

Climate change and deforestation drive the displacement and contraction of tropical montane cloud forests

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1 **Climate change and deforestation drive the displacement**
2 **and contraction of tropical montane cloud forests**

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19

20 **Global climate change and habitat loss are displacing tropical montane**
21 **forests along mountain slopes^{1–4}. Cloud forests are one of the most diverse**
22 **and fragile of these montane ecosystems^{5–8}, yet little is known about the**
23 **historical and ongoing impacts of anthropogenic disturbances on these**
24 **forests. Here we assess historical (1901–2016) changes in the altitudinal**
25 **range of vascular plant species in Mesoamerican cloud forests and evaluate**
26 **the relative impacts of climate change and land-use alterations. By analysing**
27 **thousands of occurrence records from public biological collections, we**
28 **uncovered common altitudinal shifts across species and suggest an overall**
29 **contraction of cloud forests starting in the late 1970s. We inferred a pervasive**
30 **and interrelated impact of rising temperatures, changing precipitation**
31 **patterns, and deforestation on the distribution of cloud forest species across**
32 **Mesoamerica. Over the last fifty years, cloud forests have declined due to**
33 **deforestation and warmer and more (seasonally) arid climates^{9–11}. This is**
34 **pushing species' to contract their altitudinal ranges and may lead to an**
35 **increasing probability of abrupt and devastating declines of population sizes,**
36 **local adaptation, and migration.**

37 Anthropogenic climate change is altering species ranges and abundances^{1,2}
38 and in some cases, is creating phenological mismatches and population declines
39 across ecosystems^{12–16}. Indeed, recent evidence suggests that many tropical
40 communities have experienced significant changes in species composition and
41 geographic range shifts in response to climate change over the last fifty years^{3,4,17–}
42 ²³. Global mean surface temperatures have increased by approximately 0.85°C
43 over the past 130 years. Two distinct periods of warming have occurred during
44 these years, 1910–1945 and from 1970 onwards²⁴, with the later being
45 characterized by a steep increase in mean temperatures¹⁰. Changes in
46 precipitation are much less consistent but rainfall has generally decreased in the
47 tropics²⁴, which combined with increasing temperatures and vapor pressure deficit
48 may lead to increasing water stress for many organisms.

49 In theory, species can respond to these rapid changes in climate through *in*
50 *situ* acclimation²⁵, adaptation of populations to changing conditions^{13,25}, or
51 migration of populations through habitat tracking (that is, range
52 displacement)^{2,13,16,25}. In the tropics, species migrations will be primarily towards
53 higher elevations due to a shallow latitudinal temperature gradient^{25–27}. Recent
54 upslope shifts of species have been recorded across tropical mountains^{3,4}, yet
55 evidence also suggests that some species have not shifted their ranges or have
56 undergone downslope shifts²³. These idiosyncratic responses of species to climate
57 change can be partially attributed to spatial (and temporal) variability of climate
58 change and its relationship with the biological and ecological attributes of different
59 species²⁸. For example, the range of some species may be determined primarily by
60 temperatures, whereas the range of others may be determined more by water
61 availability, soil characteristics, or competition. Therefore, although downslope
62 shifts run counter to the general expectations of global warming^{16,25}, they may be
63 explained by local/regional alterations to rainfall patterns²⁹, biological interactions
64 (e.g., predation, competition)³⁰, life history traits²⁵, or to other anthropogenic
65 pressures such as land-use change^{31,32}.

66 Altitudinal shifts have already led to changes in species composition in
67 many tropical forests, including the emergence of novel communities and the
68 increase in the relative abundance of lowland, warm-adapted species in mid to
69 high altitude montane forest communities^{17,33}. If species fail to adequately respond
70 to climate change, or their response is hindered by land-use changes³⁰ or changing
71 community structure, local extinction of populations will likely increase^{1,13,25}.
72 Neotropical montane cloud forests are one of the most biodiverse ecosystems in
73 the world and are highly vulnerable to climate alterations^{5–8}. In Mesoamerica alone,
74 these forests harbour more than 6,500 plant species³⁴ in less than 1% of the
75 region's terrestrial surface. Today, less than 20% of the region's cloud forests are
76 under protection, but this has not halted human disturbance on the few remaining
77 and protected cloud forests³⁵. Historical climate data shows a rapidly changing
78 climate across the geographic range of cloud forests¹¹, which agrees with the

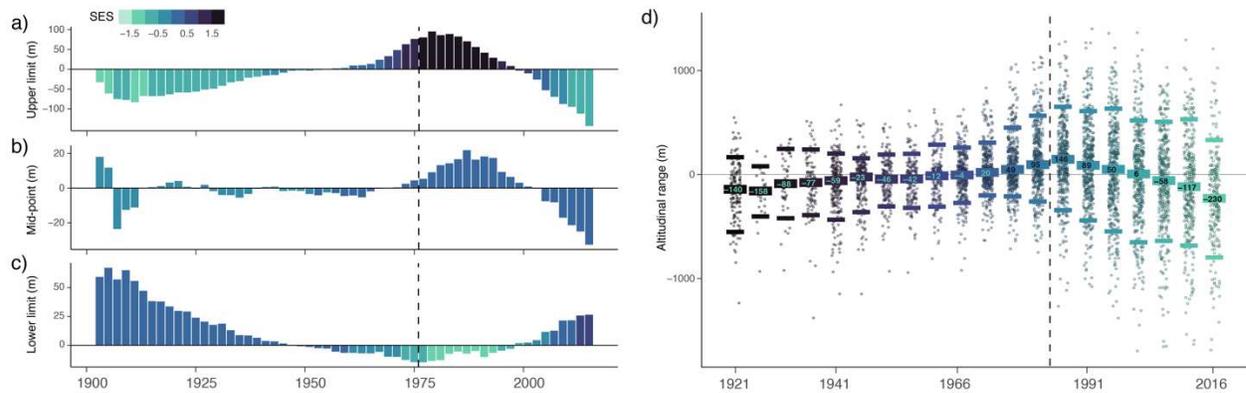
79 global trend of accelerated change since the 1970s¹⁰. In addition, historic land-use
80 change suggest that human encroachment has led to more than half of the original
81 forest cover (that is, pre-industrial) being lost to deforestation^{7,35}. Disentangling the
82 relative impacts of climate change and land-use changes on the range dynamics of
83 cloud forest species remains a major challenge and a top conservation priority.

