma_s(i) \quad (16)$$

273

274 The MMK test for the significance of trends in SPI time series in this study was analyzed at a level of  
275 significance of 0.05 using the *modifiedmk* library (Marx, 2012) of R v3.5.1 (R Core Team, 2018).

276

### 277 **3. Results and discussion**

278

#### 279 *3.1. Characterization of drought events that occurred between 1985 and 2015.*

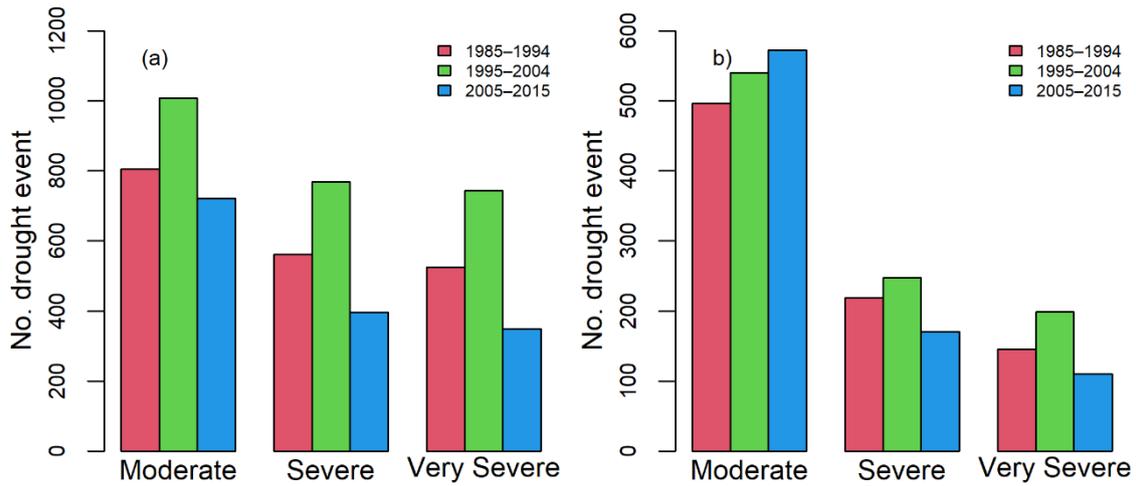
280

281 The time series analysis of total monthly rainfall in the 49 meteorological stations for 1985–2015 revealed  
282 the temporal variability of drought events that occurred in the department of La Guajira during this period.  
283 Fig. 3a and Fig. 3b show the distribution of the number of identified drought events in the meteorological  
284 stations during the subperiods of 1985–1994, 1995–2004, and 2005–2015 for different drought categories  
285 according to SPI-3 and SPI-6, respectively. Note that the subperiod of 1995–2004 had the highest number  
286 of drought events for all categories, according to SPI-3. Similar results were obtained for SPI-6, except for  
287 moderate droughts, for which the highest number of drought events occurred in the subperiod of 2005–  
288 2015.

289

290 These results correspond to those obtained by IDEAM (2016) in a study which concluded that the highest  
291 frequency of droughts with important and lasting territorial effects occurred in the Colombian-Caribbean  
292 region, specifically, in La Guajira during 1991–92, 1997–98, 2001, and 2009. Additionally, there were  
293 other, smaller drought events of lower intensity or duration during 1982–83, 1986–1987, and 2002–2003.

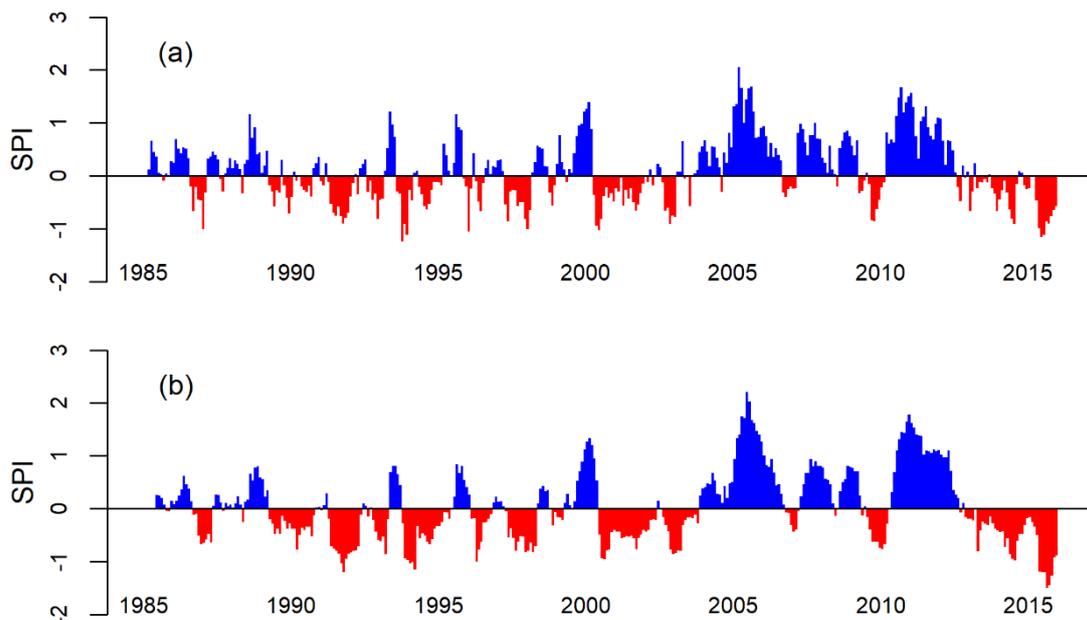
294



296 **Fig. 3** Occurrence of drought events in the subperiods of 1985–1994, 1995–2004, and 2005–2015  
 297 according to: (a) SPI-3 and (b) SPI-6  
 298

