

Spatial landslide susceptibility mapping using integrating an adaptive neuro-fuzzy inference system (ANFIS) with two multi-criteria decision-making approaches

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1 **Spatial landslide susceptibility mapping using integrating an adaptive neuro-fuzzy**
2 **inference system (ANFIS) with two multi-criteria decision-making approaches**

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21

22 **Abstract**

23 Landslide is a type of slope processes causing a plethora of economic damage and loss of lives
24 worldwide every year. This study aimed to analyze spatial landslide susceptibility mapping in
25 the Khalkhal-Tarom Basin by integrating an adaptive neuro-fuzzy inference system (ANFIS)
26 with two multi-criteria decision-making approaches, i.e. the stepwise weight assessment ratio
27 analysis (SWARA) and the new best-worst method (BWM) techniques. For this purpose, the
28 first step was to prepare a landslide inventory map, which were then divided randomly by the
29 ratio of 30/70 for model training and validation. Thirteen conditioning factors were used as
30 slope angle, slope aspect, altitude, topographic wetness index (TWI), plan curvature, profile
31 curvature, distance to roads, distance to streams, distance to faults, lithology, land use, rainfall
32 and normalized difference vegetation index (NDVI). After the database was created, the BWM
33 and the SWARA methods were utilized to determine the relationships between the sub-criteria
34 and landslides. Finally, landslide susceptibility maps were generated by implementing ANFIS-
35 SWARA and ANFIS-BWM hybrid models, and the ROC curve was employed to appraise the
36 predictive accuracy of each model. The results showed that the areas under curves (AUC) for
37 the ANFIS-SWARA and ANFIS-BWM models were 73.6% and 75% respectively, and that
38 the novel BWM yielded more realistic relationships between effective factors and the
39 landslides. As a result, it was more efficient in training the ANFIS. Evidently, the generated
40 landslide susceptibility maps (LSMs) can be very efficient in managing land use and preventing
41 the damage caused by the landslide phenomenon.

42 **Keywords:** landslide susceptibility; machine learning; GIS; ANFIS; SWARA; BWM

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45 **Introduction**

46 Causing great losses of lives and properties, landslides are dangerous processes that occur
47 repeatedly in mountainous and hilly areas worldwide, (Juliev et al. 2019; Gutiérrez et al. 2015).
48 This mass movement occurs whenever the loading of an earth material exceeds its shear
49 strength (Lin et al. 2017). Although this geological phenomenon is often triggered by
50 earthquakes and heavy rainfalls, the expansion of anthropogenic activities in susceptible areas
51 has always played an important factor in its occurrence (Baena et al. 2019) Despite the

52 increased human knowledge regarding landslide occurrence and factors controlling this
53 phenomenon, it is believed that the damage caused by landslides will increase due to
54 deforestation, climate change and urban development (Pham and Prakash 2018). Therefore, it
55 is essential to acquire accurate and realistic information about the spatial distribution and
56 degrees of susceptibility for landslide-prone regions (Colkesen et al. 2016). To achieve this
57 goal and to mitigate the destructive impacts of this phenomenon, landslide susceptibility maps
58 can serve as an appropriate tool for increasing awareness and predicting future hazards
59 (Feizizadeh et al. 2017). Based on previous landslides and identical physical features in similar
60 areas, a landslide susceptibility map provides important signs regarding the locations where
61 future landslides are likely to occur (Pradhan et al. 2017).

62 The Alborz Mountain has always been subject to the natural disasters such as landslide due
63 to its being on the seismic belt of the Himalayas (Farrokhnia et al. 2011). In a study of
64 identifying high-risk regions of the world with respect to landslide hazard, Nadim et al. (2006)
65 reported that the Alborz and Zagros Mountains of Iran were among the areas with moderate to
66 high landslide risks. In addition, according to the National Committee on Natural Disaster
67 Reduction of the Iranian Ministry of Interior, the annual damage caused by landslides in Iran
68 amounts to about 500 billion Rials (Arab Amiri et al. 2019). Consequently, if the loss of human
69 life is taken into account, it is evident that zoning of the study area is necessary.

70 In the recent years, researchers have used different methods and their combinations to zone
71 the areas susceptible to landslide in different areas worldwide that can generally classify into
72 two quantitative and qualitative groups. The qualitative approaches, also known as knowledge-
73 driven approaches, are the techniques of assigning weights to and rank criteria and sub-criteria
74 based on experts' knowledge and experience and also on mathematical relations defined for
75 each method. Some of these methods, which have been used in various studies and have yielded
76 acceptable results, include the AHP (Zhang et al. 2016; Pourghasemi and Rossi 2017; Yan et
77 al. 2019; Du et al. 2019) and hybrid methods such as MCDA and MCE (Kavzoglu et al. 2014;
78 Erener et al. 2016; Kumar et al. 2017) and the WLC (Kouli et al. 2014; Ahmed 2015). The
79 second group, also known as data-driven approaches, consists of the techniques which are not
80 influenced by experts' opinions in the computational process. Instead, the relationship between
81 the landslides and the effective parameters is determined by using probabilistic equations.
82 These methods, which have been used repeatedly in various studies on landslides, include
83 bivariate and multivariate probability models such as frequency ratio (FR) (Hong et al. 2017;
84 Sharma and Mahajan 2018; Berhane et al., 2020), weight of evidence (WoE) (Kayastha et al.
85 2013; Ding et al. 2017; Cui et al. 2017) and logistic regression (LR) (Wang et al. 2015; Oh et

86 al. 2018; Pham et al. 2019) as well as soft computing methods such as artificial neural network
87 (ANN) (Polykretis et al. 2014; Bui et al. 2016; Pham et al. 2017; Zhu et al. 2018; Moayedi et
88 al. 2018), fuzzy logic (Pradhan et al. 2011; Ramesh and Anbazhagan 2015; Turan et al. 2020),
89 adaptive neuro-fuzzy inference system (ANFIS) (Aghdam et al. 2017; Polykretis et al. 2019;
90 Paryani et al. 2020), random forest (RF) (Kim et al. 2018) and support vector machine (SVM)
91 (Pradhan 2013; Oh et al. 2018). According to the literature review, it is clear that although these
92 models have effective performance in computational process, hybrid methods performs a
93 strong structure to achieve more accurate results. Youssef et al. (2015) combined logistic
94 regression and frequency ratio methods to landslide susceptibility mapping in Saudi Arabia.
95 They concluded that ensemble of FR-LR model combining models has better results than using
96 them alone. In another study, Aghdam et al. (2017) combined FR and WoE statistical methods
97 with ANFIS algorithms to produce landslide susceptibility map of the Zagros Mountains in
98 Iran. Their results indicated that FR-ANFIS and WoE-ANFIS have better performance
99 compared with FR and WoE. It seems that using new methods is necessary to achieve more
100 accurate results. Therefore, achieving optimal results relies on (a) the quality of the input data,
101 (b) the structure of the model, used (Adineh et al. 2018).

102 In comparison with other studies, the research innovation of this study was to train the
103 ANFIS method through the new multi-criteria decision-making method BWM (ANFIS-BWM)
104 as a new hybrid model and compare it with the hybrid ANFIS-SWARA model. In other words,
105 the weights obtained from BWM and SWARA methods were used as ANFIS input data.
106 Landslide susceptibility maps were then generated for both models, and their performance
107 accuracy was estimated by using the ROC curves. Finally, the results were compared to
108 determine which of the two BWM and SWARA approaches were able to train the ANFIS
109 network better and generate more realistic maps.

110

111 **Study Region**

112 With an area of 8604 km², the Khalkhal-Tarom Basin is located on the southern slopes of the
113 Alborz mountain range along from 47° 42' 44" to 49° 10' 34" and 36° 37' 22" to 37° 56' 35"
114 (Fig. 1). Approximately 92% (7967 km²) of the Basin consists of highlands and the remainder
115 of plains. The highest and lowest elevations are 3314 m and 288 m, respectively. The data from
116 the climatological stations of Iran Meteorological Organization and Ministry of Energy were
117 utilized to estimate temperature and precipitation. The average annual temperature in the region
118 is about 10.5° C; while, the coldest month is February, and the warmest is August. In addition,
119 the average annual rainfall is about 375 mm. The difference in precipitation levels in the

120 highlands on the two sides of the main river (the Ghezel Ozan River) results from the
121 differences in the prevailing climatic conditions in the areas adjacent to the study area.
122 Although the study area has diverse lithology, pyroclastic rocks of Karaj Formation cover most
123 of its surface area. Regarding land use, range lands are the largest land cover class; however,
124 agricultural lands and orchards cover nearly 16% of the land area in the basin and exhibit an
125 increasing trend given the population increase. Given the existence of economic infrastructures
126 and the growing residential areas in future, zoning of landslide-prone regions seems to be
127 vitally important.

128

129 **Database Development and Data Preparation**

130 It is necessary to create database in any study of geographical information system. The
131 landslide zoning is no exception, and database creation including inventory map and
132 conditioning factors is considered as the first and the most important step in this process. The
133 landslide inventory map shows the locations and spatial distribution of landslides that happened
134 in the past. Since it is crucial to pinpoint the locations of the past and present landslides in order
135 to predict future high-risk areas, preparation of a landslide inventory map is a requisite to any
136 study on landslides (Regmi et al. 2013). Information on the locations of past landslides and
137 their spatial distributions was obtained from the Forest, Rangeland and Watershed
138 Organization of Iran (Fig. 1). According to Figure 1, the inventory map was employed to
139 randomly select 172 (or 70%) of the 242 landslides that have occurred in the region for training
140 the data and the 30% for model validation.

141 Various factors including geology, hydrology, geomorphology, climate and topography
142 affect slope instability. Determination of these factors is among the basic, and initial steps in
143 landslide zoning. In this study, thirteen conditioning factors including slope angle, slope aspect,
144 altitude, topographic wetness index (TWI), plan curvature, profile curvature, distance to roads,
145 distance to streams, distance to faults, lithology, land use, rainfall, normalized difference
146 vegetation index (NDVI) were selected based on the characteristics of the study area, experts
147 opinion and previous studies for the spatial modeling of the landslides (Table 1). According
148 to Table 1, these thirteen factors were determined by using the information obtained from the
149 related organizations and the reference data. Following that, ArcGIS was employed to generate
150 and digitalize the maps (30- \times 30-m pixels). Raster data models of the layers were then prepared
151 by using the selected methods.

