

# Explicitly Analyze the Soil Erosion in the Karst and Non-Karst Area of Different Morphological Types in Guizhou of China

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## Research Article

**Keywords:** Soil erosion, karst area, morphological types, spatiotemporal change intensity, quantitative analysis, Guizhou Province

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1                   **Explicitly analyze the soil erosion in the karst and**  
2                   **non-karst area of different morphological types in**  
3                   **Guizhou of China**

4                   **Anjun Lan<sup>1</sup>. Zemeng FAN<sup>1,2,3,4</sup>. Qingsong Zhao<sup>1</sup>. Xuyang Bai<sup>2,3</sup>**

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6  
7                   **Abstract:** How to explicitly understanding the soil erosion intensity change in different  
8 geomorphological types is one of key issues in the field of soil and water conservation. According  
9 to classification criterion of soil erosion intensity of China, the spatial soil erosion data with the  
10 resolution of 10 m×10m in Guizhou Province were obtained by combing with the multi-resolution  
11 remote sensing data of ALOS, ZY-3, GF-1, Landsat and GDEM V2, and 2762 field sampling data in  
12 2010 and 2015, respectively. a spatial analysis model of soil erosion was improved to analyze the  
13 spatiotemporal change of soil erosion intensity in karst and non karst area of Guizhou province,  
14 which involved the spatial soil erosion data and different geomorphological type data of Guizhou  
15 province. The results show that the soil erosion intensity decreased by 6468.13km<sup>2</sup> in Guizhou  
16 Province from 2010 to 2015. The dynamic change intensity in the high-altitude area is larger than  
17 in the low-altitude area. The soil change intensity in karst area is higher than in non karst area,  
18 especially in the high and middle elevation area in Guizhou province. Moreover, the decreasing ratio  
19 of soil erosion intensity in karst area is generally larger than in non karst area, which can be used to  
20 explain that the ecological restoration projects and water soil conservation polices carried out in  
21 karst area has a good effect, especially in western of Guizhou province from 2010 to 2015, one the  
22 other hand, the soil erosion in non karst area should also be focused by local government in the  
23 future.

24                   **Key words:** Soil erosion; karst area; morphological types; spatiotemporal change intensity;  
25 quantitative analysis; Guizhou Province

26  
27                   **1. Introduction**

28                   Soil Erosion and its initiated environmental risk is one of the most serious issue in the  
29 world(Borrelli et al. 2017a; Jiang et al. 2014; Fang and Fan, 2021; Zhao et al., 2020 ). The interacting  
30 relationship between natural environment and human activities is the key component of soil erosion  
31 study (Poesen 2018). There is no C layer between soil and it is very serious that bedrock and the  
32 soil erosion caused by denudation, dissolution and cross distribution of denudation and dissolution,

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33 especially in the wide distributed area where karst has characteristic of the vertical development(He  
34 et al. 2009).

35 Guizhou Province is a typical representative area of karst landforms in Southwest China where  
36 the landform pattern is complex( Yang et al. 2019). The increasing human activities have caused  
37 extremely serious negative interference to this area, which makes the soil erosion particularly  
38 serious (Huang and Cai 2007; Wang and Li 2007). At present, a series of studies on soil erosion in  
39 Guizhou Province have been carried out by the related scholars. For example, to address the  
40 relationship between land use and response of soil erosion, Wang Huan et al. (2019) carried out  
41 quantitative attribution on soil erosion in different geomorphic areas of Sanchahe watershed, which  
42 showed that land use had the highest explanatory power for soil erosion, and the soil erosion risk of  
43 cultivated land was higher than that of other land use types(Borrelli et al. 2017b; Dai et al. 2017; Li  
44 et al., 2018). Based on the fixed-point monitoring of slope surface, the influence of different rainfall  
45 intensity on soil erosion mode is significant(Fullhart et al. 2020; Wu et al. 2018), and the spatial  
46 distribution of soil erosion directly affects the spatial change of soil nutrient loss (Zeng et al. 2018).  
47 In addition, relevant researches have been carried out on the mechanism of soil erosion(Li et al.  
48 2017; Peng and Wang 2012; Shen et al. 2008; Wang et al. 2013a) , the construction of soil erosion  
49 model (Geissen et al. 2008; Kheir et al. 2010; Liu 2016; Xu et al. 2008) and the temporal and spatial  
50 variation of soil erosion(He et al. 2018; Wang et al. 2016). The model for scenario analysis of climate  
51 change and land use shows that different response scenarios of land use change to climate change  
52 will also cause different soil erosion intensity (Fan et al. 2015; Perović et al. 2019). According to  
53 the relationship between soil erosion intensity and terrain characteristics, a model of soil erosion  
54 degree was established to reveal the influence of different slope and surface cover types on soil  
55 erosion intensity (Bonetti et al. 2019; Vanacker et al. 2019). The MPSIAC model shows that  
56 topography is the main factor of soil erosion (Noori et al. 2018). The different type of landforms  
57 directly affects the surface water and light intensity, and then affects the spatial and temporal  
58 distribution of ecological environmental factors such as soil and vegetation, thus affecting the  
59 intensity of regional soil erosion to a certain extent (Cao et al. 2019; Cheng et al. 2009; Feng et al.  
60 2016; Wang et al. 2013b). Thus, different geomorphic types are one of the key factors to analyze  
61 the spatial and temporal distribution of soil erosion and its intensity and direction. At present, the  
62 research mainly focuses on the driving effect of land use types on soil erosion, as well as soil erosion  
63 mechanism or single terrain factor at small watershed scale. However, at the provincial level,  
64 especially in karst areas with fragile ecological environment, there is a lack of spatiotemporal  
65 variation intensity of soil erosion in different geomorphic types, especially in karst and non karst  
66 areas quantitative analysis of driving factors.

