

Melting phase relations in Fe-Si-H at high pressure: Implications for Earth's inner core crystallization and core light elements

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2 **Implications for Earth's inner core crystallization and core light**
3 **elements**

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16 **ABSTRACT**

17 Hydrogen could be an important light element in planetary cores, but its effect on phase
18 diagrams of iron alloys is not well known because the solubility of H in Fe is minimal at
19 ambient pressure and high-pressure experiments on H-bearing systems have been
20 challenging. Considering that silicon can be another major light element in planetary
21 cores, here we performed melting experiments on the Fe-Si-H system at ~50 GPa and
22 obtained the ternary liquidus phase relations and the solid/liquid partition coefficient, D
23 of Si and H based on *in-situ* high-pressure X-ray diffraction measurements and *ex-situ*
24 chemical and textural characterizations on recovered samples. Liquid crystallized
25 hexagonal close-packed (hcp) $(\text{Fe}_{0.93}\text{Si}_{0.07})\text{H}_{0.25}$, which explains the observed density
26 and velocities of the Earth's solid inner core. The relatively high $D_{\text{Si}} = 0.94(4)$ and $D_{\text{H}} =$
27 $0.70(12)$ suggest that in addition to Si and H, the liquid outer core includes other light
28 elements such as O, which is least partitioned into solid Fe and can thus explain the
29 density difference between the outer and inner core. H and O, as well as Si, are likely to
30 be major core light elements, supporting the sequestration of a large amount of water in
31 the Earth's core.

32 Introduction

33 Since Birch¹ reported the density deficit and velocity excess of the Earth's outer core
34 with respect to pure iron (Fe), light elements in the core have long been explored but
35 still remain controversial². Recent planet formation theories suggested that a large
36 amount of water could have been delivered to the growing Earth³⁻⁵. The chemical
37 reaction of water with Fe metals in a magma ocean led to the incorporation of hydrogen
38 (H) along with silicon (Si) and oxygen (O) into the core⁶⁻⁹. While O is least partitioned
39 into solid Fe and should therefore be negligible in the inner core¹⁰⁻¹², both Si and H are
40 likely to be present in both the outer and inner core. Indeed, measurements of the
41 density and sound velocity of solid Fe and Fe alloys supported that Si^{13,14} and H^{15,16} are
42 important impurity elements in the inner core. Recent theoretical calculations¹⁷
43 suggested that the Earth's solid inner core is an hcp Fe₆₀Si₄H₁₋₁₅ alloy, depending on the
44 temperature at the inner core boundary (ICB), $T_{\text{ICB}} = 5500 \text{ K to } 6500 \text{ K}$.

45 The liquidus phase diagrams of Fe alloys and the solid metal-liquid metal
46 partitioning of these light elements are of great importance to constrain possible ranges
47 of the outer and inner core compositions². Those of Fe-Si alloys have been repeatedly
48 examined at high pressures up to inner core conditions¹⁸⁻²². In contrast, the phase
49 relations of H-bearing Fe alloys are little known²³⁻²⁵. This is because 1) only a
50 negligible amount of H can be present in liquid/solid Fe at 1 bar and therefore H
51 escapes from Fe lattice during decompression^{6,9,26} and 2) high-pressure experiments on
52 H-bearing systems often involve technical problems such as the failure of a diamond-
53 anvil. Recent neutron diffraction measurements are able to reveal the position and
54 abundance of H atoms in solid Fe alloys under high pressure but so far limited to less
55 than 12 GPa^{27,28}. The properties of hcp Fe-Si-H ternary alloys, including the
56 simultaneous solubilities of Si and H, have not been examined at high pressure and high
57 temperature (P - T)²⁹.

58 In this study, we have investigated the liquidus relations and the solid/liquid
59 partitioning of Si and H in the Fe-Si-H ternary system at $\sim 50 \text{ GPa}$ based on melting
60 experiments in a laser-heated diamond-anvil cell (DAC). The results demonstrated that
61 an hcp Fe-Si-H alloy, similar in composition to those predicted for the inner core
62 solid¹⁷, crystallized from liquid with D_{Si} (solid/liquid) = 0.94(4) and $D_{\text{H}} = 0.70(12)$ on a
63 weight basis. The relatively high D_{Si} and D_{H} require the inclusion of other light
64 elements in the outer core such as O, in order to account for a density contrast between

65 the outer and inner core. We explored the possible liquid core composition and found
66 that H and O, in addition to Si, are likely to be important light elements in the Earth's
67 core.

68 **Methods**

69 High P - T experiments. Laser-heated DAC techniques were used to generate high P -
70 T conditions. The starting materials were ~ 10 μm thick Fe-Si foils containing 4.0 and
71 6.5 wt% Si. The culet size of the diamond anvils used was either 200 μm or 300 μm
72 depending on the hydrogen source, and the sample chamber was prepared by drilling a
73 hole with a diameter of ~ 80 μm or ~ 100 μm , respectively, in a rhenium gasket pre-
74 indented to ~ 25 μm thickness. Hydrogen was introduced to the system by sandwiching
75 the Fe-Si sample between $\text{C}_n\text{H}_{2n+2}$ paraffin layers (run #4) or cryogenically loading
76 liquid hydrogen into the sample chamber³⁰ (runs #1–#3, #5). In the latter case, an NaCl
77 ring prepared with a focused ion beam (FIB) was employed inside the rhenium gasket
78 as an H_2 insulator and a pressure medium. In addition, the culet of the diamond anvils
79 was coated with a thin layer of Ti by sputtering to avoid anvil failure³¹. After cooling
80 the DAC to 20 K with the sample chamber open, the insulating chamber was flooded
81 with liquid H_2 . The sample was then sealed, weakly compressed and brought back to
82 room temperature. Subsequently, it was further compressed to ~ 20 GPa and weakly
83 annealed at < 1000 K to prompt hydrogenation of the alloy. In order to remove excess
84 hydrogen, the DAC was then again cooled to ~ 90 K using liquid N_2 and decompressed
85 back to 1 atm, so that any hydrogen in excess was released from the sample chamber
86 while preventing the decomposition of hydrogenated iron.

87 Heating was conducted at the beamline BL10XU, SPring-8 synchrotron radiation
88 facility using a couple of 100 W single-mode Yb fiber lasers (YLR-100, IPG Photonics)
89 with flat-top beam shaping optics³². The samples were heated to melting temperatures
90 for a limited duration of ~ 3 sec in an attempt to avoid temperature fluctuation. This time
91 scale has been shown to be long enough to reach chemical equilibrium due to the short
92 length scale of DAC samples; in previous DAC melting experiments, liquid and
93 coexisting solid compositions did not change after 1 sec on the Fe-S system³³. We
94 obtained sample temperature profiles using spectroradiometry and estimated the
95 temperature at the liquid-solid boundary (liquidus temperature) in combination with a
96 sample cross section^{12,22} (see below) (Table 1).