152 In order to prepare the different information layers, the digital elevation model (DEM) was
153 prepared first using ASTER satellite images. DEM is one of the most important databases in

154 any landslide study because preparation of some important thematic maps depends on it. The
155 slope angle, slope aspect, altitude, TWI and plan and profile curvature layers were extracted
156 from the DEM (Fig. 2 a-f). The other considered factors (distance to roads, distance to streams,
157 distance to faults, lithology, land use, rainfall and the Normalized Difference Vegetation Index
158 were then determined, respectively (Fig. 2 g-m). In addition, the conditioning factors were
159 categorized based on experts' opinions, previous studies and study area characteristics.

160 The slope degree is always considered as an essential factor in analyzing the areas susceptible
161 to landslide (Umar et al. 2014), because it is the major cause of mass movements. Exposure to
162 sunlight, dry winds, and increased relative humidity due to rainfall are all factors associated
163 with slope aspect that trigger landslides (Kavzoglu et al. 2013). Therefore, slope aspect has
164 always been consideration by researchers. This factor is divided into 9 classes. Altitude is not
165 directly involved in the occurrence of landslides; however, other factors related to it such as
166 tectonic activity, weathering and climate change influence the entire process (Rozos et al.
167 2008). The topographic wetness index is a useful tool for estimating moisture conditions at
168 basin scale (Grabs et al. 2009). This factor was used due to the varying humidity conditions in
169 the study area. The values obtained from the slope curvature show the morphology of the
170 different elevation points (Erener et al. 2010). In this paper, both the profile curvature curve
171 and the plan curvature were taken into account. The former indicates the velocity and process
172 of sediment transport and the second the divergence and convergence of the flow passing
173 through the surface (Dehnavi et al. 2015). Road construction, especially when engineering
174 principles are ignored, reduces slope stability and consequently triggers landslides (Moosavi
175 and Niazi 2015). Therefore, the distance from the road has always attracted the interest of
176 researchers (Xiao et al. 2019; Bui et al. 2012). Streams decrease shear strength by eroding the
177 materials from the toe of the slope. Consequently, the factor of distance from the stream is very
178 important in relation to slope stability (Achour et al. 2017). Faults, especially in seismic zones,
179 play a significant role in triggering mass movements. They either act directly as a triggering
180 factor for landslides or indirectly by causing fractures in slope layers that lead to the penetration
181 of water into joints and fissures, thereby reducing the shear strength of materials constituting
182 the slope that results in the occurrence of landslides. lithology as a geological factor has always
183 been important role in predicting landslide occurrence probability because different lithologic
184 units with varying degrees of permeability influence slope stability (Chalkias et al. 2014). Due
185 to their impacts on slope instability, different types of land use have always attracted many
186 researchers in their research on landslides (Conforti et al. 2013; Dou et al. 2014). The
187 precipitation factor was used in this research because the amount of precipitation varies with

188 changes in elevation and precipitation directly and indirectly influences landslide occurrence.

189 The NDVI index was calculated to analyze the effect of vegetation on slope instability:

$$190 \quad NDVI = \frac{NIR-R}{NIR+R} \quad (1)$$

191

192 The NDVI benefits from the ratio of near-infrared (NIR) reflection to red (R) reflection to
193 estimate vegetation density (Hall et al. 1995).

194

195 **Methodology**

196 *Step-wise weight assessment ratio analysis (SWARA) model*

197 This is a multi-criteria decision-making method with an ultimate objective like that of other
198 similar approaches: assigning weights to criteria and sub-criteria. Since its introduction by
199 Keršulien *et al.* in 2011, researchers have used it to analyze various areas (Mardani et al. 2015).

200 An advantage of this method is its flexibility that allows experts to prioritize the criteria based
201 on the existing conditions. The main feature of this approach is its capability in estimating
202 experts' opinions in relation to the relative importance of the criteria in order to determine their
203 weights (Keršulien et al. 2011). This procedure consists of the following steps:

- 204 1. Selecting the required criteria and ranking them according to their degrees of
205 importance (the most important criteria take the highest position of ranking and the
206 least important ones the lowest).
- 207 2. Calculating the coefficient K_j , which is a function of the relative importance of each
208 criterion.
- 209 3. Determining the initial weight of each criterion.
- 210 4. Calculating the final normalized weight.

211

212 The final weight for each criterion is calculated through the following equations (Keršulien et
213 al. 2011):

214

$$215 \quad S_j = \frac{\sum_i^n A_i}{n} \quad (2)$$

216 In this equation, j and n represent the criterion number and the number of experts, respectively.

217 The value of A_i also indicates the suggested rating of each criterion.

$$218 \quad K_j = S_j + 1 \quad (3)$$

$$219 \quad Q_j = \frac{x_j - 1}{K_j} \quad (4)$$

220 Here, K_j and Q_i are functions of the relative importance and initial weight of each criterion,
221 respectively.

$$222 \quad W_j = \frac{Q_j}{\sum_{j=1}^m Q_j} \quad (5)$$

223 In this formula, j represents the criterion number, and m shows the number of criteria when W_j
224 indicates the final weight.

225

226 The final weight (W_j) obtained for each sub-criteria in this study indicates the relationship
227 between landslides and conditioning factors. The larger/smaller the final weight of a sub-
228 criterion is, the higher/lower its importance in landslide occurrence will be.

229

230 ***Best-worst multi-criteria decision making (BWM) model***

231 The best-worst method (BWM) is one of the newest and most efficient multi-criteria decision-
232 making approaches introduced in 2015 by Rezaei to calculate the final weights of criteria and
233 sub-criteria in decision-making problems. As in other MCDM methods such as AHP, pair-wise
234 comparisons are used in BWM. However, the differences in the final weight calculation in this
235 method have made the final result much more realistic and consistent than methods such as
236 AHP. The advantages of BWM over AHP are that fewer pair-wise comparisons are used, the
237 numbers used for pair-wise comparisons are integers ranging between 1 and 9, and there is no
238 need for fractional numbers. It is also possible to integrate the BWM with other MCDM
239 methods (Ahmad et al. 2017). The various steps in this method and its algorithms for problem
240 solving are as follows (Rezaei 2015):

- 241 1. Specifying the decision-making criteria for evaluation. The set of criteria is defined as
242 $\{C_1, C_2, \dots, C_n\}$.
- 243 2. Determining the best (B) and worst (W) criteria by the experts. The best criterion
244 include most important or the most desirable criterion, whereas the worst ones include
245 those with the least desirability and/or lowest importance.
- 246 3. Determining the priority of the best criteria compared to all the others (the numbers 1
247 to 9 are used for this purpose). This preference is represented in the form of the
248 following vector:

$$249 \quad A_B = (\alpha_{B1}, \alpha_{B2}, \dots, \alpha_{B3}) \quad (6)$$

250 Here, α_{B1} represents the preference of the best criterion (B) over the criterion j ($\alpha_{BB} = 1$) (Fig.
251 4).

252 4. Determining the priority of all the criteria over the worst one (W). The preference vector
 253 for this phase is as follows:

$$254 A_W = (\alpha_{1W}, \alpha_{2W}, \dots, \alpha_{nW})^T \quad (7)$$

255

256 Here, α_{jW} is the preference of the j criterion over the worst one (W) ($\alpha_{WW}=1$) (Fig. 4).

257 5. Calculating the final weights of the criteria. The following equations are used for this
 258 purpose:

$$259 \text{Min } \varepsilon_i$$

$$260 \text{s.t.}$$

$$261 \left| \frac{W_B}{W_j} - \alpha_{Bj} \right| \leq \varepsilon_i, \forall j = 1, 2, \dots, n$$

$$262 \left| \frac{W_j}{W_W} - \alpha_{jW} \right| \leq \varepsilon_i, \forall j = 1, 2, \dots, n \quad (8)$$

263

$$264 \sum_j^W W_j = 1$$

$$265 W_j \geq 0, \forall j = 1, 2, \dots, n$$

266

267 The values of the final optimum weights ($W_1^*, W_2^*, \dots, W_n^*$) and ε_i^* are obtained by Equations 8.

268 In addition, the consistency ratio for each criterion can be estimated by using the consistency
 269 index table (Table 3) and the ε_i^* value. The following equation states that:

$$270 \text{Consistency Ratio} = \frac{\varepsilon_i^*}{\text{Consistency Index}} \quad (9)$$

271 It is evident that the closer the value of the consistency index is to zero, the more realistic the
 272 results will be. Refer to [Rezaei et al. \(2015\)](#) for more details of this method.

273

274 *Adaptive neuro fuzzy inference system (ANFIS)*

275 Although a fuzzy inference system (FIS) using “if-then” rules can analyze complex processes,
 276 it is unable to perform the learning process. The adaptive neuro fuzzy inference system
 277 (ANFIS) ([Jang 1993](#)) is one of the most widely used fuzzy systems for modeling nonlinear
 278 problems. This approach, developed by combining a FIS and an artificial neural network
 279 (ANN), utilizes the advantages of both approaches to solve problems. The ANN model is able
 280 to optimize the fuzzy logic solution through the learning process ([Oh and Pradhan 2011](#)). The
 281 details of the ANFIS model structure are as follows:

282

283 The ANFIS structure was developed by using the Takagi-Sugeno fuzzy rule base (the details
284 are presented in equations 10 and 11).

285

286 Rule 1: if x is A_1 and y is B_1 then $f_1=p_1x + q_1y + r_1$ (10)

287 Rule 2: if x is A_2 and y is B_2 then $f_2=p_2x + q_2y + r_2$ (11)

288

289 Here, x and y are the system inputs and A_1, A_2, B_1 and B_2 are fuzzy membership functions.

290 In addition, p_i, q_i and r_i ($\forall i = 1, 2$) are the parameters of the output function (Jang 1993). In

291 general, the ANFIS structure is made of five layers described below (Fig. 5):

292

293 Layer 1: This layer is responsible for the fuzzification of the variables, and the nodes in this

294 layer are adaptive nodes.

295 $O_{A_i}^1 = \mu_{A_i}(x), \quad i=1, 2$ (12)

296 $O_{B_i}^1 = \mu_{B_i}(y), \quad i=1, 2$ (13)

297 Here, i represents the related node and x and y its input variables, A_i and B_i are linguistic

298 terms and $\mu_{A_i}(x)$ and $\mu_{B_i}(y)$ the membership functions of the node i.

