

Quantifying nitrogen fixation by heterotrophic bacteria in sinking marine particles

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Article

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1 **Quantifying nitrogen fixation by heterotrophic bacteria in sinking marine particles**

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12 Running title: Nitrogen fixation in sinking marine Particles

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20 **Abstract**

21 Nitrogen (N₂) fixation by heterotrophic bacteria associated with sinking particles contributes to
22 marine N cycling, but a mechanistic understanding of its regulation and significance are not
23 available. Here we develop a mathematical model for unicellular heterotrophic bacteria growing on
24 sinking marine particles and can fix N₂ under suitable environmental conditions. We find that the
25 interactive effects of polysaccharide and polypeptide concentrations, sinking speed of particles, and
26 surrounding O₂ and NO₃⁻ concentrations determine the N₂ fixation rate inside particles. N₂ fixation
27 inside sinking particles is mainly fueled by SO₄²⁻ respiration rather than NO₃⁻ respiration. Our
28 model suggests that anaerobic processes, including heterotrophic N₂ fixation, can take place in
29 anoxic microenvironments inside sinking particles even in fully oxygenated marine waters. The
30 modelled rates are similar to bulk rates measured in the aphotic ocean, and our study consequently
31 suggests that particle-associated heterotrophic N₂ fixation contributes significantly to oceanic N₂
32 fixation.

33 **Keywords**

34 N₂ fixation; Sinking marine particles; Heterotrophic bacteria; Anaerobic respiration; Anoxic
35 interior; Trait-based optimization model

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41 **Introduction**

42 Nitrogen (N) is an essential element for all living organisms but its availability often limits the
43 growth and productivity of terrestrial and aquatic ecosystems. Although molecular dinitrogen gas
44 (N_2) is highly abundant in the marine water column, only specific prokaryotes that can fix N_2
45 (diazotrophs) using the nitrogenase enzyme complex¹ assimilate this form of nitrogen.
46 Nevertheless, nitrogen fixation maintains the inventory of biologically available nitrogen in the
47 open oceans, which fuels primary production^{1,2}, and thereby affects the biogeochemical cycling of
48 both nitrogen and carbon³.

49 N_2 fixation was thought to exclusively be carried out by cyanobacteria in the oligotrophic and
50 sunlit upper layers of the tropical and subtropical oceans (reviewed in Zehr⁴). However,
51 accumulating evidence shows that N_2 fixation is surprisingly widespread, for example in the deep
52 sea⁵, nutrient-rich coastal waters⁶, and cold Arctic waters⁷. Moreover, analyses of genes (*nifH*)
53 encoding the enzyme complex used for N_2 fixation document that non-cyanobacterial diazotrophs
54 are almost ubiquitous across the world's oceans, often dominate *nifH* gene libraries over
55 cyanobacteria, and occasionally express nitrogenase⁸⁻¹⁰. Hence, the emerging picture shows N_2
56 fixation as a global marine process partially carried out by non-cyanobacterial diazotrophs, but their
57 ecology and contribution to total N_2 fixation remain enigmatic¹⁰.

58 Nitrogenase is irreversibly inactivated by O_2 ¹¹. Cyanobacteria adopt several strategies to
59 protect nitrogenase from inactivation by O_2 ¹². Heterotrophic diazotrophs may under rich culture
60 conditions surround cells with extracellular polymers¹³ to lower the permeability to extracellular O_2
61 to protect the nitrogenase, but since this is highly energy-demanding¹⁴ it is an unlikely strategy in
62 the relatively nutrient-poor marine water column. Recent work, inspired by the pioneering work of
63 Paerl and co-workers^{15,16}, has suggested that heterotrophic N_2 fixation takes place in low-oxygen or

64 anaerobic microzones associated with marine particles (reviewed in Riemann et al.¹⁷ and in Bombar
65 et al.¹⁰). Indeed, anaerobic microzones are occasionally associated with marine particles^{18,19}. Recent
66 studies show N₂ fixation is stimulated by the presence of particles^{20,21} and that heterotrophic
67 diazotrophs are associated with plankton specimens^{22,23} and marine aggregates^{24,25}. Hence, beyond
68 doubt, marine particles provide, at least ephemeral, conditions suitable for N₂ fixation by
69 heterotrophic bacteria.

70 Cellular O₂ removal by diazotrophs is considered highly energy-demanding, even more
71 energetically expensive than N₂ fixation *per se*¹⁴. Hence, considerable amounts of labile carbon
72 (e.g. carbohydrate and amino acids) are required to sustain particle-associated microbial respiration
73 beyond the specific energy requirements for diazotrophy. Interestingly, preferential microbial
74 utilization of N-rich organics on particles²⁶ and release of NH₄⁺²⁷ may increase particle C:N ratios
75 over time²⁶, gradually making N acquisition by N₂ fixation increasingly advantageous.

76 While synthesizing ATP during respiration, O₂ is used as the most common and favorable
77 form of electron acceptor by prokaryotes. In the absence of O₂, other electron acceptors (e.g. NO₃⁻
78 and SO₄²⁻) may be used in a stepwise manner according to their free energy yields²⁸, with a rather
79 small drop-off in the theoretical energy yield for NO₃⁻ respiration, followed by SO₄²⁻ respiration
80 with almost tenfold less energy yield per electron donor²⁹. However, SO₄²⁻ respiration is likely the
81 primary form of anaerobic respiration supporting N₂ fixation since SO₄²⁻ reducing diazotrophs
82 have been widely found^{30,31}, also on marine particles^{32,33}.

83 Another important factor regulating N₂ fixation in sinking marine particles is the particle size,
84 which varies from micrometers to several millimeters³⁴. The particle size spectrum follows a power
85 law relationship showing a decrease in particle abundance with increasing size³⁵. Because of
86 smaller surface to volume ratios, large particles are more likely to develop an anoxic interior

87 suitable for N₂ fixation. Moreover, particles face changing O₂ and NO₃⁻ concentrations while
88 descending in the water column. The rate of change depends on particle sinking speed, but no
89 universal size-sinking speed relationship exists³⁶. Although all these external factors can have huge
90 influences, the extent by which they affect heterotrophic N₂ fixation inside sinking particles is
91 currently unclear.

92 To quantitatively analyze the conditions when heterotrophic N₂ fixation occurs on sinking
93 particles, we present a trait-based model of heterotrophic bacteria associated with sinking particles.
94 This effort aims to encapsulate an understanding of the dynamics between (micro)environmental
95 conditions and the requirements and constraints of heterotrophic N₂ fixation. Specifically, the model
96 captures basic cellular processes determining growth and N₂ fixation in an individual cell, and then
97 scales up to the population level to address particle dynamics and the contribution to total N₂
98 fixation in the water column. We also examine how the size of particles, initial concentrations of
99 polysaccharide and polypeptide, and environmental O₂ concentration influence heterotrophic N₂
100 fixation inside sinking particles, and the succession of aerobic and anaerobic respiration as support
101 for N₂ fixation. In doing so we identify potentially testable hypothetical consequences: (H1) Excess
102 acquired N released by cells and hydrolysis products diffuse away from the particle and contribute
103 to an organic solute trail in the water column as the particle sinks. (H2) N₂ fixation by heterotrophic
104 diazotrophs depends on the generation of particle-associated low-oxygen microenvironments. (H3)
105 During the “life span” of a sinking marine particle there is an ephemeral window of opportunity
106 where environmental conditions are conducive for heterotrophic N₂ fixation. (H4) SO₄²⁻ reduction
107 is more important for N₂ fixation within sinking particles than NO₃⁻ reduction. (H5) The particle
108 sinking speed and concentrations of O₂ and NO₃⁻ in the water column affects N₂ fixation rates.
109 Although the model is developed to investigate N₂ fixation, it also provides critical insights on
110 biochemistry and microbial respiratory processes inside sinking particles.

111 **Results and discussion**

112 *Overview of the model*

113 The overall model consists of a ‘cell model’ and a ‘particle model’. The cell model describes basic
114 cellular processes, like uptake of resources, respiration, growth, and N₂ fixation rate. The cell model
115 is embedded in a dynamic model, called ‘the particle model’, that deals with interactions of cells
116 with the available abiotic factors (polysaccharide, polypeptide, O₂, NO₃⁻, SO₄²⁻) over time.

117 *The cell model:* The model describes a population of facultative N₂ fixing heterotrophic bacteria
118 growing inside a particle sinking through a water column. A schematic representation of the
119 processes inside a single cell is presented in Figure 1 and the full description of mathematical forms
120 and equations are provided in the Methods section. The cell uses ectoenzymes to degrade polymers
121 (polysaccharides and polypeptides) to oligomers or monomers (glucose and amino acids) that it can
122 efficiently take up to fulfill its C and N requirements. The uptake of glucose and amino acids
123 follows Michaelis-Menten kinetics. The model accounts for acquired C and N to ensure that the cell
124 satisfies its needs for both. While glucose uptake provides only C, amino acids provide both C and
125 N (equations (1) and (2)).

126 C obtained from glucose and amino acids is respired to carry out resource uptake, cellular
127 maintenance (Fig. 1), and standard metabolism (equation (6)). O₂, NO₃⁻, and SO₄²⁻ are used as
128 electron acceptors in a stepwise manner in order of their free energy yield to perform respiration³⁷.
129 In the absence of sufficient O₂, the cell uses NO₃⁻ to continue respiration, although all N necessary
130 for growth comes from organic sources and N₂ fixation (whenever possible). The further need of
131 electron acceptor is fulfilled by SO₄²⁻.

132 The cell can carry out N_2 fixation to supplement its N requirement. It regulates the rate of N_2
133 fixation to optimize its growth rate. As nitrogenase is irreversibly inhibited by O_2 ¹¹, the cell needs
134 low O_2 conditions inside particles or increased respiration to make the cell O_2 free and thereby
135 enable N_2 fixation.

136 The synthesis of biomass using available C and N from resource uptake and electron
137 acceptors follows Liebig's law of the minimum and is constrained by the cellular C:N ratio ($\rho_{CN,B}$)
138 (equation (21)). Any excess assimilated C or N is excreted from the cell. The cell division rate μ is
139 found from the mass-specific synthesis rate.

140 *The particle model:* We consider a sinking particle consisting of polysaccharides and polypeptides
141 and colonized by facultative nitrogen-fixing bacteria (Supplementary Fig. S1). Only fractions of
142 these polymers are considered labile, i.e. accessible by bacteria. Bacterial enzymatic hydrolysis
143 converts labile polysaccharides and polypeptides into monosaccharides (glucose) and amino acids
144 that are efficiently taken up by bacteria. Excess glucose and amino acids diffuse out of the particle
145 to the surrounding environment, while O_2 and NO_3^- diffuse into the particle from the surrounding
146 water. Due to high concentrations of SO_4^{2-} and N_2 in ocean waters^{38,39}, these elements are assumed
147 not to be diffusion-limited inside particles, their uptakes are, therefore, limited by the cellular
148 maximum uptake capacities. Depending on the concentrations of glucose, amino acids, O_2 , NO_3^- ,
149 and SO_4^{2-} inside the particle, bacteria carry out N_2 fixation (equation (23)). Fe, an essential
150 component in the nitrogenase complex, is considered nonlimiting as sinking particles contain high
151 levels of Fe^{40} . The interactions between particle, cells, and the surrounding environment are
152 explained in Fig. S1, equations are provided in Table 1, and a full description of the particle model
153 is provided in the Methods section.

154

155 *Biochemical dynamics inside a particle under static environmental conditions*

156 The dynamics inside a particle of radius 0.125 cm with initial polysaccharide and polypeptide
157 concentrations of $2.6 \times 10^8 \mu\text{g G L}^{-1}$ and $1.6 \times 10^8 \mu\text{g A L}^{-1}$, with relative lability of 0.238 and 0.5,
158 respectively, are depicted in Fig. 2. We simulate a population of bacterial cells of radius $0.29 \mu\text{m}$
159 ($50 \text{ fg C cell}^{-1}$) growing inside a particle where the surrounding glucose, amino acids, O_2 , NO_3^- ,
160 and SO_4^{2-} concentrations are kept fixed at $50 \mu\text{g G L}^{-1}$, $5 \mu\text{g A L}^{-1}$, $50 \mu\text{mol O}_2 \text{ L}^{-1}$, $15 \mu\text{mol NO}_3$
161 L^{-1} , and $29 \times 10^3 \mu\text{mol SO}_4 \text{ L}^{-1}$. Here, concentrations inside the particle are given as per liter of
162 particle and outside as per liter of water. A full description of the included parameters and their
163 values is available in Supplementary Material S1 and Table S1. The bacteria hydrolyze labile
164 polysaccharides and polypeptides into glucose and amino acids using ectoenzymes (Fig. 2a, b). As a
165 result, glucose and amino acid concentrations increase (Fig. 2c, d), which causes a high growth rate
166 of cells ($\sim 3.6 \text{ d}^{-1}$; Fig. 2j), an increase in bacterial abundance (Fig. 2g), and a decrease in labile
167 polysaccharide and polypeptide concentrations. The occurrence of such high growth rate⁴¹ and
168 bacterial abundance^{25,42,43} is not seldom inside natural sinking particles. The increased community
169 respiration (Fig. 2h) decreases O_2 concentration and eventually leads to anoxia in the particle
170 interior (Fig. 2e). This is consistent with ephemeral anoxia inside marine aggregates¹⁹ and the
171 anoxia observed inside suspended cyanobacterial colonies of comparable size¹⁸. The gradual
172 formation of low-oxygen or anoxic conditions and depletion of organic N (amino acids) facilitates
173 N_2 fixation (Fig. 2i), supported by aerobic respiration followed by NO_3^- respiration (Fig. 2e, f). The
174 lesser energetic yield of NO_3^- respiration leads to a reduced growth rate ($\sim 0.6 \text{ d}^{-1}$). Furthermore,
175 when NO_3^- becomes exhausted (Fig. 2f), cells respire SO_4^{2-} (not shown in the figure) and the
176 growth rate becomes very low ($\sim 0.2 \text{ d}^{-1}$). Because of the energetic constraints, N_2 fixation during
177 this phase becomes low and eventually ceases due to increased O_2 levels as the exhaustion of labile

178 carbon in the particle decreases cell concentration and O_2 influx exceeds O_2 consumption (aerobic
179 respiration).

180

181 *Cellular mechanisms of N_2 fixation*

182 To explore the cellular mechanism of N_2 fixation, we examine concentrations and rates over time at
183 a radial distance of 0.027 cm from the particle center (Fig. 3). We identify four phases based on
184 limitations of either C or N or electron acceptor where the growth rate is determined by the
185 minimum availability of these three substances (equations (21) and (22); Fig. 3a). Available C for
186 growth is the C remaining from total C uptake after paying the respiratory costs, whereas the N
187 available for growth comes from uptake and N_2 fixation.

188 I (C limited phase): Cells are limited by C, which is seen by the light blue line coinciding with the
189 magenta line in Fig. 3a. Growth rates are high, up to 3.6 day^{-1} . Excess N and hydrolysis products
190 not taken up by cells²⁶ will diffuse away from the particle and contribute to an organic solute trail in
191 the water column as the particle sinks⁴⁴, supporting our hypothesis H1.

192 II (High respiration phase): The large bacterial population causes high community respiration,
193 matching or exceeding the diffusive influx of O_2 , and anoxia forms in the particle interior (Fig. 3e).
194 Cells start respiring NO_3^- and even SO_4^{2-} reduction happens at the end of this phase (not shown in
195 the figure) when NO_3^- is depleted. Now growth is limited by the availability of electron acceptor
196 (the yellow line coincides with the magenta in Fig. 3a). An organic solute trail rich in both C and N
197 is predicted during this phase. The amino acid concentration decreases rapidly during the final part
198 of this phase (Fig. 3c).

199 III (N₂ fixing phase): Because of our initial choice of polysaccharide and polypeptide
200 concentrations, amino acids are exhausted. In real life, bacterial preferential degradation of N-rich
201 organics results in similar early exhaustion of amino acids²⁶. Glucose remains available as C source
202 (Fig. 3c), so cells start fixing N₂ to maintain growth (Fig. 3d). However, since the available O₂ in
203 the particle is insufficient to support respiration, NO₃⁻ and SO₄²⁻ also act as electron acceptors
204 during this phase. Cells become co-limited by N from N₂ fixation and electron acceptor (green and
205 yellow lines coincide; Fig. 3a) and we predict that the expected solute trail consists only of C during
206 this phase. Because of the lower free energy yield, N₂ fixation decreases when using SO₄²⁻ as an
207 additional electron acceptor. Towards the end of this phase, the respiratory cost of glucose uptake
208 decreases with the decrease in glucose concentration to such an extent that there is excess O₂ after
209 performing respiration. At that point, cells increase respiration to burn excess O₂ to perform N₂
210 fixation for a very short interval of time, resulting in a peak in N₂ fixation rate (Fig. 3d) and the
211 respiratory cost for O₂ removal (Fig. 3b).

212 IV (Fading phase): Cells have insufficient C to deal with excess O₂ and consequently stop N₂
213 fixation. N becomes the limiting factor for cell growth and growth ceases. Later, the growth rate
214 becomes negative as there is no glucose left needed for basal respiration. Throughout this phase, the
215 bacterial concentration decreases (Fig. 3c).

