

Relationship Between Induced Polarization Relaxation Time and Hydraulic Characteristics of Water-Bearing Sand

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44 **Relationship between induced polarization relaxation time and hydraulic characteristics** 45 **of water-bearing sand**

46 **Abstract:** Induced polarization method has become a popular method for evaluating formation
47 permeability characteristics in recent years because of its sensitivity to water body and water-
48 bearing pore structure. Especially, the induced polarization relaxation time can reflect the
49 macroscopic characteristics of the pore structure of rock and soil. Therefore, in order to study
50 the relationship between relaxation time and permeability, eight different sizes of quartz sand
51 were used to simulate water-bearing sand layers under different working conditions, and the
52 induced polarization experiment and Darcy seepage experiment were carried out on the same
53 sand sample in this paper, respectively. The experimental results show that the relation time
54 and the evolution of the permeability are closely correlated with the sizes of quartz sand.
55 According to the experimental data, with the particle size of the quartz sand as the link, the
56 power function equation is fitted to better describe the relationship between the permeability
57 and the relation time. It is worth noting that the equations obtained are only empirical
58 equations for quartz sand and are not suitable for general applications.

59 **Keywords:** Induced polarization method; Darcy seepage experiment; Permeability; Relaxation time

60 **1. Introduction**

61 With the increase of number of deep and long tunnels, disaster-causing water-bearing structures
62 have caused water and mud bursts in front of tunnels, it has become an important problem in some tunnels.
63 The water permeability of the water-bearing structure, that is, the permeability characteristics, directly
64 determines the magnitude and scale of water inrush. It is essential to estimate the magnitude and scale
65 of water inrush with proper evaluation of the permeability characteristics of water-bearing structures
66 (Ni et al., 2010). In general, laboratory measurement of borehole sampling, field experiment methods
67 such as pumping experiment or pressure experiment are being used to understand the permeability
68 characteristics of water-bearing structures in front of tunnel face (Attwa et al., 2013). However, these
69 methods are expensive and due to the limitation of the number of samples and ex-situ experiments, their
70 results often have hysteresis and one-sidedness. The induced polarization method has the advantage of
71 being sensitive to the pore structure of the formation and not being affected by topographical factors, and

72 it is more and more used to predict the hydraulic characteristics of the formation (Slater, 2007).

73 The induced polarization method is a geophysical method based on the induced polarization effect
74 of the geological body, which is based on the difference of induced polarization parameters between
75 different rock and soil media. Observing the complex conductivity and polarizability of the rock-soil
76 medium, we can get information such as the real part characterizing the charge conduction characteristics
77 and the imaginary part characterizing the charge storage characteristics in the complex conductivity.
78 Permeability determines the resistance of porous materials to fluid flow (Lesmes et al., 2001; Revil et al.,
79 2010). In the detection of disaster-causing water-bearing structures in tunnels, the fluid is usually
80 groundwater and its properties are relatively stable. Therefore, the permeability (k) can be used to
81 evaluate the permeability of the formation. In 1957, the induced polarization method was first proved to
82 be applicable to groundwater detection, and pointed out that the permeability of shallow aquifers can be
83 evaluated by the relevant parameters of the induced polarization attenuation curve (Vacquier et al., 1957).
84 In the past ten years or so, more and more articles have shown that the imaginary component of the
85 complex conductivity of rock and soil is positively correlated with the specific surface area of the
86 medium. Combined with the Kozeny-Carman formula, the relationship between permeability, porosity
87 and specific surface area can be obtained (Borner et al., 1991; Schon, 2015; Slater L et al., 2002; Slater
88 L et al., 2006). However, some scholars found that the correlation between the imaginary part of the
89 complex conductivity and the specific surface area is weak when using British sandstone for experiments,
90 and it is difficult to meet the accuracy requirements for predicting permeability (Binley et al., 2005). The
91 above method essentially uses indirect parameters such as porosity and specific surface area to evaluate
92 permeability, and the estimation results have certain instabilities.