299

300 Layer 2: In this section, every node is a fixed node and each one is responsible for multiplying

301 signals entering it. The nodes are named by the Π label and their outputs are as follows (Oh

302 and Pradhan 2011):

303 $O_{2,i} = \mu_{A_i}(x) \cdot \mu_{B_i}(y) = W_i, \quad for \quad i = 1, 2$ (14)

304

305 Here, W_i (the so called firing strength of each fuzzy rule) represents each node's output.

306

307 Layer 3: This layer has the task of normalizing the output of the second layer. Therefore, the

308 nodes, which are fixed ones and named by the N label, normalize the input values (Equation

309 15). The numerator of the fraction includes the firing strength of each fuzzy rule, and the

310 denominator includes the total firing strength of each rule.

311 $O_{3,i} = \frac{W_i}{W_1+W_2} = \bar{W}_i \quad for \quad i = 1, 2$ (15)

312

313 Layer 4: This is considered the second adaptive layer in the ANFIS structure and each node's

314 output is obtained from the following equation:

315 $O_{4,i} = \bar{W}_i \cdot f_i = W_i \cdot (p_i x + q_i y + r_i) \text{ for } i = 1,2$ (16)

316

317 In this equation, \bar{W}_i is the normalized firing strength of the third layer. p_i , q_i and r_i are the
 318 variable parameters (also referred to as the result parameters) of the node i .

319

320 Layer 5: The only node existing in this layer is fixed node labeled Σ . This node sums up all the
 321 input signals and calculates the resulting output (Equation 17).

322 $O_{5,i} = \sum_i \bar{W}_i f_i = \frac{\sum_i W_i f_i}{\sum_i W_i}$ (17)

323

324 For more details on the layers and the algorithms, refer to Jang (1993) and Jang and Sun
 325 (1995). Figure 3 shows the process of the study including methods and type of combination
 326 used.

327

328 **Results and Validation**

329 Table 2 shows the weights obtained from the SWARA model and BWM. The values for the
 330 slope factor indicate that most of the landslides that occurred in the study area were of the 5-
 331 15° class with weights of 0.405 and 0.409, respectively. Among the different slope aspects, the
 332 north-east aspect, with the values of 0.486 (SWARA) and 0.249 (BWM), had the highest effect
 333 on landslide occurrence. In relation to the altitude factor, the 1500-1700 m class had the highest
 334 impact on landslide (with values of 0.454 and 0.212 for SWARA and BWM, respectively). The
 335 outputs of the SWARA model for the TWI showed that the 5.65-7.31 and 7.31-9.87 classes
 336 with values of 0.482 and 0.249, respectively, had the highest probabilities of landslide
 337 occurrence. For the BWM, the 5.65-7.31 and 7.31-9.87 classes with the weight of 0.371 had
 338 the greatest impact on landslide occurrence. For the plan curvature factor, according to Table
 339 2, the maximum weights obtained from the SWARA and BWM were for the convex class with
 340 weights of 0.410 and 0.769, respectively. For the profile curvature factor, the highest SWARA
 341 weight (0.489) was that of the concave class and for BWM the highest value (0.470) was that
 342 of the concave and convex classes. Results obtained from SWARA indicated that the distance
 343 to road, distance to stream and distance to fault in the 0-100 m, 0-100 m and 1200 - 1500 m
 344 classes with weights of 0.311, 0.404 and 0.386, respectively, had the highest influence on
 345 landslide occurrence. As in the SWARA method, in the BWM also the same classes had the
 346 highest weights with the values of 0.397, 0.297 and 0.330, respectively. Concerning lithology,
 347 the highest value in the SWARA method (0.344) was that of the J1 class and the highest in the

348 BWM (0.2) those of the JI and PIQc classes. For the land use factor, the agriculture class in
349 both models had the strongest relationship with landslide occurrence with values of 0.27 and
350 0.505, respectively. Landslides were more likely to occur with increases in rainfall. For the
351 precipitation factor, 332.9 - 387.65 mm of rainfall had the highest weights in the SWARA
352 model and BWM (0.352 and 0.647, respectively). In relation to the NDVI factor, the likelihood
353 of landslide occurrence was greatest for the class >0.5 with the weights of 0.356 and 0.574 for
354 the SWARA method and BWM, respectively.

355

356 *Integration of the ANFIS with SWARA and BWM*

357 In this study, MATLAB was employed to construct the ANFIS model and the SWARA method
358 and BWM to feed it for training the network. For this purpose, all the data were first divided
359 into the training and testing sets. As mentioned earlier 70% of the data (172 landslide locations)
360 were allocated for training and 30% (70 landslide locations) for testing, and they were assigned
361 the value of 1. Using the training data and the SWARA model and BWM, the weights of the
362 sub-criteria were calculated (Table 2). In the next step, 242 non-landslide points, showing the
363 total number of data, were created in the non-landslide areas. Then 0 was allocated to each of
364 them. Out of these non-landslide points, 70% (172) points were selected randomly and
365 considered for training the network. Next, 172 landslide and non-landslide points (with values
366 of 1 and 0) were overlaid upon the conditioning factors and the value of each one was
367 determined. This process was carried out once for the SWARA model and once for the BWM.
368 The values obtained from the overlaying were used as input data for ANFIS training. After
369 ANFIS training using the BWM method and SWARA, all the pixels were entered into
370 MATLAB and the final value of each pixel was determined using the created network. Finally,
371 landslide susceptibility maps were prepared for the hybrid ANFIS-BWM and ANFIS-SWARA
372 models (Fig. 7). The prepared maps were divided into five classes with sensitivity degree of
373 very low, low, moderate, high, and very high. Figure 8 shows the percent area for each class in
374 the ANFIS-SWARA and ANFIS-BWM hybrid models. It is quite clear that the class with very
375 high landslide susceptibility had the lowest area in both LSMs with values of 16.21% and
376 18.65%, respectively (Table 4). In addition, the classes with high and low landslide
377 susceptibility had the largest areas with the values of 23.01% and 20.50% for ANFIS-SWARA
378 and ANFIS-BWM, respectively.

379

380 *Models validation and comparison*

381 Validation is a very important step in estimating the accuracy of a method in producing
382 landslide susceptibility maps. In this study, validation was performed by using 30% of landslide
383 and non-landslide locations (72 points with values of 0 and 1) in two stages. In the first stage,
384 the mean square error was calculated to estimate the accuracy of ANFIS trained network using
385 SWARA and BWM methods. MSE and RMSE are defined as follows:

$$387 \quad MSE = \frac{1}{n} \sum_{j=1}^n (T_j - \bar{T}_j)^2 \quad (18)$$

$$388 \quad RMSE = \sqrt{\frac{1}{n} \sum_{j=1}^n (T_j - \bar{T}_j)^2} \quad (19)$$

389
390 where, T_j is the target values and \bar{T}_j is the output values and n is the total number of samples.
391 RMSE is the square root of MSE.

392
393 The lower MSE value is (closer to zero) the lower the amount of error in the final prediction
394 and hence the more accuracy the modeling will be. Fig. 6c shows MSE and RMSE for the test
395 dataset. The results showed that the MSE values for the ANFIS-SWARA and ANFIS-BWM
396 models are 0.299 and 0.242, respectively. As the results indicate, the new BWM method
397 outperformed the SWARA model in training the ANFIS.

398 In the second stage the LSMs were evaluated using the ROC curve. The ROC curve is a
399 graphical representation of the balance between negative and positive error values that can
400 quantitatively estimate the model accuracy. The area under the curve (AUC) illustrates the
401 predicted value of the system by describing its ability in correctly estimating the occurrence of
402 the event (landslide) and the non-occurrence of the event (non-landslide) (Yan et al. 2018).
403 Therefore, the larger the area under (AUC) the curve is the more accurate the model will be
404 and the lower AUC show weak performance of the model. Further details on this curve for
405 validating landslide susceptibility maps are provided in articles by Pourghasemi et al. 2013 and
406 Fan et al. 2017.

407 In this study, 72 landslide and non-landslide points were overlaid upon the conditioning
408 factors to plot the ROC curves. The values obtained for each point were then used as input data.
409 Figure 9 shows the ROC curves for the methods. In this figure, the areas under the curves for
410 the ANFIS-SWARA and ANFIS-BWM hybrid models are 0.736 and 0.75, respectively. The
411 results obtained from the evaluation of the zoning suggested that both models were able to

412 predict the landslide prone areas well; however, the ANFIS-BWM model was more accurate
413 and, hence, yielded more reliable outputs.

414

415

416 **Discussion**

417 Landslide spatial modeling is a nonlinear and complex problem because it is affected by
418 various parameters. Although there are different methods for landslide susceptibility zoning,
419 the important point is which method or combination of methods should be used, depending on
420 the features of the study area, in order to obtain the best results. The novelty of this study is to
421 produce a new hybrid ANFIS-BWM model and compare it with ANFIS-SWARA to determine
422 the best combination. For this purpose, a database was first developed to contain the inventory
423 map and factors influencing landslides. In the next step, the SWARA and BWM approaches
424 were then executed by using the database and experts' opinions, and the weights of the sub-
425 criteria were determined. The weights obtained from each of the two approaches were entered
426 separately in MATLAB to train ANFIS. After developing the network and calculating the final
427 value of each pixel, the LSMs were generated, and the accuracy of each map was assessed by
428 using the ROC curve. The resultant validation output indicated that the new BWM produced
429 more realistic results than the SWARA method which trained the ANFIS model well and
430 obtained an acceptable output from it. Since bringing the prediction closer to reality is the most
431 important objective in complex environmental issues such as landslides, it is necessary to
432 compare newly introduced methods with the previous ones in order to achieve more optimal
433 results.