97 Sample analyses. *In-situ* angle-dispersive X-ray diffraction (XRD) measurements
 98 were conducted before, during and after heating (melting) using an X-ray beam with an
 99 energy of ~ 30 keV³². The beam was focused on a sample using compound refractive
 100 lenses and collimated so that the full-width at half maximum was 6 μm . The 2D
 101 diffraction patterns were collected on a flat panel detector (Perkin Elmer), and
 102 integrated to 1D patterns using the IPAnalyzer software and analyzed with the
 103 PDIndexer software³⁴. Pressure was estimated on the basis of the unit-cell volume of B2
 104 (CsCl)-type NaCl³⁵ considering its temperature following ref. 36, except for run #4
 105 where it was obtained from the Raman shift of diamond³⁷ and corrected for a
 106 contribution of thermal pressure^{12,38}. The overall uncertainty may be $\pm 5\%$ (Table 1).

107 Textural and chemical characterizations were carried out on all the samples after
 108 decompression and recovery (Fig. 1). A Ga-ion FIB (FEI VersaTM 3D DualBeamTM)
 109 was used to mill the sample parallel to the compression axis and prepare a cross section
 110 of the laser-heated portion. The sample was then examined by a field-emission (FE)-
 111 type scanning electron microscope (SEM) and energy dispersive X-ray spectrometry
 112 (EDS). Furthermore, a field-emission-type electron probe micro-analyzer (FE-EPMA,
 113 JEOL JXA-8530F) was used for quantitative chemical analyses with an accelerating
 114 voltage of 12 kV and a current of 15 nA. We used Fe, Si, Fe-0.84wt%C and Fe₃C as
 115 standards, and LIF (Fe), PETH (Si), LDE2H (C) and LDE1 (O) as analyzing crystals.
 116 Carbon (C) concentration was quantified from a calibration curve obtained by a C-free
 117 copper mesh, Fe-0.84wt%C (JSS066-6, the Japan Iron and Steel Federation) and Fe₃C.

118 Since hydrogen escapes from iron lattice at low pressure (less than ~ 3 GPa) and
 119 room temperature^{6,9,39}, we determined the H content x of (Fe,Si)H _{x} based on the volume
 120 expansion of the Fe-Si crystal lattice at room temperature²³;

$$121 \quad x = \frac{V_{(\text{Fe-Si})\text{H}_x} - V_{\text{Fe-Si}}}{\Delta V_{\text{H}}} \quad (1)$$

122 where V is lattice volume (Fig. 2) and ΔV_{H} is the volume increase caused per H atom⁴⁰.
 123 The reference lattice volume of Fe-Si at identical high pressure was estimated from
 124 those of pure Fe^{41,42} and Fe-6.5wt%Si²⁰. The H abundance in liquid was estimated from
 125 the volume of face-centered cubic (fcc) or hcp crystals formed upon quenching
 126 temperature. The validity of such estimation has been demonstrated in refs. 9 and 43;
 127 thermal annealing of such quench crystals changed their volume to a minor extent,
 128 indicating that their H concentration represents that of liquid. The EPMA analyses and

129 the H contents are summarized in [Table 1](#). The error in H concentration may be $\pm 8\%$,
130 which is derived mainly from uncertainty in $\Delta V_H^{9,28}$.

131 **Results**

132 Liquidus phase relations in Fe-Si-H. We performed five separate experiments in
133 the Fe-Si-H ternary system at ~ 50 GPa, in which C and O concentrations in liquids were
134 less than 0.6 and 0.7 wt%, respectively ([Table 1](#)). The coexisting liquid and solid
135 (liquidus phase) compositions are plotted in [Fig. 3a](#), which constrains the liquidus fields
136 of hcp Fe, fcc FeH and B2 Fe-Si.

137 In the first two runs, liquids coexisted with (Si, H)-bearing hcp Fe. In run #1, the
138 XRD pattern showed hydrogenated hcp Fe+4wt%Si together with the NaCl pressure
139 medium before heating at 56 GPa ([Fig. 2a](#)). During heating to 2100 K, we observed
140 diffuse scattering signal from liquid around two-theta angle of 12° , coexisting with
141 weak reflections from the hcp phase (liquidus phase). Upon quenching to 300 K, the fcc
142 phase appeared. Combined with microprobe analyses of the cross section of the
143 recovered sample, these observations indicated that liquid Fe-4.0wt%Si-0.67wt%H
144 coexisted with hcp Fe-3.7wt%Si-0.47wt%H during heating. Similarly in run #2
145 performed at 57 GPa, the XRD pattern collected at 2150 K included the peaks from the
146 hcp phase and the NaCl pressure medium ([Fig. 2b](#)). The diffuse signal was not clear,
147 which is consistent with the microprobe observation of a small melt pocket at one side
148 of the recovered sample surface. Liquid Fe-5.9wt%Si-0.56wt%H coexisted with hcp Fe-
149 4.3wt%Si-0.30wt%H at high P - T in this experiment. It is, however, uncertain that this
150 solid composition represents that of the liquidus phase, because the small liquid pool
151 suggests that the coexisting liquidus phase was volumetrically even smaller and the
152 determination of its chemical composition was difficult.

153 On the other hand, liquids coexisted with fcc (nearly) stoichiometric FeH in other
154 runs. In run #3, hcp (Fe,Si)H_{0.81} was found before heating at 61 GPa ([Fig. 2c](#)). Upon
155 heating to 2350 K, the peaks from the hcp phase were lost, and alternatively those of fcc
156 appeared together with the diffuse scattering signal from liquid. The XRD pattern
157 obtained after quenching temperature and the EPMA analyses on the recovered sample
158 showed that liquid Fe-5.8wt%Si-0.55wt%H crystallized Si-free fcc stoichiometric FeH
159 during melting. In run #4, an H-rich liquid (Fe,Si)H_{0.97} coexisted with both fcc
160 stoichiometric FeH and B2 Fe-13.4wt%Si-H ([Fig. 1](#)). The unit-cell volume of B2 Fe-Si

161 found in this experiment (18.5 \AA^3 at 42.6 GPa and 300 K) was 4% smaller than that of
162 H-free B2 Fe+13.4wt%Si (19.2 \AA^3) at the equivalent condition that was estimated from
163 previous volume measurements of B2 Fe-Si⁴⁴. It is difficult to estimate its H content
164 from the smaller unit-cell volume of B2; alternatively, we approximated it from the
165 relative proportions of bubbles and cracks between the B2 and neighboring
166 stoichiometric FeH in the cross section of this sample (Fig. 1a). It gives an approximate
167 amount of H in the B2 phase to be (Fe,Si)H_{0.4}. The recent experiments by ref. 45 also
168 reported a reduction in the volume of B2 stoichiometric FeSi upon hydrogenation and
169 argued that it can be caused by the substitutional incorporation of H into the B2
170 structure. In run #5, we observed that H-rich liquid (Fe,Si)H, which was similar in
171 composition to that in run #4, coexisted with fcc stoichiometric FeH.