67 Since the 21st century, due to the input and implementation of ecological protection projects  
68 and soil and water conservation policies in Guizhou Province, soil erosion has undergone significant  
69 spatiotemporal differentiation and change. Therefore, the purpose of this paper is to construct the  
70 spatio-temporal analysis model of soil erosion dynamic degree, breadth and relative change rate of  
71 different geomorphic types by combining the data of Guizhou landform type with the soil erosion  
72 intensity data of  $10 \times 10$  m resolution. Based on the quantitative analysis of soil erosion change  
73 intensity of different geomorphic types and karst and non karst areas, the spatial and temporal  
74 variation intensity difference of soil erosion in karst and non karst areas is revealed, and the driving  
75 effects of ecological protection projects and soil and water conservation policies on Soil and water  
76 erosion changes are comprehensively analyzed, so as to further implement in Guizhou Province.

77 **2. Data and methods**

78 **2.1 Data sources**

79 In the study area, ALOS 10 m resolution remote sensing images in 2009 and 2010, and ZY-3  
 80 and GF-1 2.5 m resolution remote sensing images in 2014 and 2015 were used to extract two phases  
 81 of spatial data of land use types in the man-machine interactive mode. Vegetation coverage and  
 82 bedrock exposure rate were extracted from Landsat 30 m resolution remote sensing image data in  
 83 2010 and 2015. Slope data was derived from the 30 m resolution remote sensing image of GDEM V2.  
 84 The data of lithology classification in karst and non karst areas are from the 1:50000 geological map  
 85 of Guizhou Province; the geomorphic data are from the agricultural geomorphic division of Guizhou  
 86 Province.

87 **2.1.2 Data processing**

88 Based on the 1:500000 agricultural geomorphic regionalization map of Guizhou Province as  
 89 the base map, vectorization is carried out in ArcGIS to generate corresponding vector data.  
 90 Combined with the research conception of this paper, the soil erosion intensity changes of karst and  
 91 non-karst areas in different geomorphic areas are analyzed, and the hills subdivided in karst areas  
 92 in vector data are appropriately merged, and finally low basin, medium basin and high basin are  
 93 obtained. There are 12 types of land, low hill, medium hill, high hill, low mountain, low middle  
 94 mountain, middle mountain, high mountain, low platform and high platform (Fig. 1a).

95 Due to the different sources and scales of soil erosion data, it is necessary to unify the data to  
 96 form a resolution of  $10 \times 10$  m. The resolution of ZY-3 and GF-1 is resampled for  $10 \times 10$  m, which  
 97 is consistent with the resolution of ALOS Image. The resolution of vegetation coverage, bedrock  
 98 exposure rate, slope and other data is increased to  $10 \times 10$  m to ensure the unity of various factors  
 99 in the process of comprehensive superposition.

100 Considering the great difference of soil erosion characteristics between karst and non-karst  
 101 areas in Guizhou, such as soil erosion intensity, the year limitation for soil resistant erosion, risk  
 102 coefficient of soil erosion, etc., the karst areas refer to the technical standard for comprehensive  
 103 management of soil erosion in karst areas, and the non-karst areas are based on the classification  
 104 and classification standards of soil erosion. Combined with land use type, vegetation coverage,  
 105 bedrock exposure rate, slope and other data, based on a large number of field sampling verification,  
 106 according to the above standards for karst and non-karst areas, the quantitative discrimination and  
 107 division of soil erosion grade (Table 1, table 2) are carried out respectively, so as to achieve the  $10\text{m}$   
 108  $\times 10\text{m}$  resolution soil in different geomorphological types of Guizhou Province in 2010 and 2015  
 109 Soil erosion spatial information data extraction. In the process of processing and analyzing the basic  
 110 data, in order to ensure the classification accuracy and quality of the data, a large number of field  
 111 sampling verification were carried out on the above classification results in the field, with a total  
 112 number of 2762, and the verification accuracy in 2010 and 2015 were 85.44% and 91.07%,  
 113 respectively (Fig. 1b).

114 Tab.1 Surface erosion (sheet erosion) classification index of the non-karst area

Land type		Slope(°)				
		5~8	8~15	15~25	25~35	>35
Coverage of non cultivated land (%)	60~75					
	45~60	Mild		Intense		

	30~45		Moderate	Intense	Extremely Intense
	<30				
Sloping Land		Mild	Moderate	Intense	Extremely Intense
					Violent

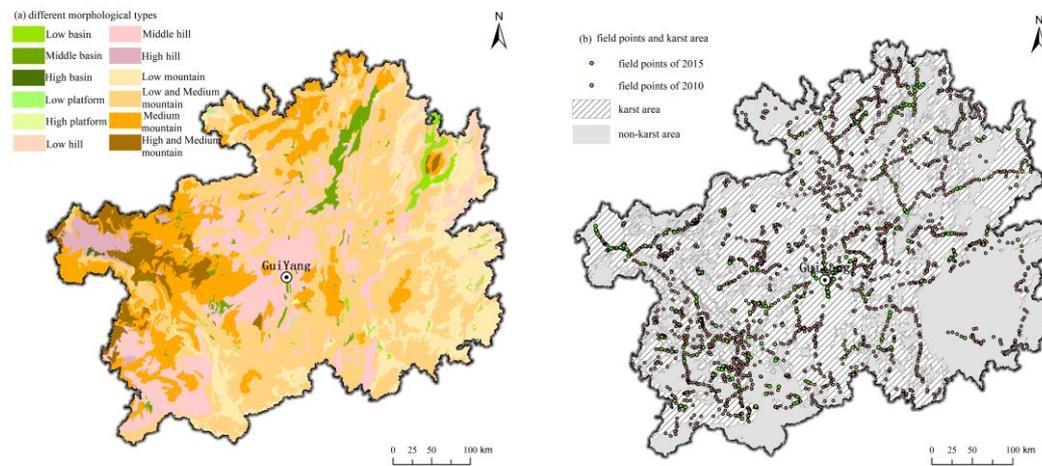
115

116

Tab.2 Sloping Farmland erosion intensity index indirect discrimination of the Karst area

Slope (°)		<5	5~8	8~15	15~25	25~35	>35
Bedrock exposure rate (%)	<5	Mild			Intense	Extremely Intense	Violent
	5~30						
	30~50				Moderate	Intense	Extremely Intense
	50~70						Intense
	>70		Slight				

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118

119

Fig.1 The distribution of different morphological types (a) and field point of karst and non-karst (b) in

120

Guizhou Province

121

## 2.2 Spatial analysis model of soil erosion

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In order to quantitatively describe and compare the spatial position, main types and regional differences of soil erosion between karst and non-karst areas in Guizhou Province, introducing the existing land use change model and combining with soil erosion breadth index (Zhu and Li 2003), dynamic degree and other soil erosion characterization indexes, the dynamic degree model, the breadth index model and the relative change rate model of soil erosion are constructed.

