

Delineation of the groundwater potential zone in Kantli River Basin, Jhunjhunu District, Rajasthan: A Geospatial Approach

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Abstract

Remote sensing and GIS are advanced techniques and tools that have been used for various researches including groundwater geology. The Landsat and the IRS satellite datasets have been utilized to extract information on the hydrogeomorphic and groundwater related features of a semi-arid Precambrian hard rock terrain in the Kantli River Basin of Jhunjhunu District, Rajasthan, India. It is the upper part of the Shekhawati River. The study area is covering about 2313.2013 km² in the Jhunjhunu District, Rajasthan. In the present study, the groundwater potential zone is identified by the GIS overlay techniques using the spatial analyst tool in ArcGIS 10.2. The developed methodology is demonstrated in the Kantli River Basin of Rajasthan, western India. Originally, nine thematic layers, viz. topographic elevation, land slope, geomorphology, geology, soil, pre and post-monsoon groundwater depths, annual net recharge, annual rainfall and proximity to surface water bodies were considered in this study. Therefore, five groundwater potential zones were identified and distinguished in the study area, viz. 'very good', 'good', 'moderate', 'poor' and 'very poor' based on groundwater potential index values, which will enable the local bodies for the future planning and management of the groundwater resource. Severe groundwater contamination has been found occasionally in the study area. Every year during the summer, the region is facing a lot of problems with portable groundwater. As the study area is semi-desert, the influences of salinity have been increasing day by day in the groundwater.

Keywords: Groundwater, Remote Sensing, Kantli River, Satellite Images, GIS, Potential zone.

Introduction

Groundwater is a precious natural resource, which provides life support to all natural beings (Agarwal et al., 2013). It is gradually declining due to rampant and unscientific use by the growing population. Over the past few decades, groundwater is being enormously exploited for drinking water in both urban and rural areas (Das et al., 2010, 2012, 2013, 2015; Nandi et al., 2017). The planning and development of any administrative area can be feasible through groundwater management (Anuradha et. al., 2010). Remote Sensing and GIS technology can facilitate to explore the groundwater sources in urban areas, especially for drinking water (Krishnamurthy et al., 2000; Khan et al., 2006). Remote sensing, GIS and field data play the most significant role in the detection and mapping of groundwater. Groundwater crisis in arid and semi-arid regions is a regular seasonal event. Over the past few decades, the international

scientific community has shown a key interest in the extensive application of remote sensing and GIS techniques in the delineation of groundwater potential zones (Singh and Prakash, 2003; Sikdar et al., 2004; Kusky and Gad, 2006; Sultan et al., 2008; Dinesh Kumar et al., 2007; Jasrotia et al., 2007; Sreedhar et al., 2009; Chowdhury et al., 2009; Chowdhury et al., 2010; Jha et al., 2010; Mohammed-Aslam et al., 2010; Agarwal et al., 2013; Abdalla, 2012; Adhikary et al. 2021; Patra et al., 2021; Mahanta et al., 2020). The present investigation is an attempt to explore the possibility and prospect of groundwater in the Kantli River Basin. Various hydrogeological parameters have been applied to identify the groundwater potential zones. The occurrence and movement of groundwater depend on several factors, namely, the depth of porous sediments, the presence of diverse morphology, surface hydrology, climate, physiography, geology, soil, and lineaments (Sreedhar et al., 2009; Abdalla, 2012; Patel et al., 2014; Goswami and Pati, 2008; Das et al., 2016, 2020; Hota and Goswami, 2018; Singh et al., 2022). The main objective of the present study is to evaluate the groundwater controlling features and identify the potential zones through geospatial techniques.

Study Area

The Kantli River basin is situated in the north-eastern parts of Rajasthan. It extends from 27° 38' 32" N to 28° 22' 14" N latitudes and 75° 21' 09" E to 76° 02' 31" E longitudes. It covers an area of 2313.2013 Km² in the central part of the Jhunjhunu District (Fig. 1). The eastern parts of the Kantli River basin are the part of Aravalli Mountain chain.

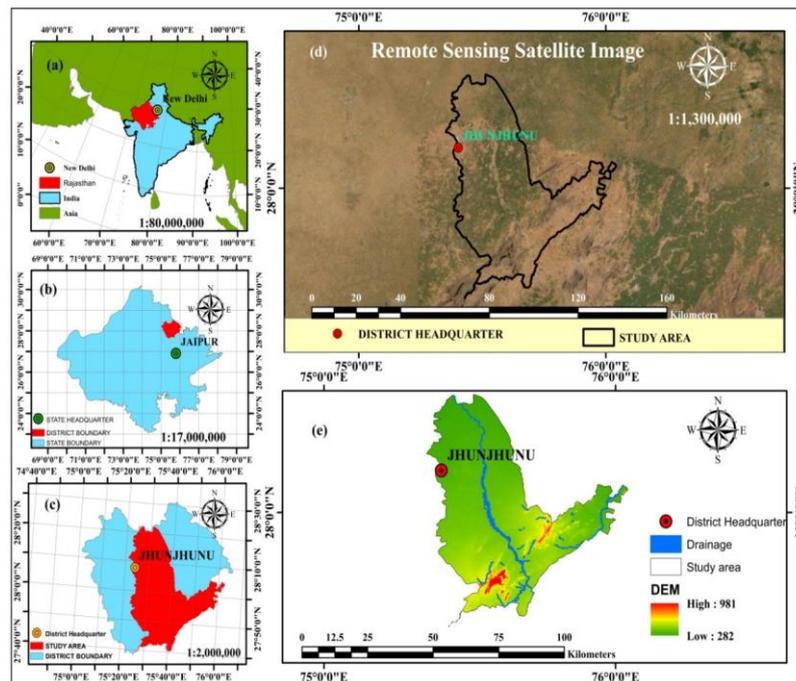


Fig.1 Study area of the upper Shekhawati River basin, India, Rajasthan, Jhunjhunu (a, b, c). Remote sensing images of the study area (d) and DEM topography of the Study area (e)

Geology of the study area

Geologically the north-northwestern part of the study area is covered by rocks of the Quaternary period and the south-eastern region is occupied by igneous rocks of the Precambrian period. The distribution of various litho-groups and their spatial extension is depicted in Table 1. The study area is occupied by the Quaternary group of rocks (89.59%), extrusive Malani Plutonic Group (0.40%) and mafic volcanic rocks of Delhi Supergroup (Alwar and Ajabgarh groups) (10.01%) (Sinha Roy, 1984; Sarkar and Dasgupta, 1980; Roy and Jakhar, 2002; Dasgupta, 1968; Pandey et al. 2010). The various geological data are derived not only from SAC and world geology USGS but also from Landsat ETM+ satellite imagery which played a significant role in map preparation. The geological map of the study area is presented in Figure 2.