84 Climate change may represent a great threat to cloud forests⁸ and, in
85 synergy with unrelenting rates of deforestation, could exacerbate their critical state
86 ^{7,35,36}. A first step to address cloud forests' vulnerability to anthropogenic
87 disturbances is to assess changes in their distribution, and that of their constituent
88 species, over the recent past. However, the extent to which climate change and
89 other anthropogenic activities have impacted Mesoamerican cloud forest
90 communities over the last century remains poorly explored. Here we address this
91 critical knowledge gap by analyzing data from biological collections with temporal
92 and spatial metadata containing invaluable information on the historical range
93 dynamics of species¹⁹. We combine these data with historical yearly data for
94 climate and land-use, and information on species' growth forms to assess whether
95 cloud forest species have undergone altitudinal shifts over the last century.

96

97 **Results and Discussion**

98 We assessed the changes in the altitudinal ranges of 419 cloud forest plant
99 species over the period 1901–2016 by using historical occurrence data. Our
100 analyses revealed a general trend of altitudinal shifts and range contractions in
101 cloud forest plant species (**figure 1**). The altitudinal shifts inferred through
102 simulation show an inflection point in range dynamics among cloud forest species
103 over the last fifty years, starting around the 1970–1980s. The observed shifts
104 coincide, albeit with a clear time lag, with the onset of an the documented
105 acceleration in the pace of climate and land-use occurring since the 1970s^{10,37}. Our
106 results suggest opposite shifts in the upper and lower species' range limits (**figure**
107 **1a,c**), which support a general contraction of species' elevation ranges (**figure 1d**).



108

109 **Figure 1. Historical altitudinal ranges for vascular plant species across Mesoamerican cloud**

110 **forests over the period 1901–2016.** a–c, mean deviations from baseline estimates (that is,

111 species' averages during 1901–1975) for the (a) upper limit, (b) mid-point, and (c) lower limit of

112 species' ranges. Values are based on the predicted values of the fitted smoothing splines. d,

113 observed altitudinal ranges (deviations from baseline estimates) over the period 1921–2016 using

114 the time-bins employed for the Piecewise Growth Model (PGM). Crossbars indicate the first and

115 third quantiles (upper and lower crossbars, respectively), and the median (middle crossbar);

116 numbers in the middle crossbar depict the values for the median for each period. Colour scale in a–

117 c represents the mean Standardized Effect sizes (SES) across species estimated through

118 simulations. Vertical dashed lines in a–d indicate the pre-definend breakpoint (that is, 1975–1976).

119

120 However, we observed some differences among species in the direction of

121 altitudinal range shifts (upslope or downslope). At least some of these

122 discrepancies appear to be associated with differences in the elevations where the

123 species occur (that is, their mean altitude prior to 1976) and differences in species'

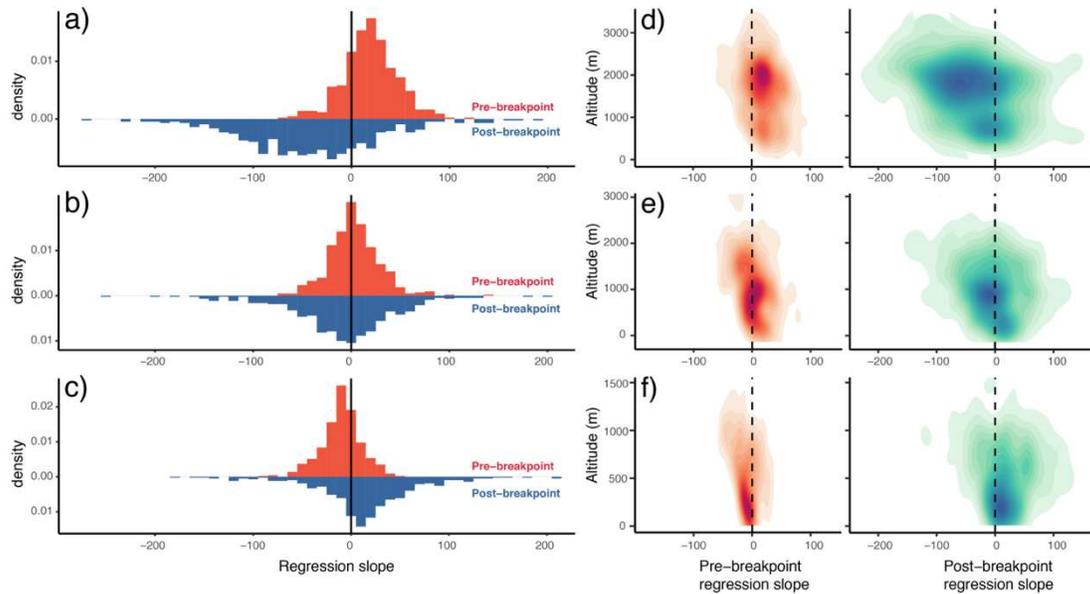
124 growth forms (**supplementary figure S1–S3**). Even when considering species with

125 no positive evidence of elevational shifts, we observe that the average altitudinal

126 trends agree with a general response of cloud forests to anthropogenic influences.