127

For further description of the spatial and temporal dynamic changes of soil erosion in Karst and non-karst regions with different geomorphic types from 2010 to 2015, a spatiotemporal dynamic index model of soil erosion in Guizhou Province was constructed, which can be expressed as:

130

$$D_i = \left\{ \sum_j^n \left( \frac{\Delta S_{i,j}}{S_a} \right) \right\} \times 100\% \quad (1)$$

131

Where  $D_i$  is the dynamic index of soil erosion, with a value of 0~1, expressed as a percentage;

132

$\Delta S_{i,j}$  represents the net area of change between soil erosion of grade  $i$  and soil erosion of grade  $j$

133

( $\text{km}^2$ );  $S_a$  is the area of a geomorphic type area ( $\text{km}^2$ ). The greater the  $D_i$  value, the more intense the

134

change in this area; the dynamic degree  $D_i$  focuses on the process of soil erosion change rather than

135 its results, which aims to reflect the intensity and overall situation of the change.

136 To reveal the main conversion types of soil erosion intensity in different geomorphic types,  
137 combined with the spatial distribution data of different geomorphic types, a soil erosion breadth  
138 index model which can be used to calculate the main conversion types of soil erosion in different  
139 geomorphic types was constructed:

$$140 \quad C_i = S_{i,j}/S \times 100\% \quad (2)$$

141 Where  $C_i$  is the breadth index of soil erosion, the value is 0~1, expressed as a percentage;  $S_{ij}$  is  
142 the change area from type  $i$  soil erosion to type  $j$  soil erosion in a certain area from 2010 to 2015  
143 ( $\text{km}^2$ );  $S$  is the total area of various soil erosion changes in the area ( $\text{km}^2$ ). The higher the  $C_i$  value,  
144 the more dominant the soil erosion from type  $i$  to type  $j$  in this region. In view of the fact that there  
145 are many types of soil erosion conversion among different geomorphic types, in order to analyze  
146 the main conversion types of different soil erosion intensities in each geomorphic type area, the  
147 cumulative sum of  $C_i$  values from large to small is greater than 70%.

148 On the basis of quantitative analysis of soil erosion extent, in order to reveal the regional  
149 differences of different soil erosion intensity changes, the relative change rate model of soil erosion  
150 was constructed:

$$151 \quad K_i = S_{i,j} / S_a \times 100\% \quad (3)$$

152 Where  $K_i$  is the relative change rate of soil erosion, the value is 0~1, expressed as a percentage;  
153 the larger  $K_i$  value in a certain area indicates that the transition from type  $i$  soil erosion to type  $j$  soil  
154 erosion is more likely to occur than that in other regions.

## 155 3 Results

### 156 3.1 Comparison of soil erosion intensity between karst and non-karst area

157 The comparative analysis of different levels of soil erosion intensity in karst and non-karst  
158 areas of Guizhou Province shows that: from 2010 to 2015, the overall erosion intensity in Guizhou  
159 Province showed a weakening trend, with the proportion of erosion area except slight erosion in the  
160 total area of Guizhou Province decreased from 31.37% in 2010 to 27.70% in 2015, and the total  
161 weakened area was 6468.13  $\text{km}^2$  (Table 3). Among them, the areas of mild, moderate, strong and  
162 severe erosion in karst area are higher than those in non-karst area, which are 566.88  $\text{km}^2$ , 794.82  
163  $\text{km}^2$ , 265.4  $\text{km}^2$  and 15.17  $\text{km}^2$  more than those in non-karst area. In addition, although the areas of  
164 extremely strong and severe erosion have weakened, the weakening extent is the smallest. The  
165 weakened areas of soil erosion in karst area and non-karst area only account for 11.03% and 22.42%  
166 of the total area of soil erosion weakening in karst area and non-karst area respectively. On the  
167 whole, the weakening area of soil erosion intensity in Guizhou karst area is larger than that in non-  
168 karst area, but the weakening trend of extremely strong and severe erosion types is contrary to the  
169 general trend of soil erosion intensity weakening (the weakening range of karst area is 11.04% and  
170 4.38% lower than that of non-karst area, respectively). The results show that the control of soil  
171 erosion in Guizhou karst area has achieved remarkable results, but the treatment of extremely strong  
172 and severe erosion types is still severe.

