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Engineering geological and petrological characterization of paleoweathered rock in the K₁/J₂ contact zone in the Ordos Basin, China

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1 Abstract

2 Various geological processes (mineral composition, structure, tectonics and weathering, etc.)
3 affect the physical-mechanical properties of rock. Petrological and engineering geological
4 characteristics of paleoweathered rock (PWR) from the K₁/J₂ contact zone are described in detail via
5 field investigation and experimental testing. This PWR exhibits mainly sandy grains and mud
6 structures, layered and massive strata, and calcareous and argillaceous cements; fissures are
7 developed and often filled with argillaceous and detrital materials; 9 minerals and 7 oxides are present,
8 and quartz is present in each sample. Long-term weathering results in a consistent bulk density and
9 high porosity due to the transformation of primary minerals into secondary clay minerals, forming
10 PWR that undergoes argillization in water. The axial point load strength (PLS) is the largest among
11 the tested PLSs, followed by the diametral PLS, and the irregular PLS. The uniaxial compressive
12 strength (UCS) varies widely, but the results are reliable. The mineralogical, physical and mechanical
13 properties of the PWR are compared to predict one parameter from another and study their mutual
14 influence. The PLS and UCS of the PWR are negatively correlated with the elastic mineral group
15 (EMG) content, weathering alteration indexes, water content, and total porosity and positively
16 correlated with the quartz content, brittle mineral group (BMG) content, bulk density, real density,
17 and longitudinal wave velocity. The UCS and the PLS, axial (diametral) PLS and irregular PLS are
18 positively correlated. These results provide a theoretical basis for physical and mechanical property
19 prediction of PWR masses and rapid estimation of UCS in engineering.

20 **Keywords:** paleoweathered rock, point load strength, uniaxial compressive strength, engineering
21 geological characteristics

22

23 1. Introduction

24 The behavior of rock under stress is studied in engineering geological research, and the
25 quantification of engineering geological and petrological characteristics is very useful in the
26 interpretation of rock behavior. Rock properties are actually affected by various geological processes,
27 such as mineral composition and content, structure, tectonics and weathering (Pappalardo et al. 2016).
28 Therefore, in geological scenarios characterized by complex geological dynamic evolution, it is very
29 meaningful to describe the engineering geology and petrological characteristics in detail.

30 The inspiration of this study is the occurrence of water (mud) inrush accidents caused by coal
31 mining in the Ordos Basin (Guo et al. 2020; Lu et al. 2018). In recent years, large-scale coal mining in
32 some mines in the Ordos Basin has caused the deformation (failure) of overlying strata, a decline in
33 impermeable capacity, the leakage of Cretaceous water bodies, water (mud) inrush accidents, and
34 damage to the ecological environment (zhu et al. 2020). Through our statistical analysis of borehole
35 data in the mining area of the Ordos Basin and a large number of field geological surveys, we found
36 that PWR in the K₁/J₂ contact zone generally exists in the central part of the Ordos Basin, serving as

1 the key aquifuge directly below the Cretaceous water body, and that the deformation and failure of
2 PWR is the root cause of the abovementioned series of problems.

3 Some physical and mechanical properties of rock depend on geological characteristics, which
4 causes the rock to show inhomogeneity, discontinuity and anisotropy (Douma et al. 2017; Vikram et
5 al. 2012). There are many studies on the geological characteristics of rocks worldwide, mainly
6 focusing on petrography, petrochemistry, structure, physical mechanics, hydrogeology and other
7 characteristics (Zhang et al. 2013; Ulyasheva et al. 2016; Marques et al. 2017; Menningen et al. 2018;
8 Vonto et al. 2020; Bai et al. 2020). Many scholars have been searching for potential relationships
9 between certain geological characteristics and rock mechanical properties (Chatterjee et al. 2002;
10 Tamrakar et al. 2007; Sousa et al. 2005). For example, the mineral composition and contents, density,
11 porosity, water content and wave speed of rock will affect its mechanical properties and lead to the
12 anisotropy of the rock (Meng et al. 2007; Cantisani et al. 2013; Undul et al. 2016; Fereidooni 2016;
13 Sun et al. 2017; Wang et al. 2019). Many scholars have carried a considerable amount of research on
14 the relationship between rock PLS and UCS (D'Andrea et al. 1965; Broch et al. 1972; Gunsallus et al.
15 1984; Hawkins 1998; Kahraman 2001; Quane 2003; Palchik et al. 2004; Sabatakakis et al. 2008; Basu
16 et al. 2010; Heidari et al. 2012; Xiang 1981; Wei 1982; Li et al. 2013; Wong et al. 2017; Liu et al.
17 2019). The above studies have improved the understanding of the macro- and microproperties of
18 rocks. However, there are no reports on related studies of PWR of the K_1/J_2 contact zone of the Ordos
19 Basin.

20 Therefore, in this work, PWR samples were collected from the K_1/J_2 contact zone in the Ordos
21 Basin to test the mineral composition and contents, chemical element composition and contents, water
22 content, density, porosity, longitudinal wave velocity, expansion rate, PLS and UCS of the PWR. The
23 aim is to study the petrological and engineering geological characteristics of the PWR in the K_1/J_2
24 contact zone. Through the regression analysis method, the relationship between each feature data type
25 and the degree of influence of each feature data type on the engineering characteristics of PWR are
26 studied. The results of this research provide a theoretical basis for the physical-mechanical properties
27 of PWR masses and the generalization of geological models in the next step of research and are of
28 great significance to the safety of coal mining and ecological environment protection in the Ordos
29 Basin.

30 2. Geological setting and sampling site

31 The Ordos Basin is a multicycle superimposed basin developed on the North China Craton and
32 currently located in the western North China Craton (Xu et al. 2013). In the middle of the Yanshan
33 tectonic movement, during the Late Jurassic to Early Cretaceous, the Ordos Basin developed
34 considerable thrust nappe structures, forming a basin pattern of uplift in the east and subsidence in the
35 west, resulting in the Jurassic Anding Formation and the overlying Cretaceous Zhidan Group (Luohe
36 Formation) being in unconformable contact (Zhang et al. 2007; Jia et al. 2005). The Upper Jurassic
37 strata were intermittently deposited for approximately 20 million years (Huang 2019), and PWR is
38 preserved on the top of the Jurassic Anding Formation. According to the drilling data of the
39 Yingpanhao mining area, PWR is widely present at the K_1/J_2 contact zone, in Shenmu County (①),
40 Jingbian County (②), the Ansai district of Yan'an city (③), and Ganquan County (④) (Fig. 1a). All
41 outcrops of PWR are found in the K_1/J_2 contact zone (Fig. 1b). The investigation shows that the PWR
42 is characterized by purple-red to gray coloring and is generally paleoweathered sandstone, sandy
43 mudstone and mudstone. These sedimentary characteristics consistently indicate the arid and hot
44 sedimentary climate in the area. This is consistent with the characteristics of the arid climate reflected

1 by the sporopollen, the specialized combination of Psilunio, and the ostracod Timiriasevia-Darwinula
2 genera found in the research area (Wang 2011).

3 In this work, the PWR samples of the K_1/J_2 contact zone were collected from core drilling in the
4 Yingpanhao mining area and field geological survey sampling. The 17 groups of samples are from
5 five regions. Each group of PWR samples was tested to determine the mineral composition, chemical
6 element composition, density, porosity, water content, longitudinal wave velocity, expansion rate, PLS
7 and UCS. Notably, the diametral and axial loading PLS test specimen diameter was nearly 50 mm,
8 and the lengths were nearly 30 mm and 70 mm, respectively; the size of the test pieces for the
9 irregular lump testing met the following requirements: the shortest side length was 30-80 mm, the
10 ratio of the distance between the loading points (D) to the average width of the smallest load surface
11 passing through the two loading points (W) was 0.5-1.0, and the distance from the loading point to the
12 free end (L) was greater than 0.5D. The UCS sample diameter and length dimensions were
13 approximately 50 mm and 100 mm, respectively. A total of 342 PWR samples were prepared for
14 strength testing, but only 216 valid samples were recorded, including 19 samples for the UCS testing
15 (2 failed), 19 samples for the diametral testing (2 invalid), 18 samples for the axial testing (1 invalid)
16 and 286 samples for the irregular lump testing (31 invalid).

17 3 Petrological features

18 This section describes how field geological surveys, scanning electron microscopy (SEM), X-ray
19 diffraction (XRD) and X-ray fluorescence (XRF) methods were used to qualitatively characterize the
20 macroscopic structure and microscopic mineral structure, quantitatively analyze the mineral
21 composition and chemical element composition, and more intuitively understand the mineralogy and
22 petrology characteristics of the PWR.

23 3.1 Methodology

24 3.1.1 SEM

25 The scanning electron microscope model used in this work is an FEI Quanta TM 250;
26 microstructure and composition characterization images of PWR samples were collected. The main
27 performance indicators of the instrument are as follows: high-vacuum mode resolutions: \leq cu nm at
28 30 kV (SE), \leq 4.0 nm at 30 kV (BSE), and \leq and nm at 3 kV (SE); magnification range: 6-100
29 million times; and acceleration voltage : 0.2-30 kV.