127 These trends could translate into a general historical tendency of elevational range

128 contractions of cloud forests and a decreased capacity for species' migrations³⁵.



129

130 **Figure 2. Historical altitudinal trends for vascular plant species across Mesoamerican cloud**
 131 **forests over the period 1901–2016.** a–c, histograms showing the distribution of the per-species
 132 slopes for the (a) upper range limit, (b) range mid-point, and (c) lower range limit, estimated using a
 133 Piecewise Growth Model with two time periods: 1901–1975 (pre-breakpoint) and 1976–2016 (post-
 134 breakpoint). d–f, two-dimensional density plots showing the distribution the pre- and post-breakpoint
 135 slopes as a function of species’ mean elevations. Darker colours indicate a higher density of slopes
 136 falling within the corresponding regions. Model coefficients for the Piecewise Growth Models are
 137 shown in Supplementary tables S1–S3.

138

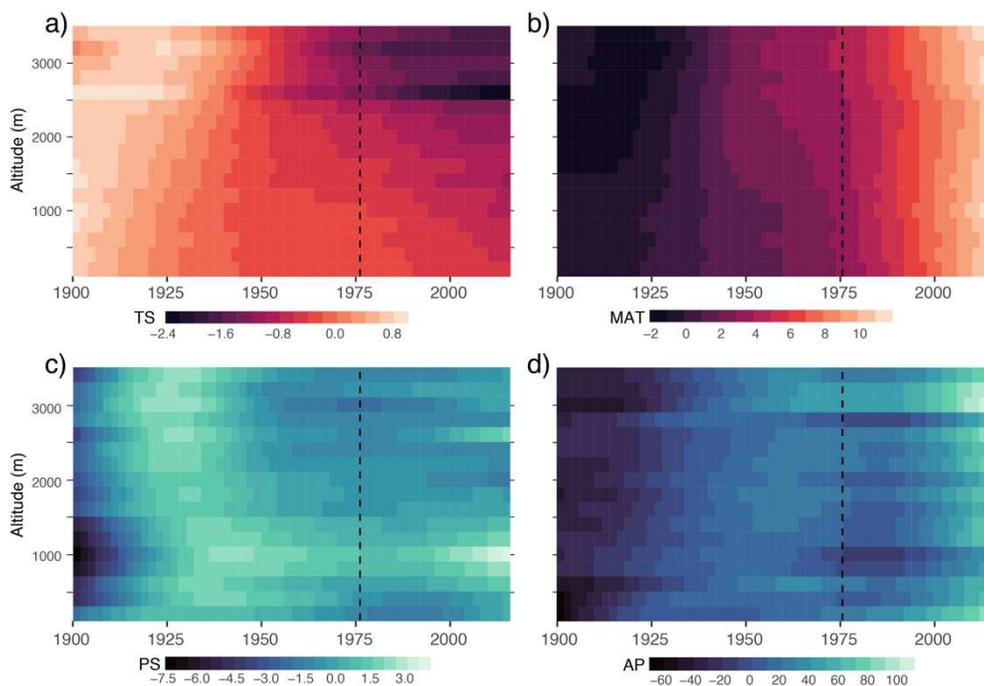
139 Species’ responses have been temporally and spatially heterogeneous
 140 across the region, highlighting specific periods with marked shifts across altitudinal
 141 belts^{38,39}. To test for time-varying changes in the altitudinal ranges of cloud forest
 142 species, we fitted a Piecewise Growth Model to estimate species’ altitudinal trends
 143 before and after a pre-defined year. Based on previous evidence of historical
 144 climate and land-use changes^{10,37}, we defined this breakpoint in the mid-1970s
 145 (that is, 1975–1976) and found clear differences between pre- and post-breakpoint
 146 trends for the lower and upper range limits (**figure 2**). For both the range limits,
 147 pre- and post-breakpoint trends show opposite signs and non-overlapping
 148 confidence intervals (**figure 2a,c; supplementary table S1–S3**). In contrast, the
 149 altitudinal mid-points of species ranges show less-marked differences between the

150 time periods (**figure 2b**), and in this case a slope of zero (that is, no altitudinal
151 shift) can explain the observed pre-breakpoint trends (**supplementary table S2**).
152 Importantly, there is no indication of an effect of elevation on the observed
153 differences between pre- and post-breakpoint trends, suggesting that most species
154 exhibit the same patterns irrespective of differences in their elevational
155 preferences.

156 The results of the Piecewise Growth Model and the simulations suggest that
157 the upper range limit and, to a lesser extent, the mid-point of species' ranges
158 underwent significant shifts after the 1970s (**figures 1,2**), with post-breakpoint
159 downslope trends (**supplementary table S1,S2**) that are reflected in significantly
160 lower altitudes with respect to baseline means. Importantly, this shift to a
161 downslope trend was accompanied by a parallel upslope shift in species' lower
162 range limits (**figures 1c,2c; supplementary table S3**). These patterns are
163 consistent with climate change during the last century (**figure 3**), which is
164 ubiquitous across Mesoamerica and Mesoamerican cloud forests but temporally
165 and spatially heterogeneous¹¹. In particular, we observed substantial deviations
166 from baseline temperature and precipitation (that is, the mean climatologies during
167 1901–1975) that have accelerated since the 1970s (**figure 3**). This is consistent
168 with previous reports of the temporal and geographic trends of changing
169 temperature and precipitation across Mesoamerica¹¹, showing a dramatic increase
170 of mean annual temperature over the last fifty years^{9–11}.