434 The research models have attracted the interest of landslide researchers. [Oh and Pradhan.](#)
435 [\(2011\)](#) employed the ANFIS model with the triangular, trapezoidal, Gaussian, 2-sided
436 Gaussian, generalized bell, sigmoid 1, sigmoid 2 and polynomial membership functions for
437 landslide spatial modeling on Penang Island in Malaysia. For this purpose, they first determined
438 the relationships between the eight effective factors and the previous landslides by using the
439 frequency ratio method. The ANFIS model was then implemented in MATLAB and to estimate
440 the values obtained from the frequency ratio method and its accuracy through the ROC curve.
441 They showed that the triangular, trapezoidal, generalized bell and polynomial membership
442 functions were slightly more accurate than the others, although all the membership functions
443 that were used yielded very good results with accuracies higher than 84%. In order landslide
444 spatial modelling for Iran, [Dehnavi et al. \(2015\)](#) integrated the SWARA multi-criteria decision-
445 making approach with the ANFIS method. In the first step, the SWARA model was used to

446 specify the weights of sub-criteria, and the LSM was then generated. In the second step, the
447 SWARA model weights were utilized to train the ANFIS and generate the map for the hybrid
448 ANFIS-SWARA landslide susceptibility model. They found that the ANFIS-SWARA hybrid
449 model with the area under the curve of 0.8 yielded a more accurate prediction than the SWARA
450 method with the area under the curve of 0.78. [Gigovic et al. \(2019\)](#) integrated the BWM with
451 the WLC and OWA methods for zoning regional landslides in western Serbia. For this purpose,
452 they first created a database that included 15 conditioning factors and 1082 previous landslides.
453 The normalized weights of the effective factors were then calculated by using the BWM. The
454 WLC and OWA methods were then employed to produce LSMs. Finally, the performance of
455 each map was assessed by through the ROC curve. They showed that the OWA method with
456 94.1% accuracy outperformed the WLC method that had the accuracy of 90.5%.

457 In this study, the important of using new models to improve the performance of machine
458 learning methods was shown. Based on the results shown in table 2, although both methods are
459 of the type of MCDM and include values between 0 and 0.5, the new BWM model performs
460 better compared to SWARA model. Therefore, the type of model that is used to determine the
461 correlation between conditioning factors and the landslide occurrence is effective in improving
462 the results. According to the research findings, the employed methods performed well in
463 estimating landslide prone areas. In addition, the use of the two ANFIS-BWM and ANFIS-
464 SWARA hybrid models with the accuracies of 75% and 73.6%, respectively, confirmed this
465 statement. It is recommended to integrate SWARA and BWM multi-criteria decision-making
466 models with machine learning methods such as ANFIS for other similar areas.

467

468 **Conclusion**

469 Known as natural destructive ground-deforming phenomena, landslides have occurred in all
470 historical periods. Evaluation of landslide susceptibility maps is a multistage and complicated
471 process and, therefore, the careful selection and execution of the stages of the model will result
472 in generation of maps yielding better results.

473 In the present study, a new comparison was drawn between two multi-criteria decision-
474 making methods to train the ANFIS model. For this purpose, it was decided to determine the
475 relationships between landslides and effective factors by using the new BWM and the SWARA
476 model. The weights obtained from both methods were then employed to train the ANFIS.
477 Finally, the landslide susceptibility maps were generated using the ANFIS-SWARA and
478 ANFIS-BWM ensemble models. The results of validation using the ROC curve showed that
479 the ANFIS-BWM model with a 75% prediction accuracy outperformed the ANFIS-SWARA

480 model with a 73.6% prediction accuracy. Although both the BWM and the SWARA technique
481 were multi-criteria decision-making models, their outputs differed in types of ranking and
482 weighting. Therefore, it is essential to decide the output of which method should be utilized to
483 train a machine learning model. Since these models yielded good results, they are
484 recommended for use in other similar areas because they can substantially help land use
485 managers and planners in making important decisions.

486

487 **Declarations**

488 **Ethics approval and consent to participate.**

489 Not applicable.

490 **Consent for publication**

491 Not applicable.

492 **Availability of data and materials**

493 All data generated or analyzed during this study are confidential.

494 **Competing interests**

495 The authors declare that they have no competing interests.

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498 **Code availability** (software application or custom code)

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500 **Authors' contributions**

501 All authors contributed to the study conception and design. Data collection and material
502 preparation were performed by SP. Development and design of methodology and creation of
503 models were performed by SP and AN. BP consulted for the methodology application. AN
504 was the supervisors of the work. The first draft of the manuscript was written by SP and AN
505 and BP commented on previous versions of the manuscript.

506 **References**

507 Achour Y, Boumezbeur A, Hadji R, Chouabbi A, Cavaleiro V, Bendaoud EA (2017) Landslide
508 susceptibility mapping using analytic hierarchy process and information value methods along
509 a highway road section in Constantine, Algeria. Arab J Geosci. 10:194.

510

511 Adineh F, Motamedvaziri B, Ahmadi H, Moeini A (2018) Landslide susceptibility mapping
512 using Genetic Algorithm for the Rule Set Production (GARP) model. J Mt Sci. 15:2013-2026.

513

514 Aghdam IN, Pradhan B, Panahi M (2017) Landslide susceptibility assessment using a novel
515 hybrid model of statistical bivariate methods (FR and WOE) and adaptive neuro-fuzzy
516 inference system (ANFIS) at southern Zagros Mountains in Iran. Environ Earth Sci. 76:237.

517

518 Ahmad WNKW, Rezaei J, Sadaghiani S, Tavasszy LA (2017) Evaluation of the external forces
519 affecting the sustainability of oil and gas supply chain using Best Worst Method. J Clean
520 Prod. 153:242-252.

521

522 Ahmed B (2015) Landslide susceptibility mapping using multi-criteria evaluation techniques
523 in Chittagong Metropolitan Area, Bangladesh. Landslides 12:1077-1095.

524

525 Arabameri A, Pradhan B, Rezaei K, Sohrabi M, Kalantari Z (2019) GIS-based landslide
526 susceptibility mapping using numerical risk factor bivariate model and its ensemble with linear
527 multivariate regression and boosted regression tree algorithms. J Mt Sci. 16:595-618.

528

529 Baena JAP, Scifoni S, Marsella M, De Astis G, Fernández CI (2019) Landslide susceptibility
530 mapping on the islands of Vulcano and Lipari (Aeolian Archipelago, Italy), using a multi-
531 classification approach on conditioning factors and a modified GIS matrix method for areas
532 lacking in a landslide inventory. Landslides 16:969-982.

533

534 Berhane G, Kebede M, Alfarrah N (2020) Landslide susceptibility mapping and rock slope
535 stability assessment using frequency ratio and kinematic analysis in the mountains of Mgulat
536 area, Northern Ethiopia. Bull Eng Geol Environ. 1-17. [https://doi.org/10.1007/s10064-020-](https://doi.org/10.1007/s10064-020-01905-9)
537 01905-9

538
539
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551
552
553
554
555
556
557
558
559
560
561
562
563
564
565
566
567
568
569
570

Bui DT, Pradhan B, Lofman O, Revhaug I, Dick OB (2012) Spatial prediction of landslide hazards in Hoa Binh province (Vietnam): a comparative assessment of the efficacy of evidential belief functions and fuzzy logic models. *Catena* 96:28-40.

Bui DT, Tuan TA, Klempe H, Pradhan B, Revhaug I (2016) Spatial prediction models for shallow landslide hazards: a comparative assessment of the efficacy of support vector machines, artificial neural networks, kernel logistic regression, and logistic model tree. *Landslides* 13:361-378.

Chalkias C, Ferentinou M, Polykretis C (2014) GIS-based landslide susceptibility mapping on the Peloponnese Peninsula, Greece. *Geosciences* 4:176-190.

Colkesen I, Sahin EK, Kavzoglu T (2016) Susceptibility mapping of shallow landslides using kernel-based Gaussian process, support vector machines and logistic regression. *J Afr Earth Sci.* 118:53-64.

Conforti M, Pascale S, Robustelli G, Sdao F (2014) Evaluation of prediction capability of the artificial neural networks for mapping landslide susceptibility in the Turbolo River catchment (northern Calabria, Italy). *Catena* 113:236-250.

Cui K, Lu D, Li W (2017) Comparison of landslide susceptibility mapping based on statistical index, certainty factors, weights of evidence and evidential belief function models. *Geocarto Int.* 329:935-955.

Dehnavi A, Aghdam IN, Pradhan B, Varzandeh MHM (2015) A new hybrid model using step-wise weight assessment ratio analysis (SWARA) technique and adaptive neuro-fuzzy inference system (ANFIS) for regional landslide hazard assessment in Iran. *Catena* 135:122-148.

Ding Q, Chen W, Hong H (2017) Application of frequency ratio, weights of evidence and evidential belief function models in landslide susceptibility mapping. *Geocarto Int.* 32:619-639.

571 Dou J, Oguchi T, Hayakawa YS, Uchiyama S, Saito H, Paudel U (2014) GIS-based landslide
572 susceptibility mapping using a certainty factor model and its validation in the Chuetsu Area,
573 Central Japan. In *Landslide science for a safer geoenvironment*. p. 419-424.
574

575 Du G, Zhang Y, Yang Z, Guo C, Yao X, Sun D (2019) Landslide susceptibility mapping in the
576 region of eastern Himalayan syntaxis, Tibetan Plateau, China: a comparison between analytical
577 hierarchy process information value and logistic regression-information value methods. *Bull*
578 *Eng Geol Environ*. 78(6):4201-4215.
579

580 Erener A, Düzgün HSB (2010) Improvement of statistical landslide susceptibility mapping by
581 using spatial and global regression methods in the case of More and Romsdal
582 (Norway). *Landslides* 7:55-68.
583

584

585 Erener A, Mutlu A, Düzgün HS (2016) A comparative study for landslide susceptibility
586 mapping using GIS-based multi-criteria decision analysis (MCDA), logistic regression (LR)
587 and association rule mining (ARM). *Eng Geol*. 203:45-55.
588

589 Fan W, Wei XS, Cao YB, Zheng, B (2017) Landslide susceptibility assessment using the
590 certainty factor and analytic hierarchy process. *J Mt Sci*. 14:906-925.
591

592 Farrokhnia A, Pirasteh S, Pradhan B, Pourkermani M, Arian M (2011) A recent scenario of
593 mass wasting and its impact on the transportation in Alborz Mountains, Iran using geo-
594 information technology. *Arab J Geosci*. 4:1337-1349.
595

596 Feizizadeh B, Roodposhti MS, Blaschke T, Aryal J (2017) Comparing GIS-based support
597 vector machine kernel functions for landslide susceptibility mapping. *Arab J Geosci*. 10:122.
598

599 Gigović L, Drobnjak S, Pamučar D (2019) The application of the hybrid GIS spatial multi-
600 criteria decision analysis best – worst methodology for landslide susceptibility
601 mapping. *ISPRS Int J Geo-Inf*. 8:79.
602

603 Gutiérrez F, Linares R, Roqué C, Zarroca M, Carbonel D, Rosell J, Gutiérrez M (2015) Large
604 landslides associated with a diapiric fold in Canelles Reservoir (Spanish Pyrenees): Detailed

605 geological – geomorphological mapping, trenching and electrical resistivity
606 imaging. *Geomorphology* 241:224-242.