216 Comparing different respiratory costs related to N₂ fixation in terms of direct respiration,
217 enzyme production, and O₂ removal, it becomes evident that when cells use respiratory protection
218 to keep nitrogenase viable and enable N₂ fixation, the related cost becomes much higher than the
219 direct cost for N₂ fixation (Fig. 3b). A similar high cost of O₂ management during N₂ fixation was
220 previously shown to exceed the costs of N₂ fixation *per se* for the heterotrophic soil bacterium
221 *Azotobacter vinelandii*¹⁴. We therefore conclude that active O₂ management by particle-associated

222 heterotrophic diazotrophs is not prevalent, but that they rather depend on the generation of low-
223 oxygen microenvironments by community respiration, supporting our initial hypothesis H2.

224

225 *Effects of particle size, O₂, and initial polysaccharide and polypeptide concentrations*

226 Marine particles are highly variable in size³⁴ and chemical composition. For example, the C:N ratio
227 (and implied polysaccharide:polypeptide availability) varies considerably, depending on
228 environmental conditions⁴⁵. Moreover, while descending in the water column, particles face a range
229 of surrounding O₂ concentrations. Therefore, we examined the implications for N₂ fixation in
230 particles of different sizes under different polysaccharide, polypeptide, and O₂ concentrations (Fig.
231 4). We considered open ocean particles with radius 5 μm to 0.25 cm⁴⁶ and estimated the total
232 amount of fixed N₂ per particle by allowing bacteria to grow inside particles for 20 days. As
233 expected, large particles provide a suitable environment for N₂ fixation. The minimum size of
234 particles where N₂ fixation is possible increases with polysaccharide concentration (Fig. 4a),
235 decreases with environmental O₂ concentration (Fig. 4c), and attains a maximum at intermediate
236 polypeptide concentration (Fig. 4b). N₂ fixation does not occur in particles with radius below ~0.03
237 cm. The increase in N₂ fixation with polysaccharide concentrations is consistent with observations
238 of stimulated N₂ fixation in seawater upon the addition of C substrate (e.g. Rahav et al.²¹).
239 Interestingly, maximum N₂ fixation occurs at intermediate polypeptide concentrations (~5 × 10⁷ μg
240 L⁻¹) in large particles (Fig. 4b). Our interpretation is that low polypeptide concentrations do not
241 allow cells to grow to high concentrations and create an anoxic interior, whereas, at high
242 concentrations, cells cover their N demand by amino acid assimilation and, therefore, refrain from
243 N₂ fixation.

244 Maximum N₂ fixation in large particles occurs under very low (~0.3 μmol L⁻¹) and
245 intermediate (~80 μmol L⁻¹) O₂ levels (Fig. 4c). Under very low O₂ concentrations, cellular
246 respiration is expected to occur using NO₃⁻ and SO₄²⁻ as electron acceptors causing low cellular
247 growth rate and a slowly increasing cell concentration. However, N₂ fixation occurs for a long time
248 period and makes the total N₂ fixation per particle relatively high (Supplementary Fig. S2d). At
249 intermediate O₂ levels, cells are not limited by O₂ and reach a high growth rate during the initial
250 phase. Therefore a high cell concentration is rapidly obtained (Supplementary Fig S2c) causing
251 reduced O₂ levels suitable for N₂ fixation for a relatively shorter time interval (Supplementary Fig
252 S2d). The combination of high cellular N₂ fixation rate and high cell concentration results in high
253 total N₂ fixation per particle. At a high O₂ level (~200 μmol L⁻¹), this low O₂ period is very short,
254 possibly because of a large diffusion loss of glucose associated with extensive polysaccharide
255 hydrolysis caused by the high cell concentration (Supplementary Fig S2c). This lowers the total
256 amount of N₂ fixed per particle. However, the concentrations of O₂, where these two maxima occur,
257 depend on the initial polysaccharide and polypeptide concentrations. With a decrease in these
258 concentrations, the intermediate O₂ concentration, where the maximum occurs, decreases
259 (Supplementary Fig. S3) and finally merges with the other maximum at low O₂ (not shown). Indeed,
260 empirical results confirm the existence of such optimal O₂ concentration for N₂ fixation and a level
261 of 6 μmol L⁻¹ has been observed for heterotrophic diazotrophs⁴⁷.

262 By simultaneously varying initial labile polysaccharide and polypeptide concentrations in
263 small and large particles at different surrounding O₂ concentrations, it appears that N₂ fixation is
264 restricted to large particles with high initial polysaccharide and polypeptide concentrations when O₂
265 concentration is high (Fig. 5a). However, under low O₂ concentrations, N₂ fixation occurs even at
266 lower concentrations of polysaccharides and polypeptides (Fig. 5b). N₂ fixation can also occur in

267 relatively smaller particles, however, only when the initial polysaccharide concentration is high and
268 the surrounding O₂ concentration is low (Fig. 5c).

269 The amount of POC present in particles can be considered a proxy for polysaccharide and
270 polypeptide concentrations. The presence of high POC in freshly formed particles from dense
271 phytoplankton blooms⁴⁸ and fecal pellets⁴⁹ would increase the likelihood of particle-associated N₂
272 fixation. Likewise, the seasonally high POC content of particles during late spring in high latitude
273 areas⁵⁰ could increase the likelihood of particle-associated N₂ fixation at this time. On the other
274 hand, the size of particles, which also has a profound impact on N₂ fixation, is closely associated
275 with the species responsible for particle formation. For instance, since particles are larger during
276 diatom blooms compared to cyanobacterial blooms⁵¹, an increased likelihood of N₂ fixation during
277 diatom blooms would be expected. The latitudinal variation in particle size spectrum with a
278 dominance of larger particles at higher latitudes compared to smaller particles in the oligotrophic
279 subtropical gyres⁵² indicates a greater opportunity for particle-associated N₂ fixation at high
280 latitudes. Hence, the potential for particle-associated N₂ fixation is highly dependent on local
281 dynamics in the size and composition of particles, together with the local O₂ conditions.

282

283 *N₂ fixation in sinking particles*

284 To explore the dynamics of N₂ fixation in sinking particles of different sizes, we use vertical
285 profiles of O₂ and NO₃⁻ in the upper 500 m of the Mauritanian upwelling zone in the North
286 Atlantic Ocean (NAO; Fig. 6)⁵³. O₂ concentration drops to hypoxic levels (~62.5–157 μmol L⁻¹) in
287 the water column between 100–600 m (Fig. 6a). N₂ fixation coincides mainly with the presence of
288 an anoxic particle interior. The amount of fixed N₂ increases with particle size and because of
289 higher sinking rates and more C to fuel respiration the existence of both anoxic interior (regions

290 within magenta lines) and N₂ fixation (within white lines) in large particles occur in deeper waters
291 and persist for longer time (Fig. 6b). We presume that N₂ fixation stops in deep water when labile
292 material in the particle is exhausted. This shows that the window of opportunity where
293 environmental conditions are conducive for heterotrophic N₂ fixation is ephemeral, supporting our
294 hypothesis H3.

295 In the anoxic particle interior, NO₃⁻ and SO₄²⁻ function as electron acceptors. The model
296 suggests that the fraction of particle volume where denitrification occurs (maximum 6%) is much
297 smaller than the fraction of volume of occurrence of SO₄²⁻ reduction (maximum 90%) (Fig. 6c, d).
298 The model predicts that despite being energetically profitable, NO₃⁻ does not play a big role in N₂
299 fixation and SO₄²⁻ reduction appears as the key anaerobic process within sinking particles. This is
300 due to very high cell concentration near the surface of the particle (Fig. 2g) that creates a high
301 respiratory demand for electron acceptors, exhausts NO₃⁻ close to the particle surface, and prevents
302 NO₃⁻ from reaching the particle interior (Fig. 2f). As a result, N₂ fixation in most of the particle
303 interior is supported by SO₄²⁻ respiration, which confirms the importance of SO₄²⁻ reduction in
304 particle-associated N₂ fixation compared to NO₃⁻ reduction, supporting hypothesis H4.

305 Diazotrophy among SO₄²⁻ reducing bacteria is well established in various marine
306 environments^{54,55}.

307

308 *Influence of sinking speed on particle-based N₂ fixation.*

309 The speed at which particles sink is a critical parameter since it determines the duration of exposure
310 to environmental conditions (e.g. O₂) that influence N₂ fixation inside particles. Sinking speed is
311 affected by a multitude of factors related to particle composition, size, and density⁵⁶⁻⁵⁸, and no

312 universal size-sinking velocity relationship exists³⁶. We, therefore, examine scenarios with three
313 types of particles: (1) natural marine snow measured *in situ* off California⁵⁹, (2) laboratory-made
314 diatom aggregates and (3) coccolithophore aggregates measured *in vitro*⁴⁸. Sinking speed is lowest
315 for natural marine snow followed by diatom aggregates and coccolithophore aggregates (Fig. 7a).
316 For the sake of simplicity, these particles are assumed to vary only in their sinking speeds and not in
317 their initial concentrations and lability of polysaccharides and polypeptides. We again use the
318 vertical profile of O₂ and NO₃⁻ at the NAO but extended to 1,500 m depth with a hypoxic region
319 between 100-600 m depth (Fig. 7b).

320 Three different aspects are evident from the analysis. Firstly, when particles sink at a speed
321 similar to natural marine snow (~15-75 m d⁻¹), anoxic microenvironments are created (regions
322 within magenta lines) and N₂ fixation happens (indicated by color) in particles within the hypoxic
323 zone (Fig. 7e). However, with higher sinking speeds similar to diatom (~2-180 m d⁻¹) and
324 coccolithophore (~20-375 m d⁻¹) aggregates, the existence of anoxia and N₂ fixation inside particles
325 can extend beyond the hypoxic strata of the water column (Fig. 7d,e); both anoxic interior and N₂
326 fixation are predicted even at 1,800 m depth for large coccolithophore aggregates (not in the figure).
327 Secondly, the depth window where N₂ fixation occurs increases with particle sinking speed (7c-e).
328 Finally, our study predicts that the highest N₂ fixation rate (1.46 μg N (cm³ particle)⁻¹ d⁻¹) is attained
329 in large particles at intermediate sinking velocities, similar to that of diatom aggregates (Fig. 7d).
330 Since, NO₃⁻ respiration yields more energy than SO₄²⁻ reduction, we speculate that the presence of
331 relatively higher concentrations of NO₃⁻ during the time of N₂ fixation helps to boost N₂ fixation
332 rate in diatom aggregates. High sinking speed, similar to that of coccolithophore aggregates,
333 transports particles quickly to deep water with high NO₃⁻ and O₂ concentrations (Fig. 7e). We
334 expect that high NO₃⁻ concentration favors relatively high N₂ fixation, but high O₂ concentration
335 decreases the N₂ fixation rate in coccolithophore aggregates. Therefore, the interplay between

336 particle sinking speed and vertical water column profiles of O_2 and NO_3^- concentrations determines
337 the N_2 fixation rate, supporting our hypothesis H5.

338

339 *N₂ fixation in contrasting oceanic environments*

340 N_2 fixation is dependent on the O_2 and NO_3^- concentrations in the water column, but those
341 are highly variable around the global ocean. We investigated N_2 fixation rates in three contrasting
342 water columns: an O_2 minimum zone in the Eastern Tropical South Pacific (ETSP)⁵³, the
343 Mauritanian upwelling zone in the NAO⁵³, and an open ocean site (OO; 30.5°N, 52.5°W)^{60,61}. The
344 ETSP has O_2 minimum zones with $< 5 \mu\text{mol } O_2 \text{ L}^{-1}$ at ~150–600 m depth (Fig. 8a), the NAO has
345 reduced O_2 levels (~62.5–157 $\mu\text{mol } O_2 \text{ L}^{-1}$) at ~100–600 m (Fig. 8d), whereas the open ocean site has
346 high O_2 concentration throughout the water column ($> 157 \mu\text{mol } O_2 \text{ L}^{-1}$; Fig. 8g). NO_3^- concentrations
347 increase gradually up to 800 m depth, with decreasing levels from ETSP to NAO and to the open
348 ocean site. We consider particles with sinking speed similar to natural marine snow and examine N_2
349 fixation rates per unit volume of particle (fig 8c,f,i). To compare with existing measurements, we
350 further calculate N_2 fixation rates per unit volume of water (Fig. 8b,e,h) and depth-integrated N_2
351 fixation rates by multiplying the number of particles with N_2 fixation per particle (equation (34))³⁵.

352 Maximum N_2 fixation rates per volume of particle and per volume of water lie within ranges
353 $0.31\text{--}1.1 \mu\text{g N (cm}^3 \text{ particle)}^{-1} \text{ d}^{-1}$ (Fig. 8c,f,i) and $0.14\text{--}0.7 \mu\text{mol N m}^{-3} \text{ d}^{-1}$ (Fig. 8b,e,h). The highest
354 rate of N_2 fixation is in the NAO followed by the ETSP. Interestingly, N_2 fixation is observed even
355 at the high O_2 concentrations of the open ocean, although, the N_2 fixation rates and the depth
356 window of N_2 fixation are smaller than for the other two scenarios. Since the abundance and size
357 spectrum of particles vary with latitude and seasonally at high latitudes⁶², we test the sensitivity of
358 N_2 fixation rates by varying the parameter determining the abundance (n_0) and the proportion of

359 large and small particles (ξ). We find that the N_2 fixation rate varies between $0.3\text{-}1.7 \mu\text{mol N m}^{-3} \text{d}^{-1}$
360 (Supplementary fig. S4). Our modelled N_2 fixation rates are comparable with bulk N_2 fixation rates
361 measured in the aphotic ocean ($0\text{-}0.89 \mu\text{mol N m}^{-3} \text{d}^{-1}$)^{9,63} where active autotrophic cyanobacterial
362 diazotrophs are not expected. Our calculated depth-integrated N_2 fixation rates lie within the range
363 $7.1\text{-}65.1 \mu\text{mol N m}^{-2} \text{d}^{-1}$. Empirical evidence from regions, where N_2 fixation is dominated by
364 heterotrophic bacteria, show similar levels of depth-integrated N_2 fixation rates; e.g. $6.27\text{-}16.6$
365 $\mu\text{mol N m}^{-2} \text{d}^{-1}$ in the equatorial and southern Indian Ocean⁶⁴ and $12.4\text{-}190.9 \mu\text{mol N m}^{-2} \text{d}^{-1}$ in the
366 South Pacific Gyre⁶⁵. In comparison, depth-integrated N_2 fixation rates by cyanobacteria in most
367 regions of the global upper ocean are in the order of $1\text{-}100 \mu\text{mol N m}^{-2} \text{d}^{-1}$ ⁶⁶. Hence, taken together,
368 our modelled rates for heterotrophic bacteria on particles are consistent with empirical bulk rates
369 from the deep sea and comparable to areal rates measured for cyanobacteria. This supports
370 empirical studies suggesting that aphotic fixation can account for a significant or even predominant
371 fraction of water column N_2 fixation^{63,67}, and substantiates the idea that aphotic N_2 fixation may be
372 important to global nitrogen budget considerations⁵.

373

374 **Conclusions and broader implications**

375 Our model suggests that particle-associated heterotrophic N_2 fixation is viable and reasonable based
376 on the known properties and physics of marine particles and reveals a significant contribution to the
377 oceanic biological N_2 fixation. The likelihood and rate of N_2 fixation associated with any individual
378 particle will depend upon numerous factors including the initial polysaccharide and polypeptide
379 concentrations (and associated C:N ratio), the size and sinking speed of the particle, and the vertical
380 profiles of O_2 and NO_3^- through which the particle sinks. We show how low- O_2 or anoxic zones
381 generated inside sinking particles by microbial respiration provide conditions suitable for

382 heterotrophic N₂ fixation, however, only in particles larger than about 0.03 cm. Moreover, we show
383 that these anoxic microenvironments can promote anaerobic respiratory processes that even extend
384 into well-oxygenated deep waters. Interestingly, our simulations suggest that even in particles that
385 favor N₂ fixation at a point in their descent, the window of time (depth) where diazotrophy occurs is
386 likely to be short. However, despite the necessity of several coinciding environmental conditions for
387 N₂ fixation in particles, the criteria are met in natural particles. Because of the huge number and
388 heterogeneity among particles sinking in the ocean⁶⁸, it is highly likely that a large number of
389 individual particles at any given time might meet these criteria and in doing so, confer a fitness
390 advantage on a subset of bacteria that retain the ability to fix nitrogen. This may explain the
391 ubiquity and persistence of the genetic signature for heterotrophic N₂ fixation throughout the
392 oceans⁸. The combined knowledge of the probability density of particle sizes, compositions, and
393 sinking speeds is suggested to predict the average rates of N₂ fixation associated with particles.

394 Our model makes several interesting and potentially testable predictions: Firstly, the C:N
395 composition of particle “trails” will reflect the interior state and might be a way to probe the
396 response of different particle types or conditions. Secondly, the model predicts a “preference” for
397 SO₄²⁻ electron acceptor over nitrate in low oxygen particles. Thirdly, heterotrophic diazotrophs
398 mostly use local and ephemeral oxygen conditions and get windows of opportunity for their N₂
399 fixation. Fourthly, N₂ fixation can occur on large particles with high concentrations of
400 polysaccharides and polypeptides in fully oxygenated marine waters, but also on older less
401 substrate-rich particles if oxygen concentration in the surrounding water is low. These predictions
402 could be promoted as perspectives for future experiments.