171 In general, the historical climate data (**figure 3a,b**) support previous
172 evidence that climate change is causing montane regions of Mesoamerica to
173 become warmer and more (seasonally) arid³³. These spatially heterogeneous
174 changes in climate may increase the vulnerability of cloud forests since species will
175 not be able to track the shifts in suitable climates through just upslope
176 migrations^{13,25,30}. In this context, the uneven changes in temperature and
177 precipitation we see across altitudinal belts can partly explain the observed shifts
178 and contraction in species' ranges. Changes to temperature and precipitation

179 seasonality have been most pronounced at both the lower (< 1,000 m) and upper
 180 elevations (> 2,500 m) of cloud forest distributions (**figure 3**). As with temperature,
 181 precipitation regimes have changed over the same time period (**figure 3c,d**), with
 182 increasingly wetter conditions through time (**figure 3d**); however, this increase in
 183 precipitation appears to be accompanied by stronger precipitation seasonality
 184 (**figure 3c**). The increase in precipitation seasonality may be of particular concern
 185 for cloud forest plant species, because many are highly sensitive to seasonal
 186 fluctuations in humidity levels and water stress^{35,36,40}.



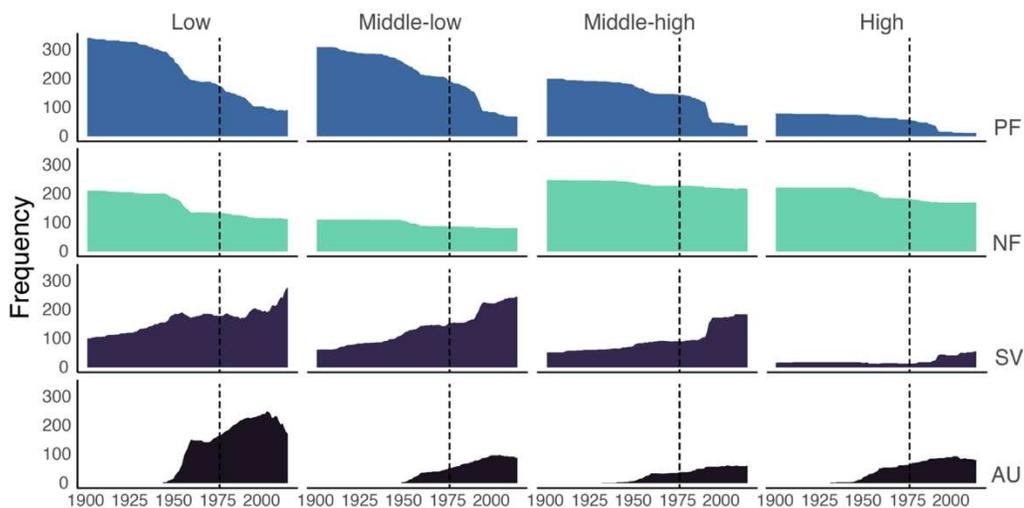
187

188 **Figure 3. Historical changes in temperature and precipitation across Mesoamerican cloud**
 189 **forests localities over the period 1901–2015.** a–d, bi-yearly deviations from baseline averages
 190 (1901–1975) in (a) temperature seasonality, (b) mean annual temperature, (c) precipitation
 191 seasonality, and (d) annual precipitation; darker colours indicate below-average conditions and
 192 brighter colours above-average conditions. Climate data is based on the CHELSAcruts high
 193 resolution temperature and precipitation time series for the 20th century⁶¹. Vertical dashed lines in
 194 a–d indicate the pre-defined breakpoint (that is, 1975–1976).

195

196 A high proportion of primary cloud forest localities have been historically
 197 subjected to human disturbance, which is indicative of land-use changes affecting

198 cloud forests throughout the region³¹. Assuming that most of the focal species are
 199 preferentially found within primary forests, or in areas with moderate to low levels
 200 of disturbance^{35,41}, the heterogeneity in land-use changes across altitudinal belts
 201 could have had a profound impact on their distributions, and hence the distribution
 202 of cloud forests, through time. We found substantial changes in the extent of
 203 primary forest cover over the period 1901–2015 (**figure 4**). Based on the historical
 204 land-use data we estimated that by 1936, ~44% (~38–59% across altitudinal belts)
 205 of the cloud forest localities in our dataset corresponded to primary forest, but by
 206 2015 only ~11% (~8–14% across altitudinal belts) of these localities remained as
 207 primary forest. However, the loss of primary forest cover has not been constant
 208 over time and a dramatic decline of 35.2% of the original forest cover (that is,
 209 classified as primary forest by 1936) is observed over the period 1976–1996.



210
 211 **Figure 4. Historical trends in land-use across Mesoamerican cloud forests localities over the**
 212 **period 1901–2015.** Area plots showing the proportional area occupied by the four land-use land-
 213 cover (LULC) categories: Agriculture-urban (AU), Primary forests (PF), Non-forest vegetation (NF),
 214 and Secondary vegetation (VG). The proportion of each land-use category was estimated across
 215 four altitudinal belts: Low (0-300 m), Middle-low (300–1,000 m), Middle-high (1,000–2,000), and
 216 High (>2,000 m). The vertical dashes lines marks the year 1976. Vertical dashed lines indicate the
 217 pre-defined breakpoint (that is, 1975–1976).