299 A total of 8,081 drought events were identified in the department of La Guajira between 1985 and 2015.  
 300 Of these, 5,380 and 2,701 correspond to droughts identified by SPI-3 and SPI-6, respectively, representing  
 301 29.5% and 14.8% of the total number of possible drought events (18,229) during the study period.  
 302

303 The temporal variations of the mean occurrence of drought events in La Guajira for the temporal  
 304 aggregations of SPI-3 and SPI-6 are illustrated in Fig. 4. It was observed that the droughts assessed by SPI-  
 305 3 exhibited higher frequencies of occurrence compared to SPI-6. However, the droughts assessed by SPI-6  
 306 were of longer duration and greater severity and intensity.  
 307



309 **Fig. 4** Temporal variability of the mean occurrence of droughts in the department of La Guajira during the  
 310 period of 1985–2015 for: (a) SPI-3 and (b) SPI-6

311

312 Comparison of the temporal aggregations of SPI-3 and SPI-6 shows that the department of La Guajira is  
313 predominantly characterized by the occurrence of frequent and moderate droughts. SPI-3 values show high  
314 fluctuations about the zero line, indicating that droughts in La Guajira assessed by SPI-3 were mostly caused  
315 by short-term rainfall variations. Conversely, SPI-6 shows greater stability, and its values show less  
316 variation about the zero line and a greater smoothness in the time series. Therefore, SPI-6 is considered a  
317 better indicator of seasonal droughts in the department of La Guajira. Although SPI-3 accurately reflects  
318 the changes in rainfall for a particular month, and, consequently, the onset and duration of a drought event,  
319 it is not the optimal temporal aggregation to characterize the persistence of droughts. In contrast, SPI-6  
320 identifies fluctuations in rainfall and the persistence of drought events; however, when compared to SPI-3,  
321 it is less reliable in identifying their onset and termination (Chang et al. 2016; Manikandan and Tamilmani,  
322 2015).

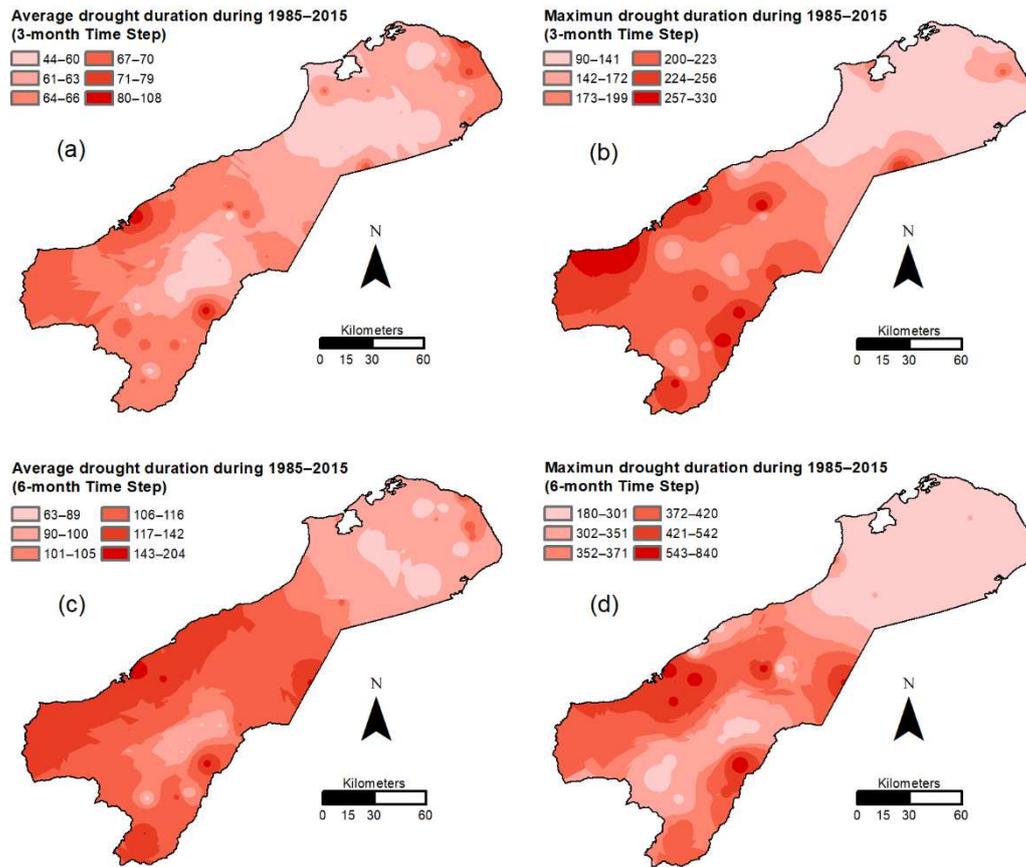
323

### 324 *3.2. Operational characteristics of droughts that occurred between 1985 and 2015.*

325

326 With regards to the main operational characteristics of droughts, the distribution of the mean and maximum  
327 values of drought duration for all events occurring in the study area during the period of 1985–2015—based  
328 on SPI-3 temporal aggregations—indicate that droughts generally affect the entire territory of the  
329 department of La Guajira (Fig. 5). However, droughts with the longest mean duration (67–108 days) were  
330 found to occur in the northern, southern, and western areas of the department (Fig. 5a). It was also evident  
331 that droughts of maximum duration (200–330 days) occurred in the southern and western areas of the  
332 department (Fig. 5b). On the contrary, the results of the SPI-6 temporal aggregation show that droughts of  
333 greater mean (106–204 days) and maximum (372–840 days) duration occurred in the central, southern, and  
334 western areas of the department of La Guajira as shown in Fig. 5c and Fig. 5d, respectively.