403

404

405 **Methods**

406 *The cell model*

407 *Growth rate of a cell:*

408 The growth rate of a bacteria cell depends on the acquisition of C (from the particle) and N (from
409 the particle and through N₂ fixation), as well as on metabolic expenses in terms of C.

410 *Uptake of C and N*

411 Bacteria get C from glucose and both C and N from amino acids. The total amount of C available
412 for the cell from monomers is (units of C per time)

413
$$J_{\text{DOC}} = f_{G,C}J_G + f_{A,C}J_A, \quad (1)$$

414 and the amount of N available from monomer is (N per time)

415
$$J_{\text{DON}} = f_{A,N}J_A, \quad (2)$$

416 where J_G and J_A are uptake rates of glucose and amino acids, $f_{G,C}$ is the fraction of C in glucose, and
417 $f_{A,C}$ and $f_{A,N}$ are fractions of C and N in amino acids.

418 The rate of obtaining N through N₂ fixation is:

419
$$J_{N_2}(\psi) = \psi M_{N_2}, \quad (3)$$

420 where ψ ($0 < \psi < 1$) regulates N₂ fixation rate and fixation can happen at a maximum rate M_{N_2} .

421 N₂ fixation is only limited by the maximum N₂ fixation rate as dissolved dinitrogen (N₂) gas in
422 seawater is assumed to be unlimited³⁹.

423 The total uptake of C and N from different sources becomes

424
$$J_C = J_{\text{DOC}} \quad (4)$$

425
$$J_N(\Psi) = J_{\text{DON}} + J_{N_2}(\Psi). \quad (5)$$

426 *Costs*

427 Respiratory costs of cellular processes together with N₂ fixation and its associated O₂ removal cost
 428 depend on the cellular O₂ concentration. Two possible scenarios can be observed:

429 *Case 1: When O₂ concentration is sufficient to maintain aerobic respiration*

430 Respiratory costs for bacterial cellular maintenance can be divided into two parts: one dependent on
 431 limiting substrates and the other one is independent of substrate concentration⁶⁹. Here we consider
 432 only the basal respiratory cost $R_B x_B$, which is independent of the limiting substrates and is assumed
 433 as proportional to the mass of the cell x_B (μg C). In order to solubilize particles, particle-attached
 434 bacteria produce ectoenzymes that cleave bonds to make molecules small enough to be transported
 435 across the bacterial cell membrane. Cleavage is represented by a biomass-specific ectoenzyme
 436 production cost R_E ⁷⁰. The metabolic costs associated with the uptake of hydrolysis products and
 437 intracellular processing are assumed to be proportional to the uptake (J_i): $R_G J_G$ and $R_A J_A$ where the
 438 R_i 's are costs per unit of resource uptake. In a similar way, the metabolic cost of N₂ fixation is
 439 assumed as proportional to the N₂ fixation rate: $R_{N_2} \rho_{\text{CN,B}} J_{N_2}$, where $\rho_{\text{CN,B}}$ is the bacterial C:N ratio.
 440 If we define all the above costs as direct costs, then the total direct respiratory cost becomes

441
$$R_D(\Psi) = R_B x_B + R_E x_B + R_G J_G + R_A J_A + R_{N_2} \rho_{\text{CN,B}} J_{N_2}(\Psi). \quad (6)$$

442 Indirect costs related to N₂ fixation arises from the removal of O₂ from the cell and the
 443 production/replenishment of nitrogenase as the enzyme is damaged by O₂. The cell can remove O₂
 444 either by increasing respiration⁷¹ or by increasing the production of nitrogenase enzyme itself⁷².
 445 Here we consider only the process of O₂ removal by increasing respiration. To calculate this
 446 indirect cost, the concentration of O₂ present in the cell needs to be estimated.

447 Since the time scale of O₂ concentration inside a cell is short, we have assumed a pseudo
 448 steady state inside the cell; the O₂ diffusion rate inside a cell is always balanced by the respiration
 449 rate¹⁴, which can be expressed as

$$450 \quad \rho_{CO} F_{O_2} = R_D(\psi). \quad (7)$$

451 Here ρ_{CO} is the conversion factor of respiratory O₂ to C equivalents and F_{O_2} is the actual O₂
 452 diffusion rate into a cell from the particle and can be calculated as

$$453 \quad F_{O_2} = 4\pi r_B K_{O_2} (X_{O_2} - X_{O_2,c}), \quad (8)$$

454 where r_B is the cell radius, X_{O_2} is the local O₂ concentration inside the particle, $X_{O_2,c}$ is the cellular
 455 O₂ concentration, and K_{O_2} is the effective diffusion coefficient of O₂ over cell membrane layers.
 456 The effective diffusion coefficient can be calculated according to Inomura et al.¹⁴ in terms of
 457 diffusion coefficient inside particles (\bar{D}_{O_2}), the diffusivity of cell membrane layers relative to water
 458 (ε_m), the radius of cellular cytoplasm (r_C), and the thickness of cell membrane layers (L_m) as

$$459 \quad K_{O_2} = \bar{D}_{O_2} \frac{\varepsilon_m(r_C + L_m)}{\varepsilon_m r_C + L_m}. \quad (9)$$

460 The apparent diffusivity inside particles (\bar{D}_{O_2}) is considered as a fraction f_{O_2} of the diffusion
 461 coefficient in seawater (D_{O_2})

$$462 \quad \bar{D}_{O_2} = f_{O_2} D_{O_2}. \quad (10)$$

463 Combining (7) and (8) gives the cellular O₂ concentration $X_{O_2,c}$ as

$$464 \quad X_{O_2,c} = \max \left[0, X_{O_2} - \frac{R_D(\psi)}{4\pi r_B K_{O_2} \rho_{CO}} \right]. \quad (11)$$

465 If there is excess O₂ present in the cell after respiration ($X_{O_2,c} > 0$), then the indirect cost of
 466 removing the excess O₂ to be able to perform N₂ fixation can be written as

467
$$R_{O_2}(\psi) = H(\psi)\rho_{CO}4\pi r_B K_{O_2} X_{O_2,c}, \quad (12)$$

468 where $H(\psi)$ is the Heaviside function:

469
$$H(\psi) = \begin{cases} 0, & \text{if } \psi = 0 \\ 1, & \text{if } \psi > 0 \end{cases}. \quad (13)$$

470 Therefore, the total aerobic respiratory cost becomes:

471
$$R_{\text{tot,A}}(\psi) = R_D(\psi) + R_{O_2}(\psi). \quad (14)$$

472 *Case 2: Anaerobic respiration*

473 When available O_2 is insufficient to maintain aerobic respiration ($R_{\text{tot}}(\psi) > \rho_{CO}F_{O_2,\text{max}}$), cells use

474 NO_3^- and SO_4^{2-} for respiration. The potential NO_3^- uptake, $J_{NO_3,\text{pot}}$, is

475
$$J_{NO_3,\text{pot}} = M_{NO_3} \frac{A_{NO_3} X_{NO_3}}{A_{NO_3} X_{NO_3} + M_{NO_3}}, \quad (15)$$

476 where M_{NO_3} and A_{NO_3} are maximum uptake rate and affinity for NO_3^- uptake, respectively.

477 However, the actual rate of NO_3^- uptake, J_{NO_3} , is determined by cellular respiration and can be

478 written as

479
$$J_{NO_3} = \min\left(J_{NO_3,\text{pot}}, \max\left(0, \frac{R_{\text{tot,A}}(\psi) - \rho_{CO}F_{O_2,\text{max}}}{\rho_{CNO_3}}\right)\right), \quad (16)$$

480 where ρ_{CNO_3} is the conversion factor of respiratory NO_3^- to C equivalents and the maximum O_2

481 diffusion rate into a cell $F_{O_2,\text{max}}$ can be obtained by making cellular O_2 concentration $X_{O_2,c}$ zero in

482 (8) as

483
$$F_{O_2,\text{max}} = 4\pi r_B K_{O_2} X_{O_2}, \quad (17)$$

484 Further, in the absence of sufficient NO_3^- , the cell uses SO_4^{2-} as an electron acceptor for

485 respiration. Since the average concentration of SO_4^{2-} in seawater is 29 mmol L^{-1} ³⁸, SO_4^{2-} is a non-

486 limiting nutrient for cell growth and the potential uptake rate of SO_4^{2-} is mainly governed by the
 487 maximum uptake rate as

$$488 \quad J_{\text{SO}_4, \text{pot}} = M_{\text{SO}_4}, \quad (18)$$

489 where M_{SO_4} is the maximum uptake rate for SO_4^{2-} uptake. The actual rate of SO_4^{2-} uptake, J_{SO_4} ,
 490 can be written as

$$491 \quad J_{\text{SO}_4} = \min \left(J_{\text{SO}_4, \text{pot}}, \max \left(0, \frac{R_{\text{tot}, \text{A}}(\psi) - \rho_{\text{CO}} F_{\text{O}_2, \text{max}} - \rho_{\text{CNO}_3} F_{\text{NO}_3, \text{pot}}}{\rho_{\text{CSO}_4}} \right) \right), \quad (19)$$

492 where ρ_{CSO_4} is the conversion factor of respiratory SO_4^{2-} to C equivalents.

493 According to formulations (16) and (19), NO_3^- and SO_4^{2-} uptake occurs only when the
 494 diffusive flux of O_2 , and both O_2 and NO_3^- are insufficient to maintain respiration. Moreover, the
 495 uptake rates of NO_3^- and SO_4^{2-} are regulated according to the cells' requirements.

496 Uptakes of NO_3^- and SO_4^{2-} incur extra metabolic costs $R_{\text{NO}_3} \rho_{\text{CNO}_3} J_{\text{NO}_3}$ and $R_{\text{SO}_4} \rho_{\text{CSO}_4} J_{\text{SO}_4}$,
 497 where R_{NO_3} and R_{SO_4} are costs per unit of NO_3^- and SO_4^{2-} uptake. The total respiratory cost can
 498 be written as

$$499 \quad R_{\text{tot}}(\psi) = R_{\text{tot}, \text{A}}(\psi) + R_{\text{NO}_3} \rho_{\text{CNO}_3} J_{\text{NO}_3} + R_{\text{SO}_4} \rho_{\text{CSO}_4} J_{\text{SO}_4}. \quad (20)$$

500 *Synthesis and growth rate*

501 The assimilated C and N are combined to synthesize new structure. The synthesis rate is
 502 constrained by the limiting resource (Liebig's law of the minimum) and by available electron
 503 acceptors such that the total flux of C available for growth J_{tot} ($\mu\text{g C d}^{-1}$) is:

$$504 \quad J_{\text{tot}}(\psi) = \min [J_{\text{C}} - R_{\text{tot}}(\psi), \rho_{\text{CN}, \text{B}} J_{\text{N}}(\psi), \rho_{\text{CO}} F_{\text{O}_2} + \rho_{\text{CNO}_3} J_{\text{NO}_3} + \rho_{\text{CSO}_4} J_{\text{SO}_4}]. \quad (21)$$

505 Here, the total available C for growth is $J_C - R_{\text{tot}}(\psi)$, the C required to synthesize biomass
 506 from N source is $\rho_{\text{CN},B}J_N$, and the C equivalent inflow rate of electron acceptors to the cell is
 507 $\rho_{\text{CO}}F_{\text{O}_2} + \rho_{\text{CNO}_3}J_{\text{NO}_3} + \rho_{\text{CSO}_4}J_{\text{SO}_4}$. We assume that excess C or N is released from the cell
 508 instantaneously.

509 Synthesis is not explicitly limited by a maximum synthesis capacity; synthesis is constrained
 510 by the C and N uptake in the functional responses (equations (28) and (29)). The division rate μ of
 511 the cell (d^{-1}) is the total flux of C available for growth divided by the C mass of the cell (x_B):

$$512 \quad \mu(\psi) = J_{\text{tot}}(\psi)/x_B. \quad (22)$$

513 The resulting division rate, μ , is a measure of the bacterial fitness and we assume that the cell
 514 regulates its N_2 fixation rate depending on the environmental conditions to gain additional N while
 515 maximizing its growth rate. The optimal value of the parameter regulating N_2 fixation ψ ($0 \leq$
 516 $\psi \leq 1$) then becomes:

$$517 \quad \psi^* = \underset{\psi}{\operatorname{argmax}}\{\mu(\psi)\}, \quad (23)$$

518 and the corresponding optimal division rate becomes

$$519 \quad \mu^* = \mu(\psi^*). \quad (24)$$

520 ***The particle model***

521 We consider a sinking particle of radius r_p (cm) and volume V_p (cm^3) (Supplementary Fig.
 522 S1). The particle contains facultative nitrogen-fixing bacterial population $B(r)$ (cells L^{-1}),
 523 polysaccharides $C_p(r)$ ($\mu\text{g G L}^{-1}$), and polypeptides $P_p(r)$ ($\mu\text{g A L}^{-1}$) at a radial distance r (cm)
 524 from the center of the particle, where G and A stand for glucose and amino acids. We assume that
 525 only fractions f_C and f_P of these polymers are labile ($C_L(r) = f_C C_p(r)$, $P_L(r) = f_P P_p(r)$), i.e.

526 accessible by bacteria. Bacterial enzymatic hydrolysis converts the labile polysaccharides and
527 polypeptides into monosaccharides (glucose) ($G \mu\text{g G L}^{-1}$) and amino acids ($A \mu\text{g A L}^{-1}$) that are
528 efficiently taken up by bacteria. Moreover, the particle contains O_2 , NO_3^- , and SO_4^{2-} with
529 concentrations $X_{\text{O}_2}(r)$ ($\mu\text{mol O}_2 \text{ L}^{-1}$), $X_{\text{NO}_3}(r)$ ($\mu\text{mol NO}_3 \text{ L}^{-1}$), and $X_{\text{SO}_4}(r)$ ($\mu\text{mol SO}_4 \text{ L}^{-1}$).
530 Glucose and amino acids diffuse out of the particle whereas O_2 and NO_3^- diffuse into the particle
531 from the surrounding environment. Due to the high concentration of SO_4^{2-} in ocean waters, we
532 assume that SO_4^{2-} is not diffusion limited inside particles, its uptake is limited by the maximum
533 uptake capacity due to physical constraint. The interactions between particle, cells, and the
534 surrounding environment are explained in Supplementary Fig. S1 and equations are provided in
535 Table 1 of the main text.

536 We assume that labile polysaccharide (C_L) and polypeptide (P_L) are hydrolyzed into glucose and
537 amino acids at rates J_C and J_P with the following functional form

$$538 \quad J_C = h_C \frac{A_C C_L}{h_C + A_C C_L} \quad (26)$$

$$539 \quad J_P = h_P \frac{A_P P_L}{h_P + A_P P_L} \quad (27)$$

540 where h_C and h_P are maximum hydrolysis rates of the carbohydrate and peptide pool, and A_C and
541 A_P are respective affinities. J_G and J_A represent uptake of glucose and amino acids:

$$542 \quad J_G = M_G \frac{A_G G}{A_G G + M_G} \quad (28)$$

$$543 \quad J_A = M_A \frac{A_A A}{A_A A + M_A} \quad (29)$$

544 where M_G and M_A are maximum uptake rates of glucose and amino acids, whereas A_G and A_A are
545 corresponding affinities. Hydrolyzed monomers diffuse out of the particle at a rate D_M .

546 μ^* is the optimal division rate of cells (equation (24)) and m_B represents the mortality rate of
 547 bacteria. F_{O_2} and $J_{NO_3^-}$ represent the diffusive flux of O_2 and the consumption rate of NO_3^- ,
 548 respectively, through the bacterial cell membrane. \bar{D}_{O_2} and $\bar{D}_{NO_3^-}$ are diffusion coefficients of O_2
 549 and NO_3^- inside the particle.

550 At the center of the particle ($r = 0$) the gradient of all quantities vanishes:

$$551 \quad \left. \frac{\partial G}{\partial r} \right|_{r=0} = \left. \frac{\partial A}{\partial r} \right|_{r=0} = \left. \frac{\partial X_{O_2}}{\partial r} \right|_{r=0} = \left. \frac{\partial X_{NO_3^-}}{\partial r} \right|_{r=0} = 0 \quad (30)$$

552 At the surface of the particle ($r = r_p$) concentrations are determined by the surrounding
 553 environment:

$$554 \quad G|_{r=r_p} = G_\infty, A|_{r=r_p} = A_\infty, X_{O_2}|_{r=r_p} = X_{O_2,\infty}, X_{NO_3^-}|_{r=r_p} = X_{NO_3^-,\infty}, \quad (31)$$

555 where G_∞ , A_∞ , $X_{O_2,\infty}$ and $X_{NO_3^-,\infty}$ are concentrations of glucose, amino acids, O_2 , and NO_3^- in the
 556 environment.

557 ***Calculation of total N_2 fixation rate***

558 The total amount of fixed N_2 in a specific size class of particle, $N_{\text{fix,P}}$ ($\mu\text{g N particle}^{-1}$), is calculated
 559 as

$$560 \quad N_{\text{fix,P}} = \int \int 4\pi r_B^2 B J_{N_2} dr_p dz, \quad (32)$$

561 where r_p (cm) is the particle radius and z (m) represents the water column depth.