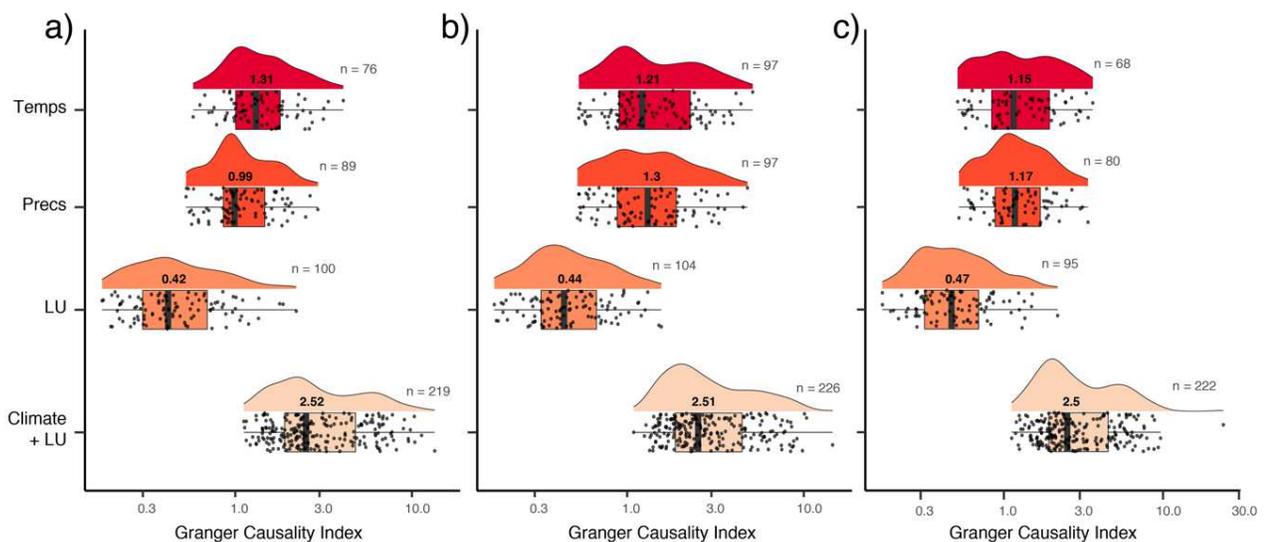
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219 Despite some uncertainty in the land-use data³⁷, our results support the
220 well-documented trend of increasing agricultural, urban, and secondary vegetation
221 land-use³¹ at the expense of primary forests and non-forest vegetation across
222 altitudinal belts (**figure 4**). One possible scenario is that the lower and upper range
223 limits of species are shifting not only because of warming and aridification, but also
224 as a result of high extirpation rates (that is, local extinction) of populations. There is
225 no indication that proportional land-use change has been more pronounced either
226 at the lowlands or highlands (**supplementary figure S4**). However, primary forests
227 have been historically more abundant at mid-elevations, so proportionally similar
228 losses across altitudinal belts would lead to more pronounced declines in absolute
229 forest cover at both ends of the altitudinal range of cloud forests.

230 We acknowledge that some of the observed patterns may be attributable to
231 biases inherent to data from biological collections, particularly the concentration of
232 collection effort in particular areas and through time⁴². As such, some of the
233 observed trends may be reflecting a ‘false’ signal and a shifting sampling bias
234 rather than species responses to climate change or land-use changes. However,
235 simulations designed to account for collector bias indicate that most of the
236 observed trends are at least in part due to actual shifts in the species’ altitudinal
237 ranges. That said, there is still the possibility that the observed patterns are being
238 influenced, by geographically heterogeneous rates of deforestation constraining
239 the areas where species can be collected and lead to apparent shifts of species’
240 ranges.

241 To test for the possible causal influence of both climate and land-use
242 changes on the observed species range shifts, we used a definition of causality
243 that is based on predictability^{43,44}. For each species, we calculated multivariate
244 Granger Causality Indices (GCI) and found that overall historical climate and land-
245 use changes –specifically the loss of primary forests and non-forested vegetation–
246 improves the predictions of species’ altitudinal shifts through time (**figure 5**). This
247 means that a significant amount of variance in species altitudinal trends can be

248 explained either by climate change, land-use transformation, or both (**figure 5**). A
 249 substantial proportion of species showed significant GCI, supporting a strong
 250 causal influence (predictability) of climate change, land-use transformation, and the
 251 interaction between the two, on the observed species' altitudinal range shifts
 252 (**figure 5**). These causal influence on species altitudinal ranges appears to be
 253 ubiquitous across species with different growth forms (**supplementary figure S5–**
 254 **S7**). Although we observed slightly higher GCI estimates for woody species (that
 255 is, trees and shrubs) than for herbaceous species, these results support an
 256 overarching influence of climate change and land-use transformation across a
 257 broad spectrum of cloud forest species.



259 **Figure 5. Granger causal influence of climate and land-use change time series on species**
 260 **historical altitudinal trends.** a–c, raincloud plots showing the distribution of the per-species
 261 Granger Causality Index (GCI) of the different climate change land-use time series for (a) upper
 262 range limit, (b) range mid-point, and (c) lower range limit. Numbers within the density plots indicate
 263 the median value across species and numbers to the left indicate the number species with a
 264 significant GCI (p-value < 0.05).