30 3.1.2 XRD

31 The X-ray diffractometer model used in this work to perform diffraction analysis of the PWR
32 samples is an X-Ray Diffraction from D8 Advance. The main performance indicators of this
33 instrument were as follows: measurement accuracy: angle reproducibility of $\pm 0.0001^\circ$; radius of the
34 goniometer: ≥ 0.0 mm; diameter of the angle measuring circle continuously changed; minimum step
35 length: 0.0001° ; angle range (2θ): $-110\sim 168^\circ$; temperature range: 20~1200urengle range (2θ 3 kW;
36 stability $\pm 0.01\%$; and tube voltage 20~60 kV (1 kV/1 step). Semiquantitative mineralogical analysis
37 of PWR diffraction data was performed using MDI Jade6 diffraction analysis software. The
38 determination of the mineral phase in each sample was achieved by comparison with the standard
39 spectrum of various mineral phases. The semiquantitative analysis of minerals was performed with
40 MDI Jade6 software according to the steps of background deduction, smoothing, peak search and
41 calculation; for each sample, the contents of various minerals were calculated by the K value method,
42 and the relevant values were provided by the PDF card in MDI Jade6 software (Zhao et al. 2019). The
43 semiquantitative K value method is as follows: suppose the mixture contains two phases, No. 1 and
44 No. 2, and the characteristic peak intensities of 100% of their relative intensities are measured as I1

1 and I₂; check the standard card library to determine their *K* values ($K = I/I_a - \text{Al}_2\text{O}_3$), which are *K*₁
2 and *K*₂, respectively; then, $1\% = \frac{I_1}{\frac{I_1+I_2}{K_1+K_2}} \times 100\%$, and $2\% = \frac{I_2}{\frac{I_1+I_2}{K_1+K_2}} \times 100\%$.

3 3.1.3 XRF

4 The XRF spectrometer model used in this paper is an X-Ray Fluorite Spectroscopy S8 Tiger and
5 was used to analyze the main and trace element types and contents of the PWR samples. The main
6 performance indicators of the instrument are as follows: scanning method: sequential scanning;
7 hardware indicators: power 4 kW, maximum voltage: 60 kV; maximum current: 170 Ma; element
8 detection range: Be(4)-U(92); and detection limit: PPM-100%. In recent years, the chemical index of
9 alteration (CIA), chemical index of weathering (CIW) and plagioclase index of alteration (PIA) have
10 been widely used to reflect the source area degree of weathering (Shaldybin et al. 2019; Elena et al.
11 2015; Farkas et al. 2018) and are defined as follows: $\text{CIA} = [\text{Al}_2\text{O}_3 / (\text{Al}_2\text{O}_3 + \text{CaO} + \text{Na}_2\text{O} + \text{K}_2\text{O})] \times 100$,
12 $\text{CIW} = [\text{Al}_2\text{O}_3 / (\text{Al}_2\text{O}_3 + \text{CaO} + \text{Na}_2\text{O})] \times 100$, $\text{PIA} = [\text{Al}_2\text{O}_3 - \text{K}_2\text{O} / (\text{Al}_2\text{O}_3 + \text{CaO} + \text{Na}_2\text{O} - \text{K}_2\text{O})]$.

13 3.2 Petrological characterization

14 Through numerous surveys of PWR outcrops in the *K*₁/*J*₂ contact zone, a preliminary
15 understanding of the macrostructure characteristics of PWR has been obtained. The color of PWR is
16 mainly light red to purple-red, light gray to gray, and light green. The structure is dominated by gravel,
17 fine grains, silt grains and pelites, and a few are cryptocrystalline (Fig. 2 e-f). The tectonic activity
18 mainly formed layered and massive strata, a few of which are earthy (Fig. 2 a-d). The cements are
19 mainly calcareous and argillaceous, with sparse iron cements (Fig. 2 a-c). Joints and fractures are
20 developed, often filled with mud and clastic material.

21 The XRD semiquantitative analysis results are shown in Fig. 3, reflecting the mineral
22 composition of PWR in the *K*₁/*J*₂ contact zone in the Ordos Basin. Clearly, all the PWR collected
23 from outcrop sections contains quartz. This is because quartz is the most stable mineral among the
24 detected mineral types; quartz is also one of the most stable minerals on Earth's surface (Sun et al
25 2013). The remaining detected minerals (calcite, dolomite, albite, muscovite, kaolinite, illite,
26 montmorillonite, and chlorite) vary depending on the sampling location. For example, the calcite of
27 the AS-1 and GQ-2 groups has been weathered and decomposed into other secondary minerals, and
28 the calcite and dolomite of the SM-1 and SM-2 groups have been weathered into other secondary
29 minerals. The rock-forming minerals listed above were identified and described according to the
30 "Geotechnical Investigation and Design Manual" (Lin et al 1996) and SEM images. Quartz is
31 kidney-shaped or irregular (Fig. 2 j), hard and brittle, and has strong resistance to weathering. Calcite
32 often forms crystal clusters and granular or nodular shapes (Fig. 2 k) and is brittle. Dolomite is often
33 granular or massive (Fig. 2 l) and brittle, with curved surfaces and stripes. Feldspar is often platy or
34 columnar (Fig. 2 m) and brittle, with stripes visible on the cleavage surface. Muscovite is flaky or
35 scaly, stacked and elastic. Kaolinite is generally lumpy, flaky, or earthy, with a slippery feel, easily
36 absorbs water, and has a plastic viscosity when wet. Illite is generally earthy, with a greasy feel, and is
37 weaker than muscovite. Montmorillonite is often earthy or cryptocrystalline, with a slippery feel, and
38 swells significantly after being immersed in water. Chlorite is scaly, has a slippery feel, is flexible,
39 and has poor elasticity.

40 XRF was used to analyze the main elemental composition of the PWR samples, and the results
41 are shown in Fig. 4. The PWR mainly contains 7 kinds of oxides; on the outcrop profile, SiO₂ is the
42 oxide with the largest proportion in most sample groups. Notably, the outcrop sections contain a large

1 amount of Fe₂O₃, but XRD analysis of some samples corresponding to these groups (JB-1 to JB-5,
 2 JB-7, AS-1, GQ-1, GQ-2) did not detect chlorite, which may have been decomposed into amorphous
 3 iron. In Fig. 4, the changes in CIA, CIW and PIA of PWR are also shown. High CIA and CIW values
 4 reflect relatively strong weathering in warm and humid climates; conversely, low CIA values reflect
 5 relatively weak weathering in cold and dry climates. PIA is suitable for judging the presence of only
 6 plagioclase in the host rock; in terms of the degree of weathering in provenance areas without
 7 potassium feldspar, the larger the PIA value is, the stronger the degree of weathering (Xu et al 2018;
 8 Zhao et al 2019). Fig. 4 shows that the values of CIA, CIW, and PIA are in the ranges of 6.83-64.64,
 9 7.01-70.15, and 3.82-20.38, respectively; the range variations indicate that the weathering degrees of
 10 the PWR samples obtained at different outcrops were quite different.

11 4 Engineering geological properties

12 4.1 Methodology

13 The physical and mechanical properties of PWR in the K₁/J₂ contact zone are tested in
 14 accordance with the “Standard for test methods of engineering rock mass” (China EC 2013). The
 15 tested physical and mechanical properties of PWR include the water content, bulk density, real density,
 16 porosity, longitudinal wave velocity, expansion rate, PLS and UCS. Rocks exhibit different
 17 engineering properties under different water content states (Majeed et al 2018). To gain a better
 18 understanding of the original properties of PWR, this paper tested the water content of the PWR
 19 samples in their natural state and calculated it by Eq. (1):

$$20 \quad w = \frac{m_o - m_s}{m_s} \times 100 \quad (1)$$

21 where w (%) is the water content of the rock, m_o (g) is the sample mass before drying, and m_s (g)
 22 is the sample mass after drying.

23 Density, defined as mass divided by volume, is the basic property of rock (Miller 1965) and is
 24 closely related to the geotechnical properties of the rock. The volume measurement includes or
 25 excludes the voids inside the rock, referred to as the bulk and real density, respectively. The bulk
 26 density is tested by the wax sealing method and calculated according to Eq. (2):

$$27 \quad \rho_{bulk} = \frac{m_s}{\frac{m_1 - m_2}{\rho_w} + \frac{m_1 - m_s}{\rho_p}} \quad (2)$$

28 where ρ_{bulk} (g/cm³) is the bulk density, m_s (g) is the drying mass of the bulk sample, m_1 (g) is the mass
 29 of the sealed test piece in air, m_2 (g) is the mass of the sealed test piece in water, ρ_w (g/cm³) is the
 30 density of water (0.998 g/cm³ at 20 e , and ρ_p (g/cm³) is the density of paraffin wax.

31 The real density is determined with the pycnometer method and is calculated according to Eq.
 32 (3):

$$33 \quad \rho_r = \frac{m_s}{m_1 + m_s - m_2} \times \rho_{WT} \quad (3)$$

34 where ρ_r (g/cm³) is the real density; m_s (g) is the drying mass of the powdered sample; m_1 (g) is
 35 the total mass of the bottle and test solution; m_2 (g) is the total mass of the bottle, test solution and
 36 rock powder; and ρ_{WT} (g/cm³) is the density of the test solution, which is the same as the test
 37 temperature.

38 Porosity is one of the important physical properties that control the mechanical properties of rock
 39 strength and deformability (Palchik and Hatzor 2002; Sabatakakis et al. 2008; Pappalardo 2015). The

1 total porosity (n) is the ratio of pore volume to sample volume and is calculated according to Eq.
2 (4):

$$3 \quad n = \left(1 - \frac{\rho_{bulk}}{\rho_r}\right) \times 100 \quad (4)$$

4 where n (%) is the total porosity, ρ_{bulk} (g/cm³) is the bulk density, and ρ_r (g/cm³) is the real
5 density.

6 The speed at which ultrasonic waves propagate in the rock depends on the density of the rock
7 and the pores and cracks in the rock (Khandelwal et al. 2010). A UTA-2000 intelligent ultrasonic
8 detector measuring instrument was used to test the longitudinal wave velocity V_p of the PWR; the
9 sampling frequency was 10 MHz, and the timing accuracy was 0.1 s. During the test, butter was
10 applied to both ends of the rock sample to better couple the acoustic sensor with the sample. V_p was
11 determined by dividing the length of the sample by the time required for the pulse to pass through the
12 sample and was calculated according to Eq. (5):

$$13 \quad v_p = \frac{L}{t_p - t_0} \quad (5)$$

14 where v_p (m/s) is the longitudinal wave velocity, t_p (s) is the propagation time of the longitudinal
15 wave in the direct transmission method, and t_0 (s) is the zero delay of the instrumental system.