562 N_2 fixation rate per unit volume of water, $N_{\text{fix,V}}(t)$ ($\mu\text{mol N m}^{-3} \text{d}^{-1}$), is calculated as

563
$$N_{\text{fix},V} = \int \int 4\pi r_B^2 \rho B J_{N_2} n(x) dr_p dx, \quad (33)$$

564 Here x (cm) represents the size range (radius) of particles, ρ is the fraction of diazotrophs of
 565 the total heterotrophic bacteria, and $n(x)$ (number of particles per unit volume of water per size
 566 increment) is the size spectrum of particles that is most commonly approximated by a power law
 567 distribution of the form

568
$$n(x) = n_0(2x)^\xi, \quad (34)$$

569 where n_0 is a constant that controls total particle abundance and the slope ξ represents the relative
 570 concentration of small to large particles: the steeper the slope, the greater the proportion of smaller
 571 particles and the flatter the slope, and the greater the proportion of larger particles³⁵.

572 Depth-integrated N_2 fixation rate, $N_{\text{fix},D}$ ($\mu\text{mol N m}^{-2} \text{d}^{-1}$), can be obtained by

573
$$N_{\text{fix},D}(t) = \int N_{\text{fix},V} dz. \quad (35)$$

574 ***Assumptions and simplification in the modelling approach***

575 According to our current model formulation, the particle size remains constant while sinking.
 576 However, in nature, particle size is dynamic due to processes like bacterial remineralization,
 577 aggregation, and disaggregation. We neglect these complications to keep the model simple and to
 578 focus on revealing the coupling between particle-associated environmental conditions and N_2
 579 fixation by heterotrophic bacteria. These factors can, however, possibly be incorporated by using *in*
 580 *situ* data or by using the relationship between carbon content and the diameter of particles⁴⁶ and
 581 including terms for aggregation and disaggregation⁵³.

582 Our model represents a population of facultative heterotrophic diazotrophs that grow at a rate
583 similar to other heterotrophic bacteria but the whole community initiates N₂ fixation when
584 conditions become suitable. However, under natural conditions, diazotrophs may only constitute a
585 fraction of the bacterial community, and their proliferation may be gradual²⁰, presumably affected
586 by multiple factors. In such case, our approach will overestimate diazotroph cell concentration and
587 consequently the N₂ fixation rate.

588 For simplicity, our approach includes only aerobic respiration, NO₃⁻ and SO₄²⁻ respiration,
589 although many additional aerobic and anaerobic processes likely occur on particles (e.g Klawonn et
590 al.¹⁸). To our knowledge, a complete picture of such processes, their interactions and effects on
591 particle biochemistry is unavailable. For example, we have assumed that when O₂ and NO₃⁻ are
592 insufficient to maintain respiration, heterotrophic bacteria start reducing SO₄²⁻. However, SO₄²⁻
593 reduction has been detected only with a significant lag after the occurrence of anaerobic conditions,
594 suggesting it as a slow adapted process⁷⁴, whereas we assume it to be instantaneous. On the other
595 hand, the lag may not be real but due to a ‘cryptic sulfur cycle’, where SO₄²⁻ reduction is
596 accompanied by concurrent sulfide oxidation effectively masking sulfide production⁷⁵. Hopefully,
597 future insights into interactions between diverse aerobic and anaerobic microbial processes can
598 refine our modelling approach and fine-tune predictions of biochemistry in marine particles.

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604 **References**

- 605 1. Gruber, N. & Galloway, J. N. An Earth-system perspective of the global nitrogen cycle.
606 *Nature* **451**, 293–296 (2008).
- 607 2. Karl, D. *et al.* Dinitrogen fixation in the world’s oceans. in *The Nitrogen Cycle at Regional*
608 *to Global Scales* 47–98 (Springer Netherlands, 2002). doi:10.1007/978-94-017-3405-9_2
- 609 3. Falkowski, P. G. Evolution of the nitrogen cycle and its influence on the biological
610 sequestration of CO₂ in the ocean. *Nature* **387**, 272–275 (1997).
- 611 4. Zehr, J. P. Nitrogen fixation by marine cyanobacteria. *Trends Microbiol.* **19**, 162–173
612 (2011).
- 613 5. Benavides, M., Bonnet, S., Berman-Frank, I. & Riemann, L. Deep into oceanic N₂ fixation.
614 *Front. Mar. Sci.* **5**, 108 (2018).
- 615 6. Mulholland, M. R. *et al.* High Rates of N₂ Fixation in Temperate, Western North Atlantic
616 Coastal Waters Expand the Realm of Marine Diazotrophy. *Global Biogeochem. Cycles* **33**,
617 826–840 (2019).
- 618 7. Blais, M. *et al.* Nitrogen fixation and identification of potential diazotrophs in the Canadian
619 Arctic. *Global Biogeochem. Cycles* **26**, (2012).
- 620 8. Farnelid, H. *et al.* Nitrogenase gene amplicons from global marine surface waters are
621 dominated by genes of non-cyanobacteria. *PLoS One* **6**, (2011).
- 622 9. Moisander, P. H. *et al.* Chasing after non-cyanobacterial nitrogen fixation in marine pelagic
623 environments. *Front. Microbiol.* **8**, (2017).

- 624 10. Bombar, D., Paerl, R. W. & Riemann, L. Marine Non-Cyanobacterial Diazotrophs: Moving
625 beyond Molecular Detection. *Trends in Microbiology* **24**, 916–927 (2016).
- 626 11. Goldberg, I., Nadler, V. & Hochman, A. Mechanism of nitrogenase switch-off by oxygen. *J.*
627 *Bacteriol.* **169**, 874–879 (1987).
- 628 12. Fay, P. Oxygen relations of nitrogen fixation in cyanobacteria. *Microbiol. Rev.* **56**, 340–373
629 (1992).
- 630 13. Bentzon-Tilia, M., Severin, I., Hansen, L. H. & Riemann, L. Genomics and ecophysiology of
631 heterotrophic nitrogen-fixing bacteria isolated from estuarine surface water. *MBio* **6**, (2015).
- 632 14. Inomura, K., Bragg, J. & Follows, M. J. A quantitative analysis of the direct and indirect
633 costs of nitrogen fixation: A model based on *Azotobacter vinelandii*. *ISME J.* **11**, 166–175
634 (2017).
- 635 15. Paerl, H. W. & Prufert, L. E. Oxygen-poor microzones as potential sites of microbial N₂
636 fixation in nitrogen-depleted aerobic marine waters. *Appl. Environ. Microbiol.* **53**, 1078–87
637 (1987).
- 638 16. Paerl, H. W. Microzone formation: Its role in the enhancement of aquatic N₂ fixation.
639 *Limnol. Oceanogr.* **30**, 1246–1252 (1985).
- 640 17. Riemann, L., Steward, G. F. & Azam, F. Dynamics of bacterial community composition and
641 activity during a mesocosm diatom bloom (Applied and Environmental Microbiology (2000)
642 66:2 (578-587)). *Appl. Environ. Microbiol.* **66**, 578–587 (2000).
- 643 18. Klawonn, I., Bonaglia, S., Brüchert, V. & Ploug, H. Aerobic and anaerobic nitrogen
644 transformation processes in N₂-fixing cyanobacterial aggregates. *ISME J.* **9**, 1456–1466

- 645 (2015).
- 646 19. Ploug, H., Kühl, M., Buchholz-Cleven, B. & Jørgensen, B. B. Anoxic aggregates - An
647 ephemeral phenomenon in the pelagic environment? *Aquat. Microb. Ecol.* **13**, 285–294
648 (1997).
- 649 20. Pedersen, J. N., Bombar, D., Paerl, R. W. & Riemann, L. Diazotrophs and N₂-Fixation
650 associated with particles in coastal estuarine waters. *Front. Microbiol.* **9**, 1–11 (2018).
- 651 21. Rahav, E., Giannetto, M. J. & Bar-Zeev, E. Contribution of mono and polysaccharides to
652 heterotrophic N₂ fixation at the eastern Mediterranean coastline. *Sci. Rep.* **6**, 1–11 (2016).
- 653 22. Scavotto, R. E., Dziallas, C., Bentzon-Tilia, M., Riemann, L. & Moisaner, P. H. Nitrogen-
654 fixing bacteria associated with copepods in coastal waters of the North Atlantic Ocean.
655 *Environ. Microbiol.* **17**, 3754–3765 (2015).
- 656 23. Yang, Q. S. *et al.* Analysis of nifH DNA and RNA reveals a disproportionate contribution to
657 nitrogenase activities by rare plankton-associated diazotrophs. *BMC Microbiol.* **19**, 188
658 (2019).
- 659 24. Farnelid, H. *et al.* Diverse diazotrophs are present on sinking particles in the North Pacific
660 Subtropical Gyre. *ISME J.* **13**, 170–182 (2019).
- 661 25. Geisler, E., Bogler, A., Rahav, E. & Bar-Zeev, E. Direct Detection of Heterotrophic
662 Diazotrophs Associated with Planktonic Aggregates. *Sci. Rep.* **9**, 1–9 (2019).
- 663 26. Smith, D. C., Simon, M., Alldredge, A. L. & Azam, F. Intense hydrolytic enzyme activity on
664 marine aggregates and implications for rapid particle dissolution. *Nature* **359**, 139–142
665 (1992).

- 666 27. Ploug, H. & Bergkvist, J. Oxygen diffusion limitation and ammonium production within
667 sinking diatom aggregates under hypoxic and anoxic conditions. *Mar. Chem.* **176**, 142–149
668 (2015).
- 669 28. Gunsalus, R. P. Control of electron flow in *Escherichia coli*: Coordinated transcription of
670 respiratory pathway genes. *Journal of Bacteriology* **174**, 7069–7074 (1992).
- 671 29. Kirchman, D. L. *Processes in Microbial Ecology*. (Oxford University Press, 2012).
- 672 30. Riederer-Henderson, M. A. & Wilson, P. W. Nitrogen fixation by sulphate-reducing bacteria.
673 *J. Gen. Microbiol.* **61**, 27–31 (1970).
- 674 31. Postgate, J. R. & Kent, H. M. Diazotrophy within *Desulfovibrio*. *J. Gen. Microbiol.* **131**,
675 2119–2122 (1985).
- 676 32. Bertics, V. J. *et al.* Occurrence of benthic microbial nitrogen fixation coupled to sulfate
677 reduction in the seasonally hypoxic Eckernförde Bay, Baltic Sea. *Biogeosciences* **10**, 1243–
678 1258 (2013).
- 679 33. Steppe, T. F. & Paerl, H. W. Potential N₂ fixation by sulfate-reducing bacteria in a marine
680 intertidal microbial mat. *Aquat. Microb. Ecol.* **28**, 1–12 (2002).
- 681 34. McDonnell, A. M. P. & Buesseler, K. O. Variability in the average sinking velocity of
682 marine particles. *Limnol. Oceanogr.* **55**, 2085–2096 (2010).
- 683 35. Jackson, G. A. *et al.* Particle size spectra between 1 µm and 1 cm at Monterey Bay
684 determined using multiple instruments. *Deep. Res. Part I Oceanogr. Res. Pap.* **44**, 1739–
685 1767 (1997).

- 686 36. Jouandet, M.-P. *et al.* Optical imaging of mesopelagic particles indicates deep carbon flux
687 beneath a natural iron-fertilized bloom in the Southern Ocean. *Limnol. Oceanogr.* **56**, 1130–
688 1140 (2011).
- 689 37. Froelich, P. N. *et al.* Early oxidation of organic matter in pelagic sediments of the eastern
690 quatorial Atlantic: suboxic diagenesis. *Geochim. Cosmochim. Acta* **43**, 1075–1090 (1979).
- 691 38. Millero, F. J. *Chemical Oceanography*. (CRC Press, 2005).
- 692 39. Maun, M. *The Biology of Coastal Sand Dunes*. (Oxford University Press, 2009).
- 693 40. de Vicente, I., Ortega-Retuerta, E., Romera, O., Morales-Baquero, R. & Reche, I.
694 Contribution of transparent exopolymer particles to carbon sinking flux in an oligotrophic
695 reservoir. *Biogeochemistry* **96**, 13–23 (2009).
- 696 41. Friedrich, U., Schallenberg, M. & Holliger, C. Pelagic bacteria-particle interactions and
697 community-specific growth rates in four lakes along a trophic gradient. *Microb. Ecol.* **37**,
698 49–61 (1999).
- 699 42. Busch, K. *et al.* Bacterial colonization and vertical distribution of marine gel particles (TEP
700 and CSP) in the arctic Fram Strait. *Front. Mar. Sci.* **4**, 166 (2017).
- 701 43. Geisler, E., Bogler, A., Bar-Zeev, E. & Rahav, E. Heterotrophic Nitrogen Fixation at the
702 Hyper-Eutrophic Qishon River and Estuary System. *Front. Microbiol.* **11**, 2012–2021
703 (2020).
- 704 44. Kiørboe, T. Formation and fate of marine snow: Small-scale processes with large-scale
705 implications. *Sci. Mar.* **65**, 57–71 (2001).

- 706 45. Martiny, A. C., Vrugt, J. A., Primeau, F. W. & Lomas, M. W. Regional variation in the
707 particulate organic carbon to nitrogen ratio in the surface ocean. *Global Biogeochem. Cycles*
708 **27**, 723–731 (2013).
- 709 46. Guidi, L. *et al.* Relationship between particle size distribution and flux in the mesopelagic
710 zone. *Deep. Res. Part I Oceanogr. Res. Pap.* **55**, 1364–1374 (2008).
- 711 47. Boström, K. H., Riemann, L., Kühl, M. & Hagström, Å. Isolation and gene quantification of
712 heterotrophic N₂-fixing bacterioplankton in the Baltic Sea. *Environ. Microbiol.* **9**, 152–164
713 (2007).
- 714 48. Iversen, M. H. & Ploug, H. Ballast minerals and the sinking carbon flux in the ocean:
715 Carbon-specific respiration rates and sinking velocity of marine snow aggregates.
716 *Biogeosciences* **7**, 2613–2624 (2010).
- 717 49. Eduardo Menschel, A. & González, H. E. Carbon and Calcium Carbonate Export Driven by
718 Appendicularian Faecal Pellets in the Humboldt Current System off Chile. *Sci. Rep.* **9**, 1–12
719 (2019).
- 720 50. Wiedmann, I., Reigstad, M., Marquardt, M., Vader, A. & Gabrielsen, T. M. Seasonality of
721 vertical flux and sinking particle characteristics in an ice-free high arctic fjord-Different from
722 subarctic fjords? *J. Mar. Syst.* **154**, 192–205 (2016).
- 723 51. Bach, L. T. *et al.* The Influence of Plankton Community Structure on Sinking Velocity and
724 Remineralization Rate of Marine Aggregates. *Global Biogeochem. Cycles* **33**, 971–994
725 (2019).
- 726 52. Cram, J. A. *et al.* The Role of Particle Size, Ballast, Temperature, and Oxygen in the Sinking

- 727 Flux to the Deep Sea. *Global Biogeochem. Cycles* **32**, 858–876 (2018).
- 728 53. Bianchi, D., Weber, T. S., Kiko, R. & Deutsch, C. Global niche of marine anaerobic
729 metabolisms expanded by particle microenvironments. *Nat. Geosci.* **11**, 1–6 (2018).
- 730 54. Desai, M. S., Assig, K. & Dattagupta, S. Nitrogen fixation in distinct microbial niches within
731 a chemoautotrophy-driven cave ecosystem. *ISME J.* **7**, 2411–2423 (2013).
- 732 55. Leloup, J. *et al.* Sulfate-reducing bacteria in marine sediment (Aarhus Bay, Denmark):
733 Abundance and diversity related to geochemical zonation. *Environ. Microbiol.* **11**, 1278–
734 1291 (2009).
- 735 56. Berelson, W. M. Particle settling rates increase with depth in the ocean. *Deep. Res. Part II*
736 *Top. Stud. Oceanogr.* **49**, 237–251 (2001).
- 737 57. De La Rocha, C. L., Nowald, N. & Passow, U. Interactions between diatom aggregates,
738 minerals, particulate organic carbon, and dissolved organic matter: Further implications for
739 the ballast hypothesis. *Global Biogeochem. Cycles* **22**, (2008).
- 740 58. Ploug, H., Iversen, M. H. & Fischer, G. Ballast, sinking velocity, and apparent diffusivity
741 within marine snow and zooplankton fecal pellets: Implications for substrate turnover by
742 attached bacteria. *Limnol. Oceanogr.* **53**, 1878–1886 (2008).
- 743 59. Alldredge, A. L. & Gotschalk, C. In situ settling behavior of marine snow. *Limnol.*
744 *Oceanogr.* **33**, 339–351 (1988).
- 745 60. Garcia, H. *et al.* *World Ocean Atlas 2018, Volume 4: Dissolved Inorganic Nutrients*
746 *(phosphate, nitrate and nitrate+nitrite, silicate)*. A. Mishonov Technical Ed.; NOAA Atlas
747 NESDIS 84. (2019).