265

266 The inferred synchronous response of species might imply that cloud forests
 267 likely respond as cohesive communities, perhaps because the high climatic
 268 sensitivity of most species. However, we observe that species' range contractions
 269 have increasingly become more variable over the last fifty years, compared to the

270 inferred responses prior to the 1970s (**figure 1d**). Thus, the impacts of an
271 accelerated environmental degradation of cloud forests could have devastating
272 effects in the near future, affecting community composition, local forest structure,
273 levels of biodiversity, and hydrological cycles^{26,40,45,46}. Improving our knowledge of
274 past range shifts on tropical montane ecosystems is imperative if we hope to
275 prevent these valuable ecosystems and species from vanishing in the near future.
276 Our study provides strong evidence for the dire status of cloud forests across the
277 region^{6,7,35,36}. Species are being forced into contracting mid-elevational refugia of
278 remaining forest cover, with an overall reduction in their elevational ranges over
279 time. This likely translates into a geographic/altitudinal decline in cloud forest
280 species abundances and may be due to the incapacity for acclimation and *in situ*
281 adaptation, limited dispersal potential, or alteration to forest structure due to
282 canopy thinning and border effects⁴⁷. In light of the temporal dynamics we
283 uncovered, we suggest a breach of a ‘climate horizon’⁴⁶ – enhanced by land-use
284 changes – beyond which there would be a dramatic increase in the probability of
285 abrupt and devastating disruptions of cloud forest communities.

286

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418 **Methods**

419 *Species occurrence data.* We obtained a list of vascular plants for cloud forests
420 from ref.⁴⁸. This is the most comprehensive and well-curated species list and
421 comprises 6,778 plant species that have been recorded in Mesoamerican cloud
422 forests. Prior to processing, we harmonized family and species names using Kew's
423 Plants of the World database⁴⁹ and obtained a list of 6,495 species belonging to
424 221 families of angiosperms, gymnosperms, ferns, and lycophytes.

425 We used historical distribution data available from the Global Biodiversity
426 Information Facility (GBIF), querying all records under 'Tracheophyta' and
427 considering only 'Preserved Specimens' in our search; this database consisted of
428 68,570,538 occurrence records⁵⁰. The occurrence records were taxonomically
429 harmonized and cleaned following the procedures described in refs.^{51–53}. Briefly,

430 we first discarded records without species identification and with missing or badly
431 formatted GPS coordinates. Then, we harmonized family and species names using
432 Kew's Plants of the World database⁴⁹, which allowed us to match the occurrence
433 database with the list of cloud forest species obtained above.

434 We used the *CoordinateCleaner* package⁵⁴ in R⁵⁵ to flag suspect
435 occurrence records meeting one (or more) of the following criteria: (1) equal
436 latitude and longitude, (2) zero latitude and longitude, (3) coordinates falling within
437 a five-kilometre radius of a country's political centroid, (4) coordinates falling within
438 a ten-kilometre radius of a country's capital, (5) coordinates falling within a two-
439 kilometre radius of biodiversity institutions, and (6) coordinates falling in the open
440 ocean using a reference landmass buffered by one degree from the coastline to
441 avoid eliminating species living on the coast. In addition, for each species we
442 applied the geographic outlier test to flag suspect records. This test estimates
443 pairwise geographic distances among all records of a species and then flags those
444 records with a minimum distance to all other records greater than a defined
445 threshold, in this case 300 km; flagged records represent single occurrences
446 isolated from the rest of the known distribution of the species. Lastly, we flagged
447 additional suspect records using the geographic information available in Kew's
448 Plants of the World database⁴⁹ that scores the 'native' or 'non-native' status of
449 species across the world using the World Geographic Scheme for Recording Plant
450 Distributions⁵⁶.

451 Since we aimed to address the historical dynamics of cloud forests in
452 Mesoamerica, we discarded species with occurrence records located outside the
453 American continent (longitudinal limits, 30°–135°W). Furthermore, the floristic list
454 includes species that are widespread and not restricted to cloud forests (or other
455 montane forest communities). To further identify these species, we used the data
456 on terrestrial eco-regions of the world to assign every occurrence record to one of
457 14 recognized biomes⁵⁷. For each species, we estimated the proportion of records
458 within each biome and discarded species with more than 2.5% of their occurrences

459 within any of the following biomes: tropical dry forests, temperate grasslands,
460 flooded grasslands, tundra, Mediterranean forests, deserts and xeric scrubland,
461 and mangroves. We considered the frequent presence of species within these
462 biomes as indicative of taxonomic problems in the original species list (that is,
463 erroneous species determinations) or indicative of occurrence records in cloud
464 forests being the result of invasive, non-native, or generalist species.

465 In sum, we obtained a database comprising 74,701 occurrence records for
466 899 cloud forest plant species (14% of the total number of species in ref.⁴⁸), which
467 belong to 148 families of angiosperms (766 species), ferns (114), lycophytes (10),
468 and gymnosperms (9). We used these occurrence records to perform all
469 subsequent analyses. For the main analyses, we limited the study area to
470 comprise grid-cells in Continental America (that is, excluding the Antilles) within
471 12–26°N, limiting our dataset to 1,895 grid-cells, and we excluded records with
472 dates prior to 1901. We selected species with more than 50 records (419 species,
473 62,973 records). Data processing was handled in R⁵⁵.

474 *Plant trait data.* We compiled data on growth form for the target species from a
475 dataset obtained from the TRY Plant Trait Database^{58,59}. For consistency, prior to
476 analysis we harmonized species names using Kew's Plants of the World
477 database⁴⁹. Given the heterogeneous categorization of growth forms in the 'raw'
478 TRY dataset, we homogenized and categorized growth-form data into four distinct
479 categories: (i) woody and non-woody climbers; (ii) epiphytes; (iii) herbs; (iv) shrubs;
480 and (v) trees. Several species in the TRY dataset have two or more records and
481 are categorized into two or more growth-form categories. Thus, we decided to
482 score species into a single category following a majority rule consensus. Although
483 the distinction among categories is not clear-cut, our scheme is aimed at
484 distinguishing among broad growth-form categories and habitats among cloud
485 forest species. Data processing was handled in R⁵⁵.