335

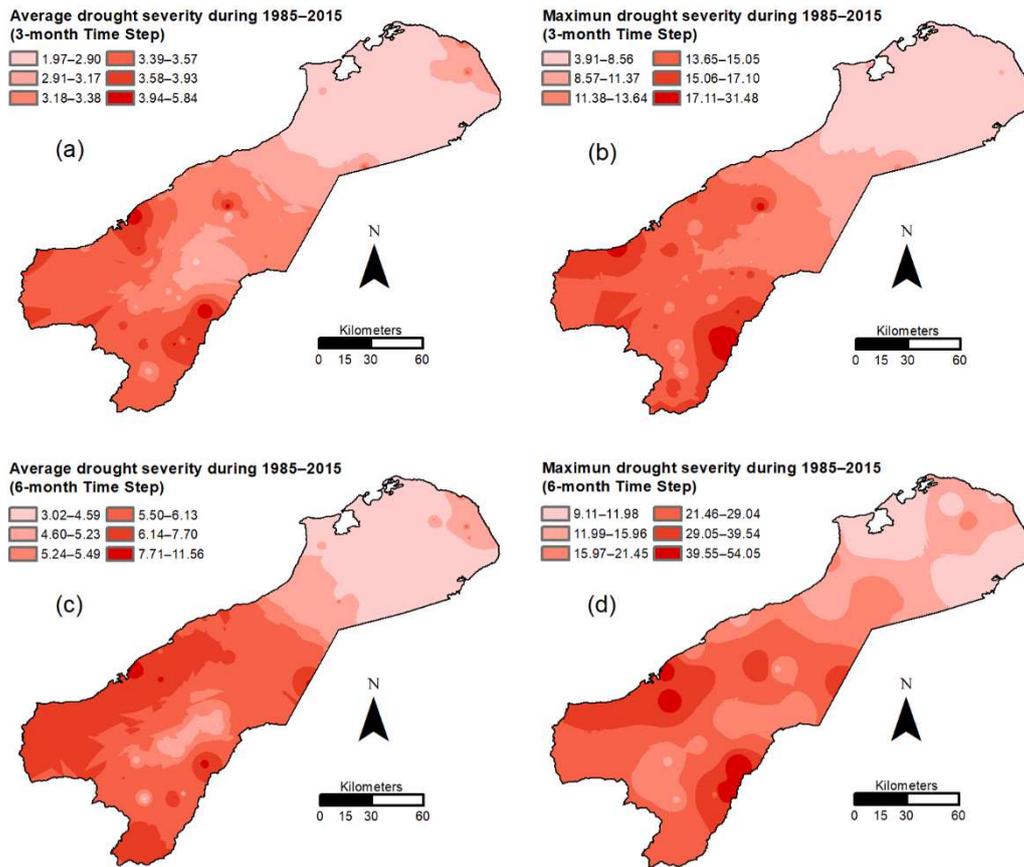


337 **Fig. 5** Spatial distribution of drought duration in the department of La Guajira during the period of 1985–  
 338 2015

339

340 In agreement with the results obtained for spatial distribution of drought duration in the department of La  
 341 Guajira during the period of 1985–2015, it was observed that the greatest values of mean (3.39–5.84) and  
 342 maximum (13.65–31.48) drought severity occurred in the central, southern, and western areas of the  
 343 department for SPI-3 (Fig. 6a and Fig. 6b). Similar results were observed for mean (5.50–11.56) and  
 344 maximum (21.46–54.05) severity, as shown in Fig. 6c and Fig. 6d, respectively. However, the SPI-6  
 345 temporal aggregation results show higher absolute mean and maximum severity values compared to those  
 346 obtained from SPI-3. Furthermore, they also show a more focused spatial distribution pattern within the  
 347 areas affected by droughts.