173

174

Tab.3 The area of soil erosion intensity in Guizhou Province in 2010 and 2015

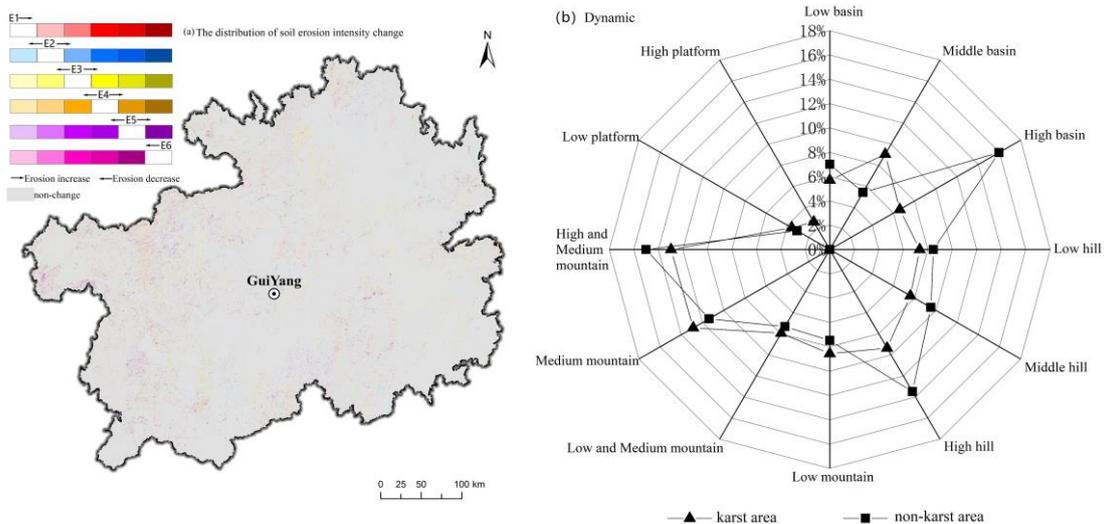
Lithology	Soil erosion intensity	2010 Year (km <sup>2</sup> )	2015 Year (km <sup>2</sup> )	2015 increased or decreased compared with 2010 (km <sup>2</sup> )(+/-)	Proportion of increase or decrease in 2010 (%) (+/-)
Non- Karst	E1(Slight erosion)	47110.03	49588.29	+2478.26	+5.26
	E2(Mild erosion)	10023.40	9550.25	-473.15	-4.72
	E3(Moderate erosion)	5739.16	4757.35	-981.81	-17.11
	E4(Strong erosion)	2452.96	1985.37	-467.58	-19.06
	E5(Extremely strong erosion)	1377.36	1132.59	-244.77	-17.77
	E6(Severe erosion)	978.87	667.92	-310.95	-31.77
Karst	E1(Slight erosion)	73772.55	77762.42	+3989.87	+5.41
	E2(Mild erosion)	17591.79	16551.76	-1040.03	-5.91
	E3(Moderate erosion)	10621.68	8844.96	-1776.73	-16.73
	E4(Strong erosion)	3592.39	2859.41	-732.98	-20.40
	E5(Extremely strong erosion)	1695.34	1581.33	-114.02	-6.73
	E6(Severe erosion)	1190.86	864.74	-326.12	-27.39
Total		176146.39	176146.39	0	

175

### 176 3.2 Comparison of soil erosion intensity changes in different geomorphic types

#### 177 3.2.1 Regional differences of different soil erosion intensity

178 The dynamic degree ( $D_i$ ) simulation results of soil erosion intensity change in different  
 179 geomorphic morphology areas of Guizhou Province from 2010 to 2015 shows that with the rising  
 180 of altitude, the  $D_i$  value of each geomorphic form area increases continuously, but the  $D_i$  value of  
 181 karst area is smaller than that of non-karst area. Therefore,  $D_i$  value in high altitude areas such as  
 182 high basin, high hill and high and middle mountain is higher than that in low altitude areas such as  
 183 low basin, low hill and low mountain. The  $D_i$  value of karst area is lower than that of non-karst area,  
 184 especially in high basin, with a difference of 9.35% (Fig. 2b). The results show that the dynamic  
 185 degree of soil erosion in high altitude area is higher than that in low altitude area, and the dynamic  
 186 degree of soil erosion in karst area is lower than that in non-karst area.



187

188

188 Fig.2 The intensity change of soil erosion distribution (a) in 2010 and 2015 and  $D_i$  value in different  
 189 morphological types (b) of Guizhou Province

190

191 3.2.2 Analysis on Transformation of soil erosion intensity in different geomorphic types

192 Based on different geomorphic types, the main conversion types of soil erosion intensity (the  
 193 sum of  $C_i$  values is more than 70%) were statistically analyzed by using the soil erosion breadth  
 194 index ( $C_i$ ). The results showed that the average  $C_i$  of the main transition types in karst area reached  
 195 40.47%, mainly from mild erosion to slight erosion, and the average  $C_i$  value of main conversion  
 196 types in non-karst areas was 33.47%. The extent of soil erosion in karst area is higher than that in  
 197 non-karst area. In addition to the high basin and high middle mountain, the other areas are mainly  
 198 transformed from mild erosion to slight erosion (Table 4). The average difference of soil erosion  
 199 extent in karst area is 10.4% in high basin, high hill and high middle mountain area, and only 1.77%  
 200 in low basin, low mountains and low hill area. In the area of low basin, low hill and low mountains,  
 201 the average difference of the first  $C_i$  value and the second  $C_i$  value between karst area and non-karst  
 202 area is 25.77% and 23.49% respectively. There are few main conversion types, and the main  
 203 transformation mode is from mild to slight erosion. In the high basin, high hill, middle mountains  
 204 and high middle mountain areas, the average difference between the first  $C_i$  value and the second  $C_i$   
 205 value between karst area and non-karst area is 7.36% and 5.63%, and the transformation from mild  
 206 to slight erosion is not obvious. The main transformation types in non-karst area are more than those  
 207 in karst area, and the conversion among all levels of erosion occurs. The above analysis shows that  
 208 mild to slight erosion is the main transition type in all physiognomy areas. With the increase of  
 209 altitude, the dominance of mild to slight erosion is weakened, and the  $C_i$  value of other erosion  
 210 levels gradually increases, and the conversion types increase. The  $C_i$  value of main transition types  
 211 in karst area is higher than that in non-karst area.