- 748 61. Garcia, H. *et al.* *World Ocean Atlas 2018, Volume 3: Dissolved Oxygen, Apparent Oxygen*
749 *Utilization, and Oxygen Saturation. A. Mishonov Technical Ed.; NOAA Atlas NESDIS 83.*
750 (2019).
- 751 62. Wiedmann, I., Reigstad, M., Sundfjord, A. & Basedow, S. Potential drivers of sinking
752 particle's size spectra and vertical flux of particulate organic carbon (POC): Turbulence,
753 phytoplankton, and zooplankton. *J. Geophys. Res. Ocean.* **119**, 6900–6917 (2014).
- 754 63. Benavides, M. *et al.* Basin-wide N₂ fixation in the deep waters of the Mediterranean Sea.
755 *Global Biogeochem. Cycles* **30**, 952–961 (2016).
- 756 64. Shiozaki, T., Ijichi, M., Kodama, T., Takeda, S. & Furuya, K. Heterotrophic bacteria as
757 major nitrogen fixers in the euphotic zone of the Indian Ocean. *Global Biogeochem. Cycles*
758 **28**, 1096–1110 (2014).
- 759 65. Halm, H. *et al.* Heterotrophic organisms dominate nitrogen fixation in the south pacific gyre.
760 *ISME J.* **6**, 1238–1249 (2012).
- 761 66. Luo, Y. W. *et al.* Database of diazotrophs in global ocean: Abundance, biomass and nitrogen
762 fixation rates. *Earth Syst. Sci. Data* **4**, 47–73 (2012).
- 763 67. Rahav, E. *et al.* Dinitrogen fixation in aphotic oxygenated marine environments. *Front.*
764 *Microbiol.* **4**, 227 (2013).
- 765 68. Boeuf, D. *et al.* Biological composition and microbial dynamics of sinking particulate
766 organic matter at abyssal depths in the oligotrophic open ocean. *Proc. Natl. Acad. Sci. U. S.*
767 *A.* **116**, 11824–11832 (2019).
- 768 69. Pirt, S. J. Maintenance energy: a general model for energy-limited and energy-sufficient

- 769 growth. *Arch. Microbiol.* **133**, 300–302 (1982).
- 770 70. Mislán, K. A. S., Stock, C. A., Dunne, J. P. & Sarmiento, J. L. Group behavior among model
771 bacteria influences particulate carbon remineralization depths. *J. Mar. Res.* **72**, 183–218
772 (2014).
- 773 71. Dalton, H. & Postgate, J. R. Effect of Oxygen on Growth of *Azotobacter chroococcum* in
774 Batch and Continuous Cultures. *J. Gen. Microbiol.* **54**, 463–473 (2009).
- 775 72. Oelze, J. Respiratory protection of nitrogenase in *Azotobacter* species: is a widely held
776 hypothesis unequivocally supported by experimental evidence? *FEMS Microbiol. Rev.* **24**,
777 321–333 (2000).
- 778 73. Inomura, K., Bragg, J., Riemann, L. & Follows, M. J. A quantitative model of nitrogen
779 fixation in the presence of ammonium. *PLoS One* **13**, e0208282 (2018).
- 780 74. Santegoeds, C. M., Ferdeman, T. G., Muyzer, G. & De Beer, D. Structural and functional
781 dynamics of sulfate-reducing populations in bacterial biofilms. *Appl. Environ. Microbiol.* **64**,
782 3731–3739 (1998).
- 783 75. Callbeck, C. M. *et al.* Oxygen minimum zone cryptic sulfur cycling sustained by offshore
784 transport of key sulfur oxidizing bacteria. *Nat. Commun.* **9**, 1–11 (2018).
- 785 76. Alldredge, A. L., Passow, U. & Haddock, S. H. D. The characteristics and transparent
786 exopolymer particle (TEP) content of marine snow formed from thecate dinoflagellates. *J.*
787 *Plankton Res.* **20**, 393–406 (1998).

788

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796 **Author contributions**

797 SC, LR, KA, and AV developed the model. SC performed all numerical simulations and analyses
798 with input from KA and AV. SC and LR wrote the manuscript. All authors provided critical
799 feedback and helped shape the manuscript.

800 **Competing interests**

801 The authors declare no competing interests.

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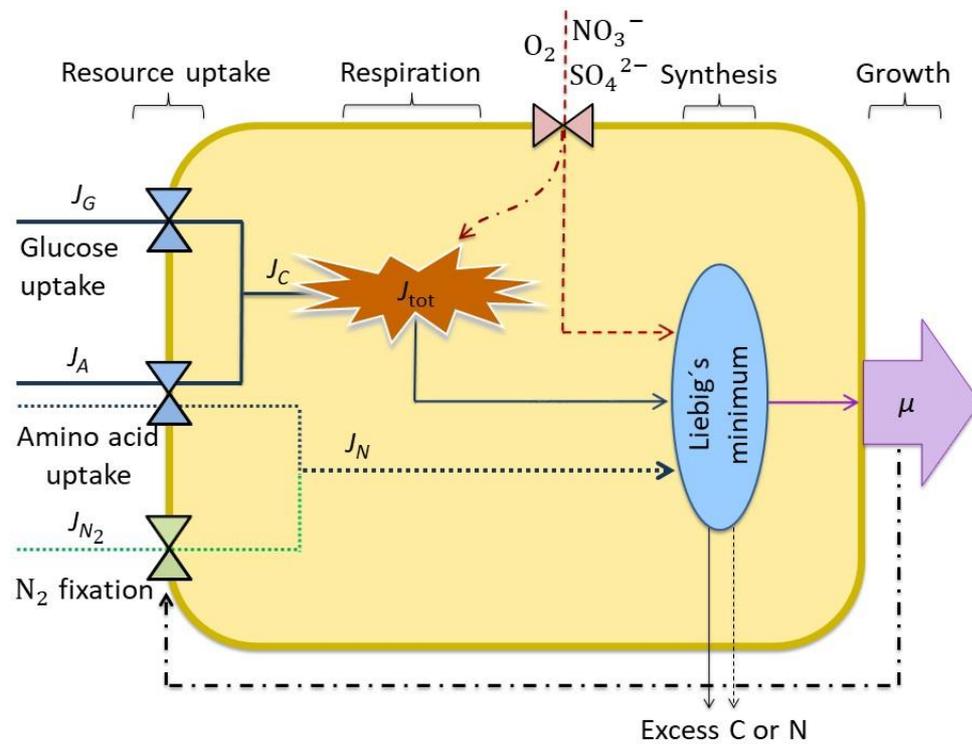
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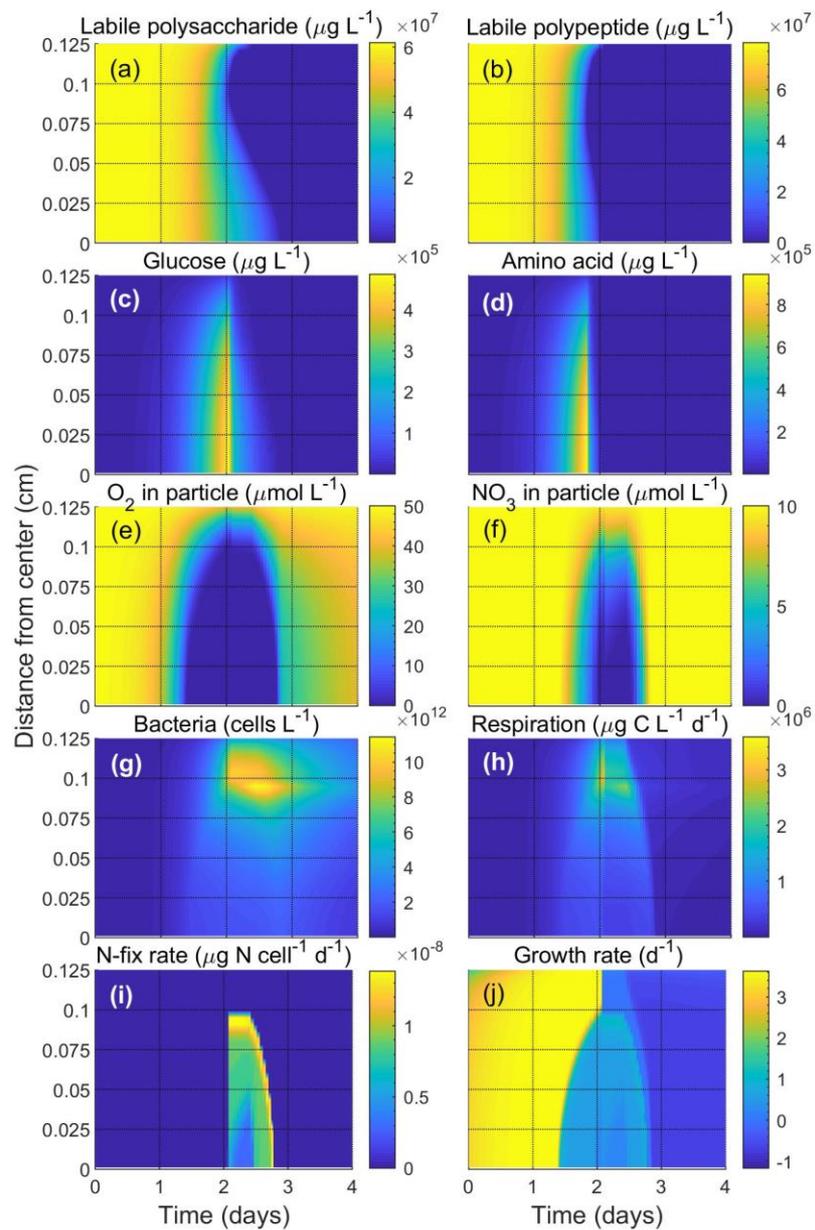
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808 **Figures:**



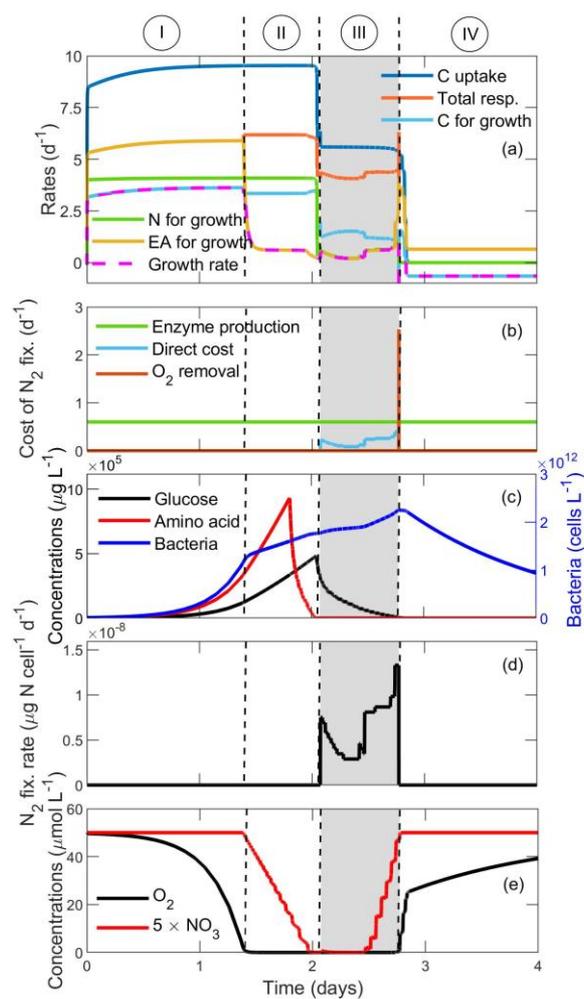
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810 **Fig. 1** Schematic representation of the cellular processes. It shows how fluxes of carbon (C), J_C
811 (solid lines), fluxes of nitrogen (N), J_N (dotted lines), and electron acceptors (O_2 , NO_3^- , and SO_4^{2-} ;
812 red dashed line) are combined (blue ellipse) to determine growth rate after paying the cost of
813 respiration (brown explosion). Triangle symbols represent the functional responses for the uptake
814 mechanisms and diffusive inflow of O_2 , NO_3^- , and SO_4^{2-} . J_{tot} represents respiration that includes
815 costs of uptake and mobilization of resources for synthesis, construction/maintenance of structure,
816 and ectoenzyme production. The ellipse represents the synthesis of biomass from the available C, N
817 and electron acceptors following Liebig's law of the minimum. Any excess assimilated C or N is
818 excreted from the cell. μ represents the division rate. The black dashed-dotted line represents the
819 regulation of N₂ fixation to optimize growth rate and the red dashed-dot line represents the
820 regulation of respiration by electron acceptors.



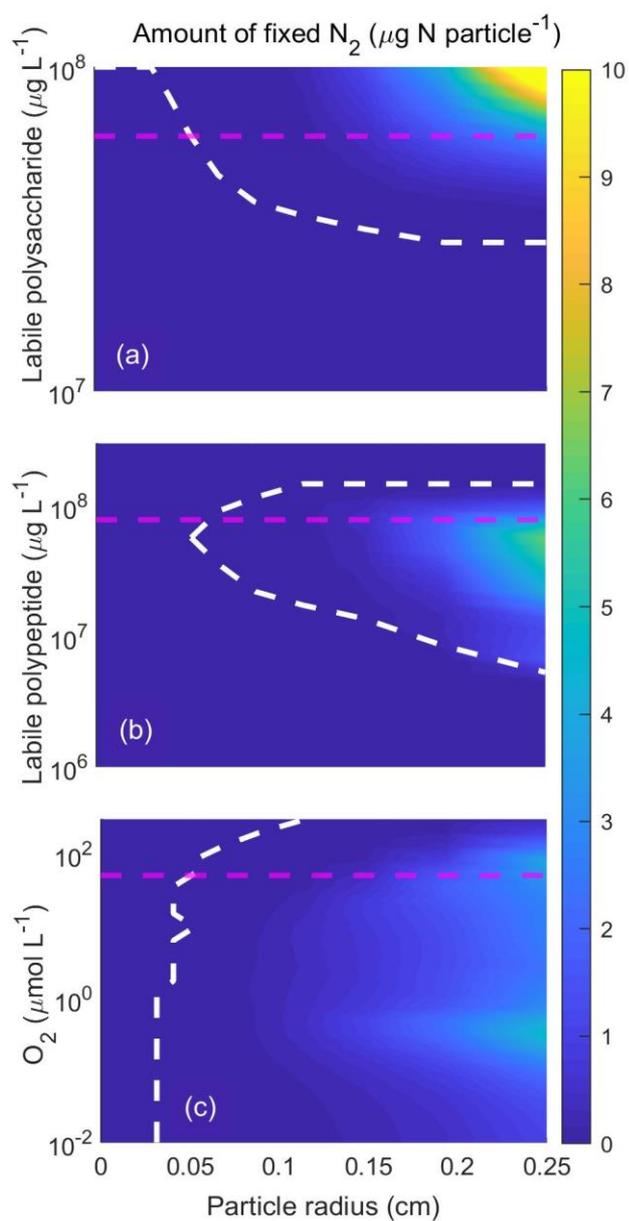
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822 **Fig. 2** Dynamics inside a particle of radius 0.25 cm with time. (a) Labile carbohydrate, (b) labile
 823 polypeptide, (c) glucose, (d) amino acid, (e) O₂ in particle, (f) NO₃⁻ in particle, (g) bacterial
 824 abundance, (h) respiration rate, (i) N₂ fixation rate, and (j) growth rate are shown along the particle
 825 radius over time. Parameters and concentrations of surrounding factors are taken from Table 2.



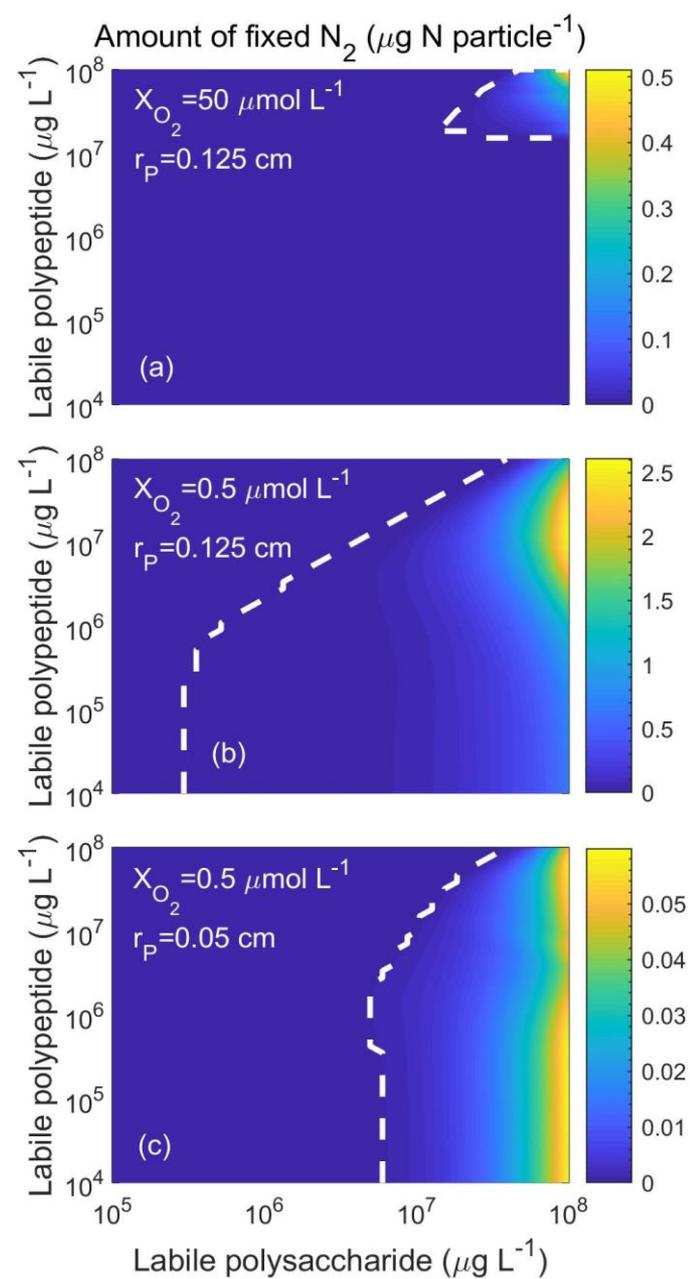
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827 **Fig. 3** Cellular rates and resource concentrations at a radial distance of 0.027 cm from the center of
 828 a 0.25 cm particle. (a) Size-specific rates of C uptake (dark blue), total respiration (orange),
 829 available C for growth (light blue), available N for growth (green), available EA (electron acceptor
 830 O_2 , NO_3^- , and SO_4^{2-} ; yellow), and growth rate of a cell (dashed magenta). Regions I, II, III, and IV
 831 represent situations when a cell is limited by C, EA, co-limited by N and EA, N, and showing
 832 negative growth rate (see text). (b) Respiratory costs related to N_2 fixation in terms of direct
 833 respiration, enzyme production, and O_2 removal. (c) Glucose, amino acid, and bacterial
 834 concentrations in the particle. (d) Cellular N_2 fixation rate. (e) O_2 and NO_3^- concentrations in the
 835 particle. The grey area represents the time interval of N_2 fixation.