607

608 Grabs T, Seibert J, Bishop K, Laudon H (2009) Modeling spatial patterns of saturated areas: A
609 comparison of the topographic wetness index and a dynamic distributed model. *J*
610 *Hydrol.* 373:15-23.

611

612 Hall FG, Townshend JR, Engman ET (1995) Status of remote sensing algorithms for estimation
613 of land surface state parameters. *Remote Sens. Environ.* 51:138-156.

614

615 Hong H, Chen W, Xu C, Youssef AM, Pradhan B, Tien Bui D (2017) Rainfall-induced
616 landslide susceptibility assessment at the Chongren area (China) using frequency ratio,
617 certainty factor, and index of entropy. *Geocarto Int.* 322:139-154.

618

619 Jang JS (1993) ANFIS: adaptive-network-based fuzzy inference system. *IEEE Trans. Syst.*
620 *Man Cybern.* 23:665-685.

621

622 Juliev M, Mergili M, Mondal I, Nurtaev B, Pulatov A, Hübl J (2019) Comparative analysis of
623 statistical methods for landslide susceptibility mapping in the Bostanlik District,
624 Uzbekistan. *Sci Total Environ.* 653:801-814.

625

626 Kavzoglu T, Sahin EK, Colkesen I (2014) Landslide susceptibility mapping using GIS-based
627 multi-criteria decision analysis, support vector machines, and logistic
628 regression. *Landslides.* 11:425-439.

629

630 Kayastha P, Dhital MR, De Smedt F (2012) Landslide susceptibility mapping using the weight
631 of evidence method in the Tinau watershed, Nepal. *Nat Hazards.* 63:479-498.

632

633 Keršulienė V, Turskis Z (2011) Integrated fuzzy multiple criteria decision-making model for
634 architect selection. *Technol. Econ Dev Econ.* 17:645-666.

635

636 Kim JC, Lee S, Jung HS, Lee S (2018) Landslide susceptibility mapping using random forest
637 and boosted tree models in Pyeong-Chang, Korea. *Geocarto Int.* 339:1000-1015.

638

639 Kouli M, Loupasakis C, Soupios P, Rozos D, Vallianatos F (2014) Landslide susceptibility
640 mapping by comparing the WLC and WofE multi-criteria methods in the West Crete Island,
641 Greece. *Environ Earth Sci.* 72:5197-5219.

642

643 Kumar S, Srivastava PK, Snehmani (2017) GIS-based MCDA-AHP modeling for avalanche
644 susceptibility mapping of Nubra valley region, Indian Himalaya. *Geocarto Int.* 32:1254-1267.

645

646 Lin L, Lin Q (2017) Landslide susceptibility mapping on a global scale using the method of
647 logistic regression. *Nat Hazards Earth Syst. Sci.* 17:1411-1424.

648

649 Mardani A, Zavadskas EK, Khalifah Z, Zakuan N, Jusoh A, Nor KM, Khoshnoudi M (2017)
650 A review of multi-criteria decision-making applications to solve energy management
651 problems: Two decades from 1995 to 2015. *Renew Sust Energ Rev.* 71: 216-256.

652

653 Moayedi H, Mehrabi M, Mosallanezhad M, Rashid ASA, Pradhan B (2019) Modification of
654 landslide susceptibility mapping using optimized PSO-ANN technique. *Eng Comput.* 35:967-
655 984.

656

657 Moosavi V, Niazi Y (2016) Development of hybrid wavelet packet-statistical models (WP-
658 SM) for landslide susceptibility mapping. *Landslides* 13:97-114.

659

660 Nadim F, Kjekstad O, Peduzzi P, Herold C, Jaedicke C (2006) Global landslide and avalanche
661 hotspots. *Landslides* 3:159-173.

662

663 Oh HJ, Pradhan B (2011) Application of a neuro-fuzzy model to landslide-susceptibility
664 mapping for shallow landslides in a tropical hilly area. *Comput Geosci.* 37:1264-1276.

665

666 Oh HJ, Kadavi PR, Lee CW, Lee S (2018) Evaluation of landslide susceptibility mapping by
667 evidential belief function, logistic regression and support vector machine models. *GEOMAT*
668 *Nat Haz Risk.* 9:1053-1070.

669

670 Paryani S, Neshat A, Javadi S, Pradhan B (2020) Comparative performance of new hybrid
671 ANFIS models in landslide susceptibility mapping. *Nat Hazards.* 1-28.
672 <https://doi.org/10.1007/s11069-020-04067-9>

673

674 Pham, B. T., Bui, D. T., Pourghasemi, H. R., Indra, P., & Dholakia, M. B. (2017). Landslide
675 susceptibility assessment in the Uttarakhand area (India) using GIS: a comparison study of
676 prediction capability of naïve bayes, multilayer perceptron neural networks, and functional
677 trees methods. *Theoretical and Applied Climatology*, 128(1-2), 255-273.

678

679 Pham BT, Prakash I (2018) Machine learning methods of kernel logistic regression and
680 classification and regression trees for landslide susceptibility assessment at part of Himalayan
681 area, India. *Indian J Sci Technol.* 11:1-11.

682

683 Pham BT, Prakash I (2019) Evaluation and comparison of LogitBoost Ensemble, Fisher's
684 Linear Discriminant Analysis, logistic regression and support vector machines methods for
685 landslide susceptibility mapping. *Geocarto Int* 34:316-333.

686

687 Polykretis C, Ferentinou M, Chalkias C (2015) A comparative study of landslide susceptibility
688 mapping using landslide susceptibility index and artificial neural networks in the Krios River
689 and Krathis River catchments (northern Peloponnesus, Greece). *Bull Eng Geol Environ.* 74(1),
690 27-45.

691

692 Polykretis C, Chalkias C, Ferentinou M (2019) Adaptive neuro-fuzzy inference system
693 (ANFIS) modeling for landslide susceptibility assessment in a Mediterranean hilly area. *B Eng*
694 *Geol Environ.* 78:1173-1187.

695

696 Pourghasemi HR, Moradi HR, Aghda SF (2013) Landslide susceptibility mapping by binary
697 logistic regression, analytical hierarchy process, and statistical index models and assessment
698 of their performances. *Nat Hazards.* 69:749-779.

699

700 Pourghasemi, H. R., & Rossi, M. (2017). Landslide susceptibility modeling in a landslide prone
701 area in Mazandarn Province, north of Iran: a comparison between GLM, GAM, MARS, and
702 M-AHP methods. *Theoretical and Applied Climatology*, 130(1-2), 609-633.

703

704 Pradhan B (2011) Manifestation of an advanced fuzzy logic model coupled with Geo-
705 information techniques to landslide susceptibility mapping and their comparison with logistic
706 regression modelling. *Environ Ecol Stat.* 18:471-493.

707
708 Pradhan B (2013) A comparative study on the predictive ability of the decision tree, support
709 vector machine and neuro-fuzzy models in landslide susceptibility mapping using
710 GIS. *Comput Geosci*. 51:350-365.
711
712 Pradhan B, Seeni MI, Kalantar B (2017) Performance evaluation and sensitivity analysis of
713 expert-based, statistical, machine learning, and hybrid models for producing landslide
714 susceptibility maps. In: *Laser scanning applications in landslide assessment*. Springer
715 international publishing, pp. 193-232. Springer, Cham.
716
717 Umar Z, Pradhan B, Ahmad A, Jebur MN, Tehrani MS (2014) Earthquake induced landslide
718 susceptibility mapping using an integrated ensemble frequency ratio and logistic regression
719 models in West Sumatera Province, Indonesia. *Catena* 118:124-135.
720
721 Ramesh V, Anbazhagan S (2015) Landslide susceptibility mapping along Kolli hills Ghat road
722 section (India) using frequency ratio, relative effect and fuzzy logic models. *Environ Earth Sci*.
723 73:8009-8021.
724
725 Regmi AD, Devkota KC, Yoshida K, Pradhan B, Pourghasemi HR, Kumamoto T, Akgun A
726 (2014) Application of frequency ratio, statistical index, and weights-of-evidence models and
727 their comparison in landslide susceptibility mapping in Central Nepal Himalaya. *Arab J*
728 *Geosci*. 7:725-742.
729
730 Rezaei J (2015) Best-worst multi-criteria decision-making method. *Omega* 53:49-57.
731
732 Rozos D, Pyrgiotis L, Skias S, Tsagaratos P (2008) An implementation of rock engineering
733 system for ranking the instability potential of natural slopes in Greek territory. An application
734 in Karditsa County. *Landslides* 5:261-270.
735
736 Sharma S, Mahajan AK (2019) A comparative assessment of information value, frequency
737 ratio and analytical hierarchy process models for landslide susceptibility mapping of a
738 Himalayan watershed, India. *Bull Eng Geol Environ* 78(4):2431-2448.
739

740 Turan, İ. D., Özkan, B., Türkeş, M., & Dengiz, O. (2020). Landslide susceptibility mapping
741 for the Black Sea Region with spatial fuzzy multi-criteria decision analysis under semi-humid
742 and humid terrestrial ecosystems. *Theoretical and Applied Climatology*, 1-14.
743

744 Wang LJ, Guo M, Sawada K, Lin J, Zhang J (2015) Landslide susceptibility mapping in
745 Mizunami City, Japan: A comparison between logistic regression, bivariate statistical analysis
746 and multivariate adaptive regression spline models. *Catena* 135:271-282.
747

748 Xiao T, Yin K, Yao T, Liu S (2019) Spatial prediction of landslide susceptibility using GIS-
749 based statistical and machine learning models in Wanzhou County, Three Gorges Reservoir,
750 China. *Acta Geochimica*, 38:654-669.
751

752 Yan F, Zhang Q, Ye S, Ren B (2019) A novel hybrid approach for landslide susceptibility
753 mapping integrating analytical hierarchy process and normalized frequency ratio methods with
754 the cloud model. *Geomorphology* 327:170-187.
755

756 Youssef, A. M., Pradhan, B., Jebur, M. N., & El-Harbi, H. M. (2015). Landslide susceptibility
757 mapping using ensemble bivariate and multivariate statistical models in Fayfa area, Saudi
758 Arabia. *Environmen. Earth Sci.* 73:3745-3761.
759

760 Zhang G, Cai Y, Zheng Z, Zhen J, Liu Y, Huang K (2016) Integration of the statistical index
761 method and the analytic hierarchy process technique for the assessment of landslide
762 susceptibility in Huizhou, China. *Catena* 142:233-244.
763

764 Zhu AX, Miao Y, Wang R, Zhu T, Deng Y, Liu J, Lin y, Qin CZ, Hong H (2018) A comparative
765 study of an expert knowledge-based model and two data-driven models for landslide
766 susceptibility mapping. *Catena* 166:317-327.