212

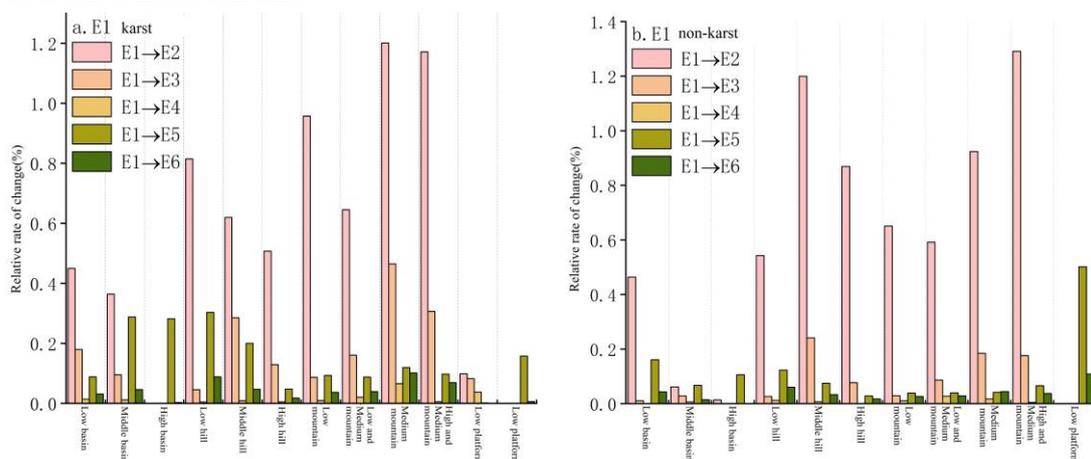
213 Tab.4 Statistical table of soil erosion intensity conversion types in different morphological types

Geomorphic form type	Non- Karst		Karst		Geomorphic form type	Non- Karst		Karst	
	Main conversion type	$C_i(\%)$	Main conversion type	$C_i(\%)$		Main conversion type	$C_i(\%)$	Main conversion type	$C_i(\%)$
Low Basin	E2→E1	42.70	E1→E2	7.89	Low mountain	E1→E2	8.71	E1→E2	11.18
	E3→E1	13.13	E2→E1	45.16		E2→E1	35.03	E2→E1	35.81
	E3→E2	18.74	E3→E1	14.89		E3→E1	14.42	E3→E1	14.08
				E3→E2		9.63	E3→E2	16.89	E3→E2
Middle basin	E2→E1	51.03	E1→E2	4.02	Low and Medium mountain	E1→E2	8.09	E1→E2	8.12
	E3→E1	9.92	E2→E1	38.15		E2→E1	33.24	E2→E1	33.89
	E3→E2	10.43	E3→E1	15.65		E3→E1	15.62	E3→E1	14.81
			E3→E2	15.99		E3→E2	13.70	E3→E2	14.21
High basin	E3→E1	24.88	E2→E1	39.10	Medium mountain	E1→E2	8.12	E1→E2	9.32
	E3→E2	10.83	E3→E1	29.30		E2→E1	22.99	E2→E1	26.52
	E4→E3	15.54	E3→E2	10.21		E3→E1	11.69	E3→E1	11.69
	E5→E4	11.21				E3→E2	10.79	E3→E2	15.70
	E6→E5	13.30				E4→E3	10.38	E4→E3	8.89
Low hill	E2→E1	41.68	E1→E2	11.10		E5→E4	8.82		
	E3→E1	12.76	E2→E1	43.75	High and Medium mountain	E1→E2	8.60	E1→E2	9.03

	E3→E2	15.79	E3→E1	18.44		E2→E1	15.64	E2→E1	22.64
Middle hill	E1→E2	12.60	E1→E2	8.14		E3→E1	15.76	E3→E1	13.97
	E2→E1	33.94	E2→E1	38.32		E3→E2	11.87	E3→E2	18.09
	E3→E1	15.31	E3→E1	16.40		E4→E3	10.03	E4→E3	8.88
	E3→E2	8.80	E3→E2	10.07		E5→E4	9.16		
High hill	E1→E2	6.44	E2→E1	31.89	Low platform	E1→E5	16.25	E2→E1	66.66
	E2→E1	21.91	E3→E1	27.63		E2→E1	44.98	E3→E1	18.50
	E3→E1	20.15	E3→E2	10.89		E3→E2	12.59		
	E3→E2	7.47			High platform			E2→E1	63.90
	E4→E2	7.92						E3→E1	9.55
	E4→E3	6.80							

### 214 3.2.3 Distribution characteristics of soil erosion intensity in different geomorphic types

215 Due to the large proportion of low and medium mountains in the whole province, the proportion  
 216 of basins and platforms is very small. Considering the distribution of the change of soil erosion  
 217 intensity only from the perspective of area, almost all the areas where the degree of soil erosion is  
 218 enhanced or weakened are distributed in low and middle mountains. Therefore, in order to further  
 219 reveal the distribution of soil erosion intensity in different geomorphic types, the relative change  
 220 rate was used for analysis and comparison. The results show that (Fig. 3), the distribution of  
 221 decreased soil erosion intensity in karst area is significantly different from that in non-karst area,  
 222 but the regional difference of enhanced soil erosion intensity is not obvious. Among them, the micro,  
 223 mild and extremely strong erosion in karst area is further enhanced, and the relative change rate  
 224 is higher in middle mountain and high mountain area, while in non-karst area, the relative change rate  
 225 of moderate and strong erosion aggravation in high basin is higher, and the difference between karst  
 226 area and non-karst area is 0.19% and 0.2%; in karst area, severe, extremely strong and strong erosion  
 227 is weakened in middle mountain, high and middle mountain and other areas. The relative change  
 228 rates of high and middle mountains are high, which are 0.63%, 0.66% and 1.15%, respectively. The  
 229 relative change rates of moderate and mild erosion are higher in high hills and middle basins  
 230 respectively; in non-karst areas, the relative change rates of severe, extremely strong, strong and  
 231 moderate erosion are mainly in high basins, with an average of 2.59%; in addition, slight erosion in  
 232 karst areas and non-karst areas turns to slight erosion in medium The relative change rate of basin  
 233 and mid hill is high, with an average of 3.19% and 3.0%, and this area is the rapid expansion area  
 234 of urban construction land.