Table 1 Distribution of Geology

Geological history	Geological group	Area (%)	Area sq.km
Undivided Precambrian	Malani Plutonic	0.4	9.25
	Delhi Group	10	231.32
	Alluvium and windblow		
Quaternary	Sand	89.6	2072.63

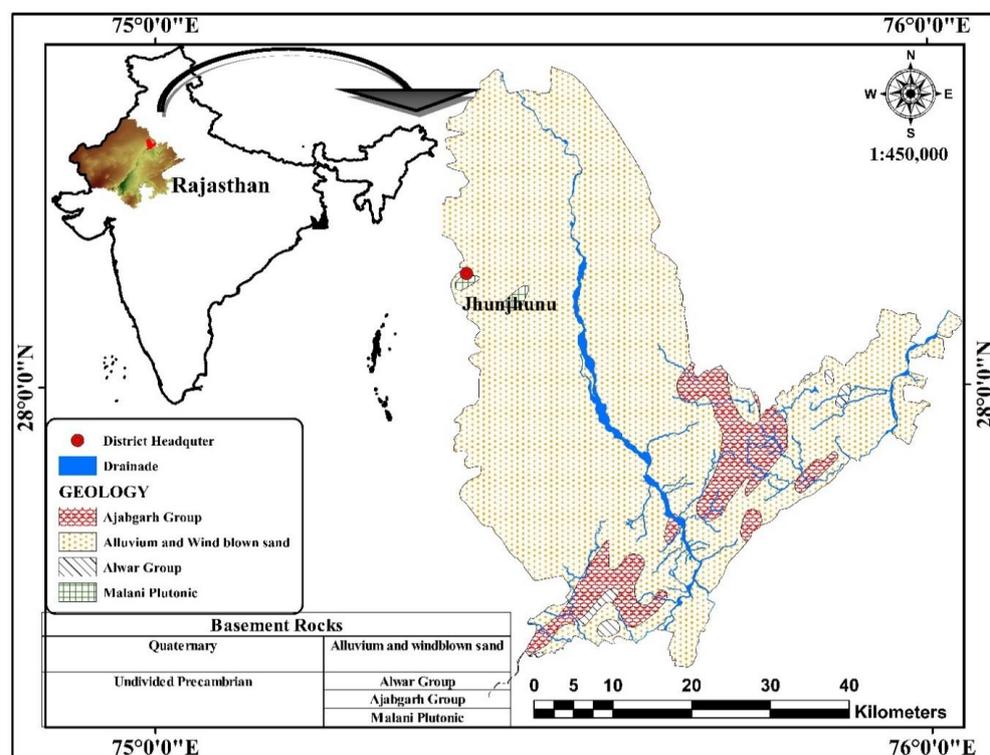


Fig. 2 Geological map of the study area

Hydrogeology of the study area

In the study area, the unconfined aquifer consists of two water-bearing formations, namely the Quaternary alluvial and the Precambrian hard rock formation. The shallow alluvial aquifer is filled with sand, silt and clay, and is observed below a depth of 50 meters in the phreatic region. Most shallow unconfined conditions are observed within the depth of 10 meters below the ground level (Ibrahim et al., 1997; Abu El-Ela et al., 1997; Sander et al, 1997). The direction of groundwater flow is decreasing from east to west. A remarkable variation in the water table has been observed in the pre-monsoon and post-monsoon periods. Most of the open wells dry out during the summer season. The hydrogeological data of the study area is given in Table 2. The confined aquifer is located in between 10-20 (m) thicknesses that occupied an area of 775.88 km² within the study area. An unconfined aquifer is found in between 0-20 (m) thicknesses that covered an area of 1923.15 km² in the investigated area. The top bedrock depth is of 80-100 m-bgl depths covering an area of about 1091.93 km² of the entire study area. The hydrogeological map of the study area is presented in Figure 3. The unconfined aquifer is available at very shallow depths (10-20 m) to the northwest of the study area. At the lowest depths, the confined aquifer is found in the north portion of the study area and the deepest depth of the bedrock is found in the north and west parts of the study area. To know the hydrology of the study area the field data has been taken into consideration during analysis.

Table 2 Hydrogeological condition of the study area

Aquifer				Depth of Bed Rock	
Confined		Unconfined		m-bgl	Area (sq.km)
Thickness (m)	Area (sq.km)	Thickness (m)	Area (sq.km)		
<10	433.27	<5	1149.15	20-40	105.86
10-20	775.88	5-10	274.80	40-60	401.62
20-30	304.01	10-20	489.20	60-80	307.30
30-40	83.42	20-30	137.36	80-100	1091.93
40-50	18.47	30-40	53.07	100-120	118.99
50-60	7.69	40-50	17.24		
60-70	1.49	50-60	3.97		
		60-70	0.93		

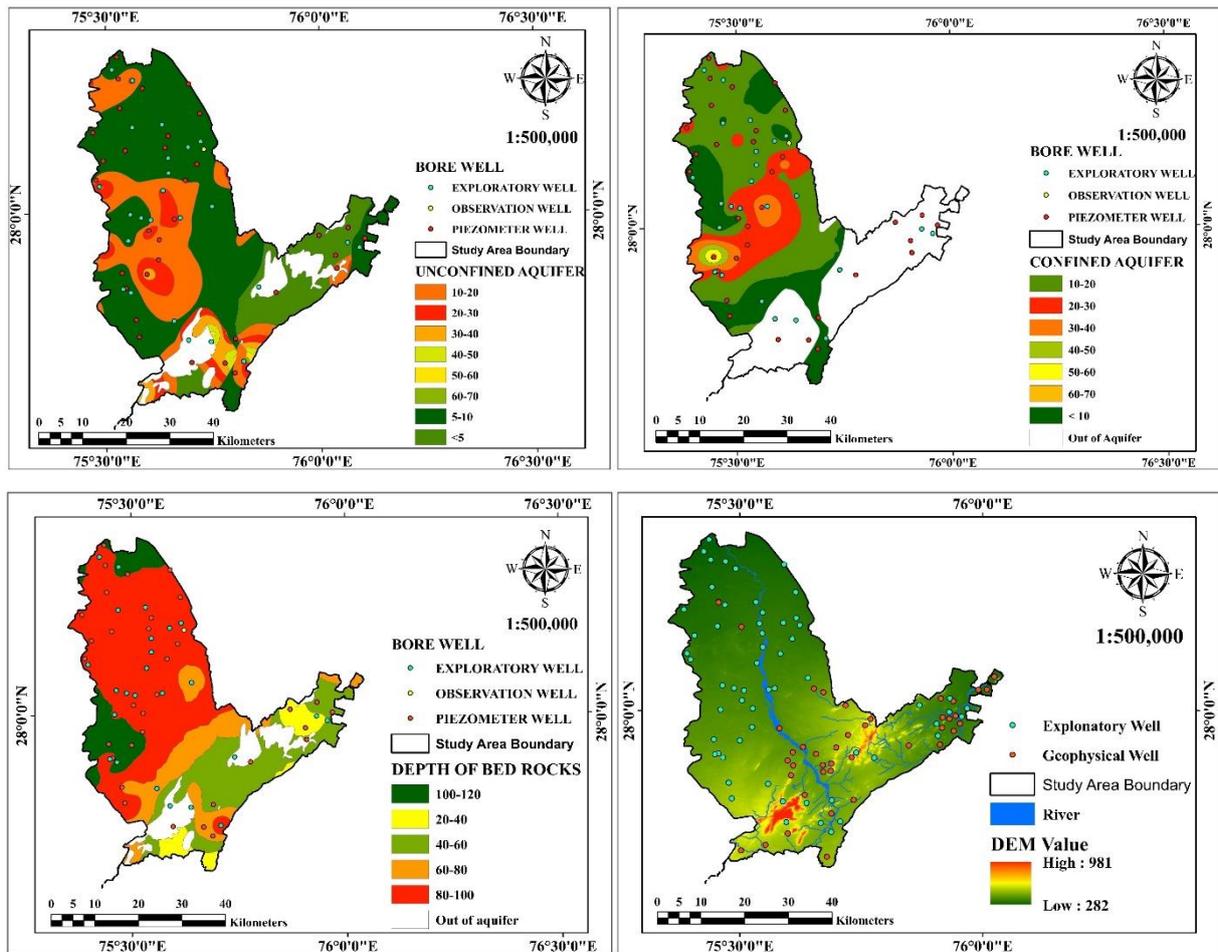


Fig. 3 Hydrogeological map of the study area, (a) unconfined aquifer, (b) confined aquifer, (c) depth of bedrock and (d) well with DEM