Figures

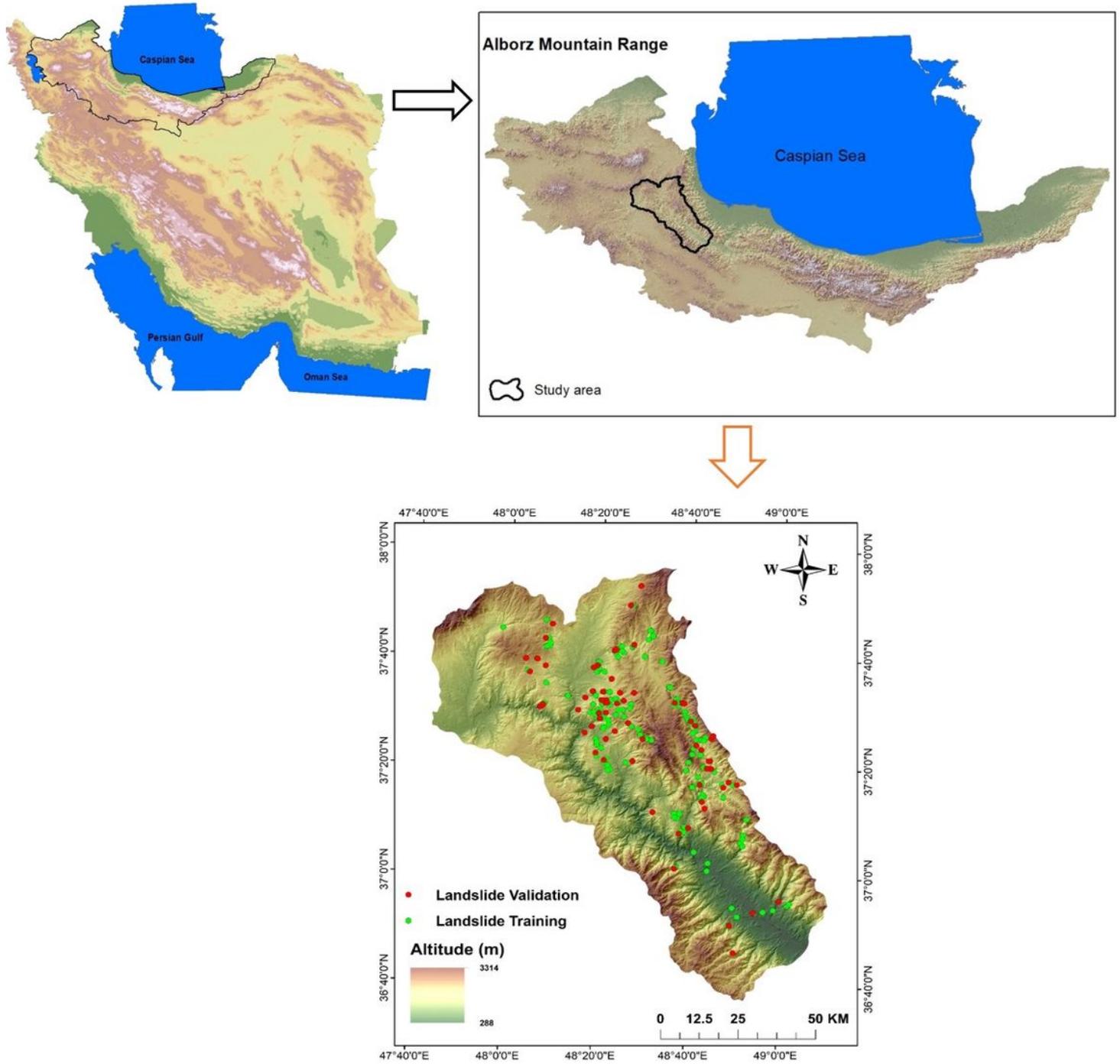


Figure 1

Location of the study area

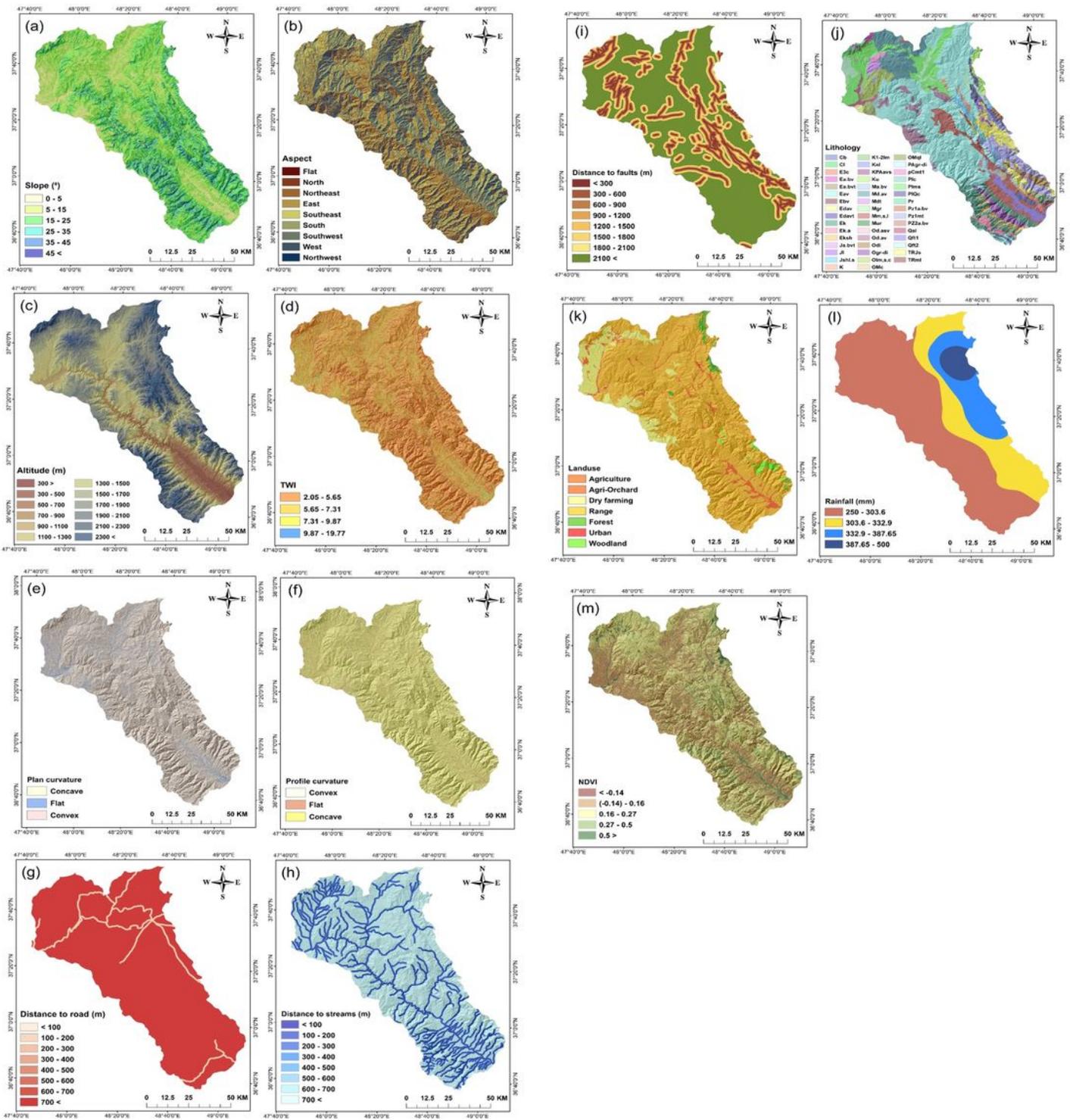


Figure 2

Produced conditioning factors of the study; a, slope angle; b, slope aspect; c, altitude; d, TWI; e, plan curvature; f, profile curvature; g, distance to roads; h, distance to streams; i, distance to faults; j, lithology; k, land use; l, rainfall; m, NDVI

