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Topographical effects on the timing of growing season in alpine grasslands

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Abstract

Context: It is important to understand the responses of alpine vegetation to recent anthropogenic climate change. The mountainous landscapes with high climatic heterogeneity are good locations to investigate the effects of microclimatic variation on alpine ecosystems.

Objectives: a) To what degree do topographical factors (aspect and elevation) affect the timing of growing season in alpine grasslands? b) Are these topographical effects different on alpine and non-alpine grasslands?

Methods: We extracted five annual growth phenology indices (Start, End, Length, Peak and Peak-NDVI) in alpine and non-alpine grasslands in the Clutha river catchment, New Zealand with a near-daily NDVI (Normalized Difference Vegetation Index) dataset through 16 years (2001-2016). The shifting rates of these phenology indices were quantified with two topographical factors (aspect and elevation).

Results: The start of season was delayed by 7.5, 5.1 and 3.7 days per 100 m higher of elevation in three grassland types (Alpine, Tall Tussock and Low Producing) respectively, and the end of season was advanced by 1.7, 1.3 days and delayed by 0.3 days per 10-degree-south on slopes individually. The longer season length was observed at lower elevation and on north-facing (sunny) slopes. The later season peak occurred at higher elevation and on north-facing slopes. The lower peak NDVI was detected at the higher elevation.

Conclusions: In the studied grasslands, aspect and elevation were correlated to different phenological indices, and they affect phenology independently. The topographical effects are more pronounced in alpine ecosystems at the elevation above 1,300 m than in non-alpine ecosystems at lower elevation.

Key words: Alpine grassland, MODIS, NDVI, Phenology, Topography.

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26 **1 Introduction**

27 Montane landscapes are characterised by their high environmental heterogeneity over a small
28 spatial extent. With this characteristic mountainous landscapes facilitate high levels of
29 biodiversity and a high ecological significance (Testolin et al. 2020; Verrall and Pickering
30 2020). Alpine ecosystems, which often live at mountains, are expected to be highly vulnerable
31 to the impacts of recent anthropogenic climate change (Rogora et al. 2018; Gottfried et al.
32 2012). It is important to understand the functional responses of vegetation to the changing
33 climatic conditions in these topographically complex environments (Opedal et al. 2015; Dufour
34 et al. 2006). The high topographical and climatic heterogeneity in montane areas are good

35 locations to investigate the effects of microclimatic variation on the key life-cycle attributes in
36 alpine vegetation ecosystems.

37 Researches showed that topographical features are substantial gradients which affect the
38 growing season of alpine species. For example, the regional difference in slope and aspect
39 affects the timing and duration of snow cover (Tennant et al. 2017; Redpath et al. 2018), and
40 in turn affects the timing of growth and ecological processes in alpine vegetation. A European
41 conifer in the Swiss Alps exhibited a delayed needle and stem growth with increasing elevation
42 (Moser et al. 2010). A shorter growing season was observed in a montane forest in Colorado,
43 the US as elevation increases or in the north-facing (shaded) sites (Barnard et al. 2017). In the
44 Australian Alps, the start of growth season in alpine species in recent years (2000-2014) was
45 detected to be sensitive to elevation change (Thompson and Paull 2017). It has been widely
46 documented that alpine plants can shift the timing of biological processes to adapt the
47 topographical variances in a montane environment.

48 There are two key effects of climate change on vegetation: range shifting and phenology
49 changing. An upward trend in range shifting of plant species has been observed in many alpine
50 ecosystems around the world, and these shifts were in the same direction that is congruent with
51 the direction expected under a global warming condition (Chen et al. 2011; Telwala et al. 2013).
52 Mountainous species have moved upwards over the last 100 years as a response to follow the
53 shifted analogous climate condition (Lenoir et al. 2008; Scherrer and Korner 2010). The
54 upwards shifting, as a result, may lead alpine species to run out of their habitats when they
55 reach mountain summits. However, other studies argue that the high topographical
56 heterogeneity in montane landscapes can offer diverse climate conditions for alpine species:
57 Instead of moving uphill, shifting sideways to a different aspect of slope, in order to find
58 analogous climate conditions, is also possible (Spasojevic et al. 2013; You et al. 2018). A study
59 in the alpine ecosystems on shaded slopes at high elevation (1,370–1,800m) in New Zealand
60 showed that microclimate factors can shape the distribution of alpine plants which adapted to
61 specific environmental drivers (Bickford et al. 2011). Is moving up mountains the same as
62 moving around mountains for alpine vegetation? To answer this question, we aim to tackle the
63 relationship between topography and growth phenology in alpine ecosystems.