16 Rock expansion refers to the expansion of rock volume with time under the action of external
17 factors. There are many external factors that affect rock swelling: water penetration, reduction in rock
18 external force, chemical substances (such as sulfide), frost action, etc. (Sun et al. 2013). The PWR in
19 the K₁/J₂ contact zone studied in this paper is located under a thick Cretaceous water body; the
20 thickness is mainly due to the expansion of the rock caused by the penetration of water. The
21 expansion rate of the PWR was investigated through free expansion rate testing of the rock and was
22 calculated according to Eq. (6):

$$23 \quad V_H = \frac{\Delta H}{H} \times 100 \quad (6)$$

24 where V_H (%) is the axial free expansion rate of the rock, ΔH (mm) is the axial deformation of the
25 specimen, and H (mm) is the height of the specimen.

26 The point load test equipment used is a YXDZ-10 rock point load tester (maximum pressure of 10
27 kN) and a Vernier caliper (Fig. 5). The sample was collected and prepared to avoid cracks. The size of
28 each irregular block specimen was 50 mm±35 mm, and the ratio of the distance between the two
29 loading points to the average width of the loading point was 0.3-1.0. In the test, the direction of the
30 minimum size of the test piece was selected as the loading direction. First, the instrument was checked
31 to ensure accurate alignment, then the center of the sample was placed close contact with the loading
32 cone. The pressure was adjusted to zero, and within 10-60 s, a load was applied at a constant speed
33 until the specimen broke. The pressure gauge reading F at the time of failure was recorded, and the size
34 of the broken surfaces of the specimen were measured.

35 According to the standard procedure of ASTM (1995), the formula for calculating the PLS of
36 rock without size correction is as follows:

$$37 \quad I_s = \frac{P}{D_e^2} \quad (7)$$

38 where P (kN) is the applied load at failure and D_e (mm) is the equivalent diameter defined. For the
39 diametral test, the calculation formula of equivalent core diameter D_e is as follows:

$$40 \quad D_e = (D \times D')^{0.5} \quad (8)$$

41 where D (mm) is the distance between two loading points and D' (mm) is the distance between
42 loading points at the moment of specimen failure after penetration of the upper and lower cone ends.
43 For axial and irregular lump tests, the calculation formula of equivalent core diameter D_e is as

1 follows:

$$2 \quad D_e = (4lW \times D' / \pi)^{0.5} \quad (9)$$

3 where W (mm) is the average width of the minimum section through two loading points.

4 When the equivalent core diameter is not 50 mm, the PLS index of rock should be corrected.
5 When there are many experimental data and the equivalent core diameter in the same group of
6 samples has multiple sizes not equal to 50 mm, the relationship curve between D_e^2 and failure load P
7 should be drawn based on the experimental results, and the corresponding P_{50} value when D_e^2 is 2500
8 mm² should be found from the curve. The PLS of the rock was calculated as follows:

$$9 \quad I_{s(50)} = \frac{P_{50}}{2500} \quad (10)$$

10 where $I_{s(50)}$ (MPa) is the PLS index of the rock with an equivalent core diameter of 50 mm and P_{50}
11 (kN) is the corresponding P value when D_e^2 is 2500 mm², as calculated from the P - D_e^2 relationship
12 curve.

13 Fig. 5 shows the definitions of L , D , D' , W_1 and W_2 (the widths at the upper and lower ends of
14 the cross-section). Note that the average width W is defined as $W = (W_1 + W_2) / 2$. The dimension L is
15 measured from the loading point to the nearest free face. The shape of the irregular lump samples
16 strictly followed the ASTM (1995) standard, which suggests that the D/W ratio should be between 0.3
17 $< D/W < 1$ and that L should be $\geq 0.5D$. If the measured L , D and W do not meet this requirement,
18 the test is invalidated, and the data are discarded. In addition, if the failure section does not pass
19 through the loading points, the specimen is invalidated. Furthermore, all the samples must be checked
20 to have an equivalent size greater than 10 times the mineral grains; otherwise, the test is also
21 invalidated.

22 The experimental instrument used in this test was a D-1000 electrohydraulic servo universal
23 testing machine, with a maximum axial load of 1000 kN (Fig. 6). During each test, the test piece was
24 placed in the center of the pressure plate of the testing machine, the ball seat was adjusted, and the
25 two ends of the test piece were in even contact with the upper and lower pressing plates of the testing
26 machine; then, loading at the rate of 0.5-1.0 MPa/s was performed until failure. The failure load (P)
27 and the phenomena that occurred during the loading process were recorded. The UCS of the rock was
28 calculated according to the failure load and the section area of the test piece. The calculation formula
29 is as follows:

$$30 \quad R_c = \frac{P}{A} \quad (11)$$

31 where R_c (MPa) is the rock UCS, P (kN) is the failure load, and A (mm²) is the cross-sectional area of
32 the rock specimen.

33 4.2 Physical-mechanical characterization

34 According to the above testing methodology, we obtained the physical and mechanical properties
35 of the PWR in the K_1/J_2 contact zone (Table 1). The water content of the PWR from the different
36 outcrops ranges between 3.86% and 10.23%, with an average of 7.66%. The bulk density is between
37 2.36 and 2.57, with an average of 2.43; the real density is between 2.61 and 2.69, which is relatively
38 uniform, with an average of 2.64. The porosity is between 4.11 and 10.27, and the average is 8.22.
39 The porosity is relatively high, which may be caused by weathering. The longitudinal wave velocity is
40 between 0.78 and 3.82, and the average is 1.95. The expansion rate is between 12.1 and 51.8, and the
41 average is 32.9.

1 According to the experimental results of the irregular sample PLS, the $P-D_c^2$ relationship curve
2 was drawn, and the $P-D_c^2$ linear regression fitting curve formula of each group of irregular samples is
3 shown in Table 2, which also shows the linear regression correlation coefficient r , the determination
4 coefficient r^2 , and the significance F value of the F test. Mathematical statistics theory (Chen et al.
5 2014) suggests that the correlation coefficient r expresses the linear correlation degree between X and
6 Y as two random variables, $-1 \leq r \leq 1$: when $-1 \leq r < 0$, X and Y are negatively correlated; when $r=0$, X
7 and Y are not correlated; and when $0 < r \leq 1$, X and Y are positively correlated. The coefficient of
8 determination r^2 indicates the degree of fit of a straight line and indicates how well the fitted straight
9 line can reflect the fluctuation in Y , $0 < r^2 \leq 1$: an r^2 value closer to 1 indicates a better fit. The F test
10 was used to determine the linear significance of all independent variables X to Y as a whole; the
11 significance F value is generally less than 0.05, and the smaller the value, the more significant it is
12 (0.05 is actually the significance level, which was artificially set). In Table 2, the JB-1 group data
13 show $r=0.95 > 0$, indicating that P and D_c^2 are positively correlated; $r^2=0.9$, which is close to 1,
14 indicating a good degree of fit; and the significance $F=5.29E-8 < 0.05$, indicating that for all the
15 samples, D_c^2 is linearly significant for P overall. Then, the P_{50} value corresponding to $D_c^2=2500$ mm
16 was calculated according to the linear regression fitting formula for each group (Table 2), and the PLS
17 index $I_{s(50)axi}$ of each group of irregular samples was calculated according to Eq. (4) (Table 1). The
18 obtained irregular sample PLS is between 1.13-6.61, with an average of 2.64.

19 According to Eqs. (7)-(9), the axial and diametral PLSs were calculated (Table 1). The axial
20 loaded PLS is between 2.65 and 13.41, with an average value of 5.97; the diametral loaded PLS is
21 between 1.78 and 8.38, and the average value is 3.96. Table 1 shows that for the same group of
22 samples with the same lithology, $I_{s(50)axi} > I_{s(50)dia} > I_{s(50)ir}$ because the direction of the axial point load is
23 approximately perpendicular to the sedimentary layer of the sample, and the loading direction of the
24 diametral point load is approximately parallel to the sedimentary bedding surface, which results in
25 $I_{s(50)axi} > I_{s(50)dia}$. This result is consistent with the research results of Basu et al (2010). The diametral
26 and axial PLS tests are performed in the dry state, and the irregular sample PLS is tested in the natural
27 moisture state. The PLS of the dry sample is higher than that of the water-containing sample (Jeng et
28 al. 2004; Deng et al. 2019; Majeed et al. 2018), and the experimental results are consistent with the
29 research results of Hawkins (1998). According to Eq. (5), the UCS of the PWR was calculated (Table
30 1) to be between 13.65 and 160.69. The degree of variation in UCS is large, which may be caused by
31 the variation in lithology, cement and degree of weathering; the average UCS is 73.61.

32 5. Discussion

33 5.1 Correlations between engineering geological and petrographic properties of PWR

34 Rocks are composed of mineral aggregates (including crystalline and amorphous minerals) with
35 a certain structure. Therefore, the mechanical properties of fresh rock mainly depend on the mineral
36 composition and relative content of the rock (Liu et al. 1999). Some scholars believe that rock
37 strength is linearly related to mineral composition and content (Fereidooni 2016; Wang et al. 2019;
38 Pappalardo et al. 2016), while others believe that rock strength is exponentially related to mineral
39 composition and content (Meng et al. 2007; Pappalardo et al. 2016). As shown in Fig. 3, each group
40 of samples contains quartz, while the other minerals present depend on the sampling location; this
41 variability is caused by the different degrees of weathering in different locations. This work attempted
42 to analyze the relationship between PLS, UCS and quartz content of PWR by linear regression and

1 exponential regression. The results show that the correlation coefficient of linear regression is higher
2 than that of exponential regression; the linear regression formula is shown in Table 3, and the PLS
3 and UCS of PWR are positively correlated with the content of quartz. According to the description of
4 their petrological characteristics, calcite, quartz, dolomite, and albite are brittle and high in hardness
5 and are collectively referred to as the BMG, while muscovite, clay minerals, and chlorite minerals are
6 elastic and low in hardness and are collectively referred to as the EMG. We analyze the correlation
7 between the strength of the PWR and the BMG and EMG contents, and the results are shown in Table
8 3. The PLS and UCS of the PWR are positively correlated with the BMG content and negatively
9 correlated with the EMG content. We also tested the expansion rate of the PWR, and the content of
10 clay minerals in the rock determines the PWR expansion rate. The expansion rate of the PWR is
11 linearly positively correlated with the content of the clay minerals (Table 3).