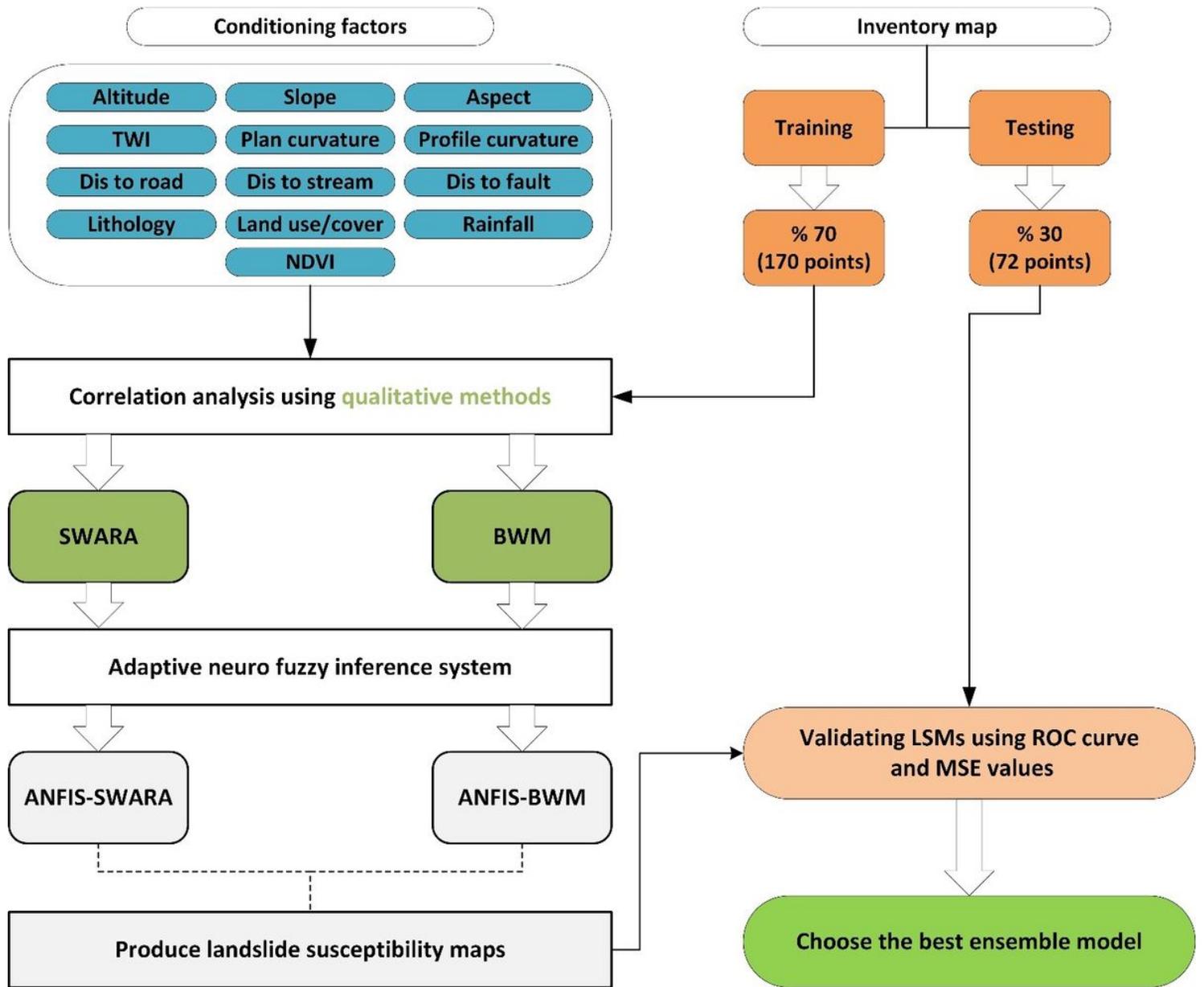


Figure 3

Flowchart of the study that showing all steps

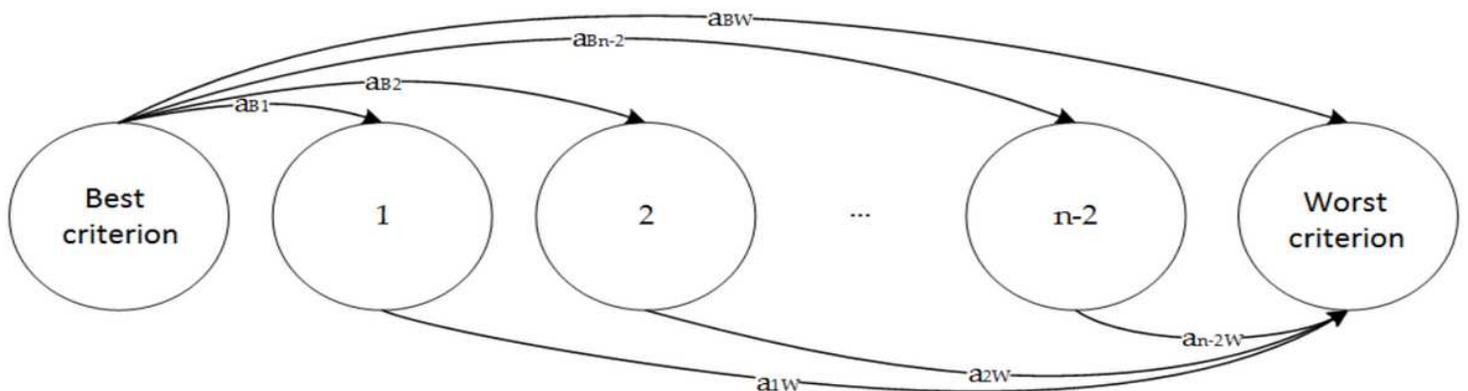


Figure 4

Reference comparison for BWM model

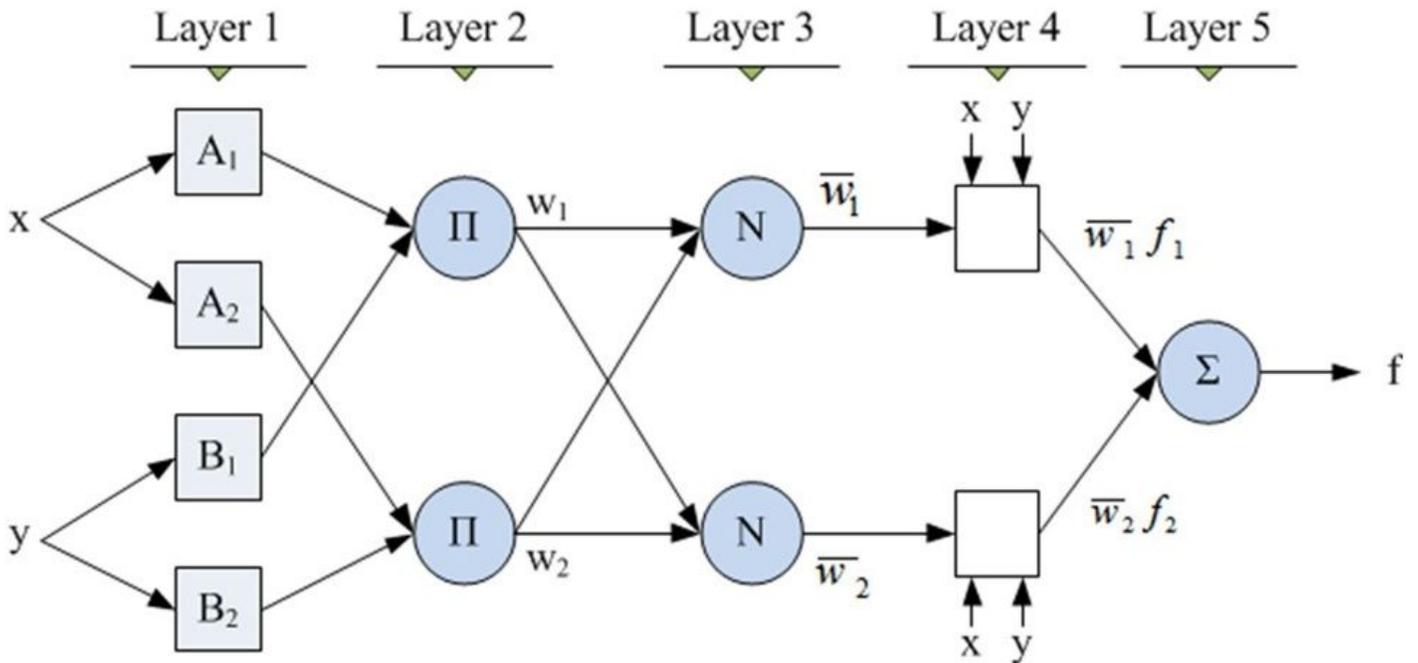


Figure 5

The structure of ANFIS model

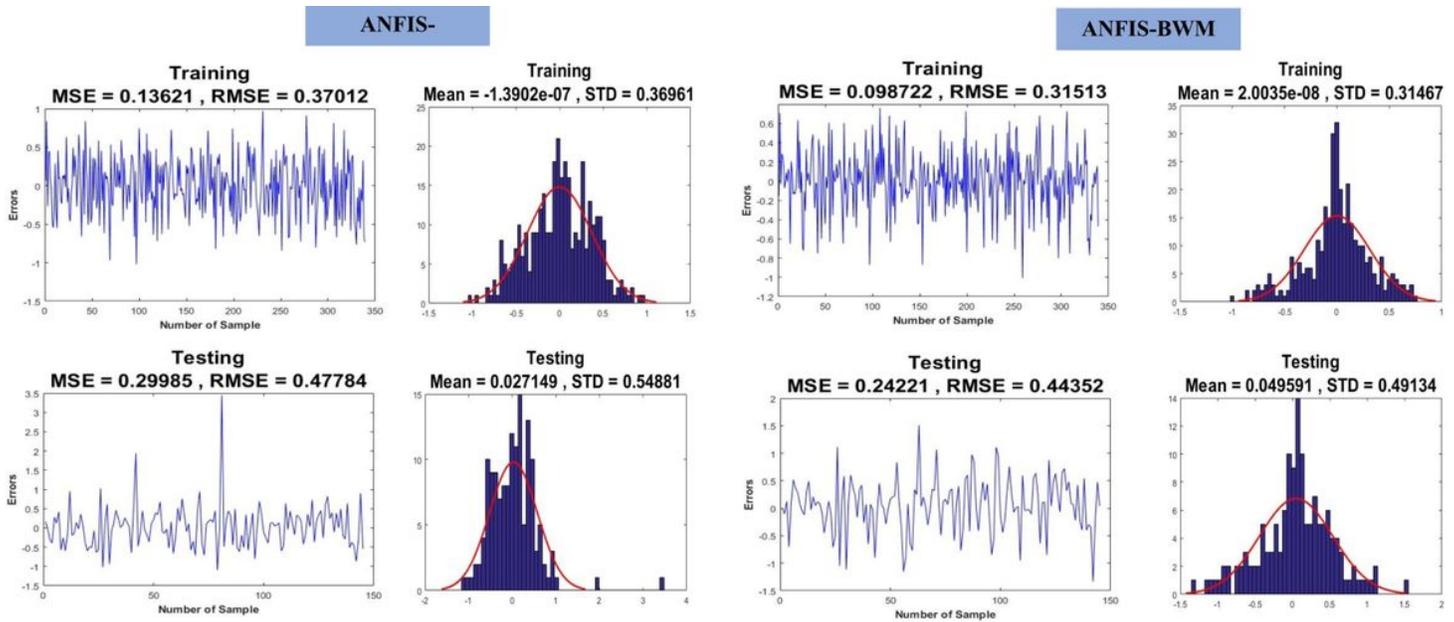


Figure 6

ANFIS-SWARA and ANFIS-BWM training and testing datasets a) MSE and RMSE value in the training phase; b) frequency errors in the training phase; c) MSE and RMSE value in the testing phase d) frequency errors in the testing phase