Methodology

In the present investigation, various parameters have been used to identify the suitable areas for groundwater prospective zones; viz. (1) Geology, (2) Geomorphology, (3) Soil, (4) Land Use and Land Cover, (5) Drainage, (6) Slope and (7) Lineament. The drainage and channel layers are generated from the survey of India toposheet in 1:50000 scale. The output layers have been updated with Landsat-4, Landsat-5, Landsat-7 and Landsat-8 satellite imagery. The vector files have been created to digitize and analyze the raster image (Strahler, 1964). The order of every stream is very systematically and sequentially observed in the study area. The digital elevation model (DEM) has been developed and used for elevation studies and comparative analysis of drainage patterns. The digitization and density mapping of all these layers have been prepared for the vector data analysis. It has been done by weighted overlay analysis with the help of ArcGIS software. This map has been scientifically used for the interpretation of groundwater occurrence in the study area and to identify the additional

recharge structure. Field observation has been undertaken to validate the map (Ibrahim et al., 1997; Abu El-Ela et al., 1997; Krishnamurthy et al. 1996; Ganapuram et al.2009). The borehole data is taken directly from the field with the help of the instrument and its validation accuracy is used to identify the groundwater potential zone. A base map of the study area is prepared with the help of Landsat series of TM, MSS, and ETM+, etc. The ERDAS Imagine 2014, Intergraph software was used for digital image processing. SRTM data (Shuttle Radar Topography Mission) of the study area were obtained and processed to construct a continuous DEM. The methodology adopted in the study area to locate the groundwater potential zone is depicted in Figure 4.

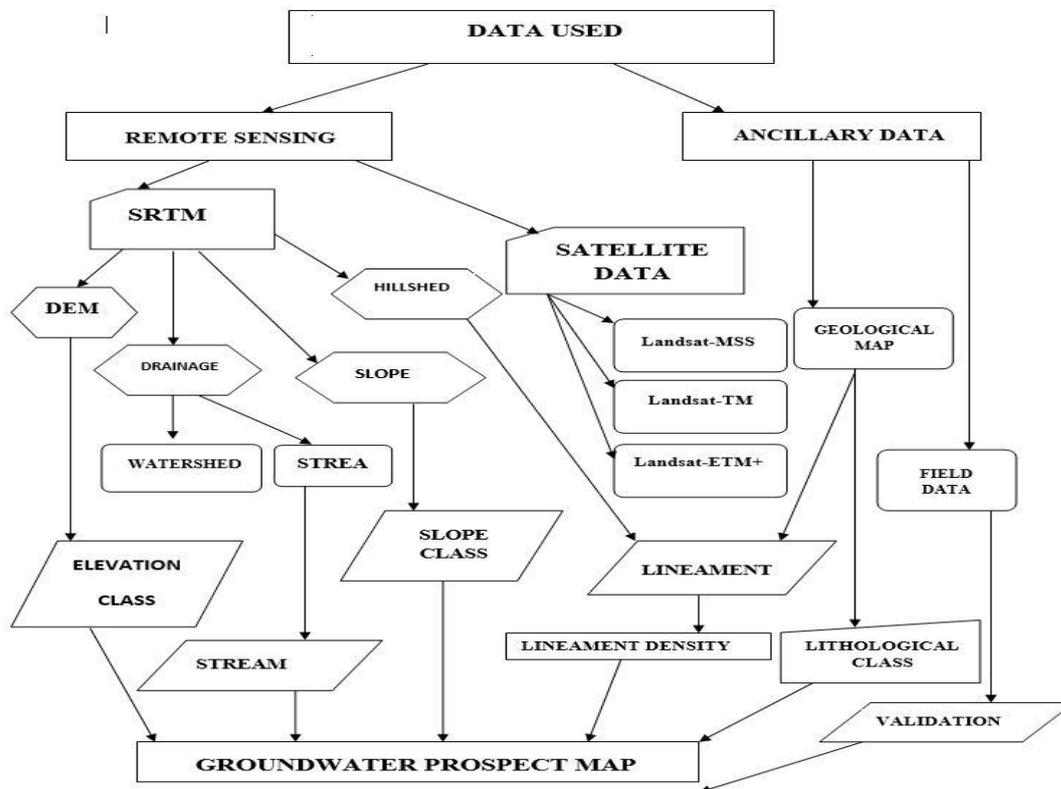


Fig. 4 Flow chart of the study area

Results and Discussions

Topographical and geological maps have been used to delineate the groundwater potential zones in the study area. The groundwater level in the study area depends on the following important factors:

- (1) Slope variation which controls the runoff of the rainwater, as a gentle slope increases the rate of infiltration of the surface water whereas a steep slope enhances the surface runoff;
- (2) Stream network which affects runoff and groundwater recharge;
- (3) Lineament as it is significantly enhancing the permeability and porosity of the rocks and affects the vertical percolation of water to recharge the local aquifers;

(4) Lithology which determines the penetration capacity of soils and opens the rock layers that manage the runoff and storage of groundwater;

(5) Topography

Each of the topographic layers is divided into different classes that are controlled by hydrogeological properties. The weighted over-analysis is done based on the relative importance of the potential of the groundwater level. The weighted over-analysis is performed by GIS for the preparation of the final groundwater potential map (Eastman et al., 1995; Voogd, 1983).

Topography Elevation

A topographical map of the study area has been created from SRTM DEM data. DEM map is depicted in Figure 5 to get an accurate idea of the surface elevation of the study area. The surface topography determines the direction of runoff over the surface. As per the DEM data, the lowest and highest elevation of the study area is 282 and 982 meters MSL respectively. Multiple irregular slopes are noticed in the study area. The slope of the northern part is relatively low and the amount of slope in the southern part is fairly high. The topographical maps are mainly classified into three numerical categories according to the characteristics of the slope (Fig. 6). The topographic features have explained the surface hydrology as well as the groundwater hydrological capabilities. The topographic elevation details are illustrated in Table 3.