12 When the rock undergoes weathering, primary minerals will transform into secondary minerals;
13 at the same time, changes in the chemical composition and structure of the rock will lead to the
14 gradual destruction of the rock. The higher the degree of weathering of the rock is, the more serious
15 the degree of destruction, and the worse its mechanical properties. The main chemical elements of the
16 PWR were analyzed by XRF, and three weathering alteration indexes were calculated (CIA, CIW, and
17 PIA) (Fig. 4). We also analyzed the correlation between the strength of PWR and CIA, CIW, and PIA,
18 and the results are shown in Fig. 7. The PLS and UCS of the PWR are negatively correlated with the
19 three weathering alteration indexes, and the correlations between the PLS and the three weathering
20 alteration indexes are higher than those between the UCS and the three indexes.

21 5.2 Correlations between the mechanical and physical properties of PWR

22 The mechanical properties of rock are affected by its water content. Jeng et al. (2004) and Deng
23 et al. (2019) studied the influence of water content on the strength of sandstone and found that water
24 saturation and wetting will cause the strength of sandstone to decrease. Majeed et al. (2018) studied
25 the influence of water content on the strength of sedimentary rocks and found that the UCS of
26 saturated rock is 40%-50% lower than that of dry rock. Wei et al. (2020) studied the influence of
27 water content on the strength of gypsum and found that an increasing water saturation had a
28 weakening effect on the strength of gypsum. The results of the correlation analysis of the strength and
29 water content of the PWR in this paper are shown in Fig. 8. The PLS and water content of the PWR
30 are negatively exponentially correlated, UCS is negatively linearly related to the moisture content,
31 and an increasing water saturation reduces the strength of the PWR. Some scholars believe that the
32 correlation between rock strength and density is linear (Chatterjee 2002; Fereidooni 2016; Aligholi et
33 al. 2019), while others believe it is exponential (Wang et al. 2019). In this paper, the correlation
34 analysis results of the PWR strength and the bulk density and real density are shown in Fig. 8. The
35 PLS is positively exponentially correlated with the bulk density and real density, and the UCS is
36 positively linearly correlated with the bulk density and real density. Porosity is one of the important
37 parameters that controls rock strength (Pappalardo 2015). The PLS and total porosity of the PWR are
38 negatively exponentially correlated, and the UCS is negatively linearly related to total porosity (Fig.
39 8). These findings are consistent with previous research results (Chatterjee 2002; Fereidooni 2016;
40 Wang et al. 2019; Dincer et al. 2004). Some scholars have studied the relationship between rock
41 strength and longitudinal wave velocity, finding that some relationships are linear (Fereidooni 2016;
42 Pappalardo et al. 2016; Aligholi et al. 2019) while others are exponential (Wang et al. 2019). In this
43 paper, there is an positive exponential correlation between the PLS and the longitudinal wave velocity,

1 and the UCS has a positive linear correlation with the longitudinal wave velocity (Fig. 8).

2 5.3 Correlation between the PLS and UCS of PWR

3 This section describes the regression analysis conducted on the PLS and UCS data of the PWR
4 in the K₁/J₂ contact zone and the relationship between them, providing a theoretical basis with which
5 engineers can use point load tests to quickly estimate the UCS of PWR in practical engineering
6 applications. There are many studies on the correlation between rock PLS and UCS. Most scholars
7 believe that UCS and PLS are linearly related (D'Andrea et al. 1965; Broch et al. 1972; Gunsallus et
8 al. 1984; Hawkins 1998; Kahraman 2001; Palchik et al. 2004; Sabatakakis et al. 2008; Basu et al.
9 2010; Xiang 1981; Wei 1982; Li et al. 2013; Liu et al. 2019), and a few scholars believe that UCS and
10 PLS are exponentially related (Quane 2003; Robina 2015). We conducted a linear regression analysis
11 on the relationships between the UCS and the PLS, axial PLS, diametral PLS and irregular PLS, and
12 the results are shown in Fig. 9; all the relationships show a positive linear correlation.

13 Whether the UCS obtained from the transformation of the PLS test is accurate and reliable is
14 discussed next. In this paper, the relative error between the UCS obtained from the transformation of
15 the PLS test and the UCS of standard laboratory rock samples is used for verification. The definition
16 of the relative error is as follows (Wang 2015):

$$17 \quad \delta : \frac{|R_c - \sigma_c|}{\sigma_c} \times 100\% \quad (7)$$

18 where R_c is the UCS obtained from PLS test conversion and σ_c is the UCS of the standard laboratory
19 rock sample.

20 The calculated relative errors of UCS obtained by the PLS relation and ISRM method are shown
21 in Table 4. Using the method described in this paper, the relative error ranges between the UCS values
22 estimated by the PLS testing of samples (irregular, diametral, and axial) and the measured values
23 obtained by standard laboratory rock UCS testing are 0.38%-10.12%, 0.91%-13.22%, and
24 0.25%-7.03%, respectively; the average values are 3.13%, 6.02%, and 2.79%, respectively. Using the
25 ISRM method, the relative error range between the UCS of PWR samples estimated by the PLS
26 testing of samples (irregular, diametral, and axial) and the measured values obtained by the laboratory
27 standard rock UCS testing are 17.91%-101.42%, 13.89%-108.67%, and 0.15%-30.31%, and the
28 average values are 48.95%, 54.21%, and 9.83%, respectively. By comparison, the method in this
29 paper is more reasonable than the ISRM method, and the error of the estimated UCS is smaller, which
30 is more suitable for the transformation between the PLS and UCS values of the PWR in the K₁/J₂
31 contact zone of the Ordos Basin.

32 6. Conclusions

33 Experimental analysis of complete rock blocks is one of the main basic problems that need to be
34 solved in rock mass research (Bieniawski 1989). Therefore, a detailed description of the geological
35 characteristics of the rock blocks affected by a project is of great significance to the study of the
36 physical and mechanical properties of the rock mass.

37 For this research, laboratory tests were conducted on PWR samples from the K₁/J₂ contact zone
38 of the Ordos Basin to study the petrological and engineering geological characteristics of PWR,
39 analyze the impact of the petrological and physical characteristics of PWR on its mechanical
40 properties, and ensure the safe implementation of PWR engineering activities in the K₁/J₂ contact
41 zone in the Ordos Basin.

42 From the hand-scale analysis, the PWR is massive, with sandy gravel, fine-grained and muddy

1 structures, occasionally black and white cryptocrystalline, and can be observed as irregular rock
2 blocks, occasionally with fine cracks.

3 From a mineralogical point of view, 9 minerals, including quartz, calcite, dolomite, albite,
4 muscovite, kaolinite, illite, montmorillonite and chlorite were detected in the studied PWR, but only
5 quartz is constantly present. The presence of other minerals vary depending on the sampling location
6 and the degree of weathering. A total of 7 types of oxides were detected in the PWR; each group of
7 samples contained Fe_2O_3 , which had been decomposed into amorphous iron, and chlorite was not
8 detected in some groups of samples. The values of the three weathering alteration indexes (CIA, CIW,
9 and PIA) vary widely, indicating that the degree of weathering of the PWR varies greatly.

10 From a physical point of view, the real density of the PWR is relatively uniform. After long-term
11 weathering, the porosity is relatively high. The existence of clay minerals causes the PWR to have a
12 certain expansion rate.

13 From a mechanical point of view, the PWR is irregular, the axial and diametral PLSs are
14 relatively close, and for the same group of samples with the same lithology, $I_{s(50)axi} > I_{s(50)dia} > I_{s(50)irr}$.
15 The UCS of the PWR has a wide range. Notably, the UCS of each group of samples is determined in
16 accordance with the “Standard for test methods of engineering rock mass”. Therefore, the
17 experimental results and relationships can be considered reliable for the characterization of rock
18 types.

19 The mineralogical, physical and mechanical properties of the PWR were compared to predict
20 one parameter from another and study their mutual influence.

21 The relationships between the PLS and UCS of the PWR and the content of quartz minerals were
22 studied, and the results show positive linear trends. According to the petrological description, calcite,
23 quartz, dolomite and albite are collectively referred to as the BMG, and muscovite, clay minerals and
24 chlorite minerals are collectively referred to as the EMG. The PLS and UCS of the PWR are linearly
25 positively correlated with the BMG and EMG, respectively. Brittle minerals have greater hardness, so
26 the higher their content is in the PWR, the greater its PLS and UCS; the higher the content of elastic
27 minerals is, the lower the PLS and UCS. This result is a new aspect of rock behavior prediction. The
28 expansion rate of PWR is linearly positively correlated with the content of clay minerals. The PWR
29 swells when in contact with water and becomes muddy, which explains the water (mud) inrush
30 accidents of the coal mines in the study area. This PWR has experienced weathering for a long time,
31 and changes in mineral composition, chemical composition and structure have led to the deterioration
32 of its rock mechanical properties; the PLS and UCS of the PWR are negatively correlated with the
33 three weathering alteration indexes (CIA, CIW, and PIA).

34 The relationships between the strength of the PWR and its physical characteristics are studied.
35 The results show that the PLS is negatively correlated with the water content and total porosity and
36 positively correlated with the bulk density, real density, and longitudinal wave velocity. The UCS has
37 a negative linear correlation with the water content and total porosity, as well as a positive linear
38 correlation with the bulk density, real density and longitudinal wave velocity.

39 The relationship between UCS and PLS was studied, and the results show that UCS and PLS
40 have a positive linear correlation; additionally, the axial and diametral PLSs and irregular PLS also
41 show a positive linear correlation. This research result provides a theoretical basis for engineers to
42 quickly estimate the UCS and PLS of regular samples by using irregular rock block point load tests in
43 practical engineering applications.