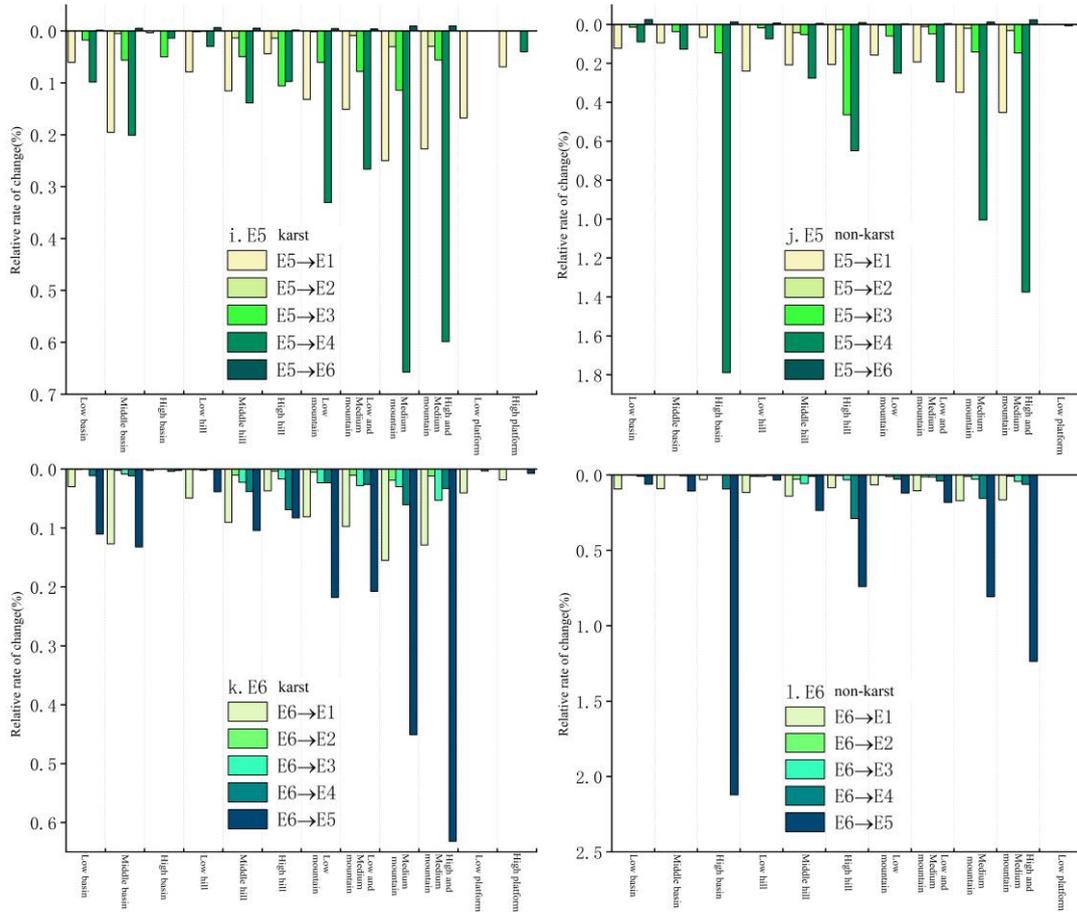


Fig.3 The intensity changes of soil erosion between karst and no-karst area in different morphological types

## 4 Discussion and conclusions

### 4.1 Discussion

The driving forces of soil erosion are directly or indirectly controlled by different geomorphic types. Different altitudes affect the vertical difference distribution of heat and moisture, and then affect vegetation types and land use patterns; the greater the surface relief, the stronger the terrain cutting, soil erosion is prone to occur; slope and aspect directly affect the intensity and direction of soil erosion. Different soil erosion intensity is the result of the synthesis of various factors. This paper analyzes the geomorphic differences of different soil erosion intensity changes, which has a certain indication for the changes of various factors.

The intensity of soil erosion in Guizhou Province showed a decreasing trend. Among them, the dynamic degree of soil erosion in the high-altitude area in the west is higher than that in the low-altitude area in the East. This result reflects the ecological effect of the project of returning farmland to forest (grass). From 2010 to 2015, the cultivated land in the middle mountain and high and middle mountain areas converted to woodland and grassland by 13.83 km<sup>2</sup>(Han et al., 2015). Due to the large-scale plantation and the implementation of rocky desertification ecological project in the west of Guizhou Province (Yan et al. 2018; Zhang et al. 2017), the overall trend of soil erosion in Guizhou Province has become better, showing a dynamic change pattern of soil erosion in the western part of Guizhou Province, which changes violently in the west and relatively stable in the east (Fig. 2a). On the one hand, due to the main distribution of forest land and cultivated land in karst area, under

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262 the drive of ecological and environmental protection policies, slope cultivated land is gradually  
263 reduced, so that the breadth of soil erosion in karst area is higher than that in non-karst area, and the  
264 conversion is mostly to weak first-order erosion. On the other hand, the severe, extremely strong  
265 and intense soil erosion in karst area is mainly concentrated in the middle and high mountains areas,  
266 and the relative change rate of soil erosion in these areas is relatively high, which reflects that the  
267 ecological restoration project has achieved good ecological effect in karst area. In addition, due to  
268 the rapid development of urbanization in the central basin and hilly areas along the Zunyi-Guiyang-  
269 Anshun line, the construction land has expanded greatly (increased by 1029.86 km<sup>2</sup> from 2010 to  
270 2015), mainly from cultivated land (Cheng et al. 2019; Cheng et al. 2018), which makes the relative  
271 change rate of soil erosion from mild to slight erosion in this area is in a high state. From the overall  
272 situation, because the slopes of middle mountains and middle and high mountains areas in Guizhou  
273 Province are basically above 25 ° and the terrain fluctuates greatly, coupled with the implementation  
274 of ecological restoration projects such as slope farmland conversion and rocky desertification  
275 control, the dynamic degree of soil erosion intensity change in this area is the strongest.