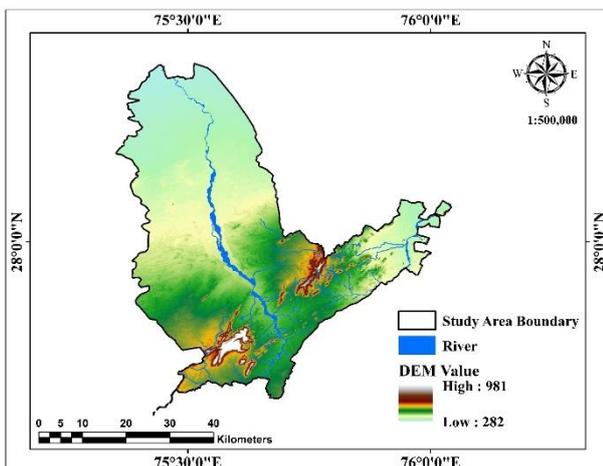


Fig. 5 Digital Elevation Model of the study area

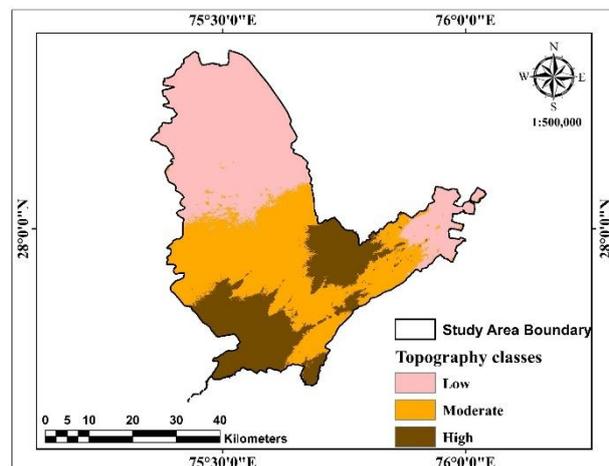


Fig. 6 DEM Topography of the study area

Slope

Critical analysis of slopes reveals the sources of groundwater infiltration and recharge and in turn enables to delineate groundwater potential zones (Dawoud et al., 2005; Vittala et al., 2005; Solomon and Quiel, 2006, Bagyaraj et al., 2013). A close contour indicates a steeper slope whereas a widely spaced contour specifies a gentle slope. A slope layer has been used

to locate the groundwater potential zones (Sreedhar et al., 2009). The groundwater percolation is higher in the gentle and flat areas. In the case of steep slopes, the runoff of rainwater is high, and infiltration is low. The slope amount and numerical values of the slope are presented in Figures 7 and 8 respectively. The major part of the region (around 73-90%) is covered by a steep slope. Slope details of the area are depicted in Table 3.

Table 3 Details of various hydrological parameters

Parameter	Classes	Groundwater Prospect	Weight	Rank
Topography (Elevation M.)	282-359	Very High	10	5
	259-422	High		4
	422-517	Moderate		3
	517-687	Low		2
	687-981	Very Low		1
Slope	15% < Plain Region	Very Good	10	1
	15%-25% Undulating Region	Good		2
	25%-40% Pediment Region	Moderate		3
	40%-55% Plateau Region	Poor		4
	55%-90% Hilly Region	Very Poor		5
Drainage Density (km/km ²)	0.0-0.6	Very Good	20	1
	0.6-1.8	Good		2
	1.8-3.2	Moderate		3
	3.2-5.2	Low		4
	5.2-9.12	Very Low		5
LULC	Crop Land	Very Good	10	5
	Lake Pond, reservoir, Tanks, river, Stream, Drainage	Good		4
	Forest	Moderate		3
	Core urban, Gullied, Hamlets Household, Peri Urban, Scrub land	Poor		2
	Barren Rocky	Very Poor		1
Geomorphology	Pediment	Very Good	10	5
	Structural/ Linear denudation, Interdunal Depression.	Good		4
	Waterbody, River, Pond, Reservoir, Plain area.	Moderate		3
	Intermontane Valley, fill valley.	Poor		2
	Obstacle Dune	Very Poor		1
Geology	Alwar groups	Very Good	20	5
	Ajabgarh groups	Very Good		5
	Malani Plutonic	Good		4
	Alluvium and windblown Sand	Moderate		3
Soil	Clay loam	Very Good	10	5
	Fine Sand	Moderate		3
	Loam	Moderate		2
	Sandy Loam	Poor		2
Lineament Density (km/km ²)	0.0-0.32	Very Good	10	1
	0.33-0.64	Good		2
	0.65-0.95	Moderate		3
	0.96-1.27	Poor		4
	1.28-1.59	Very Poor		5

Geomorphology

In the study area, twelve geomorphic units have been identified with the help of remote sensing and Landsat ETM+ satellite imagery. The plain area is about 1463.98 km² (63.35%)

of the investigated terrain. All the geomorphic units are shown in Table 4. Since the study area is part of the semi-desert region, the predominance of sandy plain lands has been observed. The geomorphic categories and geomorphic numerical values are depicted in Figures 9 and 10 respectively. Groundwater scarcity is noticed in the sandy plains. The geomorphic units like plains, valleys, pediments, and structural hills yield moderate groundwater.

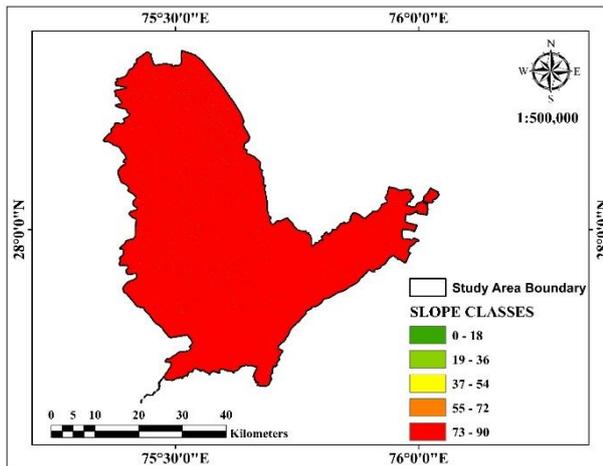


Fig. 7 Map depicting different slope classes

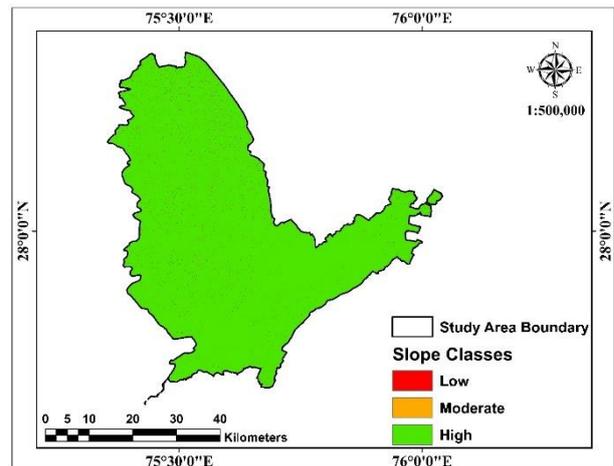


Fig. 8 Map showing slope zoning (numerical value)

Table 4 Geomorphic units and their areal extension in the study area

Geomorphology Category Description	Area sq.km	Area (%)
Alluvial Plain	538.73	23.34
Buried Pediment	305.37	13.21
Aeolian Plain	65.55	2.8
Flood Plain	21.28	0.9
Interdunal Depression	149.82	6.52
Intermontane Valley	28.89	1.2
Obstacle Dune	53.51	2.3
Pediment	30.43	1.3
River/Pond/Reservoir	45.48	2
Sandy Plain	860.70	37.21
Structural/Linear/Denudational	197.91	8.55
Valley Fill	15.52	0.67
Total Area	2313.20	100