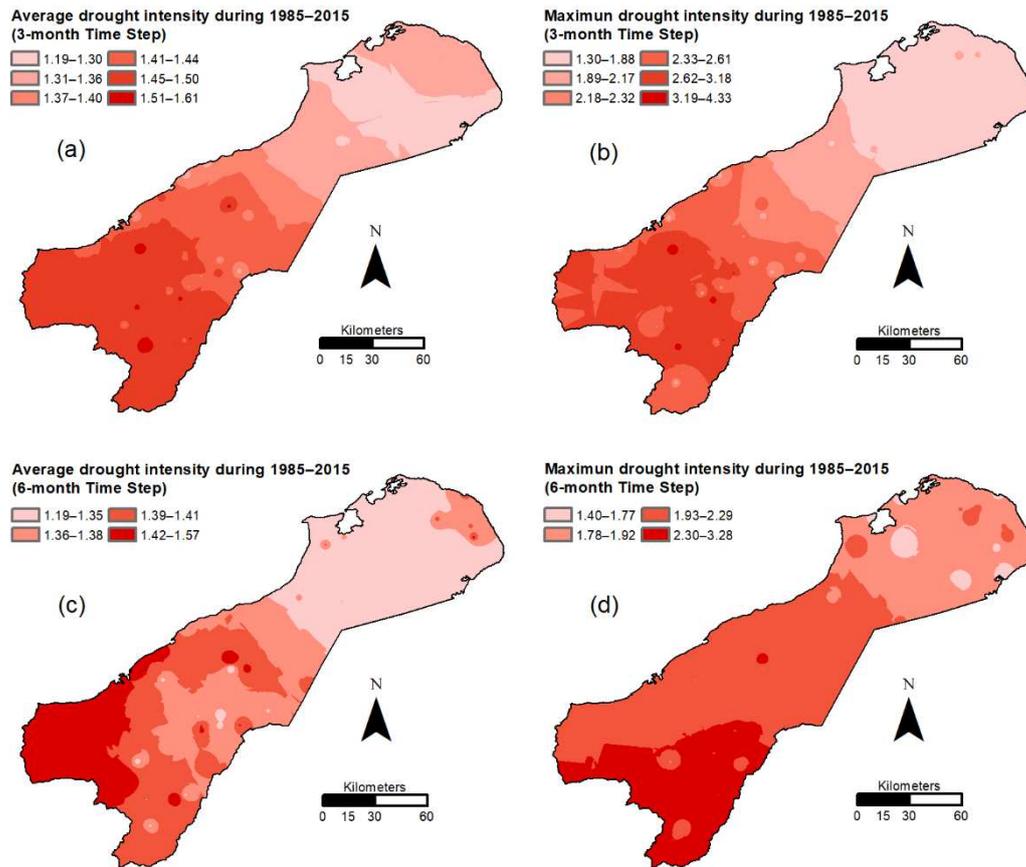
348



350 **Fig. 6** Spatial distribution of drought severity in the department of La Guajira during the period of 1985–  
 351 2015

352  
 353 Finally, when considering the intensity of droughts occurring in the department of La Guajira during the  
 354 study period (Fig. 7), it was observed that the highest mean (1.41–1.61) and maximum (2.33–4.33)  
 355 intensities for SPI-3 were predominantly distributed in the central, southern, and western areas of the  
 356 department (Fig. 7a and Fig. 7b, respectively). Similar results, albeit in a more focused manner, were  
 357 obtained for the greatest mean (1.39–1.57) and maximum (1.93–3.28) values for SPI-6 temporal  
 358 aggregation, as shown in Fig. 7c and Fig. 7d, respectively. It is notable that, despite the notorious spatial  
 359 association, SPI-3 showed slightly higher values of mean and maximum intensity than those obtained for  
 360 SPI-6. As this characteristic is derived from the previously discussed results regarding duration and  
 361 severity, and has an inverse relationship with drought duration, it reaffirms the conclusion that these areas  
 362 of the department were the most affected by droughts of high duration and severity.

363



365 **Fig. 7** Spatial distribution of drought intensity in the department of La Guajira during the period of 1985–  
 366 2015

367

368 The previous results are indicative of a spatial correlation between the operational characteristics of the  
 369 droughts that occurred in the department of La Guajira in the period of 1985–2015 for the analyzed temporal  
 370 aggregations (SPI-3 and SPI-6). This indicates that areas with droughts of longer duration correspond to  
 371 those areas that experience droughts of greater severity and intensity. Therefore, these are the areas of the  
 372 department where water-dependent socioeconomic activities, such as agriculture and livestock activities,  
 373 are more vulnerable (Contreras et al., 2020).

374

### 375 *3.3. Temporal drought trends SPI-3 and SPI-6*

376

377 The trend analysis of the time series of drought events that occurred in the department of La Guajira in the  
 378 period of 1985–2015 through the MMK test at different levels of significance  $\alpha$  (0.1, 0.05, 0.01, and 0.001)  
 379 showed that the majority of meteorological stations in the department did not generally exhibit significant  
 380 trends in the occurrence of drought events of greater magnitude in successive years of the study period,  
 381 with percentages of 75.51% and 81.63% for SPI-3 and SPI-6, respectively (Fig. 8). However, the two SPI  
 382 temporal aggregations revealed that some meteorological stations located in the central and southern areas  
 383 of the department showed negative (decreasing) temporal trends at levels of significance between 0.1 and  
 384 0.01, with a cumulative proportion of stations of 6.12% for SPI-3 and SPI-6.