93 Previous studies have shown that permeability calculations can be carried out by using the
94 correlation between relaxation time and pore size, and the calculated results are in good agreement with
95 the directly measured permeability (Niu et al., 2016). However, some scholars have proposed that the
96 permeability calculation method based on the relationship between relaxation time and pore size has
97 certain limitations, and it is difficult to obtain a universally applicable prediction model for mixed particle
98 size samples (Joseph et al., 2016; Kruschwitzv et al., 2010; Titov et al., 2010). The relaxation time
99 spectrum inversion method is to obtain the excitation polarization relaxation time distribution of the
100 sample from the time-domain attenuation curve, and use the excitation polarization relaxation time

101 spectrum to characterize the pore structure distribution and estimate the permeability (Tong et al., 2006a;
102 Tong et al., 2006b; Titov et al., 2010). In addition, some scholars have studied the relationship between
103 the two in terms of models, using the average relaxation time (τ) of the complex resistivity Cole-Cole
104 model to estimate the permeability of sandstone. Revil and Florsch proposed a model showing that
105 permeability is linearly related to relaxation time (Binley et al., 2005; Revil et al., 2010). The above
106 method research shows to varying degrees that there is a correlation between relaxation time and
107 permeability.

108 In this paper, 8 kinds of quartz sands are used to simulate the water-bearing sand layer under
109 different working conditions, and the indoor and outdoor experiments are carried out through the
110 designed multiple experiment devices. The first is the induced polarization experiment, which uses
111 square experiment tanks of different sizes to fill quartz sand samples of different particle sizes to obtain
112 the observation data of the induced polarization experiment, and then obtain the relaxation time. The
113 second is the measurement of permeability. Using the same sample, the Darcy percolation experiment is
114 performed to determine the permeability parameters to obtain the permeability. Considering the particle
115 size of the water-bearing sand sample as the intermediate quantity, the relationship between relaxation
116 time and permeability is established through the compound mapping relationship of relaxation time-sand
117 sample size-permeability. The relationship between the two will be further clarified to verify the
118 effectiveness of the method for estimating the induced polarization permeability.

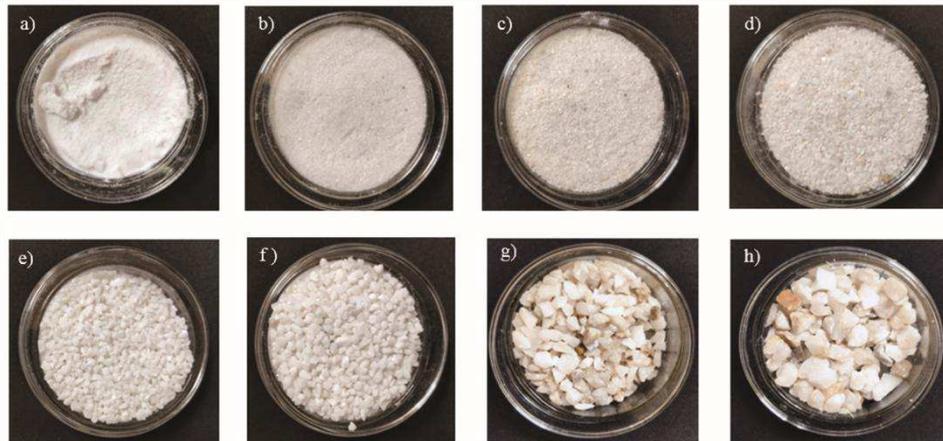
119 **2. Experimental Principles and Method**

120 2.1 Quartz sand sample preparation

121 In order to simulate different pore structures of various formations, at the same time, considering
122 the ease of control and preparation of loose samples, quartz sand is used as the formation simulation
123 material. Fine quartz sand is selected, in order to ensure the differentiation of particle size, this
124 experiment uses custom-made screens with different apertures to screen the quartz sand. The quartz sand
125 is divided into 8 particle sizes: 0.1~0.2mm, 0.2~0.5mm, 0.5~1mm, 1~2mm, 2~3mm, 3~4mm, 4~6mm,
126 6~8mm.

127 The sieved quartz sand still has some impurities. In order to avoid clay or other impurities having a
128 significant impact on the induced polarization effect, the sieved sand sample is cleaned, dried, and then

129 poured with water to saturation for testing. The water sample used was the site water of the tunnel project
130 to simulate the actual engineering conditions. The quartz sand after screening and cleaning is shown in
131 Fig.1. At the same time, in order to reduce the influence of temperature on the induced polarization effect,
132 the measurement needs to be carried out in an environment with small temperature changes.