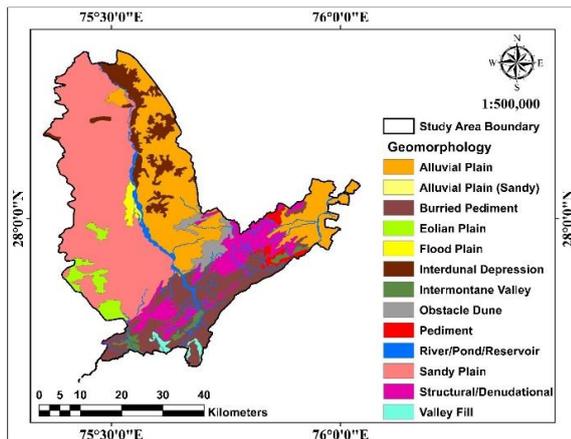


Fig. 9 Map showing different geomorphological types

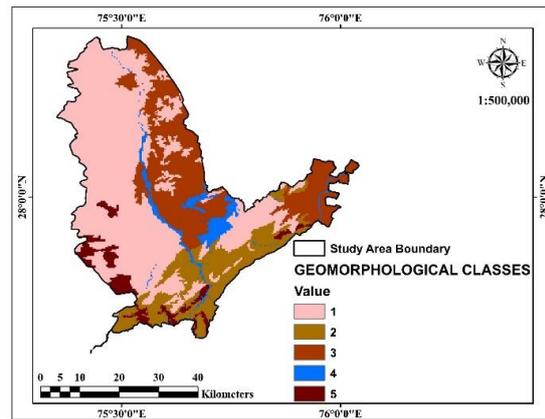


Fig. 10 Map showing various geomorphological classes

Geology

The occurrence and movement of groundwater depend on the porosity and permeability of the rocks that are closely associated with geology (Sreedhar et al., 2009). The Quaternary and Precambrian deposition are found in the study area. The geology and numerical value of the litho units are depicted in Figures 11 and 12 respectively. There are many pediments and structural areas between the Ajabgarh group and its adjoining areas. Groundwater has existed in fractured and weathered Precambrian rocks in the study area.

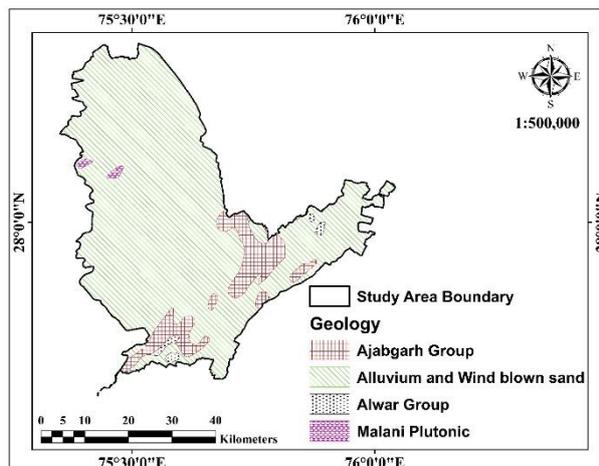


Fig. 11 Geological map of the study area

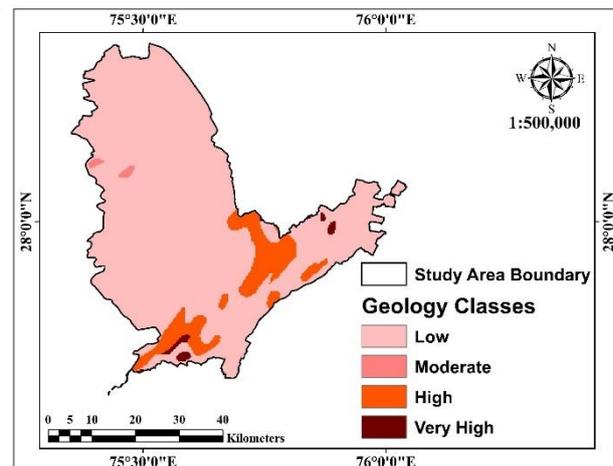


Fig. 12 Map depicting geological classes

Soil

The role of soil is important in groundwater research. Low permeability and limited water infiltration are observed in fine-grained soils. Four major soil units were observed in the study area: Clay loam soils (92.53 km², 4.0%), Fine sand soils (1201.29 km², 51.9%), Loamy soils (25.52 km², 1.1%), and Sandy loam soils (994.20 km², 43 %). Soil map of the study area and map showing soil classes are depicted in Figures 13 and 14. Soil data from the study area is presented in Table 5. Soil data reveals that there is inadequate water holding capacity and a shortage of groundwater storage as a larger part of the study area is covered by sandy soils.

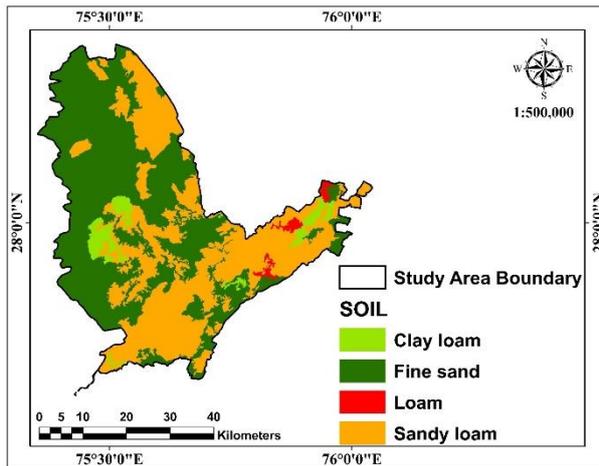


Fig. 13 Soil map of the study area

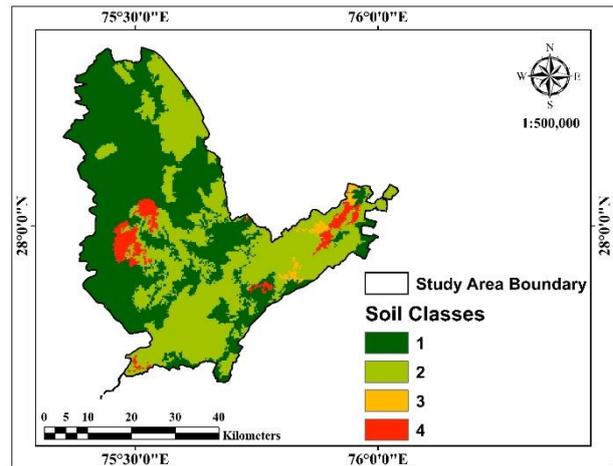


Fig. 14 Map showing soil classes

Table 5 Distribution of Soil in the study area

SOIL Category Description	Area sq.km	Area (%)
Clay loam	92.53	4.0
Fine sand	1201.29	51.9
Loam	25.52	1.1
Sandy loam	994.20	43.0
Grand Total	2313.20	100.0

LULC

Land use land cover mapping has an immense role in delineating the groundwater potential zone as hydrogeological processes are particularly controlled by land use and land cover. Generally, forest areas have high infiltration and runoff is very low. Table 6 demonstrates that the forest land is around 253.04 km² (10.93 %) where there is more infiltration; cropland spreads around 1610.31 km² (69.61%), where infiltration is moderate. Similarly, scrub land spreads over an area of 255.12 km² (11.02%), where infiltration is moderate to low (Figs. 15, 16).