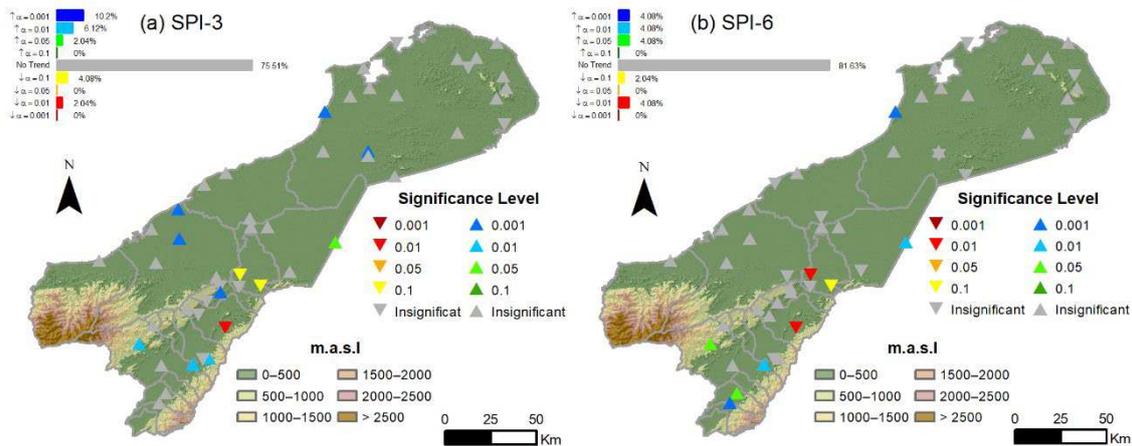
385

386 In contrast, positive (increasing) trends were observed in different areas of the department for events of  
387 greater magnitude that occurred in each year of the study period, especially in the northern, central, and  
388 southern areas at levels of significance between 0.05 and 0.001 and an accumulated proportion of  
389 meteorological stations of 18.36% for SPI-3 (Fig. 8). Similarly, positive trends were observed in the  
390 northern, eastern, and southern zones of the study area for SPI-6 in 12.24% of the meteorological stations  
391 at levels of significance between 0.05 and 0.001.

392

393 This indicates that, although droughts are a natural and recurring element of the climate in the department  
394 of La Guajira due to the arid and semi-arid climate prevailing across much of its area, there are some periods  
395 of extreme drought and others during which water excesses can occur. This is governed by a climate  
396 variability phenomenon known as ENSO. The warm phase of ENSO (*El Niño*) reduces rainfall and  
397 increases air temperature, thus inducing droughts; whereas the cold phase (*La Niña*) strengthens the rainy  
398 season, thus significantly increasing the region's water supply (Bedoya-Soto et al. 2019; Hoyos et al. 2013).

399



401 **Fig. 8** Trend analysis at different levels of significance for all meteorological stations in the department of  
402 La Guajira during 1985–2015 for: (a) SPI-3 and (b) SPI-6

403

#### 404 4. Conclusions

405

406 The findings of this study support the conclusion that droughts are recurrent events in the department of La  
407 Guajira due to the biophysical characteristics of the arid and semi-arid climate prevalent across most of its  
408 area. However, drought analysis during the period of 1985–2015 reveals that 1995–2004 was characterized  
409 by the highest occurrence of drought events, governed by climate variability phenomena such as ENSO in  
410 its warm phase (*El Niño*).

411

412 Both the temporal aggregations, SPI-3 and SPI-6, used for assessing droughts in this study were able to  
413 accurately identify variations in rainfall for a given month, especially when a rainfall deficit considering  
414 the climatological mean constituted a drought event. SPI-6 was shown to be a better indicator of seasonal  
415 droughts in the department as it identifies rainfall fluctuations and the persistence of deficit conditions  
416 during a drought event.

417

418 In terms of the operational characteristics of droughts identified in the department of La Guajira for the  
419 period of 1985–2015, the results show that the southern, central, and western areas of the department  
420 experienced droughts of greater duration, severity, and intensity, and were also the areas where water-  
421 dependent socioeconomic activities such as agriculture and livestock were most vulnerable.

422

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427

#### 428 **Statements and Declarations**

429 All authors certify that they have no affiliations with or involvement in any organization or entity with any  
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431

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#### References

433

434 Abramowitz M, Stegun IA (1965) Handbook of Mathematical Functions With Formulas, Graphs and  
435 Mathematical Tables (National Bureau of Standards Applied Mathematics Series No. 55). Journal of  
436 Applied Mechanics, Vol. 32, p. 239. <https://doi.org/10.1115/1.3625776>

437 Asadi Zarch MA, Sivakumar B, Sharma A (2015) Droughts in a warming climate: A global assessment of  
438 Standardized precipitation index (SPI) and Reconnaissance drought index (RDI). Journal of Hydrology,  
439 526, 183–195. <https://doi.org/10.1016/j.jhydrol.2014.09.071>

440 Ashraf M, Routray JK (2015) Spatio-temporal characteristics of precipitation and drought in Balochistan  
441 Province, Pakistan. Natural Hazards, 77(1), 229–254. <https://doi.org/10.1007/s11069-015-1593-1>

442 Bedoya-Soto JM, Poveda G, Trenberth KE, Vélez-Upegui JJ (2019) Interannual hydroclimatic variability  
443 and the 2009–2011 extreme ENSO phases in Colombia: from Andean glaciers to Caribbean lowlands.  
444 Theoretical and Applied Climatology, 135(3–4), 1531–1544. <https://doi.org/10.1007/s00704-018-2452-2>

445 Begueria S, Serrano V, Sawasawa H (2017) SPEI: Calculation of Standardised Precipitation-  
446 Evapotranspiration index. R package version 1.7. A Case Study Birkoor Kortigiri Mandals.  
447 <https://doi.org/10.1175/2009JCLI2909.1>