172 Figure 3a illustrates the compositions of these liquids and coexisting solids (liquidus
173 phase) plotted in the Fe-Si-H ternary diagram. The recent experiments²² reported the Fe-
174 FeSi binary eutectic liquid composition to be Fe + 11.5 wt% Si at 50 GPa, which is
175 helpful to obtain the ternary liquidus phase relations. While the liquids formed in runs
176 #2 and #3 were similar to each other, they coexisted with hcp Fe and fcc FeH,
177 respectively. It indicates that these two liquid compositions are close to the Fe + FeH
178 cotectic line (showing liquids coexisting two solid phases). The liquid in run #1
179 coexisted with the hcp phase, which also constrains the location of this Fe + FeH
180 cotectic line. The H-rich liquid found in run #4 crystallized both fcc FeH and B2 Fe-Si
181 (Fig. 1) and is therefore on the FeH + Fe-Si cotectic line. It is consistent with the result
182 of run #5.

183 These observations show each liquidus field (showing a compositional range of
184 liquids that first crystallize a given solid phase) of Fe, FeH and Fe-Si at ~50 GPa (Fig.
185 3a). Considering that we often obtained liquids with compositions close to binary
186 eutectic point or ternary cotectic lines in similar melting experiments on Fe alloy
187 systems such as Fe-Fe₃S³³ and Fe-S-O¹², the results of runs #1–#3 suggest the Fe-FeH
188 binary eutectic composition to be around FeH_{0.5} at 50 GPa.

189 The liquidus temperatures (temperatures at solid/liquid boundary) found in runs #1–
190 #3 were 2100–2350 K, which are much lower than the Fe-FeSi binary eutectic
191 temperature of 2900 K at 50 GPa¹⁹, indicating a large effect of hydrogen to reduce
192 melting temperature^{24,43}. The ternary invariant point at which liquid coexists with three
193 solid phases (Fe, FeH and Fe-Si) may be a peritectic (not eutectic) point at ~2400 K.

194 Solid-liquid partitioning of Si & H. The solid/liquid partition coefficients of Si and H
195 between (Si, H)-bearing hcp Fe and coexisting liquid were obtained in run #1, where
196 strong diffuse signal from liquid was observed in the high-temperature XRD pattern
197 (Fig. 2a); $D_{\text{Si}} = 0.94(4)$ and $D_{\text{H}} = 0.70(12)$ (weight basis) (Table 1). The $D_{\text{Si}} = 0.94(4)$ is
198 consistent with those previously observed in the H-free Fe-FeSi binary system at
199 ambient and high pressures^{18,21}.

200 Fcc stoichiometric FeH included the least amounts of Si when coexisting with Fe-
201 Si-H liquids (Fig. 3a), suggesting that Si atoms do not substitute Fe when octahedral
202 interstitial sites are fully occupied by H atoms²⁸. On the other hand, the chemical
203 composition of the B2 phase formed in run #1 was estimated to be $(\text{Fe}_{0.77}\text{Si}_{0.23})\text{H}_{0.4}$,
204 which is enriched in both Si and H than hcp $(\text{Fe}_{0.93}\text{Si}_{0.07})\text{H}_{0.25}$ found in run #4. The D_{H}
205 (B2/liquid) could be 0.45, certainly smaller than the D_{H} (hcp/liquid) = 0.70(12).

206 Discussion

207 Crystallization of hcp Fe-Si-H at Earth's inner core. The solid inner core of our
208 planet consists of hcp Fe containing some light elements; the inner core density deficit
209 with respect to pure Fe has been estimated to be about 4%^{41,46}. While the least amounts
210 of O and C are incorporated into solid Fe in the inner core^{10,12,22}, Si, H and S are known
211 to form solid solution with Fe to some extent^{18,19,23,33} and likely present in the solid
212 core. Nevertheless, interactions among Si, H and S atoms in hcp Fe could be strong and
213 affect their simultaneous solubilities as well as solid-liquid partitioning^{47,48}. Indeed, the
214 liquid immiscibility, a typical consequence of the strong interaction, has been observed
215 between Fe-H and Fe-S liquids to >100 GPa⁴⁹.

216 Recent *ab initio* simulations performed by refs. 17 and 50 emphasized the presence
217 of C and/or H in the solid inner core, in order to account for not only the density but the
218 low P- and S-wave velocities observed. As mentioned above, C is unlikely to be an
219 important impurity element in the inner core because of its low D_{C} (solid-hcp/liquid)
220 ~ 0.1 ²²; otherwise the liquid core should be enriched in C, which is not compatible with
221 its density and velocity observations^{51,52}. Alternatively the inner core alloy may be H-
222 bearing hcp $\text{Fe}_{60}\text{Si}_4\text{H}_{1-15}$ when T_{ICB} ranges from 5500 K to 6500 K (more H is necessary
223 for lower T_{ICB})¹⁷.

224 The present high-pressure experiments demonstrated that liquid crystallized hcp
225 $(\text{Fe}_{0.93}\text{Si}_{0.07})\text{H}_{0.25}$ ($\text{Fe}_{60}\text{Si}_{4.5}\text{H}_{16}$), which is almost equivalent to the inner core solid

226 proposed by ref. 17 when T_{ICB} is 5500 K. Although ref. 29 examined the compression
227 behaviors of hcp $(\text{Fe}_{0.88}\text{Si}_{0.12})\text{H}_{0.61-0.79}$ at 300 K, this study first confirmed that the hcp
228 Fe-Si-H alloy proposed for the inner core is stable to melting temperatures. While D_{H}
229 (solid-Fe/liquid) has never been reported in the literature, $D_{\text{Si}} = 0.94(4)$ obtained in this
230 study is similar to those previously found in the Fe-Si system, indicating no remarkable
231 dependence on H concentration in liquid^{18,21}. It contrasts the large effects of sulfur
232 (S)^{47,48} and C²², which remarkably enhance the solid/liquid D_{Si} . $D_{\text{H}} = 0.70(12)$ observed
233 in this study will be independent from Si concentration in liquid Fe; the little interaction
234 between Si and H in Fe is suggested from the fact that Si atoms substitute Fe, while H
235 atoms occupy the interstitial sites⁵².