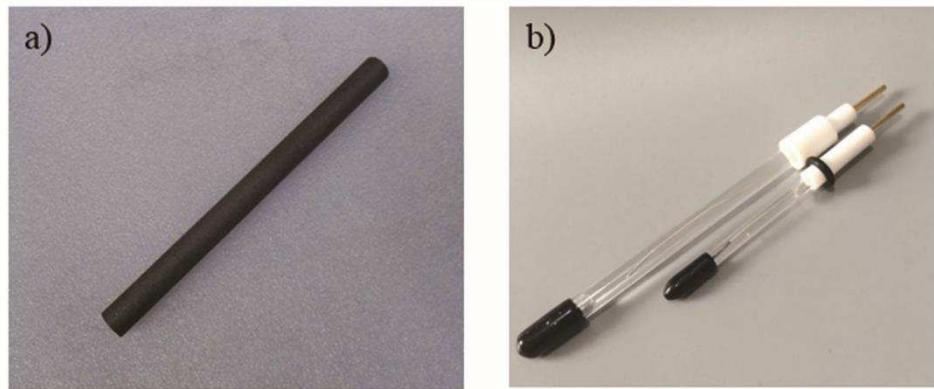
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134 **Fig. 1** sand of samples with different particle size grades. a) 0.1-0.2 mm, b) 0.2-0.5 mm, c) 0.5-1
135 mm, d) 1-2 mm, e) 2-3 mm, f) 3-4 mm, g) 4-6 mm, h) 6-8 mm

136 2.2 Induced polarization experiment

137 The commonly used detection methods of the induced polarization method are time domain and
138 frequency domain. Time domain measurement has the advantages of high efficiency and convenient
139 operation in the field work. Through the measurement of a charge and discharge process, all time domain
140 induced polarization observation data can be obtained, which can greatly save field work time and may
141 reduce work efficiency. Therefore, the time-domain measurement method is used for the experimental
142 study of induced polarization. In the observation device of the experiment, four-electrode arrangement
143 is used for measurement. The outer side is two current electrodes. In order to reduce the influence of the
144 polarization of the current electrodes and the potential electrodes on the measurement results, the current
145 electrode A/B uses graphite electrodes, as shown in Fig. 2(a). There are two potential electrodes on the
146 inner side. In order to avoid the influence of the polarization of the electrodes on the observation data,
147 there are two potential electrodes M/N on the inner side, using Ag/AgCl non-polarized electrodes, as
148 shown in Fig. 2(b). The electrolyte adopts saturated KCL solution and is replaced regularly. The purpose
149 is to prevent the electrolyte from failing due to long-term use, resulting in the generation of over-potential
150 of the electrode itself, and to minimize the influence of the polarization of the electrode on the

151 measurement result.



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Fig. 2 The current electrode and potential electrode. a) Graphite current electrode, b) Ag/AgCl unpolarized potential electrode

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The experiment transmitter uses Pentium WDFZ-10T excited polarization transmitter. The signal transmitter is connected to the current electrode, which can transmit DC square wave signals with different duty ratios and different power supply durations, and can measure the current size of the transmitted signal. The receiver adopts Horn3D full-function IP instrument, its sampling frequency is up to 250Hz, and it can collect full waveform time series data.

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The used transmission period and reception period are both 64s. In order to eliminate the influence of DC offset caused by unidirectional power supply, each cycle uses square wave pulses of opposite polarity to supply power with a duty cycle of 50%, that is, the power supply duration is the same as the power failure duration. At the same time, in order to reduce accidental errors in the data process, multiple cycles of power supply and power failure measurements were performed, so the power supply current signal of the transmitter is shown in Fig. 3(a).

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In terms of data collection, the receiver adopts a late synchronization method for data collection, without manual synchronization of the transmitted signal. The obtained observation signal is shown in Fig. 3(b). It can be seen that the secondary field voltage slowly decays within a certain period of time after the power supply signal is turned off.