- 448 Belal AA, El-Ramady HR, Mohamed ES, Saleh AM (2012) Drought risk assessment using remote sensing  
449 and GIS techniques. *Arabian Journal of Geosciences*, 7(1), 35–53. [https://doi.org/10.1007/s12517-012-](https://doi.org/10.1007/s12517-012-0707-2)  
450 0707-2
- 451 Benitez, JB, Domecq RM (2014) Analysis of meteorological drought episodes in Paraguay. *Climatic*  
452 *Change*, 127(1), 15–25. <https://doi.org/10.1007/s10584-014-1260-7>
- 453 Bonaccorso B, Cancelliere A, Rossi G (2015) Probabilistic forecasting of drought class transitions in Sicily  
454 (Italy) using Standardized Precipitation Index and North Atlantic Oscillation Index. *Journal of Hydrology*,  
455 526, 136–150. <https://doi.org/10.1016/j.jhydrol.2015.01.070>
- 456 Byun HR, Wilhite DA (1999) Objective quantification of drought severity and duration. *Journal of Climate*,  
457 12(9), 2747–2756. [https://doi.org/10.1175/1520-0442\(1999\)012<2747:OQODSA>2.0.CO;2](https://doi.org/10.1175/1520-0442(1999)012<2747:OQODSA>2.0.CO;2)
- 458 Caloiero T (2017) Drought analysis in New Zealand using the standardized precipitation index.  
459 *Environmental Earth Sciences*, 76(16). <https://doi.org/10.1007/s12665-017-6909-x>
- 460 Chang J, Li Y, Ren Y, Wang Y (2016) Assessment of precipitation and drought variability in the Weihe  
461 River Basin, China. *Arabian Journal of Geosciences*, 9(14). <https://doi.org/10.1007/s12517-016-2638-9>
- 462 Contreras D, Voets A, Junghardt J, Bhamidipati S, Contreras S (2020) The Drivers of Child Mortality  
463 During the 2012–2016 Drought in La Guajira, Colombia. *International Journal of Disaster Risk Science*,  
464 11(1), 87–104. <https://doi.org/10.1007/s13753-020-00255-0>
- 465 Corporación Autónoma Regional de La Guajira (CORPOGUAJIRA), I. de I. M. y C. "José B. V. de A.  
466 (INVEMAR) (2012) Atlas Marino Costero de la Guajira. Retrieved from  
467 <https://colaboracion.dnp.gov.co/CDT/Inversiones y finanzas pblicas/La Guajira 15-Ajustada.pdf>
- 468 Dahal P, Shrestha NS, Shrestha ML, Krakauer NY, Panthi J, Pradhanang SM, Jha A, Lakhankar T (2016)  
469 Drought risk assessment in central Nepal: temporal and spatial analysis. *Natural Hazards*, 80(3), 1913–  
470 1932. <https://doi.org/10.1007/s11069-015-2055-5>
- 471 Ghosh S, Srinivasan K (2016) Analysis of Spatio-temporal Characteristics and Regional Frequency of  
472 Droughts in the Southern Peninsula of India. *Water Resources Management*, 30(11), 3879–3898.  
473 <https://doi.org/10.1007/s11269-016-1396-5>
- 474 Gómez J (2016) Estudio de Sequías en Colombia. In Instituto de Meteorología, Hidrología y Estudios  
475 Ambientales - IDEAM.
- 476 Guajira, G. de La. (2012). Plan Departamental de Gestión del Riesgo de Desastres: La Guajira. Retrieved  
477 from <http://hdl.handle.net/20.500.11762/381>
- 478 Guenang GM, Mkankam-Kamga F (2014) Computation of the standardized precipitation index (SPI) and  
479 its use to assess drought occurrences in Cameroon over recent decades. *Journal of Applied Meteorology*  
480 *and Climatology*, 53(10), 2310–2324. <https://doi.org/10.1175/JAMC-D-14-0032.1>
- 481 Hamed KH, Rao AR (1998) A modified Mann-Kendall trend test for autocorrelated data. *Journal of*  
482 *Hydrology*, 204(1–4), 182–196. [https://doi.org/10.1016/S0022-1694\(97\)00125-X](https://doi.org/10.1016/S0022-1694(97)00125-X)
- 483 Hayes M, Svoboda M, Whilite DA, Wilhite DA (1999) Monitoring drought using the standardized  
484 precipitation index. *Drought: A Global Assessment Volume 1*, 168–180.  
485 <https://doi.org/http://dx.doi.org/10.1108/17506200710779521>
- 486 He B, Lü A, Wu J, Zhao L, Liu M (2011) Drought hazard assessment and spatial characteristics analysis in  
487 China. *Journal of Geographical Sciences*, 21(2), 235–249. <https://doi.org/10.1007/s11442-011-0841-x>
- 488 He J, Yang XH, Li JQ, Jin JL, Wei YM, Chen XJ (2015) Spatiotemporal variation of meteorological  
489 droughts based on the daily comprehensive drought index in the Haihe River basin, China. *Natural Hazards*,  
490 75(2), 199–217. <https://doi.org/10.1007/s11069-014-1158-8>