236 Possible range of the outer core composition. The liquidus phase relations, in
237 particular the liquidus field of Fe in the Fe-Si-H ternary system at the ICB pressure
238 constrain the liquid core composition. Those determined at ~ 50 GPa (Fig. 3a) may be
239 extrapolated to higher pressures based on the pressure evolutions of the Fe-FeSi and Fe-
240 FeH binary eutectic liquid compositions (Fig. 3b). Si concentration in the Fe-FeSi
241 eutectic liquid has been shown to decrease from 11.5 wt% at 50 GPa to 8 wt% at 330
242 GPa²². On the other hand, the Fe-FeH eutectic composition (Fe + 0.8 wt% H) at ~ 50
243 GPa would remain similar at higher pressures because the temperature/pressure slope of
244 the melting curve of stoichiometric FeH is comparable to that of Fe at >40 GPa⁵³. The
245 outer core composition should be within the liquidus field of Fe—the (Si, H)-depleted
246 hcp phase—at 330 GPa (blue area in Fig. 3b) to form the dense inner core when both Si
247 and H are important impurity elements. It is noted that the presence of other light
248 elements such as O and S diminishes the Si and H variations in the liquidus field of Fe.

249 The possible range of the inner core composition proposed by ref. 17 for $T_{\text{ICB}} =$
250 6000–6500 K is $\text{Fe}_{60}\text{Si}_4\text{H}_1$ – $\text{Fe}_{60}\text{Si}_4\text{H}_8$. If $T_{\text{ICB}} = 5500$ K, it can be $\text{Fe}_{60}\text{Si}_4\text{H}_{15}$. With D_{Si}
251 and D_{H} , the compositions of liquids in equilibrium with these possible inner core solids
252 are calculated to be Fe + 3.5(2) wt% Si + 0.04(1)–0.61(11) wt% H. Such liquid
253 compositions are almost fully within the liquidus field of Fe at 330 GPa (Fig. 3b),
254 which ensures that they crystallize hcp Fe-Si-H under inner core conditions. However,
255 both $D_{\text{Si}} = 0.94(4)$ and $D_{\text{H}} = 0.70(12)$ are close to 1.0 and do not make large differences
256 in Si and H concentrations between the outer and inner core. Indeed, such Fe-Si-H
257 liquid compositions require additional light elements to explain the observed outer core

258 density and velocity according to the *ab initio* calculations by ref. 52. Recent
259 experiments on the metal-silicate partitioning of C during core formation suggested that
260 the core contains at most 0.2 wt% C^{54,55}. If O is an additional impurity element in the
261 outer core (note that O is not soluble into the inner core and thus does not alter the
262 possible range of the inner core composition considered here) and $T_{ICB} = 6000$ K, the
263 outer core liquid may include 1.7–4.4 wt% O⁵² along with 3.5 wt% Si and 0.04–0.32
264 wt% H (Fe₆₀Si_{4.4–4.5}O_{3.8–9.9}H_{1.5–12}). When T_{ICB} is 5500 K, we found liquid Fe + 3.5 wt%
265 Si + 0.61 wt% H + 0.15 wt% O (Fe₆₀Si_{4.3}O_{0.3}H₂₁) for the outer core.

266 While the inner core temperature is still uncertain², these results suggest that H and
267 O, in addition to Si, are likely to be important light elements in the core (note that the
268 liquid core constitutes 95% of the bulk core by mass). It supports recent arguments^{3–5}
269 on the delivery of a large amount of water to the accreting Earth and its sequestration in
270 metals during core formation for the most part^{6–9}.

271 **Conclusions**

272 Recent theoretical calculations¹⁷ found that hcp Fe₆₀Si₄H_{1–15} alloys, depending on core
273 temperatures, account for the density and velocities observed in the Earth's solid inner
274 core. Our experiments demonstrated that liquid metal crystallized hcp Fe₆₀Si_{4.5}H₁₆
275 which is close in composition to the predicted inner core alloys, indicating that hcp Fe
276 can simultaneously include both Si and H unlike fcc FeH that does not incorporate Si.
277 We determined the liquidus phase relations in the Fe-Si-H ternary system, suggesting
278 the Fe-FeH eutectic composition to be around FeH_{0.5}. We also obtained the solid hcp-
279 Fe/liquid partition coefficients for Si and H, $D_{Si} = 0.94(4)$ and $D_H = 0.70(12)$. While
280 these experiments were carried out at ~50 GPa, the D_H value is likely not sensitive to
281 pressure because the size of H atom is substantially smaller than those of Fe and Si
282 atoms even at 330 GPa⁵². Similar temperature/pressure slopes between the melting
283 curves of Fe and stoichiometric FeH⁵³ suggest that the Fe-FeH liquidus phase relations
284 change little with increasing pressure. Therefore, the Fe-Si-H ternary liquidus diagram
285 may be extrapolated to ICB conditions by primarily considering the change in the Fe-
286 FeSi liquidus phase relations reported in previous studies^{19,22}.

287 Such liquidus phase relations, in particular the liquidus field of Fe as well as D_{Si} and
288 D_H between the solid and liquid cores help constrain the Earth's core composition. We
289 explored the possible compositional range of the outer core liquid that is in equilibrium

290 with the predicted solid inner core alloy¹⁷. The relatively high D_{Si} and D_{H} close to 1.0
291 do not make much differences in Si and H concentrations between the outer and inner
292 core, requiring other light elements such as O that is least partitioned into solid Fe^{10–12}.
293 Depending on the ICB temperature that is still uncertain, we found H and O, as well as
294 Si, are important core light elements, which support recent arguments on the
295 sequestration of a large amount of water in the Earth's core^{6–9}.

296 **Data availability**

297 All data supporting the findings of this study are available in the paper or from the
298 corresponding author upon request.