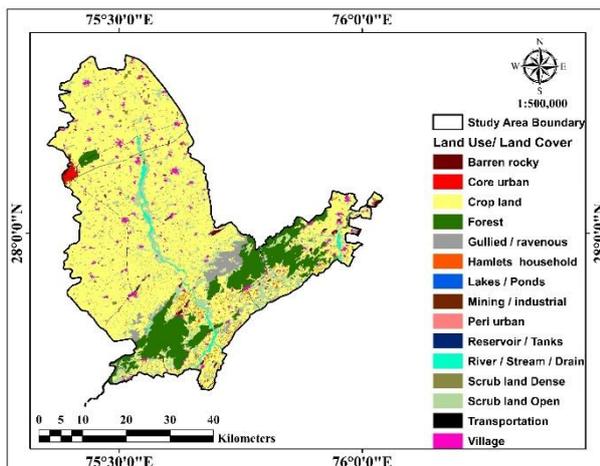


Fig. 15 LULC map of the study area

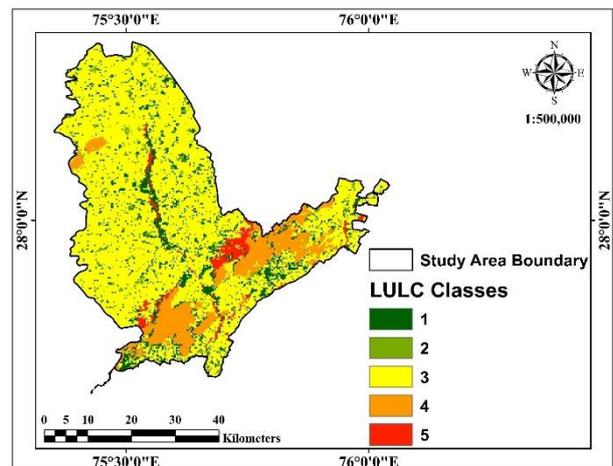


Fig. 16 Map showing LULC classes

Table 6 Land use /Land cover distribution of the study area

LULC Category Description	Area sq.km	Area (%)
Barren rocky	10.43	0.45
Core urban	8.00	0.35
Crop land	1610.31	69.61
Forest	253.04	10.93
Gullied / ravenous	51.31	2.21
Hamlets and dispersed household	29.50	1.27
Lakes / Ponds	0.20	0.008
Mining / industrial	7.10	0.306
Peri urban	0.40	0.017
Reservoir / Tanks	0.08	0.003
River / Stream / Drain	24.40	1.054
Scrub land Dense	8.30	0.358
Scrub land Open	255.12	11.028
Transportation	2.21	0.095
Village	52.80	2.28
Grand Total	2313.20	100

Pre and Post monsoon Groundwater depth

The groundwater level in the Indian subcontinent is controlled by monsoon rainfall. The variation of groundwater depth is observed between the pre-monsoon and post-monsoon periods. Groundwater recharge is good after post-monsoon periods. The aquifer became rich in the later stage of monsoon rainfall. The pre-monsoon and post-monsoon groundwater depth and the numerical value of the groundwater level are depicted in Figures 17, 18, 19 and 20.

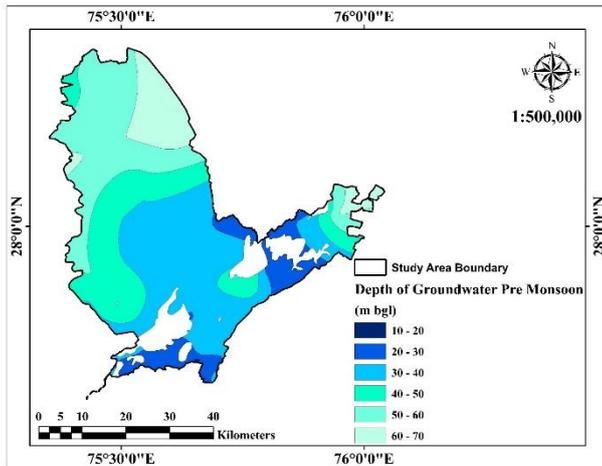


Fig. 17 Groundwater depth of pre-monsoon map

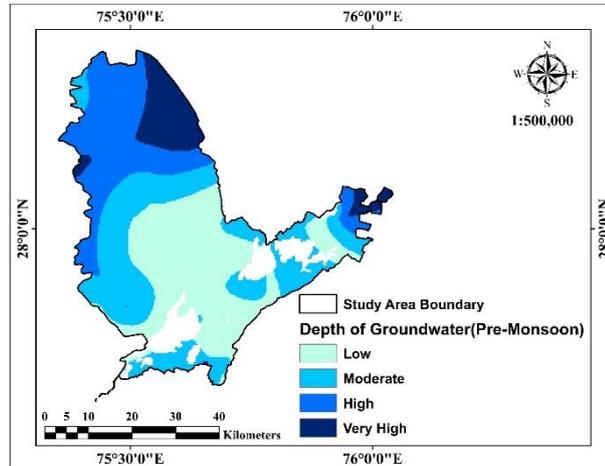


Fig. 18 Map showing groundwater depth classes of pre-monsoon

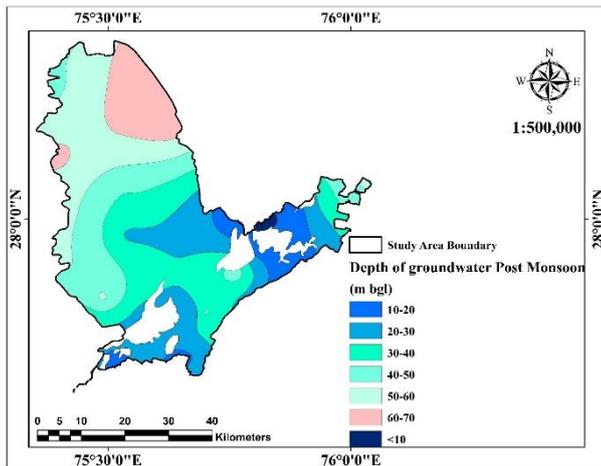


Fig. 19 Groundwater depth of the post-monsoon map

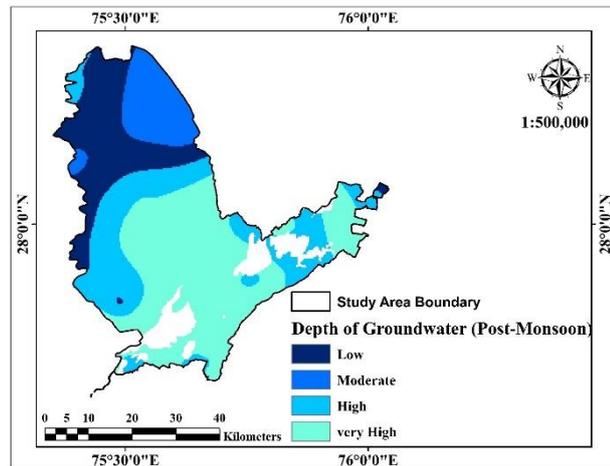


Fig. 20 Map showing groundwater depth classes of post-monsoon

Rainfall

Rainfall directly recharges the groundwater. The isohyet line of maximum rainfall was detected in the north-eastern and south-western parts of the study area. The rainfall distribution and rainfall numerical values of the study area are presented in Figures 21 and 22. The amount of annual rainfall in various weather monitoring stations located in the study area is presented in Table 7. The maximum amount of rainfall is in the Khetri block of Jhunjhunu district, Rajasthan.