- 491 Hong X, Guo S, Xiong L, Liu Z (2015) Spatial and temporal analysis of drought using entropy-based  
 492 standardized precipitation index: a case study in Poyang Lake basin, China. *Theoretical and Applied*  
 493 *Climatology*, 122(3–4), 543–556. <https://doi.org/10.1007/s00704-014-1312-y>
- 494 Hoyos I, Baquero-Bernal A, Jacob D, Rodríguez BA (2013) Variability of extreme events in the Colombian  
 495 Pacific and Caribbean catchment basins. *Climate Dynamics*, 40(7–8), 1985–2003.  
 496 <https://doi.org/10.1007/s00382-012-1487-9>
- 497 Hurtado, G. (2012). *Sequía Meteorológica Y Sequía Agrícola En Colombia: Incidencia Y Tendencias*.
- 498 Husak G, Michaelsen J, Funk C (2007) Use of the gamma distribution to represent monthly rainfall in  
 499 Africa for drought monitoring applications. *International Journal of Climatology*, 27(December 2007), 935-  
 500 944. <https://doi.org/10.1002/joc.1441>
- 501 IDEAM (2016) *Estudio de sequias en colombia*.
- 502 Influence C, Henrique B, Michiko C, Okawa P (2017) Application of a Standardized Precipitation Index  
 503 for Meteorological Drought Analysis of the Semi-Arid Climate Influence in Minas Gerais, Brazil.  
 504 *Hydrology*, 4(2), 26. <https://doi.org/10.3390/hydrology4020026>
- 505 Ionita M, Scholz P, Chelcea, S (2016) Assessment of droughts in Romania using the Standardized  
 506 Precipitation Index. *Natural Hazards*, 81(3), 1483–1498. <https://doi.org/10.1007/s11069-015-2141-8>
- 507 Karabulut M (2015) Drought analysis in Antakya-Kahramanmaraş Graben, Turkey. *Journal of Arid Land*,  
 508 7(6), 741–754. <https://doi.org/10.1007/s40333-015-0011-6>
- 509 Livada I, Assimakopoulos VD (2007) Spatial and temporal analysis of drought in Greece using the  
 510 Standardized Precipitation Index (SPI). *Theoretical and Applied Climatology*, 89(3–4), 143–153.  
 511 <https://doi.org/10.1007/s00704-005-0227-z>
- 512 Lloyd-Hughes B, Saunders MA (2002) A drought climatology for Europe. *International Journal of*  
 513 *Climatology*, 22(13), 1571–1592. <https://doi.org/10.1002/joc.846>
- 514 Loaiza-Cerón W, Carvajal-Escobar Y, De Souza RVA, Kayano MT, López NG (2020) Spatio-temporal  
 515 analysis of the droughts in Cali, Colombia and their primary relationships with the El Niño-Southern  
 516 Oscillation (ENSO) between 1971 and 2011. *Atmosfera*, 33(1), 51–69.  
 517 <https://doi.org/10.20937/ATM.52639>
- 518 Loaiza-Cerón W, Carvajal-Escobar Y, Baquero-Montoya OL (2015) Índice estandarizado de precipitación  
 519 (SPI) para la caracterización de sequías meteorológicas en la cuenca del río Dagua-Colombia. *Estudios*  
 520 *Geográficos*, 76(279), 557–578. <https://doi.org/10.3989/estgeogr.201520>
- 521 Logan KE, Brunsell NA, Jones AR, Feddema JJ (2010) Assessing spatiotemporal variability of drought in  
 522 the U.S. central plains. *Journal of Arid Environments*, 74(2), 247–255.  
 523 <https://doi.org/10.1016/j.jaridenv.2009.08.008>
- 524 Luetkemeier R, Stein L, Drees L, Liehr S (2017) Blended Drought Index: Integrated Drought Hazard  
 525 Assessment in the Cuvelai-Basin. *Climate*, 5(3), 51. <https://doi.org/10.3390/cli5030051>
- 526 Manatsa D, Mukwada G, Siziba E, Chinyanganya T (2010) Analysis of multidimensional aspects of  
 527 agricultural droughts in Zimbabwe using the Standardized Precipitation Index (SPI). *Theoretical and*  
 528 *Applied Climatology*, 102(3), 287–305. <https://doi.org/10.1007/s00704-010-0262-2>
- 529 Manikandan M, Tamilmanni D (2015) Spatial and Temporal Variation of Meteorological Drought in the  
 530 Parambikulam-Aliyar Basin, Tamil Nadu. *Journal of The Institution of Engineers (India): Series A*, 96(3),  
 531 177–184. <https://doi.org/10.1007/s40030-015-0121-3>
- 532 Mann HB (1945) Nonparametric Tests Against Trend. *The Econometric Society*, 13(3), 245–259.
- 533 Mansouri-Daneshvar MR, Bagherzadeh A, Khosravi M (2013) Assessment of drought hazard impact on  
 534 wheat cultivation using standardized precipitation index in Iran. *Arabian Journal of Geosciences*, 6(11),  
 535 4463–4473. <https://doi.org/10.1007/s12517-012-0695-2>