299 **References**

- 300 1. Birch, F. Elasticity and constitution of the Earth's interior. *J. Geophys. Res.* **57**,
301 227–286 (1952).
- 302 2. Hirose, K., Wood, B. & Vočadlo, L. Light elements in the Earth's core. *Nat. Rev.*
303 *Earth & Environ.* **2**, 645–658; <https://doi.org/10.1038/s43017-021-00203-6> (2021).
- 304 3. Raymond, S. N., Quinn, T. & Lunine, J. I. High-resolution simulations of the final
305 assembly of Earth-like planets. 2. Water delivery and planetary habitability.
306 *Astrobiology* **7**, 66–84 (2007).
- 307 4. Walsh, K. J., Morbidelli, A., Raymond, S. N., O'Brien, D. P. & Mandell, A. M. A
308 low mass for Mars from Jupiter's early gas-driven migration. *Nature* **475**, 206–209
309 (2011).
- 310 5. Sato, T., Okuzumi, S. & Ida, S. On the water delivery to terrestrial embryos by ice
311 pebble accretion. *Astron. Astrophys.* **589**, A15 (2016).
- 312 6. Okuchi, T. Hydrogen partitioning into molten iron at high pressure: implications for
313 Earth's core. *Science* **278**, 1781–1784 (1997).
- 314 7. Li, Y., Vočadlo, L., Sun, T. & Brodholt, J. P. The Earth's core as a reservoir of
315 water. *Nat. Geosci.* **13**, 453–458 (2020).
- 316 8. Yuan, L. & Steinle-Neumann, G. Strong sequestration of hydrogen into the Earth's
317 core during planetary differentiation. *Geophys. Res. Lett.* **47**, e2020GL088303
318 (2020).

- 319 9. Tagawa, S. *et al.* Experimental evidence for hydrogen incorporation into Earth's
320 core. *Nat. Commun.* **12**, 2588; <https://doi.org/10.1038/s41467-021-22035-0> (2021).
- 321 10. Alfè, D., Gillan, M. J. & Price, G. D. Composition and temperature of the Earth's
322 core constrained by combining ab initio calculations and seismic data. *Earth*
323 *Planet. Sci. Lett.* **195**, 91–98 (2002).
- 324 11. Ozawa, H., Hirose, K., Tateno, S., Sata, N. & Ohishi, Y. Phase transition boundary
325 between B1 and B8 structures of FeO up to 210 GPa. *Phys. Earth Planet. Inter.*
326 **179**, 157–163 (2010).
- 327 12. Yokoo, S., Hirose, K., Sinmyo, R. & Tagawa, S. Melting experiments on liquidus
328 phase relations in the Fe-S-O ternary system under core pressures. *Geophys. Res.*
329 *Lett.* **46**, 5137–5145 (2019).
- 330 13. Mao, Z. *et al.* Sound velocities of Fe and Fe–Si alloy in the Earth's core. *Proc.*
331 *Natl. Acad. Sci. USA* **109**, 10239–10244 (2012).
- 332 14. Antonangeli, D. *et al.* Sound velocities and density measurements of solid hcp-Fe
333 and hcp-Fe-Si (9 wt.%) alloy at high pressure: constraints on the Si abundance in
334 the Earth's inner core. *Earth Planet. Sci. Lett.* **482**, 446–453 (2018).
- 335 15. Shibazaki, Y. *et al.* Sound velocity measurements in dhcp-FeH up to 70 GPa with
336 inelastic X-ray scattering: implications for the composition of the Earth's core.
337 *Earth Planet. Sci. Lett.* **313–314**, 79–85 (2012).
- 338 16. Sakamaki, T. *et al.* Constraints on Earth's inner core composition inferred from
339 measurements of the sound velocity of hcp-iron in extreme conditions. *Sci. Adv.* **2**,
340 e1500802; [10.1126/sciadv.1500802](https://doi.org/10.1126/sciadv.1500802) (2016).
- 341 17. Wang, W., Li, Y., Brodholt, J. P., Vočadlo, L., Walter, M. J. & Wu, Z. Strong shear
342 softening induced by superionic hydrogen in Earth's inner core. *Earth Planet. Sci.*
343 *Lett.* **568**, 117014 (2021).
- 344 18. Kuwayama, Y. & Hirose, K. Phase relations in the system Fe–FeSi at 21 GPa. *Am.*
345 *Mineral.* **89**, 273–276 (2004).
- 346 19. Fischer, R. A. *et al.* Phase relations in the Fe–FeSi system at high pressures and
347 temperatures. *Earth Planet. Sci. Lett.* **373**, 54–64 (2013).
- 348 20. Tateno, S., Kuwayama, Y., Hirose, K. & Ohishi, Y. The structure of Fe-Si alloy in
349 Earth's inner core. *Earth Planet. Sci. Lett.* **418**, 11–19 (2015).

- 350 21. Ozawa, H., Hirose, K., Yonemitsu, K. & Ohishi, Y. High-pressure melting
351 experiments on Fe–Si alloys and implications for silicon as a light element in the
352 core. *Earth Planet. Sci. Lett.* **456**, 47–54 (2016).
- 353 22. Hasegawa, M., Hirose, K., Oka, K. & Ohishi, Y. Liquidus phase relations and
354 solid-liquid partitioning in the Fe-Si-C system under core pressures. *Geophys. Res.*
355 *Lett.* **48**, e2021GL092681 (2021).
- 356 23. Fukai, Y. Some properties of the Fe-H system at high pressures and temperatures,
357 and their implications for the Earth’s core. in *High-Pressure Research:*
358 *Applications to Earth and Planetary Sciences* (eds. Syono, Y. & Manghnani, M.
359 H.) **67**, 373–385 (AGU, 1992).
- 360 24. Sakamaki, K. *et al.* Melting phase relation of FeH_x up to 20 GPa: implication for
361 the temperature of the Earth’s core. *Phys. Earth Planet. Inter.* **174**, 192–201 (2009).
- 362 25. Shibazaki, Y. *et al.* High-pressure and high-temperature phase diagram for
363 Fe_{0.9}Ni_{0.1}-H alloy. *Phys. Earth Planet. Inter.* **228**, 192–201 (2014).
- 364 26. Iizuka-Oku, R. *et al.* Hydrogenation of iron in the early stage of Earth’s evolution.
365 *Nat. Commun.* **8**, 14096; <https://doi.org/10.1038/ncomms14096> (2017).
- 366 27. Machida, A. *et al.* Site occupancy of interstitial deuterium atoms in face-centred
367 cubic iron. *Nat. Commun.* **5**, 5063; <https://doi.org/10.1038/ncoms6063> (2014).
- 368 28. Ikuta, D. *et al.* Interstitial hydrogen atoms in face-centered cubic iron in the Earth’s
369 core. *Sci. Rep.* **9**, 7108; <https://doi.org/10.1038/s41598-019-43601-z> (2019).
- 370 29. Tagawa, S., Ohta, K., Hirose, K., Kato, C. & Ohishi, Y. Compression of Fe-Si-H
371 alloys to core pressures. *Geophys. Res. Lett.* **43**, 3686–3692 (2016).
- 372 30. Chi, Z. *et al.* Cryogenic implementation of charging diamond anvil cells with H₂
373 and D₂. *Rev. Sci. Instrum.* **82**, 105109 (2011).
- 374 31. Ohta, K. *et al.* Phase boundary of hot dense fluid hydrogen. *Sci. Rep.* **5**, 16560;
375 <https://doi.org/10.1038/srep16560> (2015).
- 376 32. Hirao, N. *et al.* New developments in high-pressure X-ray diffraction beamline for
377 diamond anvil cell at SPring-8. *Matter Radiat. at Extremes* **5**, 1–10 (2020).
- 378 33. Mori, Y. *et al.* Melting experiments on Fe–Fe₃S system to 254 GPa. *Earth Planet.*
379 *Sci. Lett.* **464**, 135–141 (2017).
- 380 34. Seto, Y., Nishio-Hamane, D., Nagai, T. & Sata, N. Development of a software suite
381 on X-ray diffraction experiments. *Rev. High Press. Sci. Technol.* **20**, 269–276
382 (2010).