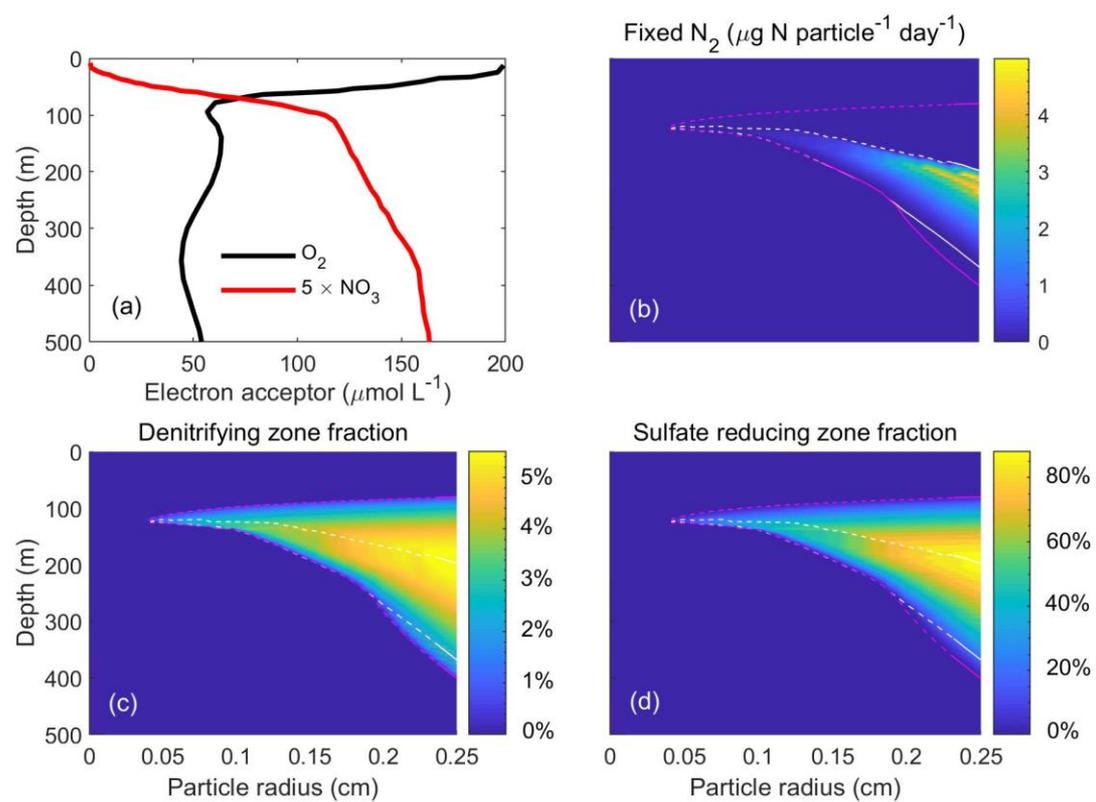
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837 **Fig. 4** The total amount of fixed N_2 per particle as a function of particle radius and initial labile
 838 polysaccharide (a), initial labile polypeptide (b), and surrounding O_2 concentration (c). The white
 839 dashed line separates regions of occurrence and non-occurrence of N_2 fixation ($N_{\text{fix}} > 10^{-3} \mu\text{g N}$
 840 particle^{-1}). The horizontal magenta dashed lines indicate the base value for other plots, e.g., in (a),
 841 magenta dashed line represents the level of polysaccharide concentration in panels b and c. The
 842 chosen concentration ranges corresponds to those in natural particles^{58,76}.



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844 **Fig. 5** The total amount of fixed N_2 at different initial labile polysaccharide and polypeptide
 845 concentrations. N_2 fixation is observed in a large particle of radius 0.125 cm and surrounding O_2
 846 concentrations (a) $50 \mu\text{mol } O_2 \text{ L}^{-1}$ and (b) $0.5 \mu\text{mol } O_2 \text{ L}^{-1}$, and (c) in a relatively smaller particle of
 847 radius 0.05 cm and surrounding O_2 concentrations $0.5 \mu\text{mol } O_2 \text{ L}^{-1}$. Line types are similar as in Fig.
 848 4. Note the different ranges of the scale of N_2 fixation.



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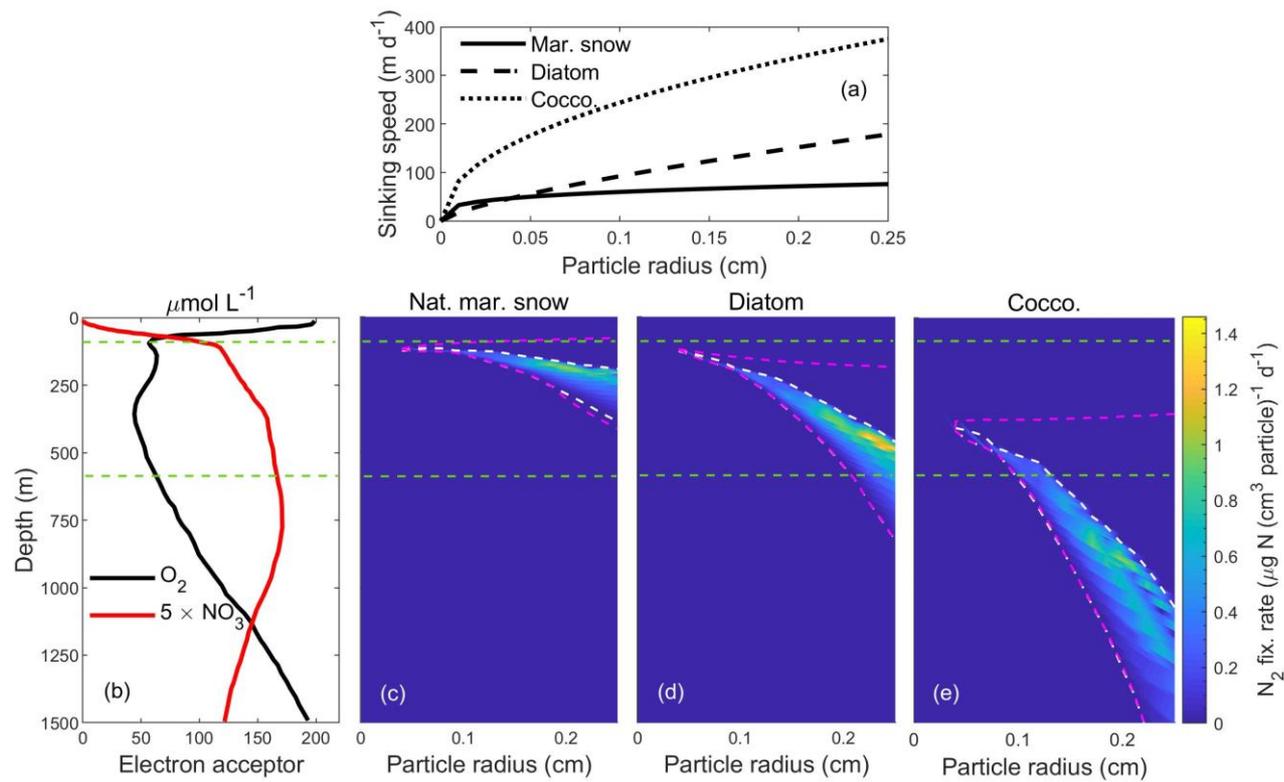
850 **Fig. 6** Dynamics inside sinking particles of different sizes in the upper 500 m of the water column.

851 (a) O_2 and NO_3^- profile in the water column⁵³. Total amount of fixed N_2 per particle (b), fraction
 852 of particle volume where respiration is fueled by denitrification (c), and fraction of particle volume
 853 where respiration is fueled by SO_4^{2-} reduction (d). Areas enclosed by white and magenta dashed
 854 lines represent the occurrence of N_2 fixation ($>10^{-3} \mu\text{g N day}^{-1} \text{ particle}^{-1}$) and anoxia ($\text{O}_2 < 10^{-3}$
 855 $\mu\text{mol O}_2 \text{ L}^{-1}$), respectively, at the center of particles.

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860 **Fig. 7** N₂ fixation rates in different types of sinking particles with different sinking speeds. (a)

861 Particle radius (cm) versus sinking speeds (m d⁻¹) of natural marine snow (continuous line)

862 measured *in situ* off California ($v = 108.95 \times r_p^{0.26}$; ⁵⁹) and of laboratory-made diatom (dashed

863 line) and coccolithophore (dotted line) aggregates measured *in vitro* ($v = 484.09 \times r_p^{0.72}$ and $v =$

864 $719.44 \times r_p^{0.47}$, respectively⁴⁸). (b) O₂ and NO₃⁻ profiles in the water column⁵³. (c-e) N₂ fixation

865 rates per unit volume of particles of different sizes and types. White and magenta dashed lines are

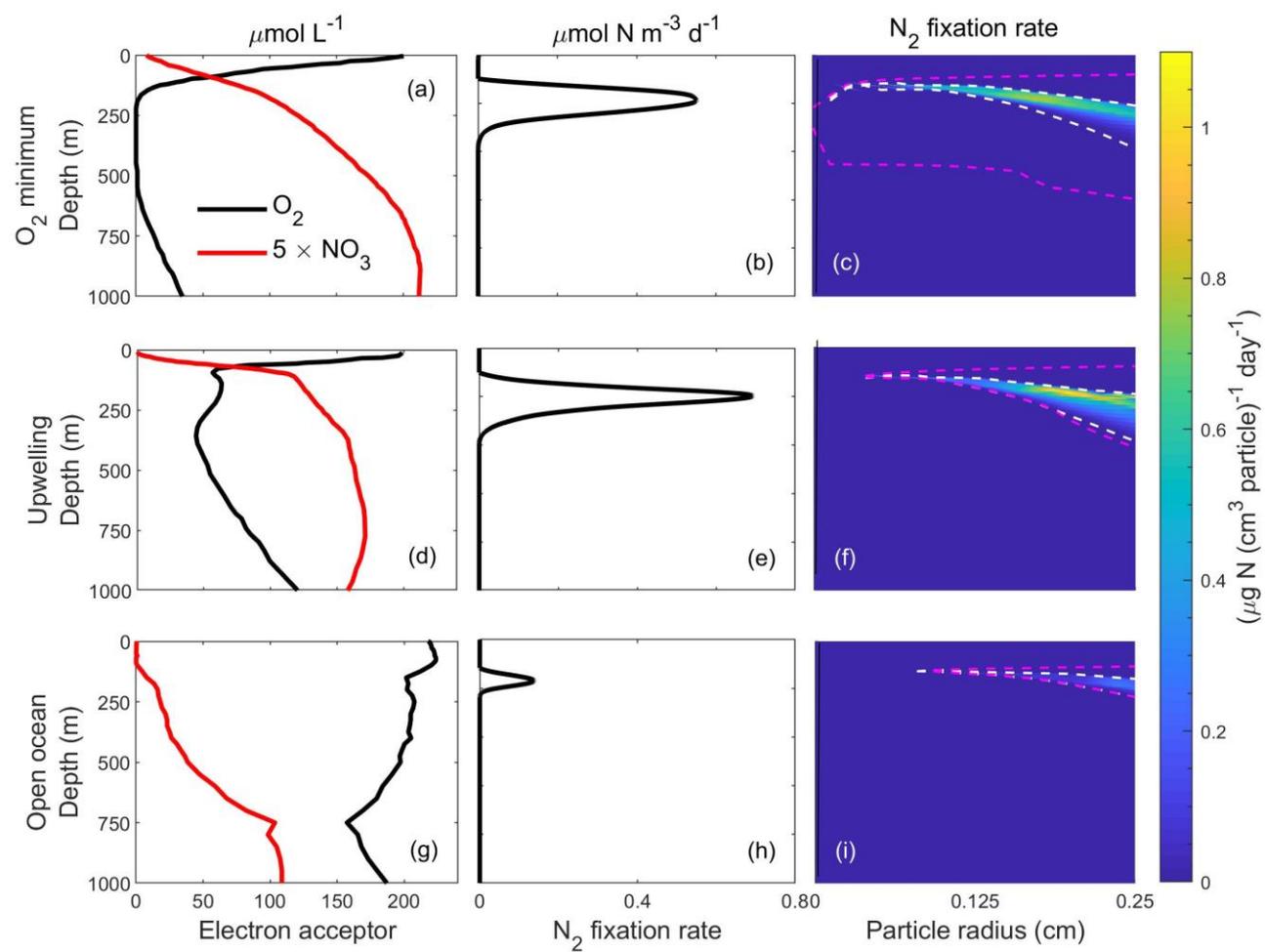
866 same as in Fig. 6. The area enclosed within the horizontal green lines represents the low-O₂ zone.

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872 **Fig. 8** Comparison of predicted N_2 fixation rates in natural marine snow at three contrasting sites in
 873 terms of vertical distributions of O_2 and NO_3 in the ocean: (a-c) O_2 minimum zone in the Eastern
 874 Tropical South Pacific (ETSP)⁵³, (d-f) the Mauritanian upwelling zone in the North Atlantic Ocean
 875 (NAO)⁵³, and (g-i) open ocean (OO; 30.5°N, 52.5°W)^{60,61}. (a,d,g) O_2 and NO_3 concentrations in
 876 the upper 1000 m vertical water column. (b,e,h) N_2 fixation rates per unit volume of water. (c,f,i) N_2
 877 fixation rates per unit volume of particle of different size classes.

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881 **Table 1.** Equations for the particle model. All quantities vary with time t and with distance from
882 the center r . The operator in the brackets represents diffusion in spherical coordinates. Definitions,
883 units, and values of each of the parameters are provided in Supplementary Table S1.