486 *Land-use series.* We obtained land use and vegetation data for the region from
487 ref.⁶⁰, who selected and transformed the data from the Land Use Harmonization

488 Project (LUH2)³⁷. The LUH2 is a global terrestrial dataset at 0.25° spatial resolution
489 that provides land-use and land-cover (LULC) data from 850 to 2300^{37,60}. We used
490 the final dataset provided by ref.⁶⁰ and retrieved LULC data for the years 1901–
491 2016. For each of the 1,895 grid-cells (resolution of 0.083°, ~9 km²) in our
492 occurrence database we extracted yearly data on land-use. The original LULC data
493 includes the following twelve categories^{37,60}: forested primary land, non-forested
494 primary land, potentially forested secondary land, potentially non-forested
495 secondary land, managed pasture, rangeland, urban land, C3 annual crops, C3
496 perennial crops, C4 annual crops, C4 perennial crops, and C3 nitrogen-fixing
497 crops. For our analyses, we summarized these categories into four broad
498 categories: (1) AU, urban and agricultural; (2) NF, non-forested primary vegetation;
499 (3) PF, primary forested vegetation; and (4) SV, secondary vegetation.

500 We visualized temporal changes across land-use categories using chord
501 diagrams, which represent ‘flows’ between multiple entities. In this case, the
502 entities represent land-use categories at two distinct time points (for example, PF-
503 1936 and PF-1956) and the flows represent the proportion of grid cells transitioning
504 between categories from past to present (for example, PF-1936 to AU-1956). The
505 size of each category is represented by a fragment of the outer circle, where all the
506 fragments for a given time period sum up to 1. This way we can estimate and
507 visualize the proportional changes through time for any given land-use category
508 and the proportional size of different transitions across categories. We estimated
509 land-use transitions between 1936 and 2015; we chose 1936 as the starting point
510 because the AU category appears first in our grid-cells during the 1930s. Data
511 processing was handled in R⁵⁵.

512 *Historical climate change series.* We used the climate data from the CHELSAcruts
513 high resolution temperature and precipitation time series for the 20th century⁶¹.
514 These time series include monthly data for maximum temperature, minimum
515 temperature, and net precipitation at a 30 arc-sec spatial resolution. We extracted
516 climate data for every occurrence record and then aggregated the monthly data

517 into four yearly variables: annual precipitation (Ann_Prec), precipitation seasonality
518 (CV_Prec), mean annual temperature (Mean_temp), and temperature seasonality
519 (CV_temp). Variable processing was handled in R⁵⁵. We extracted climate and
520 elevation data for every occurrence record, which we then aggregated for the
521 1,895 grid-cells comprising our study area.

522 We used the aggregated climate data to visualize climate change series
523 within cloud forest localities, extracting yearly climate data for the period 1901–
524 2016 for each locality. For each climatic variable, we estimated a rolling mean
525 using a moving window of ten time points centred on the corresponding year (that
526 is, the time window for 1955 is 1951–1960), with partial estimates allowed (that is,
527 the time window for 1901 is 1901–1906). We then fitted cubic smoothing splines of
528 climate as a function of time for each locality, using a smoothing parameter that
529 was automatically set between 0.45 and 0.7 to avoid over-fitting; smooth model fit
530 was performed as implemented in the ‘*npreg*’ package⁶² in R⁵⁵. We used the
531 resulting smoothing spline models to estimate average ‘predicted’ climate values
532 every two years. We set 1975–1976 as a breaking point in historical climate trends
533 based on empirical evidence¹⁰ and estimated climate averages before this point
534 (baseline means) across 200-meter altitudinal belts. We then calculated deviations
535 from these baseline estimates across the full climate series for every altitudinal
536 belt, separately. Data processing was handled in R⁵⁵.

537 *Historical altitudinal series.* We used the geographic occurrences for the selected
538 species to evaluate altitudinal series for each species for the period 1901–2016.
539 For this, we used the GTOPO30 global digital elevation model (US Geological
540 Service’ Earth Resources Observation and Science Center) to assign altitudinal
541 data to every occurrence record; this global elevation model has a grid resolution
542 of 30 arc-seconds (~1 km²).

543 To characterize shifts in species’ altitudinal ranges, we used information on
544 the year of collection of every record to estimate a rolling median of altitude (mid-
545 point) for each species as described above for climate series. We only considered

546 rolling estimates calculated from a minimum of ten occurrence records, which
547 ensures that every rolling estimates is based on an average of one record per year.
548 As described above for climate series, we fitted cubic smoothing splines of mid-
549 point altitude as a function of time for each species; smoothing splines were only
550 fitted for species with at least ten data points (rolling estimates). For each species,
551 we estimated the average ‘predicted’ altitude every two years since the species’
552 first historical occurrence.

553 We also assessed shifts in the lower and upper limits of species’ altitudinal
554 ranges as described above. To minimize the possible influence of outlying records,
555 we used the 0.025 and 0.975 quantiles of the altitudinal range to approximate the
556 lower and upper limits, respectively. Finally, we estimated the breadth of the
557 altitudinal range of species by taking the difference between the rolling estimates
558 for the lower and upper range limits. This allowed us to visualize trends of range
559 contractions/expansions through time.