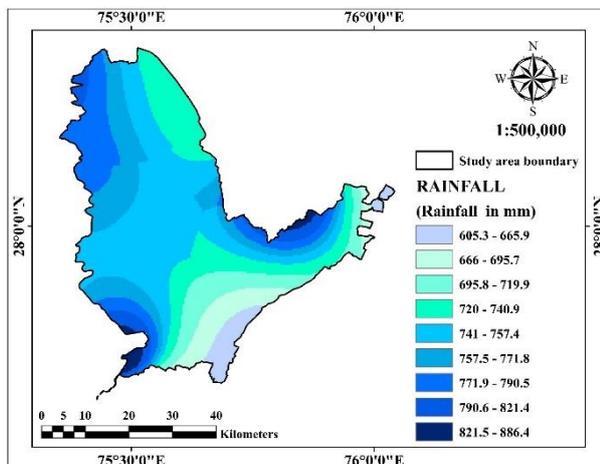


Fig. 21 Rainfall distribution map

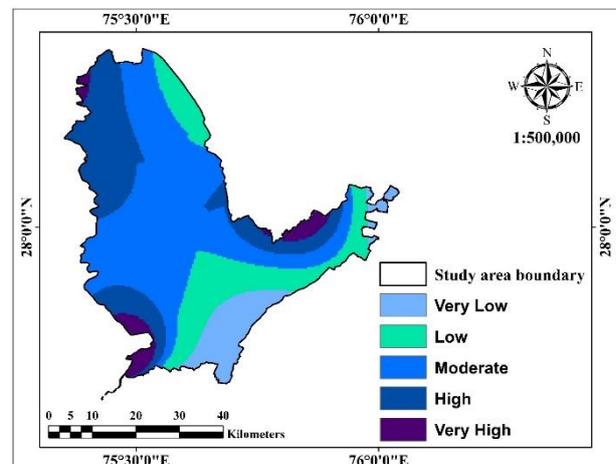


Fig. 22 Rainfall classes map

Table 7 Rainfall distribution and rainfall survey station.

District	Rainfall Station (Near the study area)	Rainfall (mm)
Jhunjhunu (Rajasthan)	Alsisar	854.7
Jhunjhunu (Rajasthan)	Buhana	650.5
Jhunjhunu (Rajasthan)	Chirawa	734
Jhunjhunu (Rajasthan)	Jhunjhunu	778.9
Jhunjhunu (Rajasthan)	Khetri	913.7
Jhunjhunu (Rajasthan)	Nawalgarh	740.7
Jhunjhunu (Rajasthan)	Surajgarh	617.5
Jhunjhunu (Rajasthan)	Udaipurwati	888
Churu (Rajasthan)	Rajgarh	700.3
Sikar (Rajasthan)	Khandela	414.25

Sikar (Rajasthan)	Neem Ka Thana	620.25
Mahendragarh (Haryana)	Narnaul	570

Drainage

Drainage plays an important role in groundwater storage and movement. The drainage density scientifically reveals the groundwater potential and recharge condition of the terrain (Prasad et al., 2008). The drainage pattern is depicted in Figure 23. The drainage density of the study area is presented in Figure 24. The maximum drainage density is located in the north-eastern region of the study area.

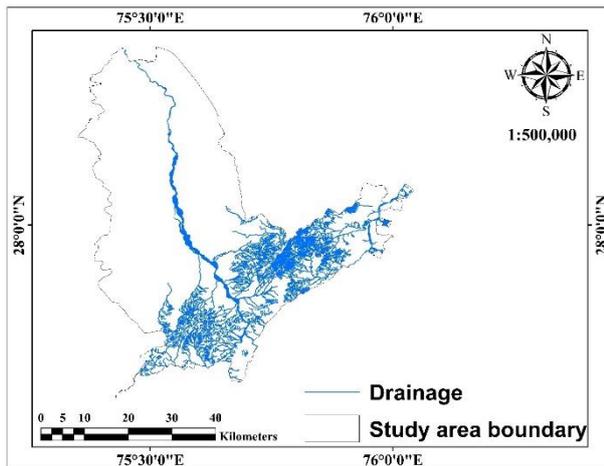


Fig. 23 Drainage map of the study area

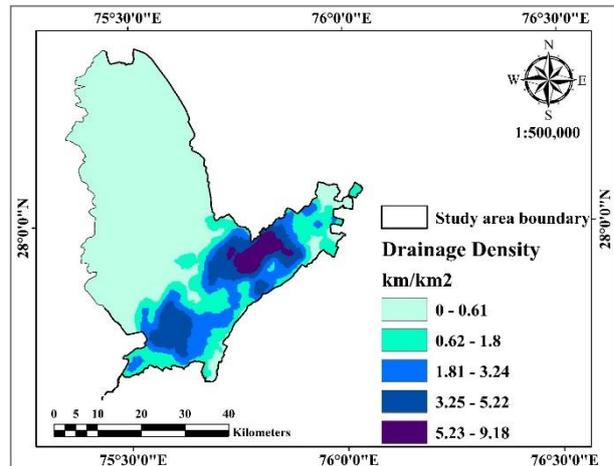


Fig. 24 Drainage Density map of the study area

Lineament

Lineament analysis reveals a close relationship between groundwater infiltration and fractured rocks. Groundwater occurrence and movement are governed by the fracture, joints, and lineament in the hard rock region. The source of lineament data is SRTM images. This layer provides important information about the fractured rocks, which control the movement and storage of groundwater (Pradeep, 1998; Sreedhar et al., 2009). Table 3 shows the lineament density of the study area, which is divided into five categories. The high lineament density region is the potential location for groundwater infiltration. Figures 25 and 26 show the location of the lineament area and the numerical value of lineament density.

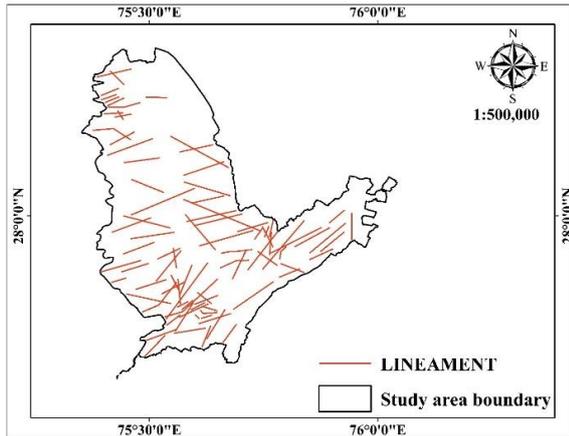


Fig. 25 Lineament map of the study area

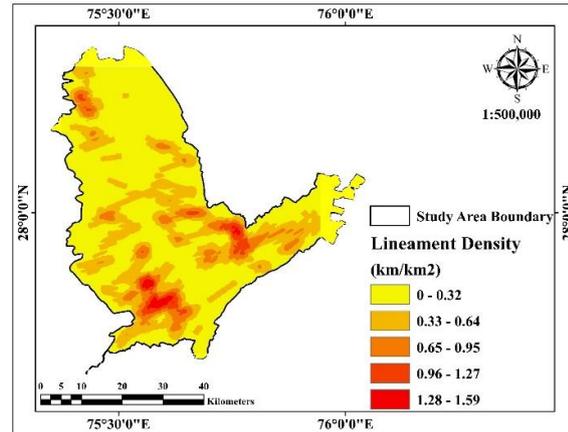


Fig. 26 Lineament Density map of the study area

Stream network

The morphometric analysis depends on the stream network. SRTM topographic data are mainly used for stream networks and for delineating watershed regions. Researchers have used stream networks to identify groundwater potential zones (Abdalla, 2012; Pradeep, 1998; Sreedhar et al., 2009; Adhikary et al. 2021). Stream network is mainly influenced by lithology, rock masses and lineament. The stream of a region refers to the main rivers, tributaries, and distributaries; and stream network refers to a watershed that flows in the direction of the river. Stream ordering has been done using the Strahler method. It is mainly divided into five numerical groups (Table 3) depending on the capability of runoff and infiltration. Figures 27 and 28 show the stream ordering and stream network density of the study area.