- 536 Marx W (2012) Tracking historical papers and their citations. *European Science Editing*, 38(2), 35–37.  
537 <https://doi.org/10.1023/B>
- 538 Mayorga R, Hurtado G (2006) La Sequía En Colombia Documento Técnico De Respaldo a La Información  
539 En La Página Web Del Ideam. Instituto de Meteorología, Hidrología y Estudios Ambientales - IDEAM,  
540 66.
- 541 Mckee TB, Doesken NJ, Kleist J (1993) The relationship of drought frequency and duration to time scales.  
542 AMS 8th Conference on Applied Climatology, (January), 179–184. [https://doi.org/citeulike-article-](https://doi.org/citeulike-article-id:10490403)  
543 [id:10490403](https://doi.org/citeulike-article-id:10490403)
- 544 Mendoza M, Puche N (2007) Evaluación de la ocurrencia de sequía en localidades de Venezuela Drought  
545 occurrence evaluation in localities of Venezuela Introducción. *Rev. Fac. Agron*, 24, 661–678.
- 546 Mishra AK, Singh, VP (2010) A review of drought concepts. *Journal of Hydrology*, 391(1–2), 202–216.  
547 <https://doi.org/10.1016/j.jhydrol.2010.07.012>
- 548 Mitra S, Srivastava P (2017) Spatiotemporal variability of meteorological droughts in southeastern USA.  
549 *Natural Hazards*, 86(3), 1007–1038. <https://doi.org/10.1007/s11069-016-2728-8>
- 550 Nam WH, Hayes MJ, Svoboda MD, Tadesse T, Wilhite DA (2015) Drought hazard assessment in the  
551 context of climate change for South Korea. *Agricultural Water Management*, 160, 106–117.  
552 <https://doi.org/10.1016/j.agwat.2015.06.029>
- 553 Palmer WC (1965) Meteorological Drought. U.S. Weather Bureau, Res. Pap. No. 45, p. 58. Retrieved from  
554 <https://www.ncdc.noaa.gov/temp-and-precip/drought/docs/palmer.pdf>
- 555 Parker DJ, Priest SJ, Tapsell SM (2009) Understanding and enhancing the public's behavioural response  
556 to flood warning information. *Meteorological Applications*, 114(January), 103–114.  
557 <https://doi.org/10.1002/met>
- 558 Rahman, M. R., & Lateh, H. (2016). Meteorological drought in Bangladesh: assessing, analysing and  
559 hazard mapping using SPI, GIS and monthly rainfall data. *Environmental Earth Sciences*, 75(12).  
560 <https://doi.org/10.1007/s12665-016-5829-5>
- 561 Rivera del Río R, Crespo-Pichardo G, Arteaga-Ramírez R, Quevedo-Nolasco A (2007) Temporal and  
562 Spatial Behavior of Drought in the State of Durango , Mexico. *Terra Latinoamericana*, 25, 383–392.
- 563 Rivera J, Penalba O (2014) Trends and Spatial Patterns of Drought Affected Area in Southern South  
564 America. *Climate*, 2(4), 264–278. <https://doi.org/10.3390/cli2040264>
- 565 Rojas LPT, Díaz-Granados M (2018) The construction and comparison of regional drought severity-  
566 duration-frequency curves in two Colombian River basins-study of the Sumapaz and Lebrija Basins. *Water*  
567 (Switzerland), 10(10). <https://doi.org/10.3390/w10101453>
- 568 Santos CAG, Brasil-Neto RM, Passos JSA, da Silva RM (2017) Drought assessment using a TRMM-  
569 derived standardized precipitation index for the upper São Francisco River basin, Brazil. *Environmental*  
570 *Monitoring and Assessment*, 189(6). <https://doi.org/10.1007/s10661-017-5948-9>
- 571 Shahid S, Behrawan H (2008) Drought risk assessment in the western part of Bangladesh. *Natural Hazards*,  
572 46(3), 391–413. <https://doi.org/10.1007/s11069-007-9191-5>
- 573 Sönmez FK, Kömüscü AÜ, Erkan A, Turgu E (2005) An analysis of spatial and temporal dimension of  
574 drought vulnerability in Turkey using the standardized precipitation index. *Natural Hazards*, 35(2), 243–  
575 264. <https://doi.org/10.1007/s11069-004-5704-7>
- 576 Thom HCS (1966) Some methods of climatological analysis. *WMO Technics/Note*, 81, 55.
- 577 Thomas T, Jaiswal RK, Nayak PC, Ghosh NC (2014) Comprehensive evaluation of the changing drought  
578 characteristics in Bundelkhand region of Central India. *Meteorology and Atmospheric Physics*, 127(2),  
579 163–182. <https://doi.org/10.1007/s00703-014-0361-1>

- 580 Touchan R, Funkhouser G, Hughes MK, Erkan N (2005) Standardized precipitation index reconstructed  
581 from Turkish tree-ring widths. *Climatic Change*, 72(3), 339–353. [https://doi.org/10.1007/s10584-005-](https://doi.org/10.1007/s10584-005-5358-9)  
582 5358-9
- 583 Vicente-Serrano SM, Beguería S, López-Moreno JI (2010) A multiscalar drought index sensitive to global  
584 warming: The standardized precipitation evapotranspiration index. *Journal of Climate*, 23(7), 1696–1718.  
585 <https://doi.org/10.1175/2009JCLI2909.1>
- 586 Yan Z, Zhang Y, Zhou Z, Han N (2017) The spatio-temporal variability of droughts using the standardized  
587 precipitation index in Yunnan, China. *Natural Hazards*, 88(2), 1023–1042. [https://doi.org/10.1007/s11069-](https://doi.org/10.1007/s11069-017-2904-5)  
588 017-2904-5
- 589 Yevjevich V (1967) An objective approach to definitions and investigations of continental hydrologic  
590 droughts. *Journal of Hydrology*, 7(3), 353. [https://doi.org/10.1016/0022-1694\(69\)90110-3](https://doi.org/10.1016/0022-1694(69)90110-3)
- 591 Yu X, He X, Zheng H, Guo R, Ren Z, Zhang D, Lin J (2014) Spatial and temporal analysis of drought risk  
592 during the crop-growing season over northeast China. *Natural Hazards*, 71(1), 275–289.  
593 <https://doi.org/10.1007/s11069-013-0909-2>
- 594 Yuan Z, Yan DH, Yang ZY, Yin J, Yuan Y (2015) Temporal and spatial variability of drought in Huang-  
595 Huai-Hai River Basin, China. *Theoretical and Applied Climatology*, 122(3–4), 755–769.  
596 <https://doi.org/10.1007/s00704-014-1332-7>
- 597 Zuo D, Cai S, Xu Z, Li F, Sun W, Yang X, Kan, G, Liu P (2018) Spatiotemporal patterns of drought at  
598 various time scales in Shandong Province of Eastern China. *Theoretical and Applied Climatology*, 131(1–  
599 2), 271–284. <https://doi.org/10.1007/s00704-016-1969-5>