276 Due to the difficulty in obtaining the refined spatial information of soil erosion, the spatial-  
277 temporal variation trend and regional difference of soil erosion intensity in karst and non- karst areas  
278 of Guizhou Province from 2010 to 2015 are only revealed at the resolution of 10m × 10m. In the  
279 further work, we will continue to obtain more refined spatial information data of soil erosion, and  
280 then realize the dynamic detection and fine identification of spatial and temporal variation patterns  
281 and regional differences of soil erosion intensity in Karst and non-karst areas of Guizhou Province,  
282 and further implement soil and water loss prevention measures and control measures in karst and  
283 non-karst areas in Guizhou Province Rocky desertification ecological restoration project provides  
284 dynamic and refined spatial data and method support.

## 285 4.2 Conclusions

286 In view of the lack of quantitative analysis of soil erosion intensity in different geomorphic  
287 types, this paper takes Guizhou Province as an example to study the spatiotemporal variation  
288 intensity of soil erosion and its driving factors in karst and non-karst regions of different geomorphic  
289 types. The spatial and temporal variation intensity of soil erosion in karst and non- karst regions was  
290 revealed by constructing spatiotemporal analysis models of soil erosion in different geomorphic  
291 types. This study can provide method support and data support for the refined study of soil erosion  
292 change intensity in complex geomorphic type area.

293 The results showed that: 1) from 2010 to 2015, the soil erosion intensity in Guizhou Province  
294 showed a decreasing trend, with a total area of 6468.13 km<sup>2</sup>. The total area of soil erosion intensity  
295 weakening in karst area is larger than that in non-karst area, but the weakening extent of soil erosion  
296 intensity of extremely strong and severe erosion in karst area is less than that in non-karst area,  
297 which indicates that the soil erosion restoration difficulty of extremely strong and severe erosion  
298 area in karst area is higher than that in non-karst area; 2) with the increase of altitude, the dynamics  
299 of various geomorphic types are also discussed. The average  $D_i$  value increased from 6% in low  
300 altitude area to 13.61% in high altitude area, and the increase of  $D_i$  value in karst area was less than  
301 that in non-karst area; 3) the main conversion type of soil erosion in each geomorphic form type  
302 area was from mild to slight erosion, and the conversion breadth index of soil erosion in karst area  
303 was higher than that in non-karst area, although  $C_i$  value in low altitude area was only 1.77%; 4)  
304 The results showed that the average effect of the project was 10.4% higher than that of the other

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305 areas in Guizhou Province.

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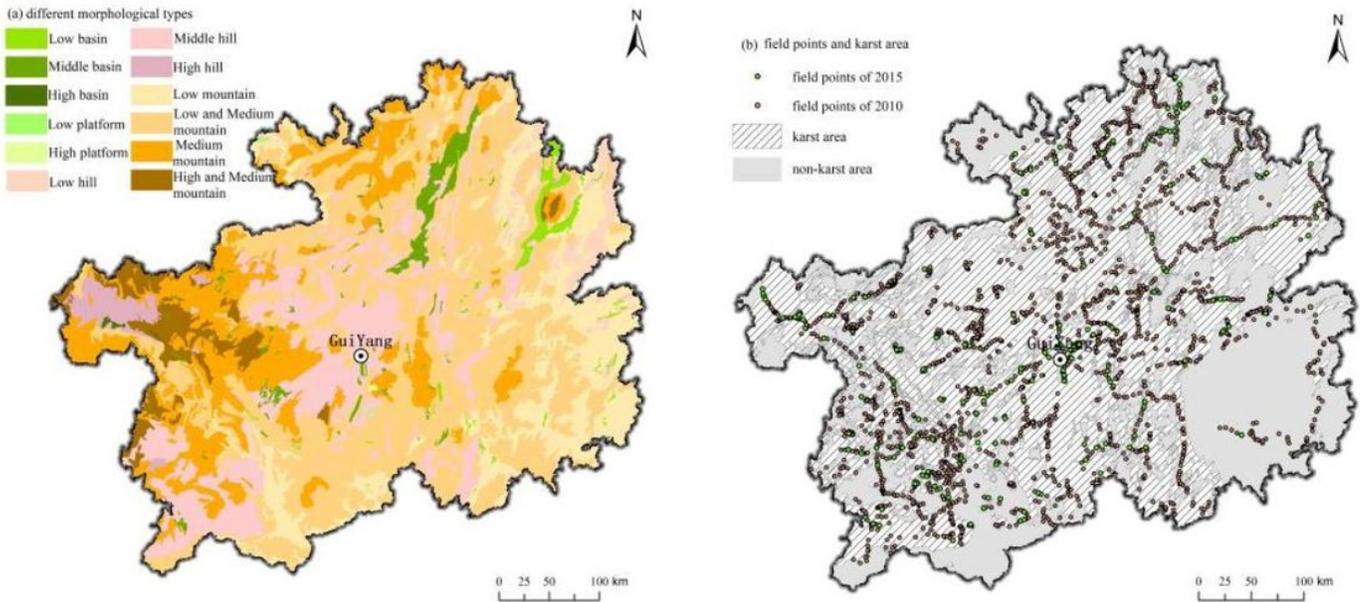
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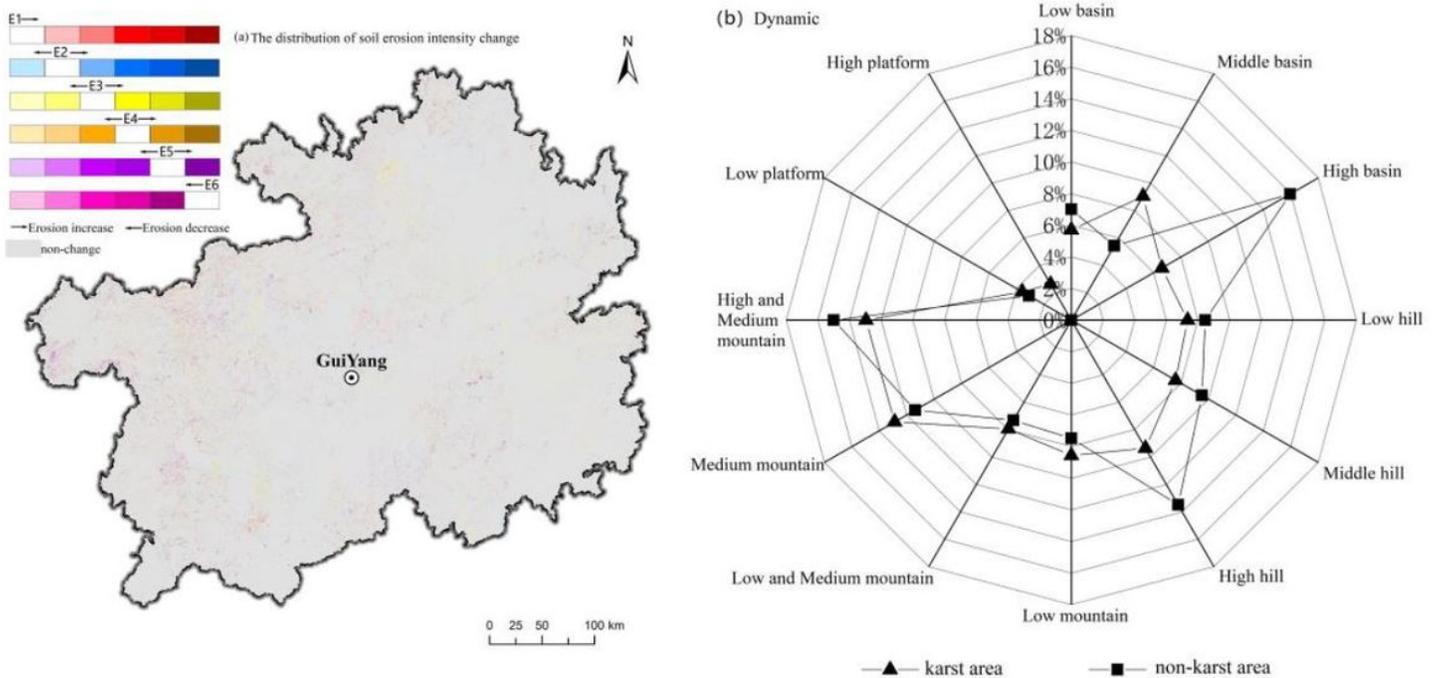
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# Figures