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- 1 **List of figure and table captions:**
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13 **Table 1 Physical and mechanical properties of the PWR in the K_1/J_2 contact zone**

Sample groups	Water content (%)	Bulk density (g/cm ³)	Real density (g/cm ³)	Total porosity (%)	V_p (Km/s)	Expansion rate (%)	$I_{s(50)}$ (Mpa)			UCS (Mpa)
	mean/times						irregular	axial	diameter	
JB-1	7.06/5	2.41	2.63	8.37	2.83	—	2.82	7.78	4.45	105.31
JB-2	8.04/6	2.36	2.62	9.92	1.35	—	1.19	4.02	2.23	15.77
JB-3	8.66/6	2.36	2.61	9.58	1.23	—	1.15	3.42	2.16	14.87
JB-4	7.51/5	2.44	2.64	7.58	2.15	12.1	2.76	7.78	4.79	105.52
JB-5	5.97/6	2.49	2.66	6.39	2.24	15.6	3.69	8.37	6.49	135.69
JB-6	9.09/7	2.36	2.63	10.27	0.88	18.4	1.07	2.65	1.78	14.35
JB-7	4.81/6	2.55	2.69	5.21	3.42	—	4.23	11.02	6.89	139.77
JB-8	3.86/5	2.57	2.68	4.11	3.82	12.4	5.28	13.41	8.38	160.69
JB-9	7.15/5	2.39	2.62	8.78	1.96	50.5	2.1	4.81	3.49	86.2
JB-10	6.91/6	2.45	2.65	7.55	2.74	29	2.56	5.36	4.09	100.99
JB-11	7.32/5	2.4	2.63	8.75	1.83	48.1	1.81	4.19	2.99	71.85
JB-12	7.86/5	2.45	2.67	8.24	2.38	41.2	2.22	4.87	3.68	90.58
AS-1	8.89/6	2.43	2.65	8.31	1.71	25.4	1.86	4.25	3.39	81.98
GQ-1	9.32/6	2.39	2.63	9.13	1.57	31.5	1.75	4.15	2.96	62.13
GQ-2	10.23/6	2.38	2.62	9.16	1.25	44.3	1.13	3.75	2.41	14.56
SM-1	9.56/5	2.35	2.60	9.62	0.97	51.8	1.06	3.68	2.26	13.65
SM-2	9.15/6	2.4	2.63	8.75	0.78	49.8	1.35	3.94	2.63	37.35

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1 **Table 2 The $P-D_e^2$ regression equations and coefficient of determination**

Sample groups	$P-D_e^2$ regression equation	r	r^2	Significance F
JB-1	$P=0.2151+0.0027 D_e^2$	0.95	0.9	5.29E-08
JB-2	$P=0.0282+0.0012 D_e^2$	0.83	0.69	6.22E-05
JB-3	$P=-0.3143+0.0021 D_e^2$	0.89	0.79	2.47E-05
JB-4	$P=0.3775+0.0026 D_e^2$	0.91	0.83	6.15E-06
JB-5	$P=0.2851+0.0036 D_e^2$	0.87	0.76	2.62E-05
JB-6	$P=0.0285+0.0011 D_e^2$	0.95	0.91	3.86E-08
JB-7	$P=0.5816+0.004 D_e^2$	0.96	0.92	5.46E-09
JB-8	$P=1.4681+0.0047 D_e^2$	0.9	0.81	4.65E-06
JB-9	$P=2.3231+0.0012 D_e^2$	0.92	0.85	9.23E-07
JB-10	$P=0.0728+0.0025 D_e^2$	0.87	0.76	2.03E-05
JB-11	$P=-2.3877+0.0028 D_e^2$	0.92	0.84	1.68E-06
JB-12	$P=0.4701+0.0020 D_e^2$	0.92	0.85	9.23E-07
AS-1	$P=0.1258+0.0018 D_e^2$	0.85	0.72	5.36E-06
GQ-1	$P=0.8564+0.0014 D_e^2$	0.91	0.83	3.24E-07
GQ-2	$P=0.7496+0.0008 D_e^2$	0.87	0.76	2.65E-06
SM-1	$P=0.0598+0.0010 D_e^2$	0.88	0.77	1.84E-07
SM-2	$P=1.0247+0.0008 D_e^2$	0.92	0.85	8.32E-05

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4 **Table 3 Regression equations and determination coefficient of PWR strength and mineral**
 5 **content**

Regression equation	r	r^2	Significance F
$I_{s(50)irr}=1.46+5.35 Qu$	0.6	0.36	2.13E-03
$I_{s(50)axi}=3.49+15.34 Qu$	0.64	0.41	3.89E-03
$I_{s(50)dia}=2.68+7.88 Qu$	0.59	0.35	1.24E-03
$UCS=45.33+193.61 Qu$	0.66	0.44	4.53E-03
$I_{s(50)irr}=3.23-1.78 EMG$	-0.67	0.45	4.68E-03
$I_{s(50)axi}=8.60-5.12 EMG$	-0.71	0.51	1.23E-04
$I_{s(50)dia}=5.34-2.70 EMG$	-0.66	0.44	4.45E-03
$UCS=107.56-60.64 EMG$	-0.64	0.41	3.12E-03
$I_{s(50)irr}=1.45+1.78 BMG$	0.67	0.45	4.67E-03
$I_{s(50)axi}=3.48+5.12 BMG$	0.71	0.51	1.31E-04
$I_{s(50)dia}=2.64+2.70 BMG$	0.66	0.44	4.59E-03
$UCS=46.93+60.64 BMG$	0.62	0.39	2.98E-03
$ER=0.03+0.70 CM$	0.95	0.9	3.85E-07

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Figures

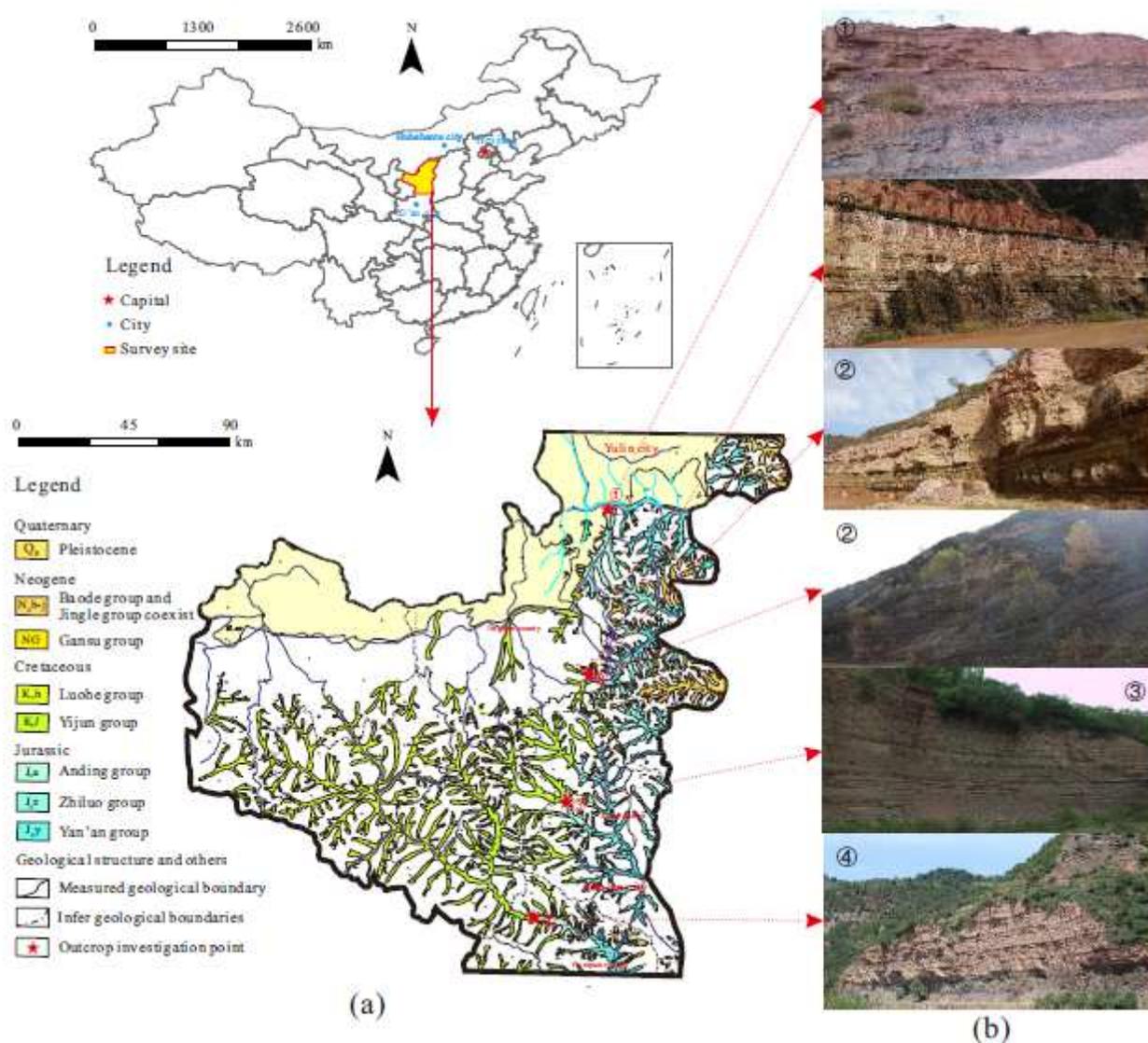


Figure 1

Location of study area: a Field outcrop survey sites, b Typical outcrop profile. Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its authorities, or concerning the delimitation of its frontiers or boundaries. This map has been provided by the authors.