- 383 35. Dorogokupets, P. I. & Dewaele, A. Equations of state of MgO, Au, Pt, NaCl-B1,
384 and NaCl-B2: internally consistent high-temperature pressure scales. *High Press.*
385 *Res.* **27**, 431–446 (2007).
- 386 36. Campbell, A. J. *et al.* High pressure effects on the iron-iron oxide and nickel-nickel
387 oxide oxygen fugacity buffers. *Earth Planet. Sci. Lett.* **286**, 556–564 (2009).
- 388 37. Akahama, Y. & Kawamura, H. High-pressure Raman spectroscopy of diamond
389 anvils to 250 GPa: method for pressure determination in the multimegabar pressure
390 range. *J. Appl. Phys.* **96**, 3748–3751 (2004).
- 391 38. Hirose, K. *et al.* Crystallization of silicon dioxide and compositional evolution of
392 the Earth’s core. *Nature* **543**, 99–102 (2017).
- 393 39. Fukai, Y. & Suzuki, T. Iron-water reaction under high pressure and its implication
394 in the evolution of the Earth. *J. Geophys. Res.* **91**, 9222 (1986).
- 395 40. Caracas, R. The influence of hydrogen on the seismic properties of solid iron.
396 *Geophys. Res. Lett.* **42**, 3780–3785 (2015).
- 397 41. Dewaele, A. *et al.* Quasihydrostatic equation of state of iron above 2 Mbar. *Phys.*
398 *Rev. Lett.* **97**, 215504 (2006).
- 399 42. Tsujino, N. *et al.* Equation of state of γ -Fe: reference density for planetary cores.
400 *Earth Planet. Sci. Lett.* **375**, 244–253 (2013).
- 401 43. Hirose, K. *et al.* Hydrogen limits carbon in liquid iron. *Geophys. Res. Lett.* **46**,
402 5190–5197 (2019).
- 403 44. Edmund, E. *et al.* Structure and elasticity of cubic Fe-Si alloys at high pressures.
404 *Phys. Rev. B* **100**, 134105 (2019).
- 405 45. Fu, S., Chariton, S., Prakapenka, V. B., Chizmeshya, A. & Shim, S.-H. Hydrogen
406 solubility in FeSi alloy phases at high pressures and temperatures. *Am. Mineral.*
407 10.2138/am-2022-8295 (2022).
- 408 46. Fei, Y., Murphy, C., Shibasaki, Y., Shahar, A. & Huang, H. Thermal equation of
409 state of hcp-iron: constraint on the density deficit of Earth’s solid inner core.
410 *Geophys. Res. Lett.* **43**, 6837–6843 (2016).
- 411 47. Tateno, S. *et al.* Melting experiments on Fe-Si-S alloys to core pressures: silicon in
412 the core? *Am. Mineral.* **103**, 742–748 (2018).
- 413 48. Tao, R. & Fei, Y. High-pressure experimental constraints of partitioning behavior
414 of Si and S at the Mercury’s inner core boundary. *Earth Planet. Sci. Lett.* **562**,
415 116849 (2021).

- 416 49. Yokoo, S., Hirose, K., Tagawa, S., Morard, G. & Ohishi, Y. Stratification in
417 planetary cores by liquid immiscibility in Fe-S-H. *Nat. Commun.* **13**, 644;
418 <https://doi.org/10.1038/s41467-022-28274-z> (2022).
- 419 50. Li, Y., Vočadlo, L. & Brodholt, J. P. The elastic properties of hcp-Fe alloys under
420 the conditions of the Earth's inner core. *Earth Planet. Sci. Lett.* **493**, 118–127
421 (2018).
- 422 51. Badro, J., Côté, A. S. & Brodholt, J. P. A seismologically consistent compositional
423 model of Earth's core. *Proc. Natl Acad. Sci. USA* **111**, 7542–7545 (2014).
- 424 52. Umemoto, K. & Hirose, K. Chemical compositions of the outer core examined by
425 first principles calculations. *Earth Planet. Sci. Lett.* **531**, 116009 (2020).
- 426 53. Tagawa, S., Gomi, H., Hirose, K. & Ohishi, Y. High-temperature equation of state
427 of FeH: implications for hydrogen in Earth's inner core. *Earth and Space Science*
428 *Open Archive*; <https://doi.org/10.1002/essoar.10508224.1> (2021).
- 429 54. Fischer, R., Cottrell, E., Hauri, E., Lee, K. K. M. & Le Voyer, M. The carbon
430 content of Earth and its core. *Proc. Natl. Acad. Sci. USA* **117**, 8743–8749 (2020).
- 431 55. Blanchard, I. *et al.* The metal–silicate partitioning of carbon during Earth's
432 accretion and its distribution in the early solar system. *Earth Planet. Sci. Lett.* **580**,
433 117374 (2022).

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439 **Author Contributions**

440 K. Hikosaka and K. Hirose designed the project. All authors were involved in high-
441 pressure experiments and data analyses. K. Hikosaka and K. Hirose wrote the paper
442 with input from all authors.

443 **Additional Information**

444 **Competing Interests:** The authors declare no competing interests.