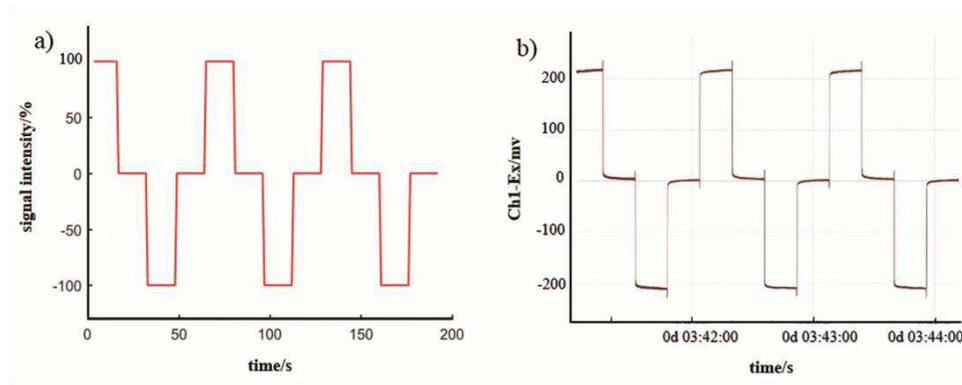


Fig. 3 a) Schematic of current signal, b) Measured full waveform signal

We try a larger outdoor square experiment tank with a size of 100cm×100cm×100cm is used, as shown in Fig. 4. The side walls and the ground are made of cement. Data collection is performed on the process of excited polarization power supply and power-off, and time-domain attenuation information is obtained.

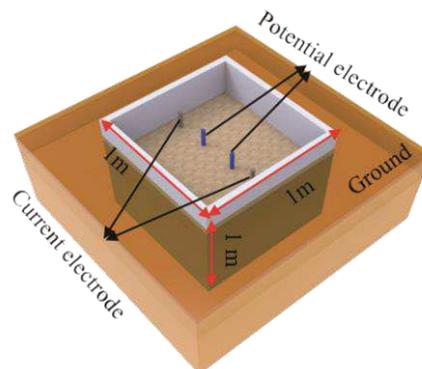


Fig. 4 Outdoor laboratory tank

2.3 Darcy flow experiment

In order to experiment the actual permeability coefficient of the quartz sand samples and calculate the permeability, the Darcy seepage experiment device was used to conduct seepage experiments on quartz sand samples with different particle sizes. Since the sample is highly water-permeable, the measurement is carried out by the constant head method. The experimental instrument is shown in Fig.5;

the current experiment system is mainly comprised of five parts:

- (1) Water supply device: it can realize continuous replenishment of experimental water and keep the water head stable during the experiment;
- (2) Permeation device: an acrylic cylinder is used to place the experiment sample, the upper end is

- 188 equipped with a water inlet, the side is equipped with a pressure measuring hole, the lower end is
189 equipped with a water outlet, and the bottom is equipped with a permeable filter plate;
- 190 (3) Pressure measuring device: connect the pressure measuring tube with the pressure measuring hole
191 to measure the pressure head on different sections;
- 192 (4) Drainage device: set a series of round holes in the piezometric tube to adjust the drainage water level;
- 193 (5) Other equipment: stopwatch, 1000mL measuring cylinder, beaker, funnel, glass rod, thermometer,
194 tube clamp, rubber tube and suction balloon, etc.
- 195



196
197 **Fig. 5** Schematic diagram of Darcy seepage tester

198 Before the formal experiment, first determine the relationship between the permeability of the quartz
199 sand sample and the head loss, and test the experimental instrument. After confirming that the
200 experimental instrument is in good condition, the permeability measurement of 8 kinds of homogeneous
201 quartz sand samples with particle size is carried out. The experimental process is as follows:

- 202 (1) Connect the instrument: check the state of the instrument, such as whether the piezometer tube and
203 the infiltration device are airtight, and record the inner diameter of the infiltration device, the
204 distance between the piezometer tubes and other parameters;
- 205 (2) Filling the sample: first install a permeable filter plate at the bottom of the infiltration device, then
206 the sample is loaded, every time a certain thickness is loaded, a certain degree of vibrating is
207 performed with a glass rod;
- 208 (3) Saturated sample: inject water from top to bottom until water film appears on the surface of the
209 sample;

210 According to Darcy's law, the seepage flow Q of the cross section through the infiltration device is

211 proportional to the cylindrical section A and the hydraulic slope I , and is related to the soil permeability
212 coefficient K . The basic relationship is as follows:

$$213 \quad Q = KAI \quad (1)$$

$$214 \quad I = (H_2 - H_1) / L \quad (2)$$

215 Where I is the hydraulic gradient, H_1 and H_2 are respectively the head of the piezometer, and L is
216 the length of the seepage path, both in m.