Variables	Equations	
Bacteria (cells L ⁻¹)	$\frac{\partial B}{\partial t} = \mu^*(G, A, X_{O_2}, X_{NO_3})B - m_B B$	25(a)
Labile polysaccharides (μg G L ⁻¹)	$\frac{\partial C_L}{\partial t} = -J_C B$	25(b)
Labile polypeptides (μg A L ⁻¹)	$\frac{\partial P_L}{\partial t} = -J_P B$	25(c)
Glucose (μg G L ⁻¹)	$\frac{\partial G}{\partial t} = J_C B - J_G B + D_M \left(\frac{\partial^2 G}{\partial r^2} + \frac{2}{r} \frac{\partial G}{\partial r} \right)$	25(d)
Amino acids (μg A L ⁻¹)	$\frac{\partial A}{\partial t} = J_P B - J_A B + D_M \left(\frac{\partial^2 A}{\partial r^2} + \frac{2}{r} \frac{\partial A}{\partial r} \right)$	25(e)
Oxygen (μmol O ₂ L ⁻¹)	$\frac{\partial X_{O_2}}{\partial t} = -F_{O_2} B + \bar{D}_{O_2} \left(\frac{\partial^2 X_{O_2}}{\partial r^2} + \frac{2}{r} \frac{\partial X_{O_2}}{\partial r} \right)$	25(f)
Nitrate (μmol NO ₃ L ⁻¹)	$\frac{\partial X_{NO_3}}{\partial t} = -J_{NO_3} B + \bar{D}_{NO_3} \left(\frac{\partial^2 X_{NO_3}}{\partial r^2} + \frac{2}{r} \frac{\partial X_{NO_3}}{\partial r} \right)$	25(g)

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Figures

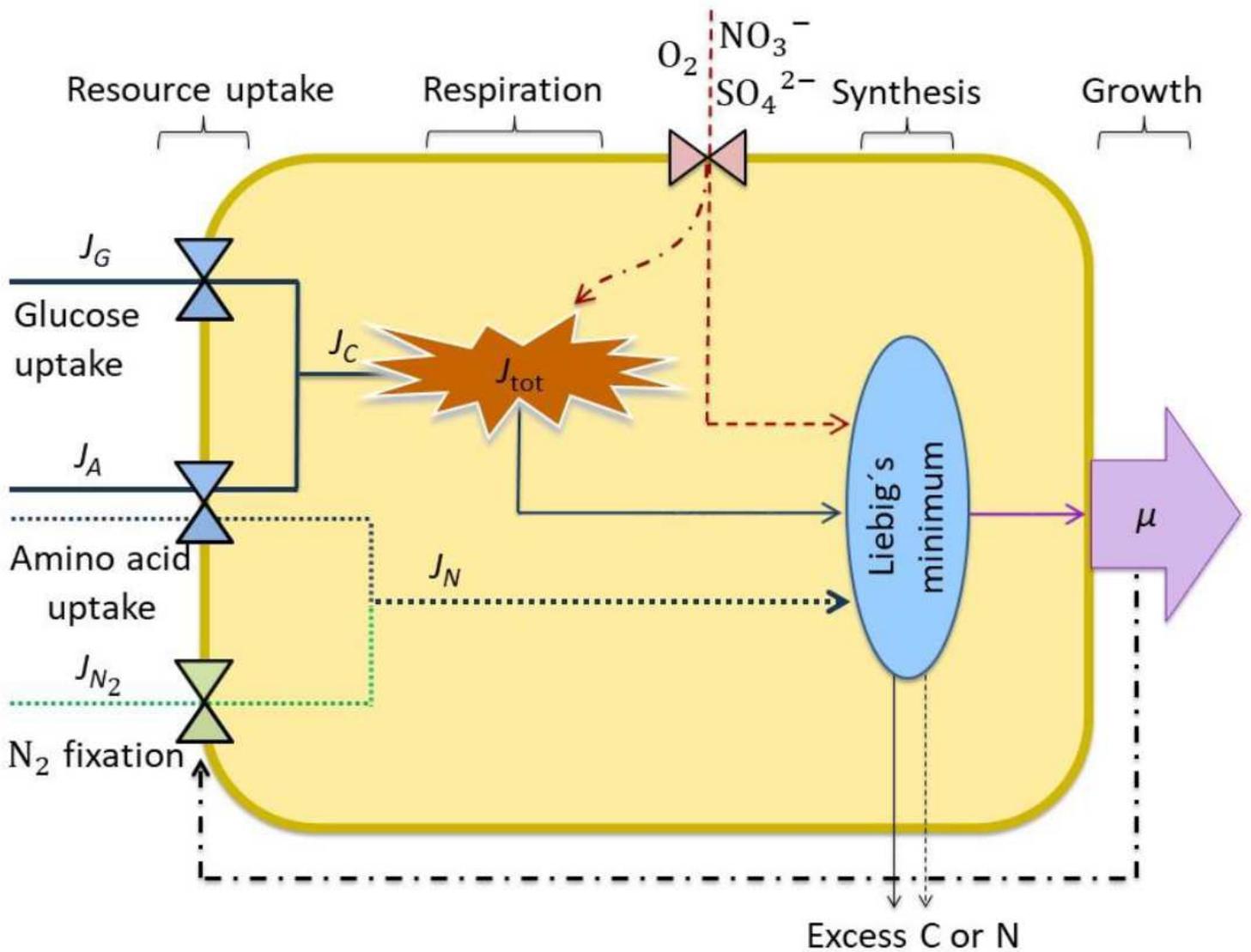


Figure 1

Schematic representation of the cellular processes. It shows how fluxes of carbon (C), C 811 (solid lines), fluxes of nitrogen (N), N (dotted lines), and electron acceptors (O_2 , NO_3^- , and SO_4^{2-} ; red dashed line) are combined (blue ellipse) to determine growth rate after paying the cost of respiration (brown explosion). Triangle symbols represent the functional responses for the uptake mechanisms and diffusive inflow of O_2 , NO_3^- , and SO_4^{2-} . J_{tot} represents respiration that in costs of uptake and mobilization of resources for synthesis, construction/maintenance of structure, and ectoenzyme production. The ellipse represents the synthesis of biomass from the available C, N and electron acceptors following Liebig's law of the minimum. Any excess assimilated C or N is excreted from the cell. μ represents the division rate. The black dashed-dotted line represents the regulation of N_2 fixation to optimize growth rate and the red dashed-dot line represents the regulation of respiration by electron acceptors.

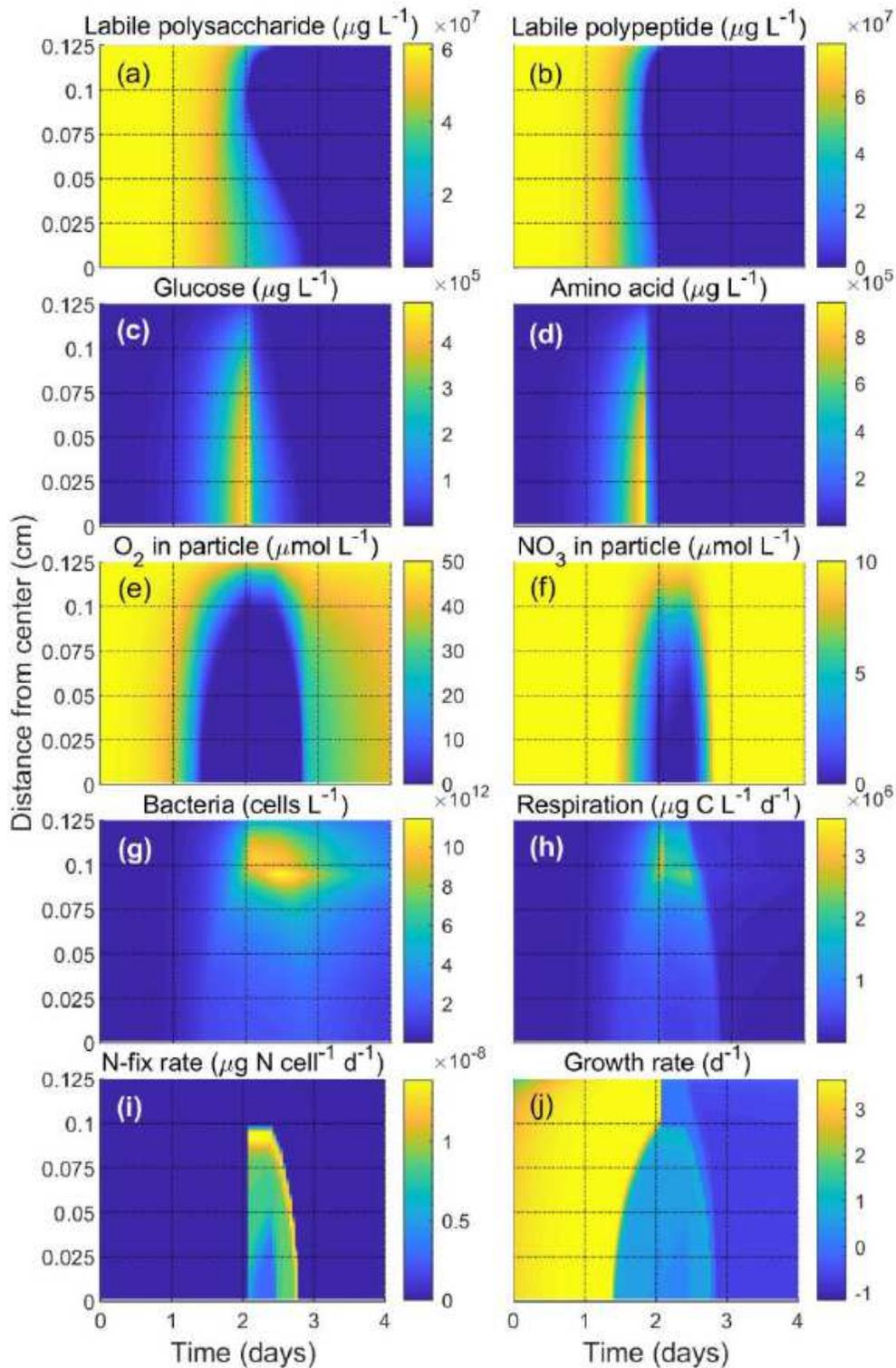


Figure 2

Dynamics inside a particle of radius 0.25 cm with time. (a) Labile carbohydrate, (b) labile polypeptide, (c) glucose, (d) amino acid, (e) O₂ in particle, (f) NO₃⁻ in particle, (g) bacterial abundance, (h) respiration rate, (i) N₂ fixation rate, and (j) growth rate are shown along the particle radius over time. Parameters and concentrations of surrounding factors are taken from Table 2.

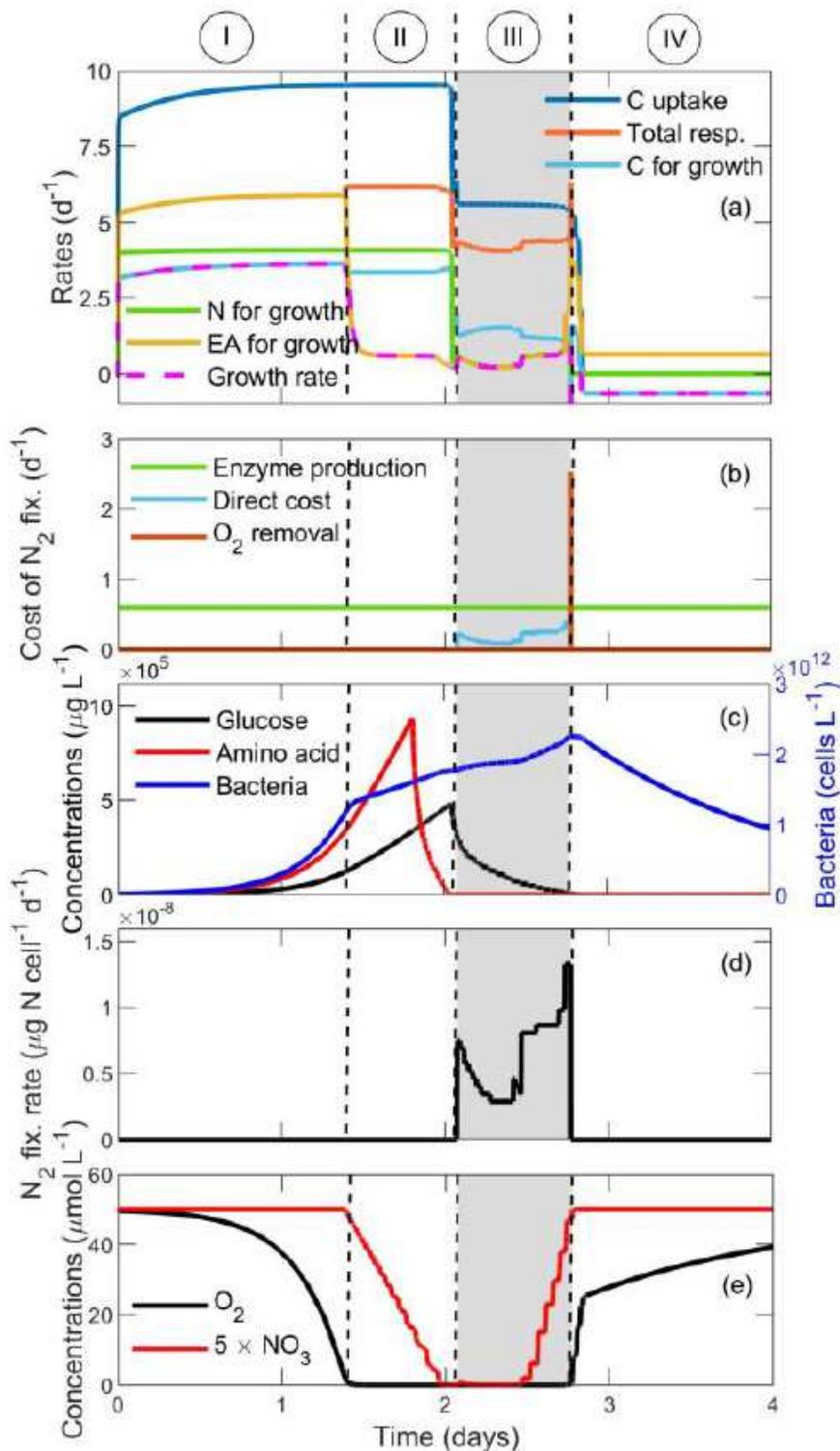


Figure 3

Cellular rates and resource concentrations at a radial distance of 0.027 cm from the center of a 0.25 cm particle. (a) Size-specific rates of C uptake (dark blue), total respiration (orange), available C for growth (light blue), available N for growth (green), available EA (electron acceptor 830 O₂, NO₃⁻, and SO₄²⁻; yellow), and growth rate of a cell (dashed magenta). Regions I, II, III, and IV represent situations when a cell is limited by C, EA, co-limited by N and EA, N, and showing negative growth rate (see text). (b)

Respiratory costs related to N₂ fixation in terms of direct respiration, enzyme production, and O₂ removal. (c) Glucose, amino acid, and bacterial concentrations in the particle. (d) Cellular N₂ fixation rate. (e) O₂ and NO₃⁻ concentrations in the particle. The grey area represents the time interval of N₂ fixation.

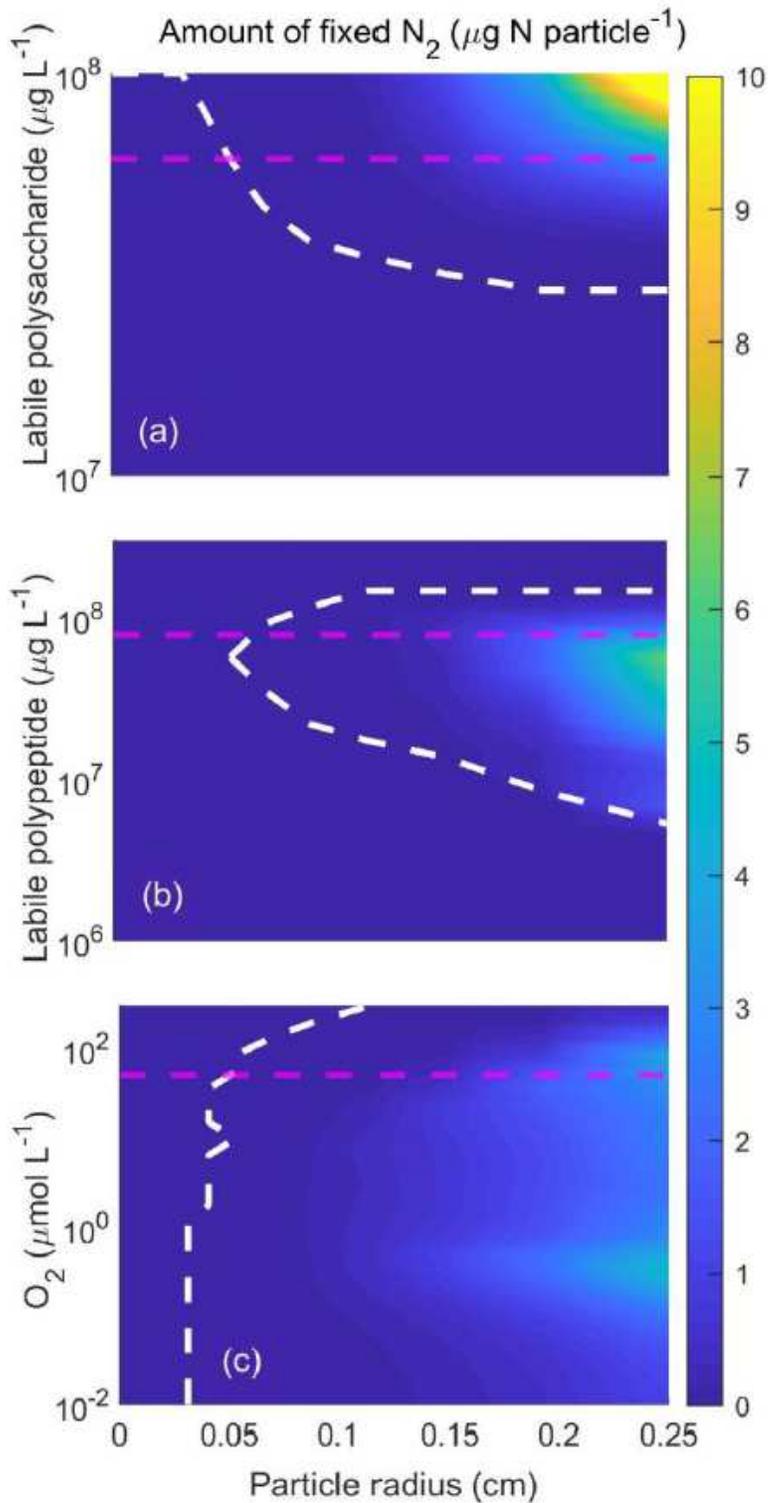


Figure 4

The total amount of fixed N₂ per particle as a function of particle radius and initial labile polysaccharide (a), initial labile polypeptide (b), and surrounding O₂ concentration (c). The white dashed line separates

regions of occurrence and non-occurrence of N₂ fixation ($N_{fix} > 10^{-3} \mu\text{g N particle}^{-1}$). The horizontal magenta dashed lines indicate the base value for other plots, e.g., in (a), magenta dashed line represents the level of polysaccharide concentration in panels b and c. The chosen concentration ranges corresponds to those in natural particles 58,76.