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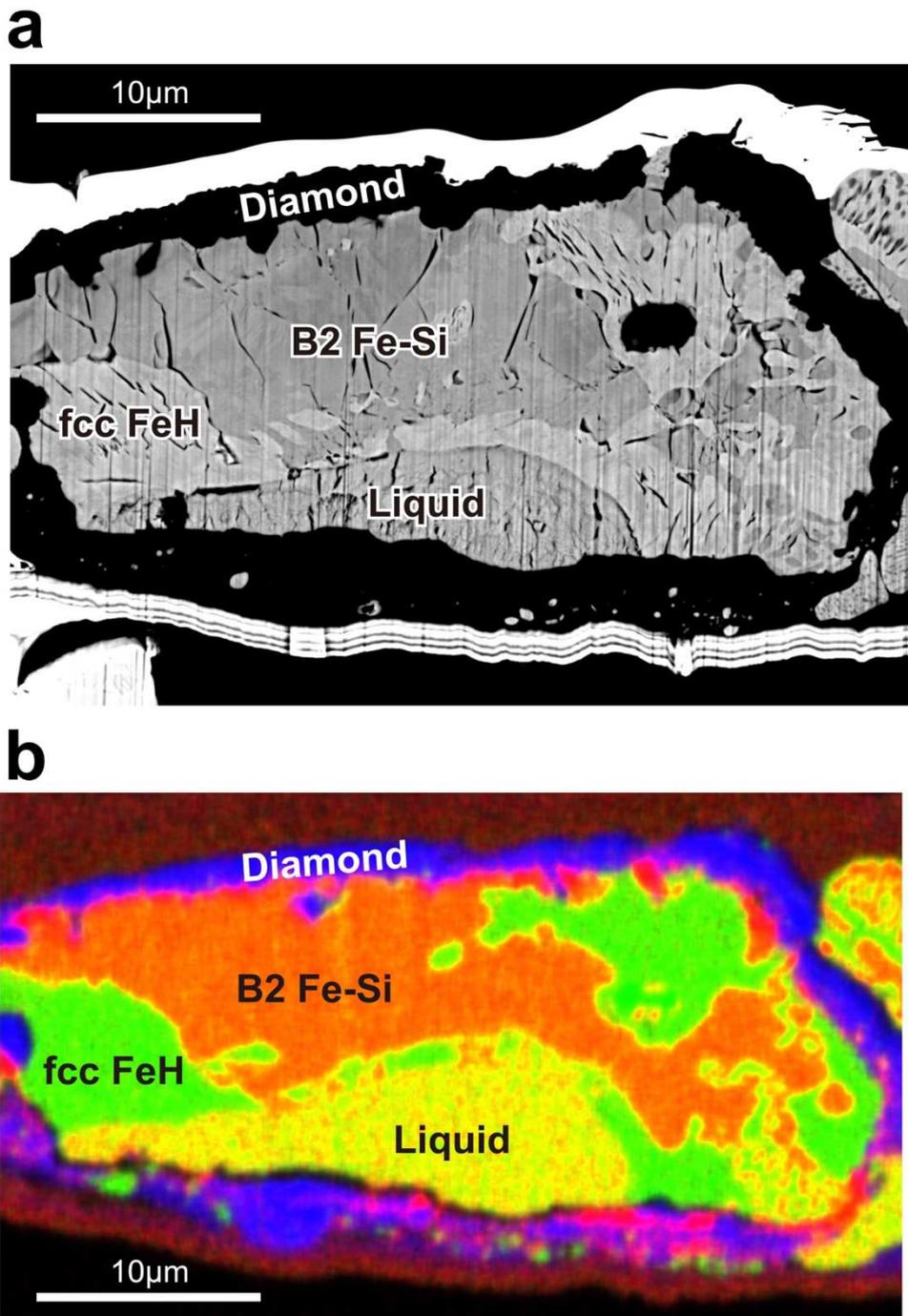
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Table 1. Experimental conditions and chemical compositions of liquid and coexisting solid. Phases observed by XRD before and after heating are given. The H content was determined from XRD data except for the B2 phase. See text for details. Numbers in parentheses indicate errors in the last digits. ^ax in (Fe,Si)H_x.

Run #	P (GPa)	T (K)	Phase	Fe wt%	Si wt%	H wt%	x ^a	C wt%	O wt%	
1	56(3)	2100(210)	before	hcp	96.0	4.0	0.19(2)	0.10	-	-
			after	hcp	97.5(9)	3.7(1)	0.47(4)	0.25	0.0(1)	0.1(1)
			after	liquid (fcc)	97.9(7)	4.0(1)	0.67(5)	0.35	0.0(1)	0.1(1)
			before	hcp	96.0	4.0	0.58(5)	0.31	-	-
2	57(3)	2150(110)	after	hcp	96.2(8)	4.3(3)	0.30(2)	0.16	0.2(1)	0.1(1)
			after	liquid (hcp)	92.6(20)	5.9(1)	0.56(4)	0.30	0.0(1)	0.5(3)
3	61(3)	2350(400)	before	hcp	96.0	4.0	1.54(12)	0.81	-	-
			after	fcc	99.7(12)	0.8(2)	1.59(13)	0.87	0.3(2)	0.3(2)
			after	liquid (hcp)	93.2(13)	5.8(2)	0.55(4)	0.29	0.4(3)	0.7(4)
			before	hcp	93.5	6.5	-	-	-	-
4	48(2)	2100(210)	after	fcc	98.9(17)	0.3(2)	1.76(14)	0.98	0.1(1)	0.2(1)
			after	liquid (dhcp)	94.2(10)	5.6(2)	1.84(15)	0.97	0.6(1)	0.5(2)
			after	B2	87.5(5)	13.4(2)	0.82	0.4	0.4(1)	0.2(1)
			before	hcp	96.0	4.0	1.86(15)	0.99	-	-
5	49(2)	2450(120)	after	fcc	100.4(5)	0.6(1)	1.78(14)	0.98	0.1(1)	0.1(1)
			after	liquid (fcc)	91.9(4)	7.1(1)	1.79(14)	0.93	0.5(1)	0.7(2)