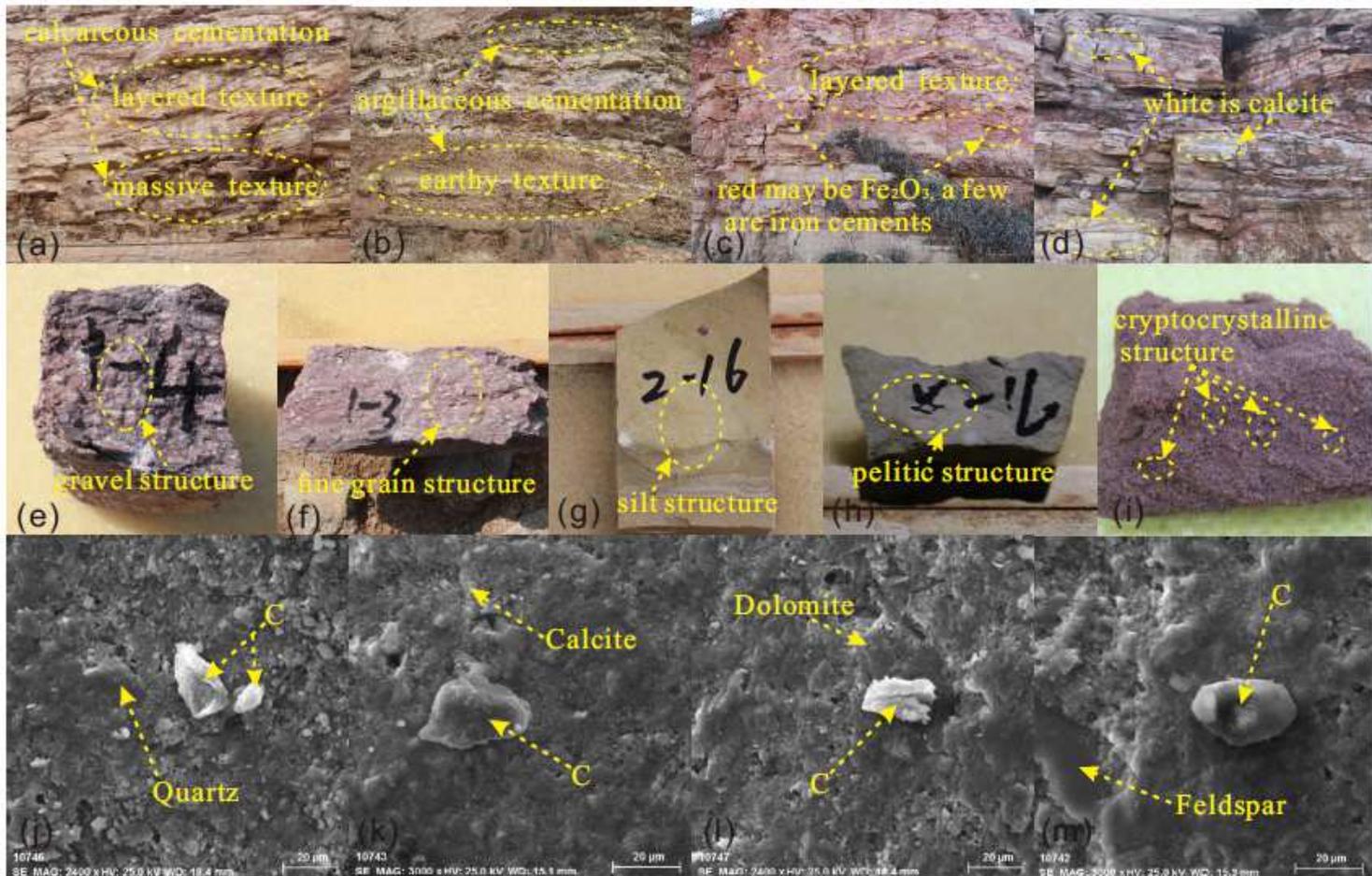


Figure 2

photos of the PWR macroscopic and microscopic features.

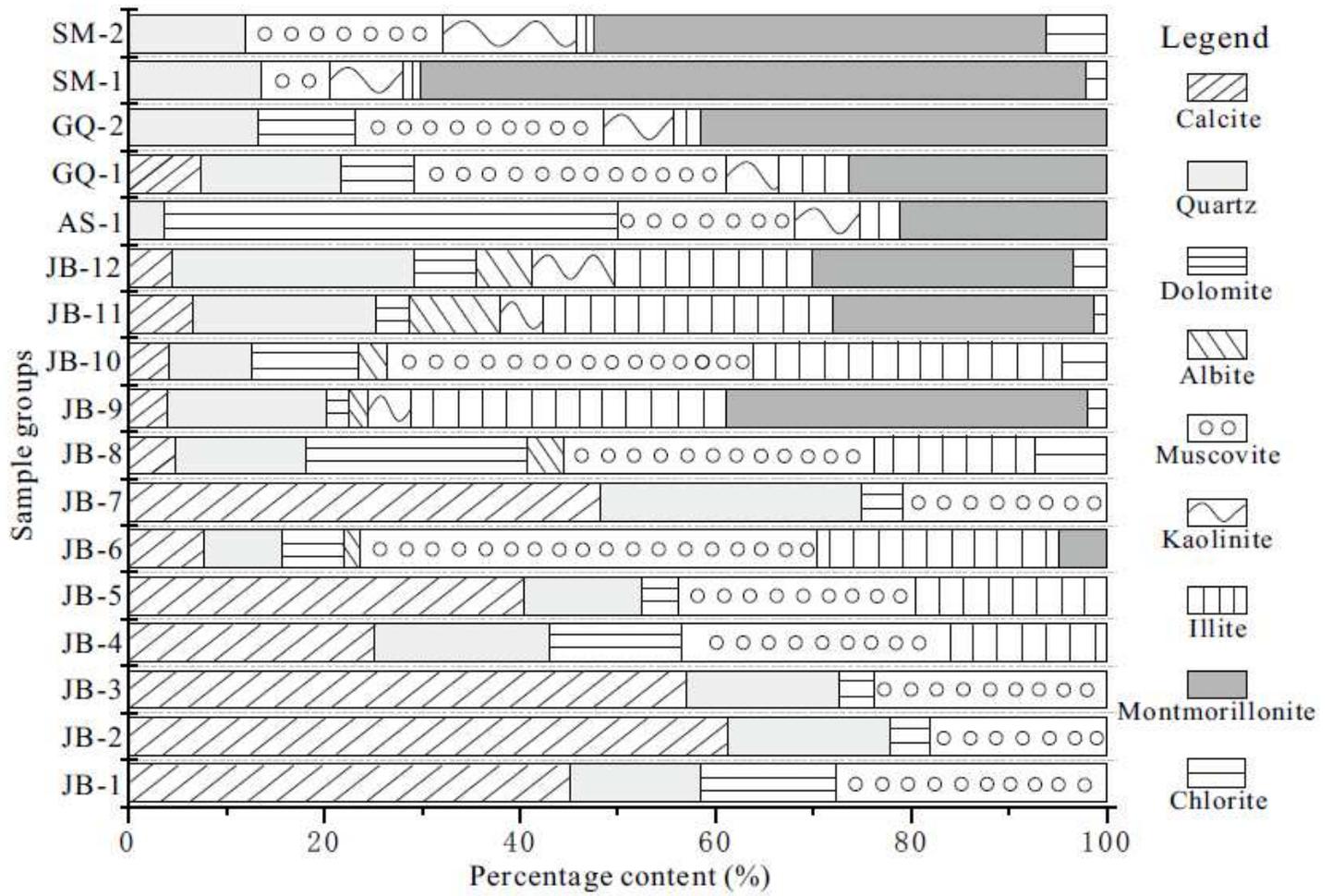


Figure 3

Mineral compositions of the PWR.