217 (4) Experimental measurement: After the water level of the piezometric pipe is stable, record the
218 piezometric water level and start to measure the seepage flow out of the permeation device within a
219 certain period of time. After repeating the measurement, change the hydraulic slope of the device
220 and repeat the above process for measurement. In order to prevent the osmotic pressure in the device
221 from changing too drastically and damaging the original structure of the sample, the hydraulic
222 gradient should be increased or decreased step by step to avoid jumping changes.

223 The permeability coefficient K can be used to evaluate the difficulty of the fluid passing through
224 the pore framework of rock and soil, and its definition is shown in Eq. (3):

$$225 \quad K = k \frac{\rho g}{\mu} \quad (3)$$

226 Where ρ is the fluid density, g is the acceleration due to gravity, and μ is the hydrodynamic viscosity
227 coefficient.

228 The Darcy seepage experiment is used to test samples of different particle sizes. According to the
229 above operation method, the flow rate is changed 2 to 3 times to obtain the flow rate, time, head and
230 other parameters, and the permeability coefficient and permeability are calculated by Eq. (3). Record and
231 analyze the experiment data.

232 **3. Results and discussions**

233 3.1 Relaxation time measurements

234 In order to verify the repeatability of the experimental data and reduce the impact of accidental
235 errors on the measurement results, the same measurement device is used to perform repeated observations
236 on the same sample. In the time domain induced polarization experiment, the attenuation curve of the
237 secondary field at the corresponding particle size is obtained, and the attenuation curve of the polarization

238 rate is obtained by Eq. (4).

239
$$\eta_s = \frac{\Delta U_2(t)}{\Delta U(T)} \quad (4)$$

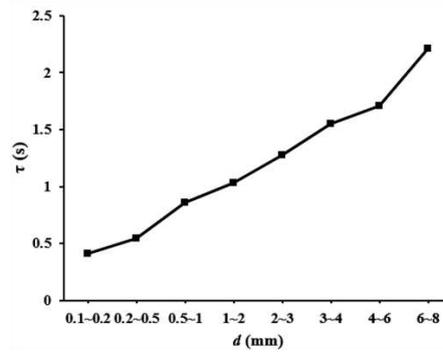
240 Where η_s is polarization rate, $\Delta U(T)$ is the total field potential difference measured before the
241 power is cut off by supplying power to the body polarized medium with a stable current for a period of
242 time T , $\Delta U_2(t)$ is the secondary field potential difference measured at time t after power failure.

243 We obtain the polarizability decay curve of the corresponding particle size under a certain current
244 from the average of multiple sets of data, and fit it with the second nonlinearity of the Cole-Cole model
245 to obtain the relaxation time curve under the corresponding particle size. Table 1 shows the relaxation
246 time distribution of different particle sizes.

247 **Table 1** Relaxation time of sand sample

Particle size (mm)	0.1~0.2	0.2~0.5	0.5~1	1~2	2~3	3~4	4~6	6~8
The relaxation time (s)	0.410	0.541	0.857	1.035	1.273	1.552	1.705	2.215

248 The relaxation time curves of eight different particle sizes of quartz sand are shown in Fig.6.



249

250 **Fig. 6** Relaxation time of sand sample

251 3.2 Permeability measurements

252 Through measurement, the inner diameter of the instrument $D=6.4\text{cm}$, the cross-sectional area of
253 the quartz sand sample $A=0.0033\text{m}^2$, and the pressure measurement interval $L=10\text{cm}$. The outdoor
254 temperature was continuously measured for three days, and the average temperature for three days was
255 15°C , so the hydrodynamic viscosity coefficient was $0.001\text{Pa}\cdot\text{s}$. Use the constant head method to test.
256 After the water flow remains stable, record the experiment data when the water flow out of the drain is

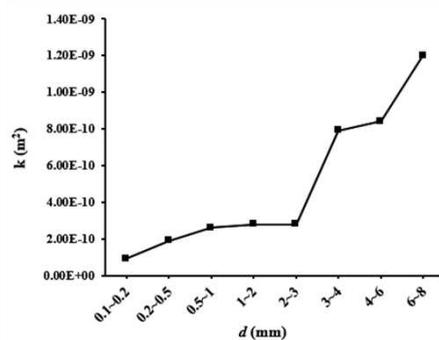
257 1000mL. The permeability coefficient of the sample can be calculated through the measurement results
 258 of the Darcy seepage experiment, and the permeability of the quartz sand sample of this particle size can
 259 be obtained by conversion by Eq. (3). The experimental data is shown in the following table 2:

260

Table 2 Darcy experiment result

Particle size (mm)	Time t(s)	Water volume w(m ³)	Flow Q(m ³ /s)	Piezometer		Water level	Hydraulic gradient I	Permeability coefficient K(m/s)	Permeability k(m ²)
				water level		difference			
				<i>h</i> ₁	<i>h</i> ₂	<i>h</i> ₁ - <i>h</i> ₂			
0.1~0.2	189	0.001	5.3E-06	20	2	18	1.8	9.1E-04	9.1E-11
0.2~0.5	92.6	0.001	1.1E-05	20	2	18	1.8	1.9E-03	1.9E-10
0.5~1	66.3	0.001	1.5E-05	20	2	18	1.8	2.6E-03	2.6E-10
1~2	62	0.001	1.6E-05	20	2	18	1.8	2.8E-03	2.8E-10
2~3	61.5	0.001	1.6E-05	20	2	18	1.8	2.8E-03	2.8E-10
3~4	22	0.001	4.5E-05	20	2	18	1.8	7.9E-03	7.9E-10
4~6	20.6	0.001	4.9E-05	20	2	18	1.8	8.4E-03	8.4E-10
6~8	14.3	0.001	7.0E-05	20	2	18	1.8	1.2E-02	1.2E-09

261 It can be seen that as the particle size of the quartz sand sample increases, the time required to reach
 262 the same seepage flow has a significant difference, and the difference between the minimum and
 263 maximum time is close to 10 times. Through data such as flow, time, pressure head, etc., through data
 264 such as flow, time and pressure head, it is possible to calculate parameters such as seepage flow, seepage
 265 velocity, water level difference and hydraulic slope, so as to obtain permeability coefficient and
 266 permeability from Darcy's law. The experimental results show that the evolution of the permeability is
 267 closely correlated with the sizes of sand sample.



268

269

Fig. 7 Permeability of sand sample

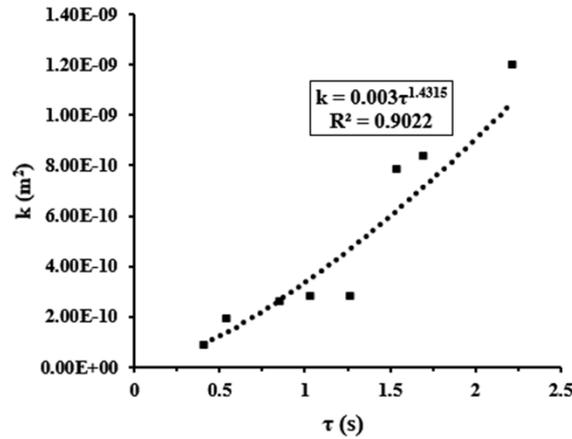
270 3.3 Fitting of curves

271 According to the experimental data, the power function equation is used to fit the relaxation time and
 272 permeability of the same particle size sand sample, and the empirical equations for the relationship

273 between the permeability and the relaxation time is obtained, and the relationship between relaxation
274 time and permeability is established. The fitting equations are given below.

275
$$k = 0.003\tau^{1.4315} \quad (5)$$

276 The relationship between relaxation time and permeability under the same particle size is shown in
277 the fig. 8 below:



278

279 **Fig. 8** Relationship between relaxation time and permeability

280 **4. Conclusion**

281 In this paper, 8 kinds of quartz sands with different particle sizes are used to measure the relaxation
282 time and permeability of sand samples with different particle sizes using the field time domain induced
283 polarization experiment system and the indoor Darcy flow experiment system. According to the
284 experimental data, the following conclusions can be made:

285 (1) The relaxation time and permeability increase with the increase of quartz sand particle size, and the
286 increasing trend gradually increases, which has a significant positive correlation.

287 (2) Using the particle size of the water-containing sand sample as the intermediate quantity, the
288 relationship curve of permeability and relaxation time under the same particle size is formed by
289 mathematical fitting. Finally, the power function equation to describe the correlation between relaxation
290 time and permeability is obtained.

291 In this paper, the obtained relationship is the experiment equation of quartz sands. They are not
292 for the universal equations application, and their coefficients may be different in different experiment
293 conditions and rock samples.

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