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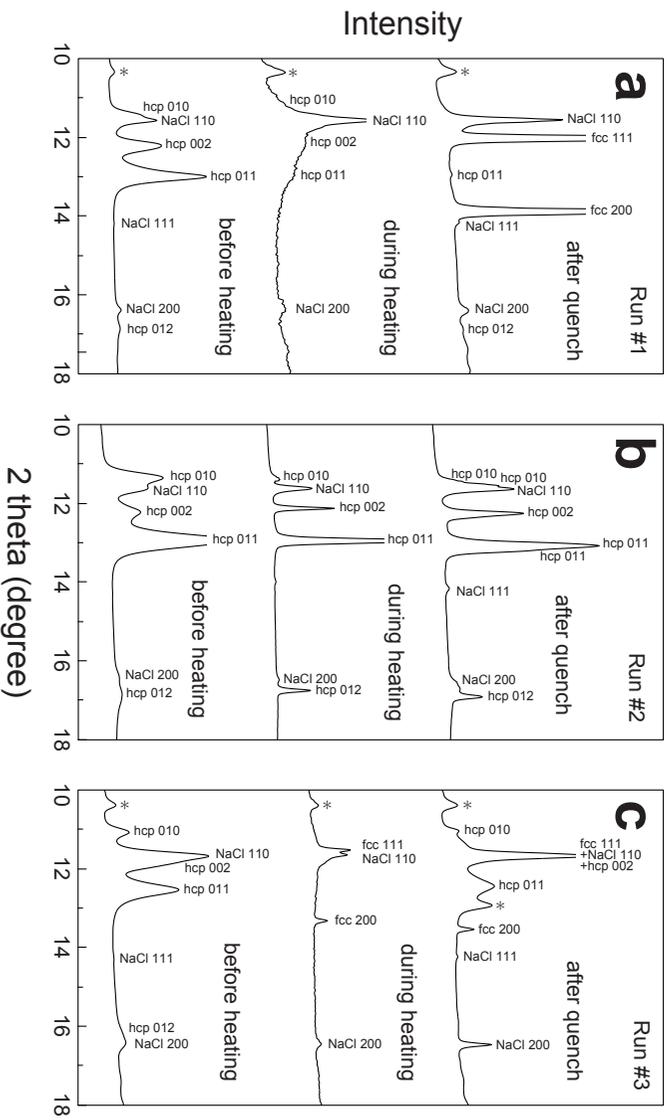
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450 **Figure 1.** Coexisting Fe-Si-H liquid and solid fcc FeH and B2 Fe-Si phases at 48 GPa
 451 and 2100 K in run #4 using Fe-Si and C_nH_{2n+2} paraffin as starting materials. (a) Back-
 452 scattered electron image and (b) combined X-ray elemental map of Fe (green) + Si (red)
 453 + C (blue). Bubbles and cracks in (a) indicate that hydrogen was present in liquid and
 454 solids at high pressure and escaped during decompression.

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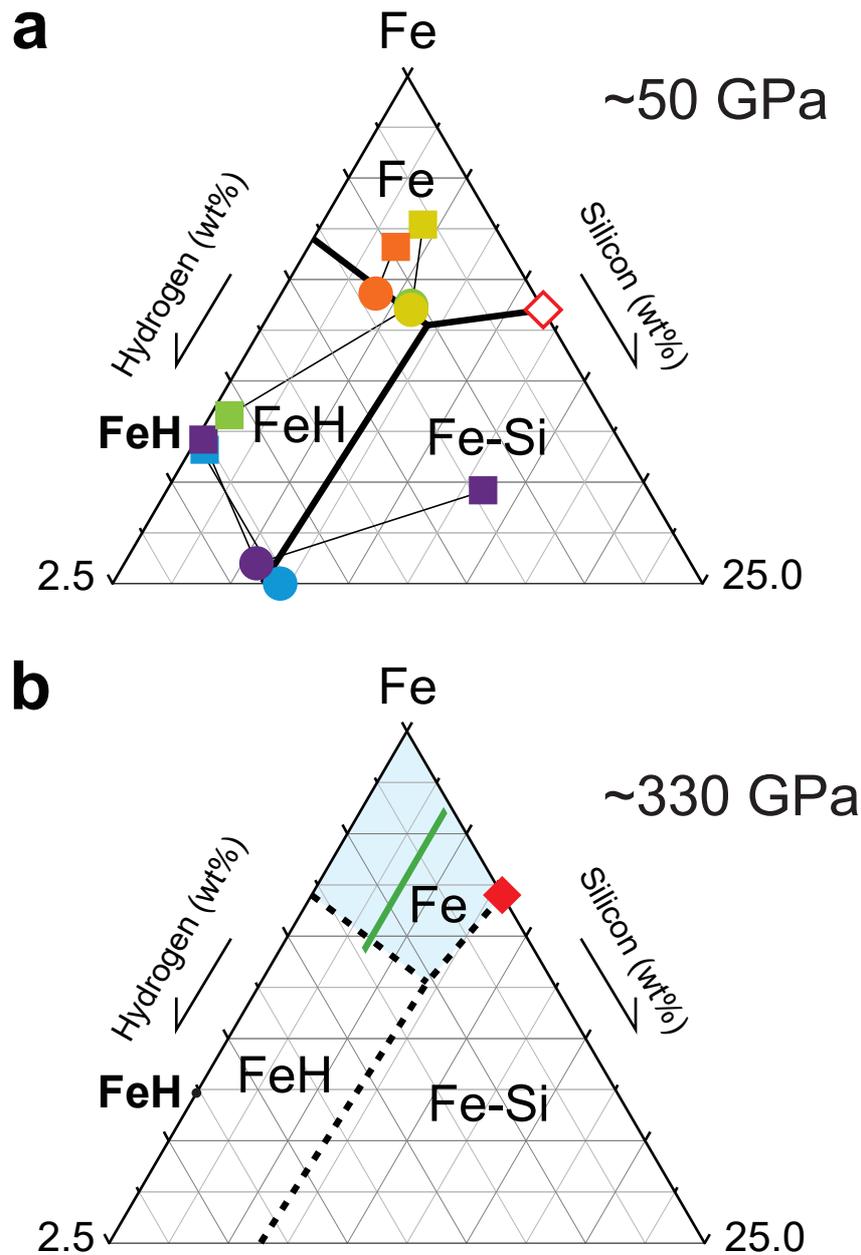
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457 **Figure 2.** Sample XRD patterns collected at high pressures in (a) run #1, (b) run #2 and

458 (c) run #3 before, during and after heating (melting). Asterisks indicate unknown peaks.

459 See text for details.

460



461

462 **Figure 3.** Liquidus phase relations in Fe-Si-H showing the liquidus fields of hcp Fe, fcc
 463 FeH and B2 Fe-Si at (a) ~50 GPa and (b) ~330 GPa. The compositions of liquids
 464 (circles) and coexisting solid phases (squares) obtained in the present experiments are
 465 plotted in (a); orange, run #1; yellow, run #2; green, run #3; purple, run #4; blue, run
 466 #5. Open and closed red diamonds in (a) and (b) show the Fe-FeSi eutectic liquid
 467 composition at respective pressure²². The liquid core compositions (green belt),
 468 calculated with D_{Si} and D_{H} considering their uncertainties from the proposed inner core
 469 solid compositions¹⁷, are mostly within the liquidus field of (Si, H)-bearing hcp Fe (blue
 470 area) at 330 GPa (b).