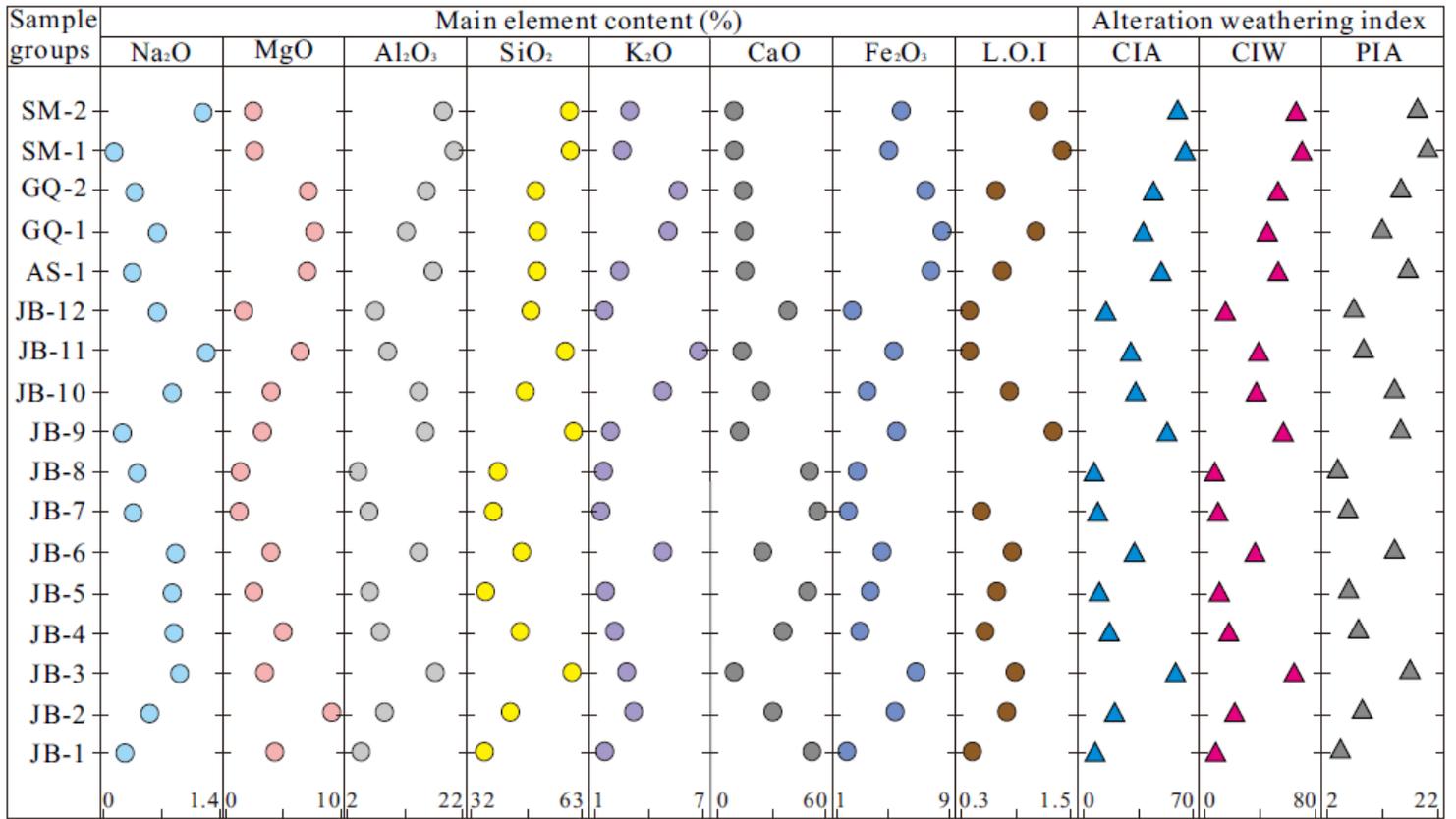


Figure 4

Main element contents and alteration weathering index of the PWR.

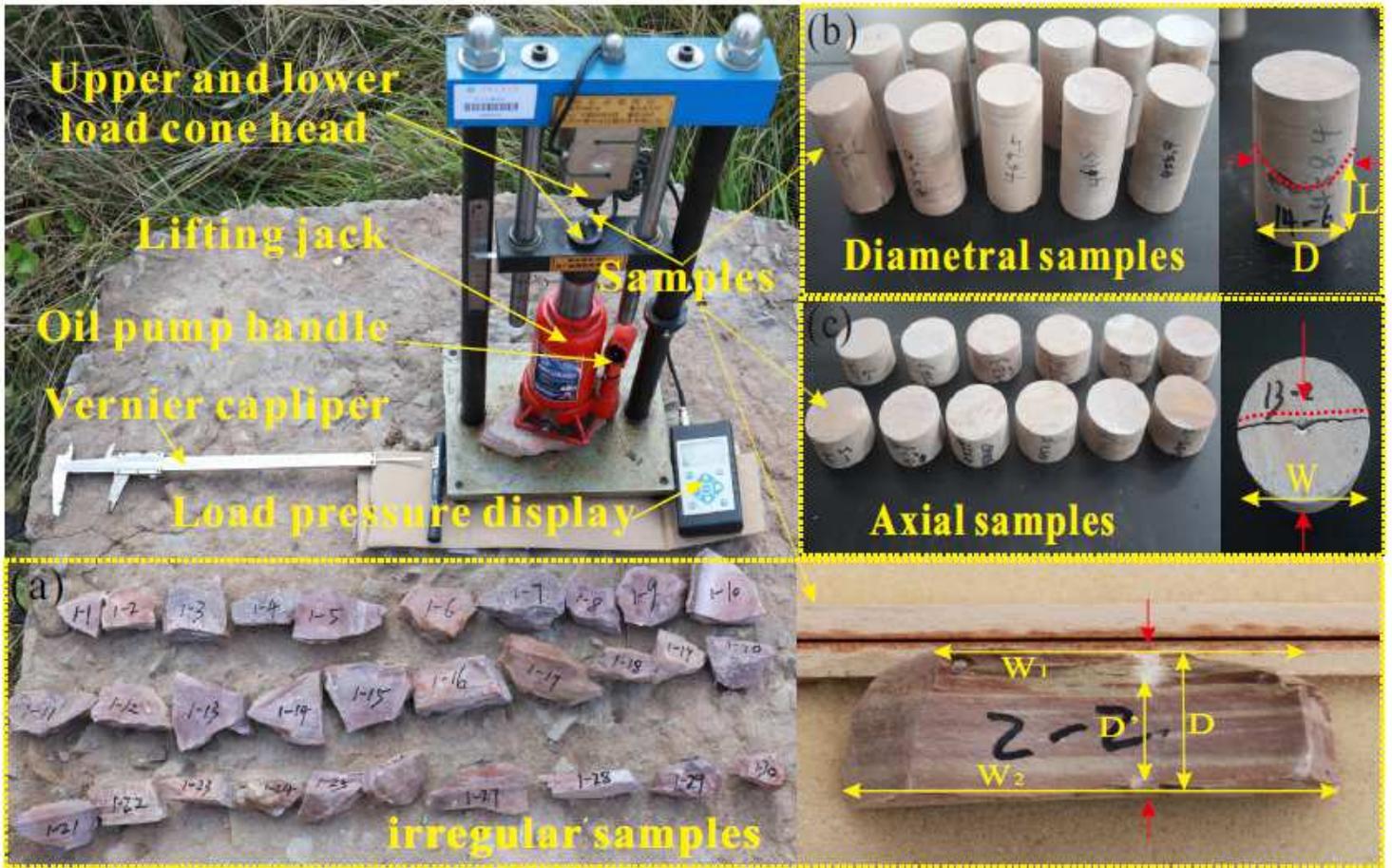


Figure 5

Point load instrument (YXDZ-10) and the definitions of L, D, D', W1 and W2.

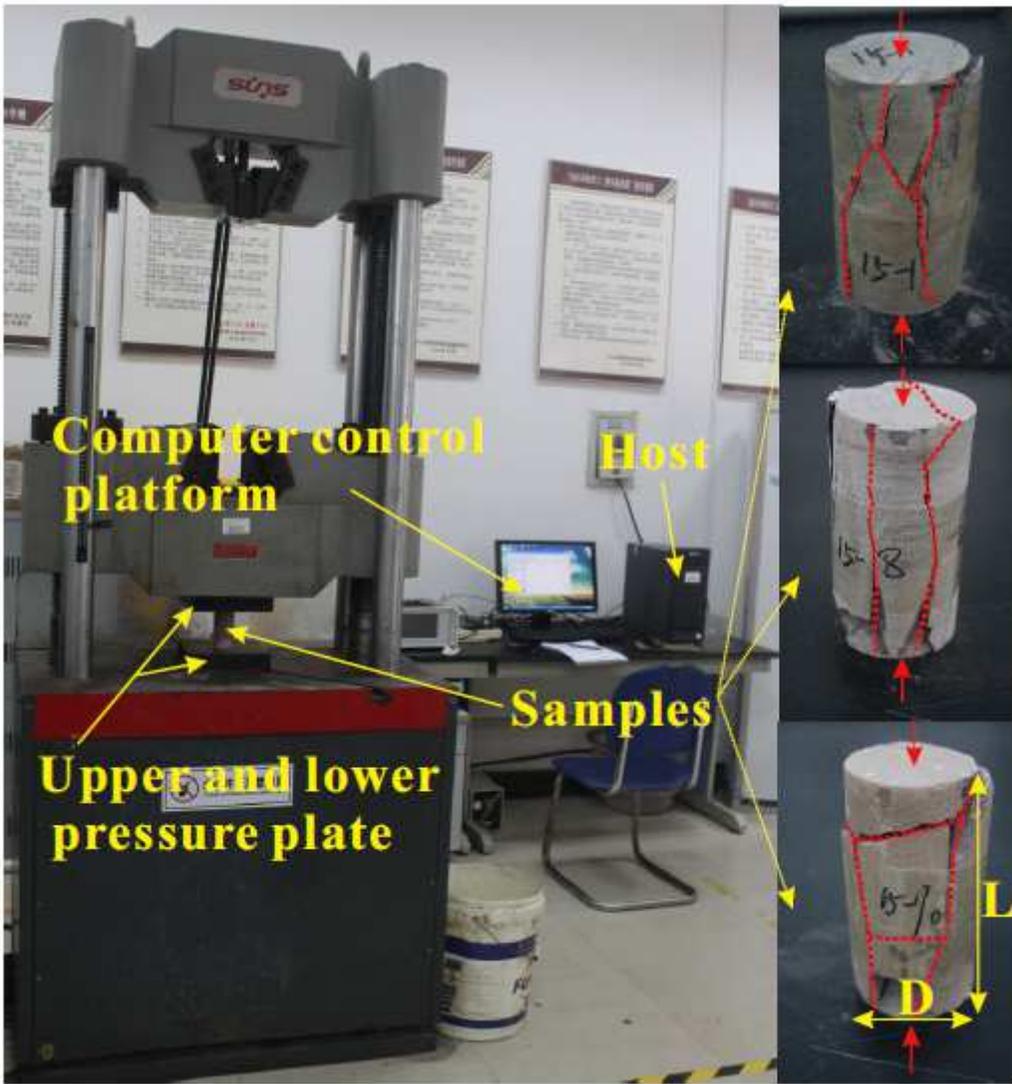


Figure 6

D-1000 electro-hydraulic servo universal testing machine.