Groundwater Potential Zone

The groundwater potential zone has been obtained with the help of a thematic map of the overlay method of the spatial analysis tool of ArcGIS 10.2 software. Each map was created by a common UTM projection and WGS84 datum. The thematic map has been changed into a raster format. Groundwater potential zones are based on the drainage density, slope, land use/land cover, geomorphology, geology, lineament density, and topography data of the study area (Voogd, 1983; Sikdar et al., 2004; Abdalla, 2012). The output prospective map describes the potential groundwater zone. The Groundwater potential zone map of Kantli River Basin (upper part of the Skekhawati River Basin) is divided into five sections namely: very poor (1051.35 km², 45.45%), poor (544.76 km², 23.55%), moderate (460.33 km², 19.90%), high (175.34 km², 7.58%), very high (81.43 km², 3.52%) (Table 8). The largest area under the study belongs to the very poor groundwater potential zones, which cover 1051.35 km² (45.45%). The groundwater potential map of the study area is depicted in Figure 29.

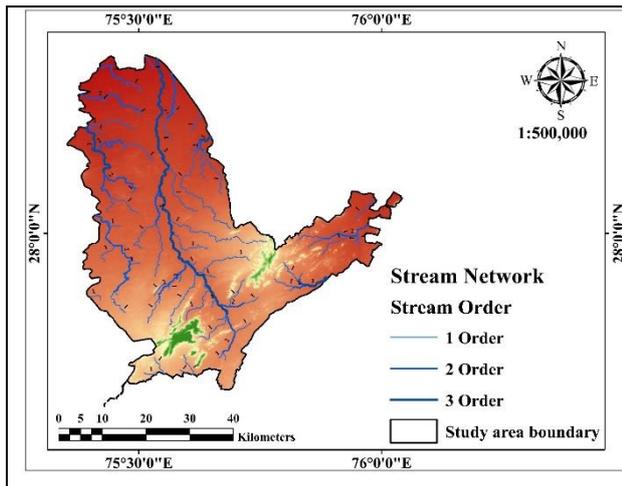


Fig. 27 Stream ordering map of the study area

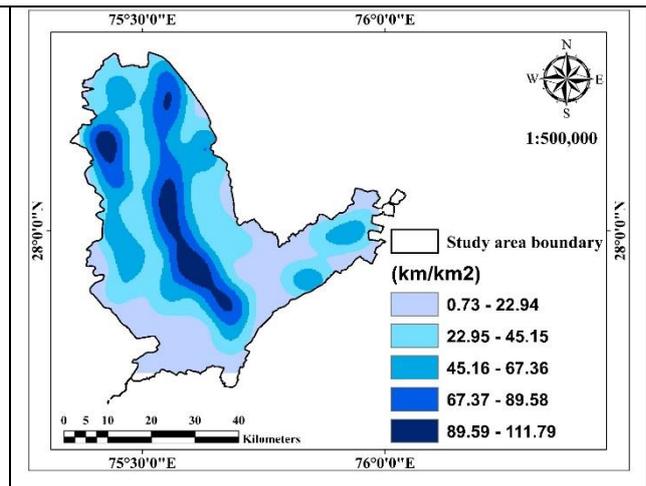


Fig. 28 Stream Network density map of the study area

Table 8 Various groundwater potential zones

Sl.no.	Value	Potential Zone	Area (km ²)	Area (%)
1	0-1	Very Poor	1051.35	45.45
2	1-2	Poor	544.76	23.55
3	2-4	Moderate	460.33	19.90
4	4-6	High	175.34	7.58
5	6-12	Very High	81.43	3.52

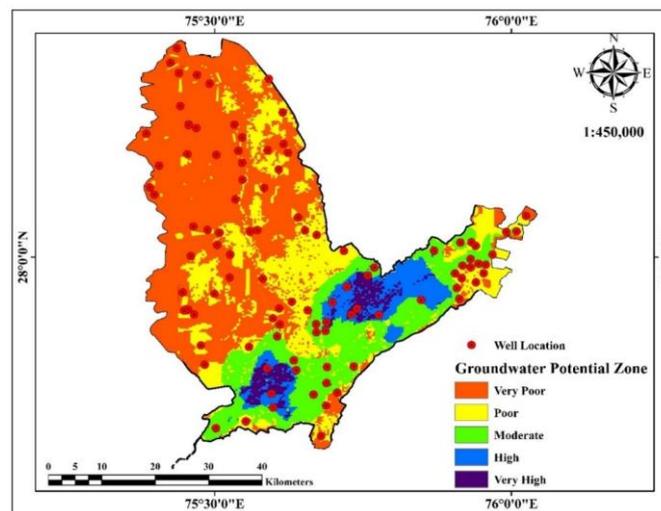


Fig. 29 Groundwater potential zone map of the study area

Results validation with borehole data

Two types of aquifers are commonly noticed in the region; confined and unconfined. The result obtained from the boreholes data is that the depth of the shallow unconfined aquifer is located at 8-70 m depth of the study area (27.1%). The confined aquifer is found in a large area, the depth of this well is above 70 meters (72.9%). The groundwater scarcity is found in most of the study area, which is derived from borewell data and pumping test data. Table 9 shows the boreholes (Geophysical Well, Exploratory Well, Observation Well, and Piezometer Well) types and depth of groundwater.

Table 9 Various borehole data of the study area

Types of Boreholes	Several Well no	Number of Borehole	Percentage of Well (%)	Depth (m)	Aquifer
Geophysical Well	15	29	27.1	8-70 M	Unconfined Shallow
Exploratory Well	01				
Observation Well	00				
Piezometer Well	13				
Geophysical Well	33	78	72.9	70 M Above	Confined Deep
Exploratory Well	24				
Observation Well	01				
Piezometer Well	20				

Conclusion

The present study reveals the delineation of potential groundwater zones in the Kantali Basin of Jhunjhunu District (upper Shekhawati basin). The various informative thematic map layers have been compiled to identify potential zones of groundwater (very poor, poor, moderate, high, very high). It is observed that only 11% of the total study area constitutes a very good to good groundwater potential zone. Sandy soil, low lineament density, medium to low rainfall, insignificant forest density, moderate to steeply slope have reduced the infiltration and increased the runoff, which has resulted in low and unhealthy groundwater prospective zone in the Kantali Basin. Steps may be taken to construct good numbers of check dam and water harvesting structures at the Government level. Moreover, extensive plantation programmes may also be undertaken in these areas by local bodies and NGOs. Hopefully, subsequently increased groundwater recharge will enhance the groundwater prospects and potentials of the area.

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