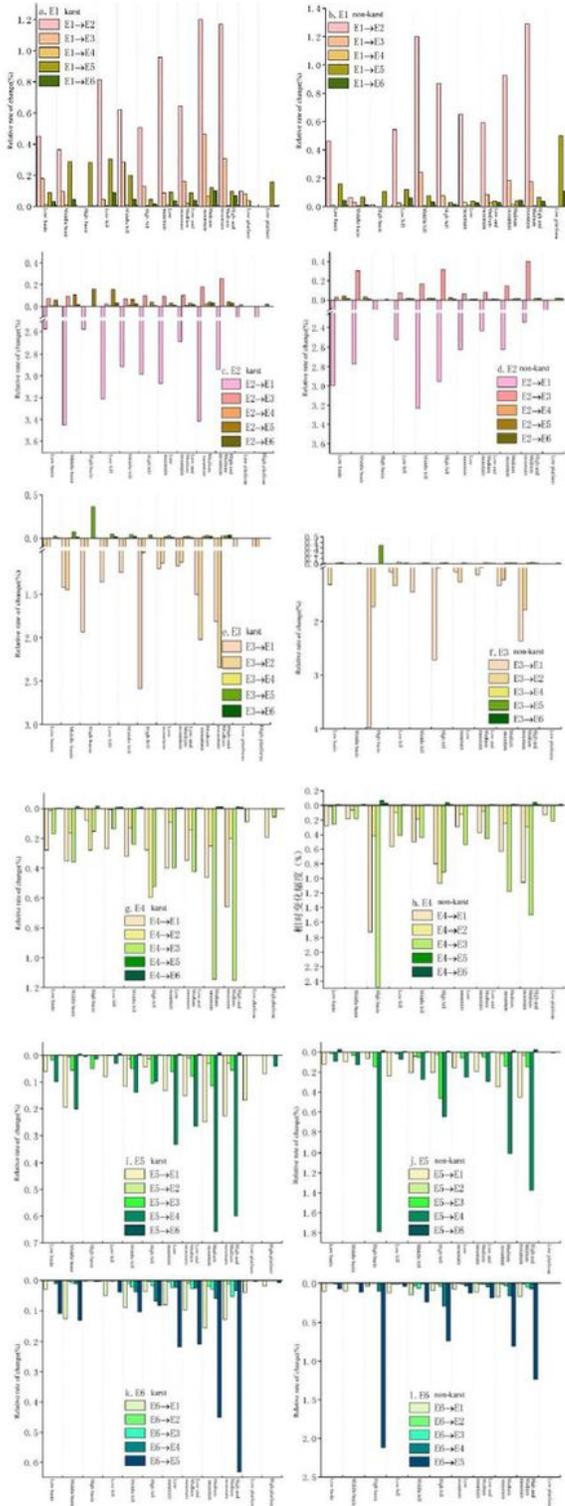
**Figure 1**

The distribution of different morphological types (a) and field point of karst and non-karst (b) in Guizhou Province. Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its authorities, or concerning the delimitation of its frontiers or boundaries. This map has been provided by the authors.



**Figure 2**

The intensity change of soil erosion distribution (a) in 2010 and 2015 and Di value in different morphological types (b) of Guizhou Province. Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its authorities, or concerning the delimitation of its frontiers or boundaries. This map has been provided by the authors.



### Figure 3

The intensity changes of soil erosion between karst and no-karst area in different morphological types