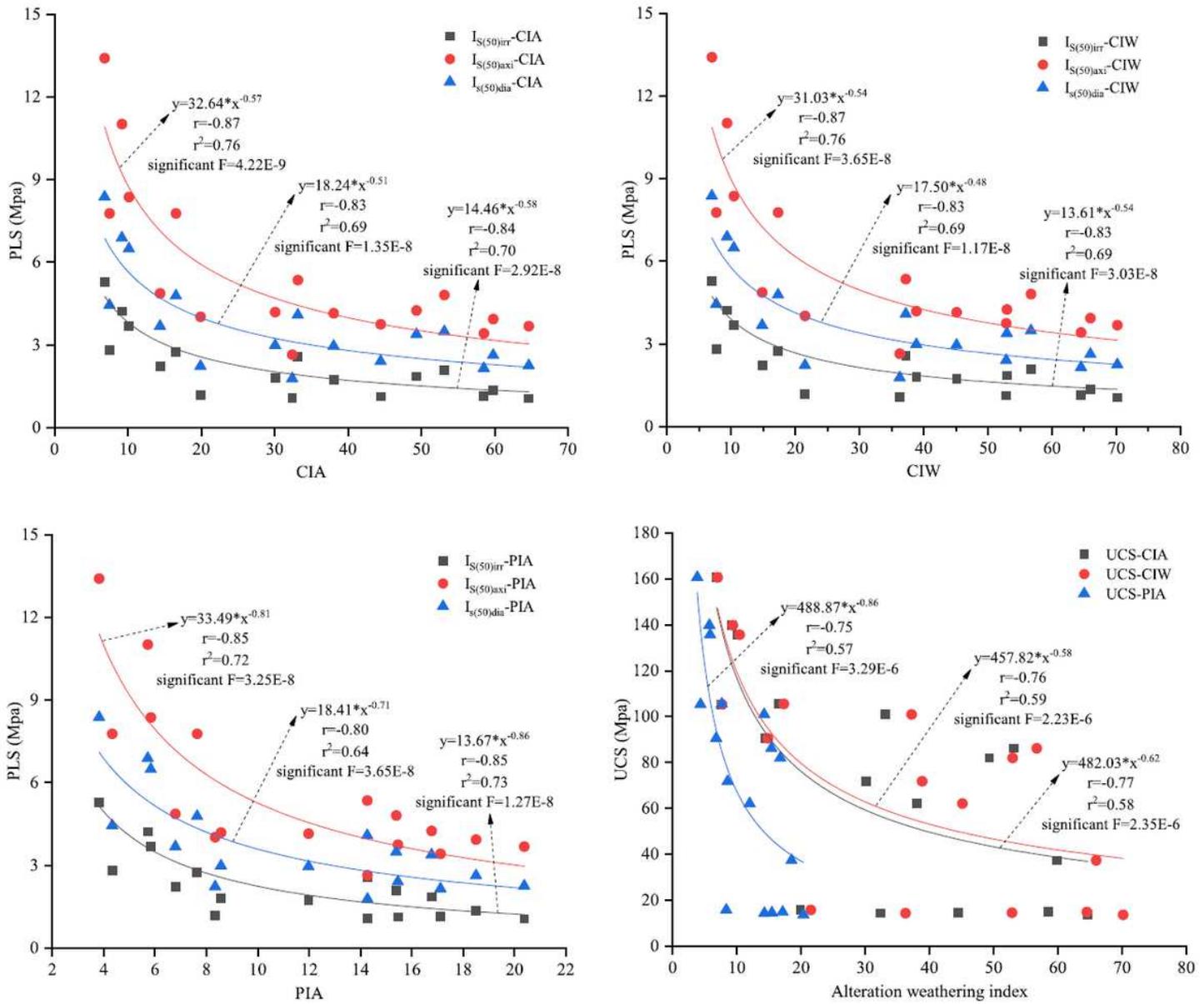


Figure 7

Relationships between strength of PWR and CIA, CIW, and PIA.

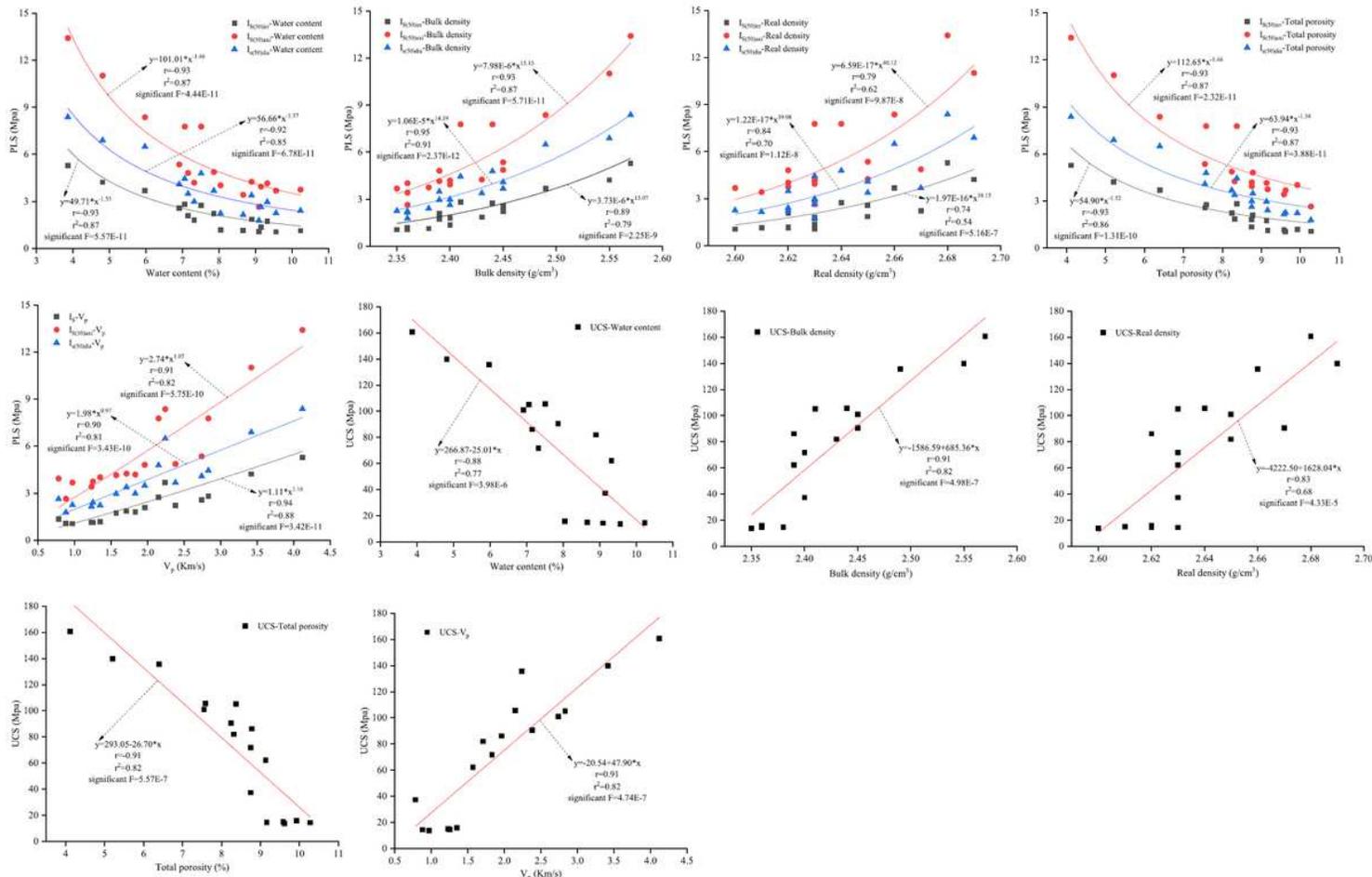


Figure 8

Relationships between strength of PWR and physical properties.

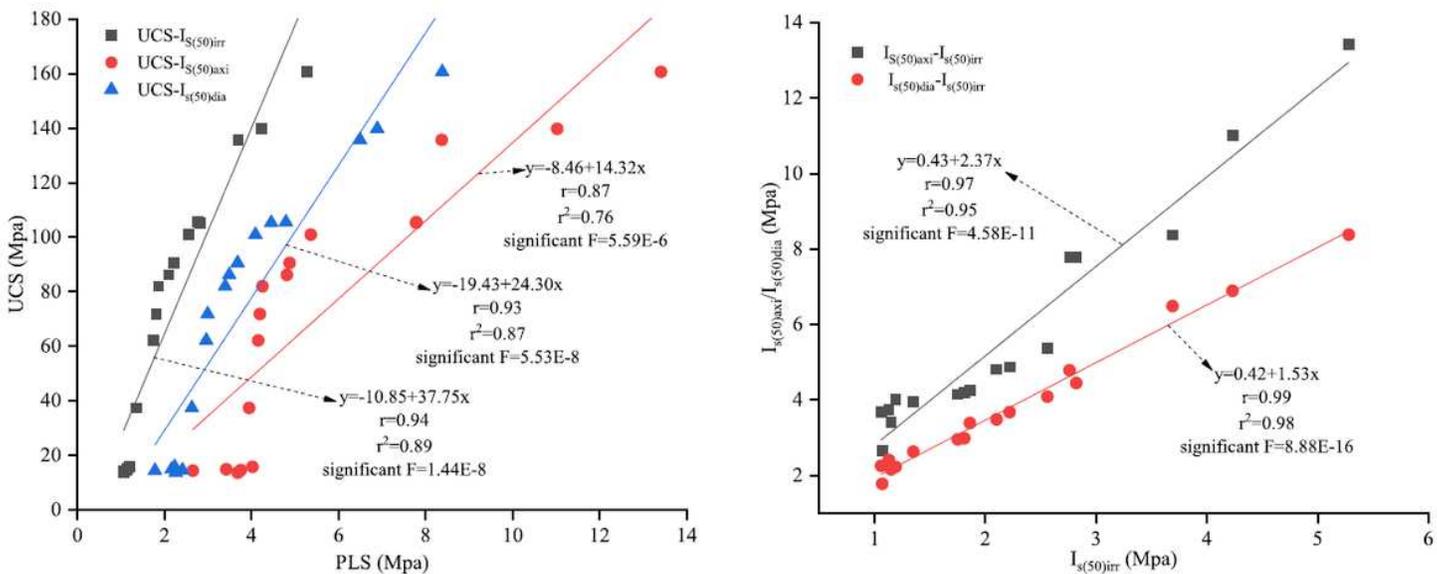


Figure 9

Relationships between the UCS and PLS, axial PLS, diametral PLS and irregular PLS.