560 *General altitudinal shifts.* We tested for a general altitudinal shift of cloud forests by
561 fitting a Piecewise Growth Model (PGM) using the ‘lme4’ package⁶³ in R⁵⁵. We
562 fitted PGM to the lower, mid-point, and upper altitude of species separately, using
563 the Nelder-Mead optimizer. To fit the PGM we summarised the altitudinal data into
564 twenty periods (5-year bins); due to the reduced data prior to 1921, we defined the
565 first time period as a 20-year bin (that is, 1901–1921). Briefly, the model estimates
566 the average slope and intercept of the altitudinal series while modelling non-linear
567 change through time. The general fitted PGM takes the form (‘lme4’ syntax),

$$568 \quad y_i \sim \text{intercept} + \text{slope}_0 + \text{slope}_1 + (\text{intercept} \mid \text{Species}) +$$
$$569 \quad (0 + \text{slope}_0 \mid \text{Species}) + (0 + \text{slope}_1 \mid \text{Species}),$$

570 where y_i is the mean altitude across species and slope_0 and slope_1 represent the
571 slope coefficients before and after a pre-defined breakpoint. This model
572 incorporates random effects for the intercept and the two slopes (that is, species
573 effects) and uncorrelated random effects for slope_0 and slope_1 . We used 1975–

574 1976 as a breaking point to segment the time series using ‘dummy’ variables. The
575 PGM was assessed against less complex nested models, that is growth models
576 with a single time trend (that is, a single slope coefficient) and fewer random
577 effects (that is, only random intercepts). We fitted PGM to the lower, mid-point, and
578 upper altitudinal limits of species, separately.

579 *Significance of species’ altitudinal series.* In order to account for the possibility of
580 collection bias impacting the observed species’ altitudinal shifts, we compared the
581 observed series against null models based on time-varying collection efforts.
582 Basically, we implemented a re-sampling approach of the complete occurrence
583 database by generating ‘simulated species’ that match the per-year sampling
584 intensity and the altitudinal range of species. To achieve this, for a given species,
585 we first estimated the recorded altitudinal range throughout the entire period (that
586 is, 1901–2016) and counted the number of occurrence records per year. We used
587 this information to identify occurrence records (across all species in the database)
588 within the altitudinal range of the focal species, and then sampled this subset of
589 occurrences using the per-year sampling intensity of the focal species. In principle,
590 this procedure generates a sample of records matching the historical altitudinal
591 range and temporal sampling effort for a given species.

592 For every ‘simulated species’, we produced rolling estimates, fitted
593 smoothing splines and estimated average ‘predicted’ values as described above.
594 We constructed 500 replicates of the null models (that is, simulated species) for
595 every species, and compared the empirical values against those obtained through
596 simulation. We estimated the mean and standard deviation across the 500
597 replicates and estimated per-species Standardized Effect Sizes (SES) at time i as:

$$598 \quad SES_{Ti} = OBS - MEAN / SD,$$

599 Where OBS , $MEAN$, and SD correspond to the empirical value, the mean and
600 standard deviation of the null model, respectively. SES_{Ti} indicates whether
601 observed altitudes are higher or lower than expected under the null model; their

602 significance is approximated using 95% confidence intervals around zero (± 1.96).
603 We constructed the null models in R⁵⁵.

604 We used the Standardized Effect Sizes (SES_{Ti}) for the average ‘predicted’
605 altitude over the period 1901–2016 to visualize the overall trend of altitudinal shifts
606 among cloud forest species. For this, we estimated per-year averages of observed
607 altitudes across species, but only for estimates with associated significant SES_{Ti} .
608 We further visualized the overall trends by grouping species into altitudinal and
609 growth-form categories; for altitude, we classified species into four altitudinal belts
610 based on their mean elevation: (i) low (148), (ii) mid-low (164), (iii) mid-high (166),
611 and (iv) high (42).

612 *Influence of climate change and land-use on altitudinal series.* Using the species’
613 altitudinal series, we assessed whether species responses are causally influenced
614 by climate and land-use changes. For this, we employed the Granger Causality
615 Index (GCI) that uses a concept of causality based on prediction^{43,44}. More
616 specifically, the Granger Causality test is a statistical hypothesis to test whether
617 one (or more) time series is useful for predicting another series. For each species,
618 we summarised climate and land-use series across the grid cells where the target
619 species has been recorded. For the climate series we estimated series for each of
620 the four yearly variables and for the land-use series we estimated the temporal
621 change in the frequency of grid-cells with forested (PF) and non-forested (NF)
622 primary vegetation. Prior to estimating the GCI, we pre-whitened all series
623 according to an ARIMA model fitted to the altitudinal series (that is, the ARIMA
624 model was fitted by selecting the best model according to the AICc value) as
625 implemented in the ‘forecast’ package^{64,65} in R⁵⁵. We then estimated GCI for
626 multivariate series using the FIAR package⁶⁶ in R⁵⁵. More specifically, we tested
627 GCI of four multivariate series: (1) temperature (that is, mean annual temperature
628 and temperature seasonality); (2) precipitation (that is, annual precipitation and
629 precipitation seasonality); (3) land-use; and (4) temperature, precipitation, and
630 land-use combined.

632 **Methods references**

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685 **Author contributions**

686 S.R.-B. conceived the research with contributions from A.C.-R., K.F., J.F.O., and
687 H.R.-C.; S.R.B. and A.C.-R compiled and curated the occurrence, climate, and
688 land-use data; S.R.-B. designed and performed the analysis with contributions by
689 K.F.; S.R.-B. wrote the manuscript with contributions with contributions from A.C.-
690 R, K.F, A.E.O.-R., A.A.V.-A., J.F.O., and H.R.-C.

691 **Competing interests**

692 The authors declare no competing interests.

693 **Data and code availability**

694 The R code and data supporting the results in the main text are available at
695 <https://github.com/spiritu-santi/CF-mig> and at Zenodo with the identifier
696 <https://doi.org/10.5281/zenodo.5587219>. The geographic occurrence data is
697 available through the Global Biodiversity Information Facility with the identifier
698 <https://doi.org/10.15468/dl.bdxzkw>.

699 **Additional information**

700 Supplementary Information is available for this paper.

701 Correspondence and requests for materials should be addressed to Santiago
702 Ramírez-Barahona.

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