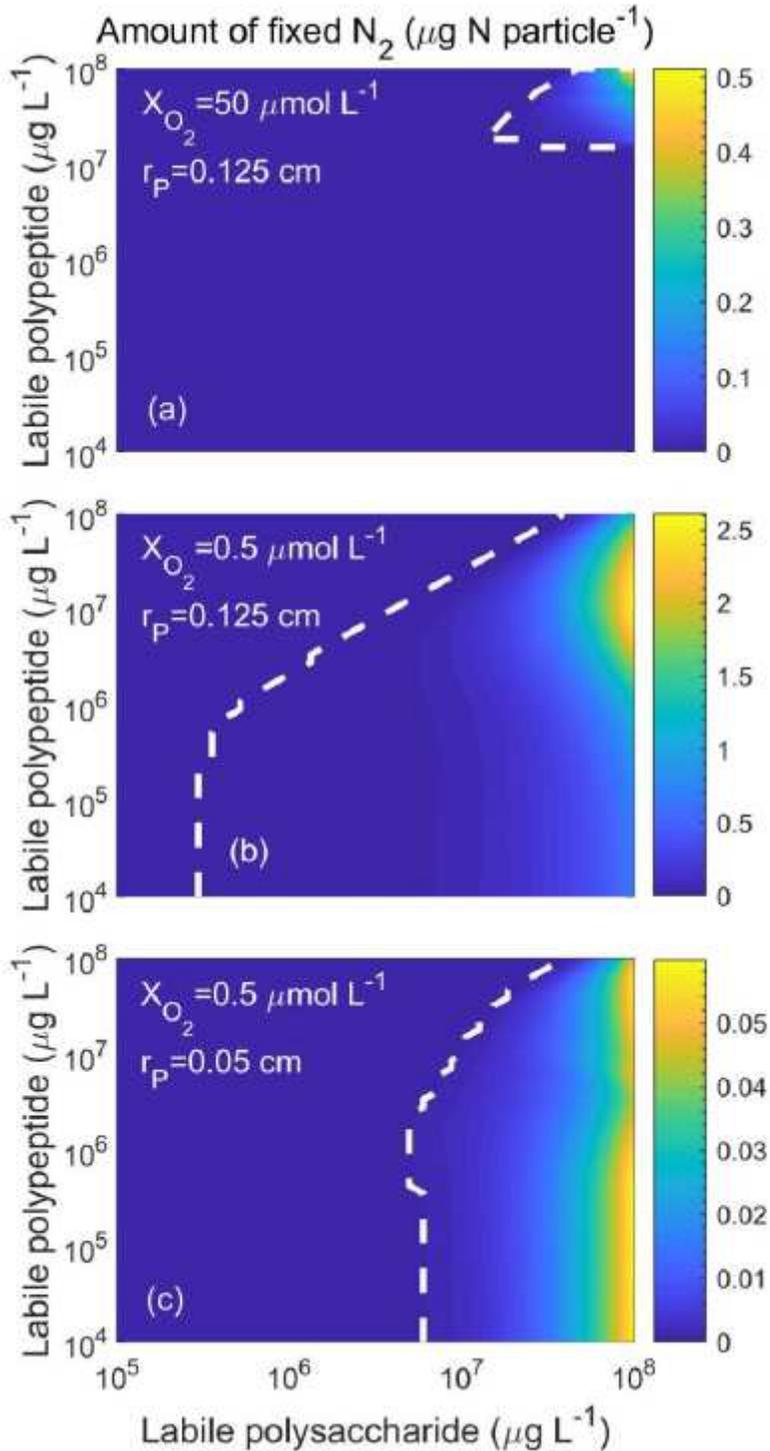


Figure 5

The total amount of fixed N₂ at different initial labile polysaccharide and polypeptide concentrations. N₂ fixation is observed in a large particle of radius 0.125 cm and surrounding O₂ concentrations (a) 50 μmol

0.5 $\mu\text{mol L}^{-1}$ and (b) 0.5 $\mu\text{mol L}^{-1}$, and (c) in a relatively smaller particle of radius 0.05 cm and surrounding O_2 concentrations 0.5 $\mu\text{mol L}^{-1}$. Line types are similar as in Fig. 4. Note the different ranges of the scale of N_2 fixation.

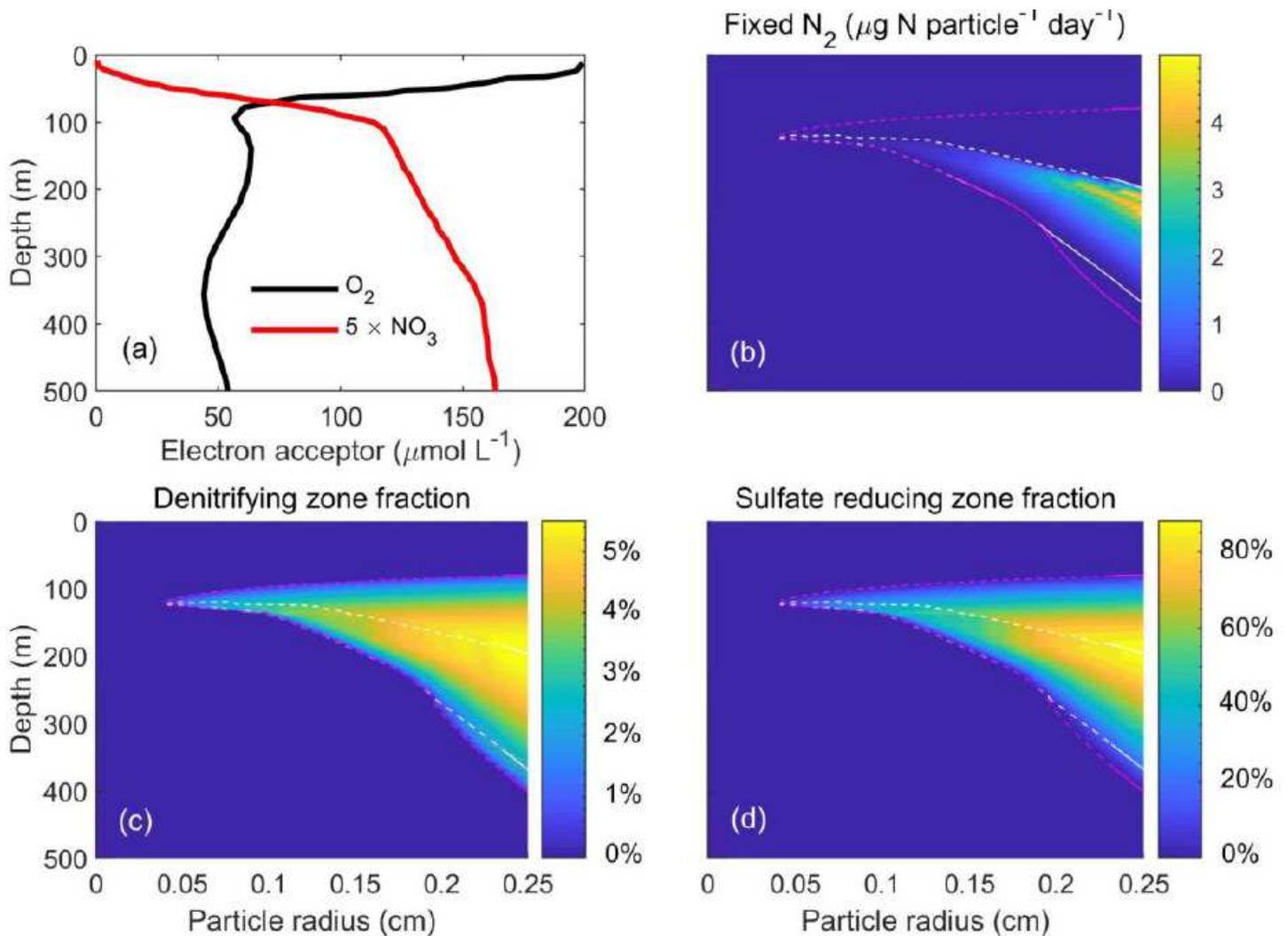


Figure 6

Dynamics inside sinking particles of different sizes in the upper 500 m of the water column. (a) O_2 and NO_3^- profile in the water column 53. Total amount of fixed N_2 per particle (b), fraction of particle volume where respiration is fueled by denitrification (c), and fraction of particle volume where respiration is fueled by SO_4^{2-} reduction (d). Areas enclosed by white and magenta dashed lines represent the occurrence of N_2 fixation ($>10^{-3} \mu\text{g N day}^{-1} \text{ particle}^{-1}$) and anoxia ($\text{O}_2 < 10^{-3} \mu\text{mol L}^{-1}$), respectively, at the center of particles.

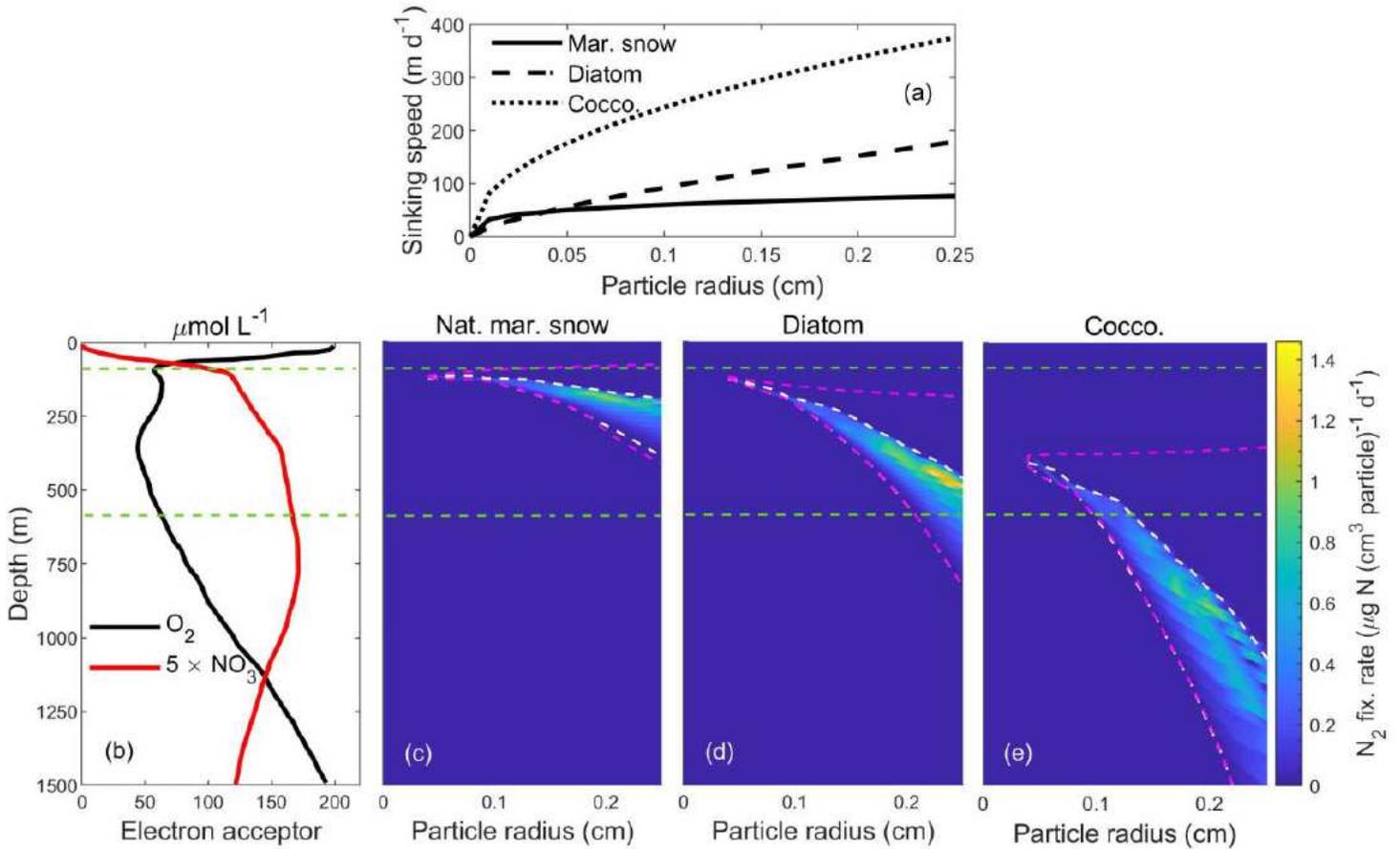


Figure 7

N₂ fixation rates in different types of sinking particles with different sinking speeds. (a) Particle radius (cm) versus sinking speeds (m d⁻¹) of natural marine snow (continuous line) measured in situ off California ($\bar{x} = 108.95 \times \bar{\sigma} 0.26; 59$) and of laboratory-made diatom (dashed line) and coccolithophore (dotted line) aggregates measured in vitro ($\bar{x} = 484.09 \times \bar{\sigma} 0.72$ and $\bar{x} = 719.44 \times \bar{\sigma} 0.47$, respectively⁴⁸). (b) O₂ and NO₃⁻ profiles in the water column⁵³. (c-e) N₂ fixation rates per unit volume of particles of different sizes and types. White and magenta dashed lines are same as in Fig. 6. The area enclosed within the horizontal green lines represents the low-O₂ zone.

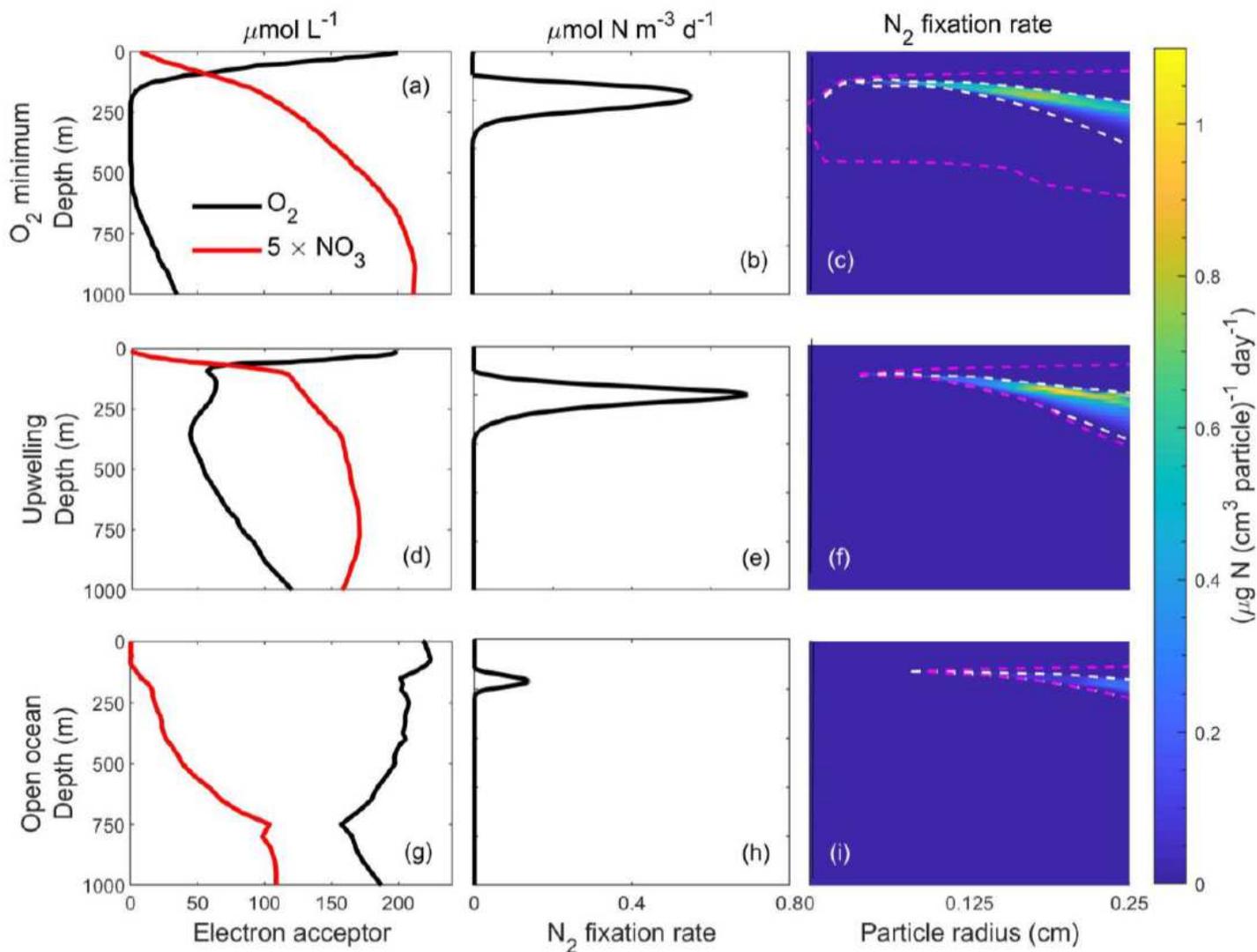


Figure 8

Comparison of predicted N₂ fixation rates in natural marine snow at three contrasting sites in terms of vertical distributions of O₂ and NO₃ in the ocean: (a-c) O₂ minimum zone in the Eastern Tropical South Pacific (ETSP)53, (d-f) the Mauritanian upwelling zone in the North Atlantic Ocean (NAO)53, and (g-i) open ocean (OO; 30.5oN, 52.5oW)60,61. (a,d,g) O₂ and NO₃ concentrations in the upper 1000 m vertical water column. (b,e,h) N₂ fixation rates per unit volume of water. (c,f,i) N₂ fixation rates per unit volume of particle of different size classes.

Supplementary Files

This is a list of supplementary files associated with this preprint. Click to download.

- [SuppChakrabortyetal04112020.pdf](#)