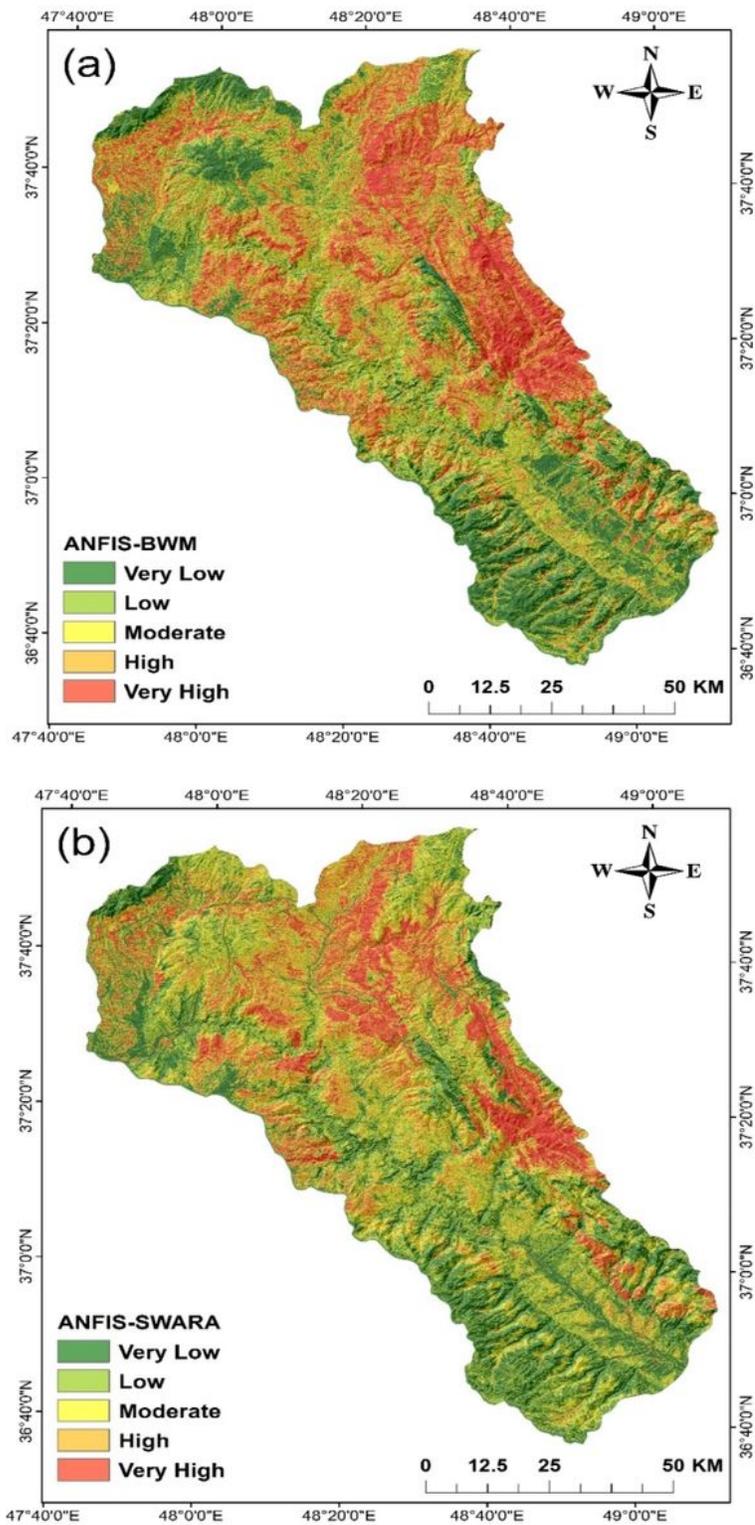


Figure 7

Landslide susceptibility map produced by ANFIS-BWM (a) and ANFIS-SWARA (b) models

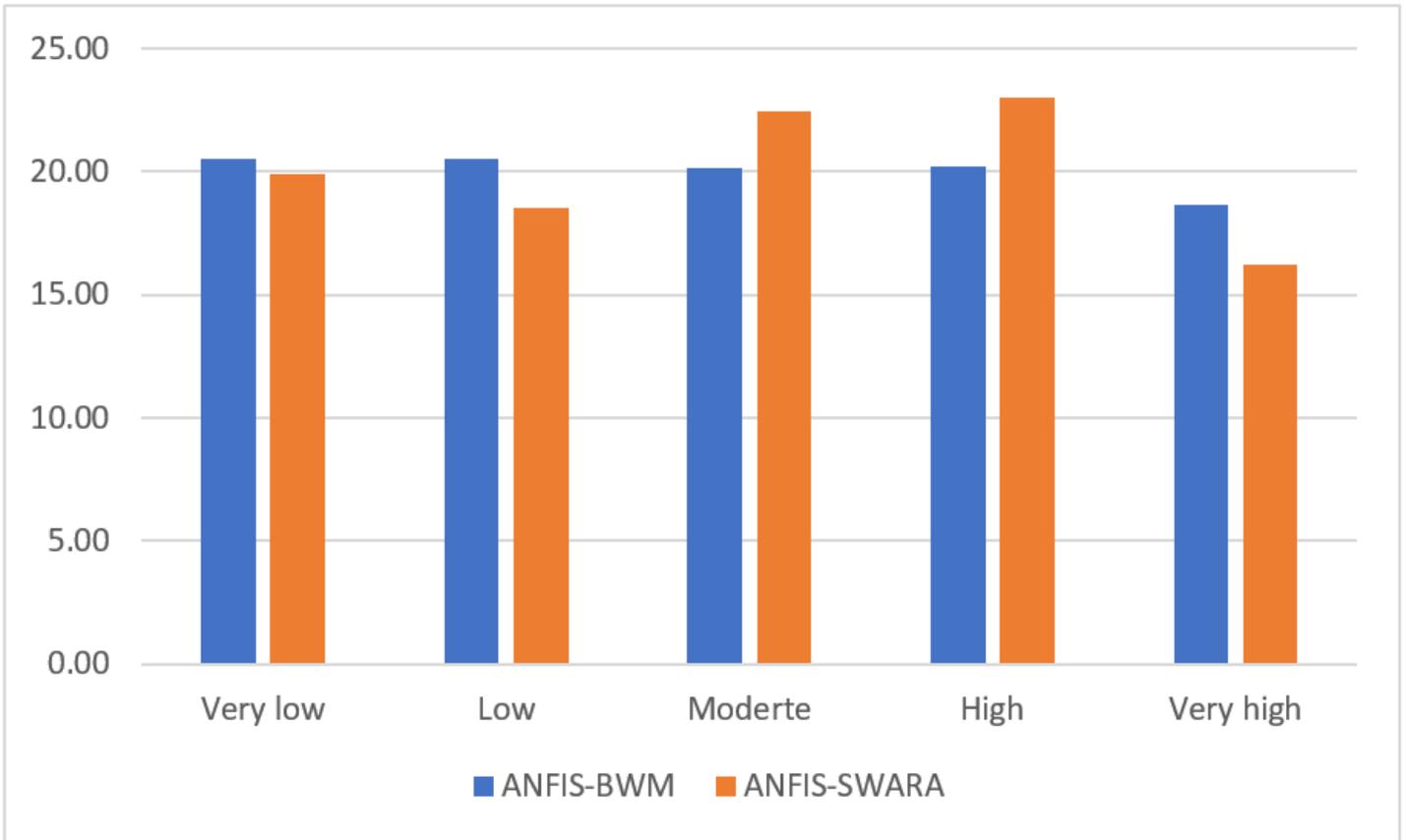


Figure 8

Percentages of landslide susceptibility classes

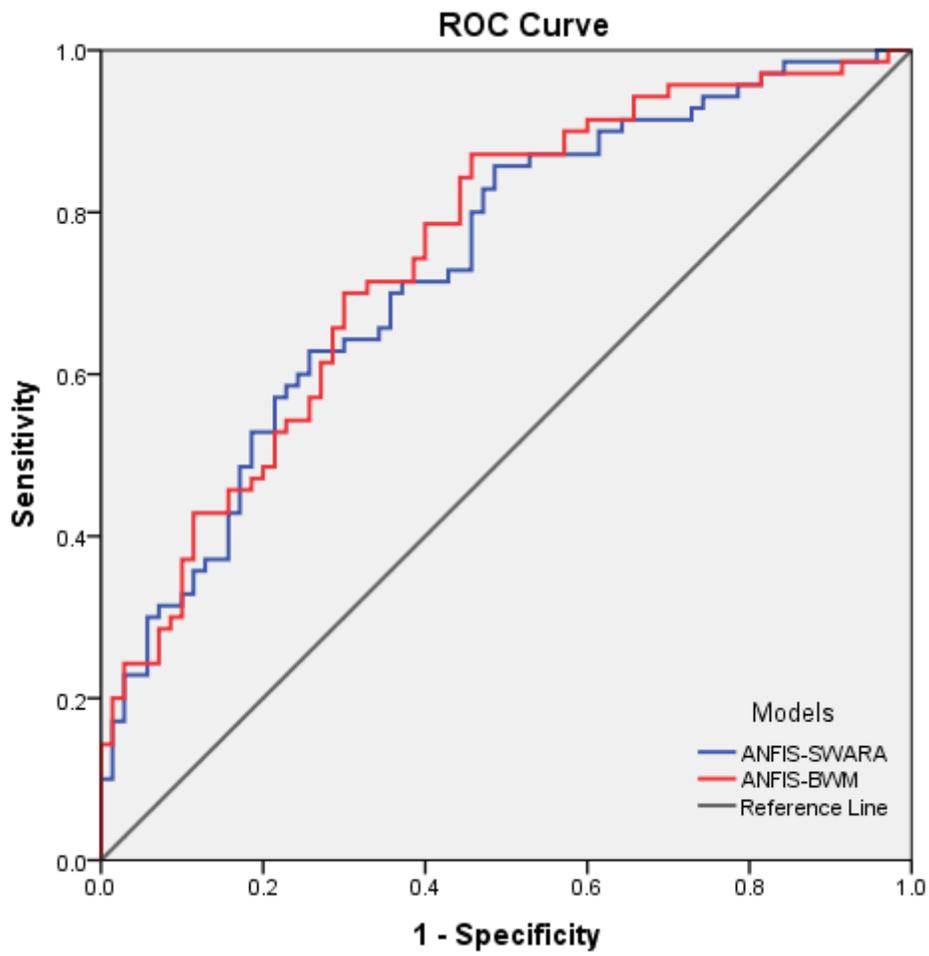


Figure 9

receiver operating characteristic curve for validation

Supplementary Files

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- [GraphicalAbstract.docx](#)
- [Highlight.docx](#)