64 In this study, we used the near daily MODIS (Moderate-resolution Imaging Spectroradiometer)
65 NDVI (Normalized Difference Vegetation Index) images through 16 years (2001-2016) to
66 derive growth phenology indices. We analysed individual and interactive effects of two
67 topographical factors (aspect and elevation) on the timing of growing season in three types of
68 grassland in the Clutha river catchment in New Zealand, to address the following questions:

69 1) To what degree do topographical factors (aspect and elevation) affect the timing of growing
70 season in New Zealand's alpine grasslands?

71 2) Are there any differences of the effects of aspect / elevation on the growth phenology
72 between in alpine and in non-alpine grassland ecosystems?

73 **2 Methods**

74 **2.1 Study Area**

75 We chose the Clutha/Mata-Au River catchment (21,400km²) in the South Island of New
76 Zealand as our study area. Clutha catchment is New Zealand's largest river catchment. It
77 harbours a large proportion of alpine ecosystems in this country and conserves valuable
78 indigenous biodiversity. Over half of Clutha catchment's land has been covered by grassland.
79 There are three types of New Zealand's indigenous grassland in the montane areas of this
80 catchment, and they were investigated here (10,562km², 49.4% of Clutha catchment), namely
81 Alpine grassland (278km², 1.3% of Clutha catchment), Tall Tussock grassland (6,406km², 29.9%
82 of Clutha catchment), Low Producing grassland (3,878km², 18.1% of Clutha catchment)
83 (LCDB-v4.1 2015). Alpine grassland situates at the average elevation of 1,574m, and about
84 80% of this type lives in the range of 1,302m-1,857m. Over 80% of Tall Tussock grassland is
85 located at the elevation of 807m - 1,600m with mean at 1,200m, and nearly half of this type
86 lives above the natural treeline (900m), where is defined as the alpine region in New Zealand
87 (Wardle 2008). Low Producing grassland can be found at mean elevation of 640m, and more
88 than 80% of this type lives between 338m and 929m elevation. This type is recognised as the
89 non-alpine region in this study (Fig.1a).

90 **2.2 Growth phenology extraction**

91 Five annual growing phenological indices (Table 1) were calculated with the Moderate-
92 resolution Imaging Spectroradiometer (MODIS) imagery. We produced a time series of
93 Normalized Difference Vegetation Index (NDVI) from 2001 to 2016 by the near-daily MODIS
94 images at 250m resolution. The 5-day maximum composite method (Fontana et al. 2008) was
95 used to eliminate the signal noises from the raw data due to high cloud/snow coverages. We
96 used the tool of TIMESAT 3.3 on the Matlab R2013b platform (Jonsson and Eklundh 2004)
97 (<http://web.nateko.lu.se/timesat/timesat.asp?cat=0>, visited on 4th July, 2020) to extract the five
98 phenological indices. The Asymmetric Gaussian function in TIMESAT was selected for
99 calculation. To better describe the growing season in the Southern Hemisphere, we created a
100 modified day of year (mDOY, 1st July = Day 1 of mDOY; 30th June in next year = Day 365 of
101 mDOY) to number the days of a year (Fig.2). The peak of season (POS) is defined as the day
102 of highest NDVI in each modified year. The difference of NDVI between the peak NDVI of a
103 year (P-NDVI) and the base level (the mean of the two lowest NDVIs before and after the POS
104 of a year) is defined as the "seasonal amplitude". We used 50% of the seasonal amplitude as
105 the threshold to decide when a growing season starts or ends: Specifically, the start of season
106 (SOS) is the day before the POS of a year and when NDVI first reaches 50% of the seasonal

107 amplitude in that year. Similarly, the end of season (EOS) is the day after POS of a year and
 108 when NDVI first drops to lower than 50% of the seasonal amplitude. The number of days
 109 between SOS and EOS is the length of season (LOS). Other parameters in TIMESAT were
 110 configured with sensitivity analyses: The cut-off point, which is used to remove spike and
 111 outliers with a median-filter method, was set to two times of standard deviation of the whole
 112 dataset (spike = 2.0 in TIMESAT); We used twice “upper envelope” weight assignments
 113 (number of envelopes = 3 in TIMESAT) and an adaptation (degree = 3.0 in TIMESAT) to
 114 reduce negative bias of NDVI. With above settings, we still have a majority of pixels (>90%)
 115 in our data left for seasonal signals extraction (Eklundha and Jönsson 2017).

116 **2.3 Topographical effect analysis**

117 The NZSoSDEM data (Columbus et al. 2011) (Fig.1b) were used to calculate aspect and
 118 elevation. This original 15m resolution DEM (digital elevation model) was resampled to 250m
 119 resolution in ArcGIS 10.1. Particularly, we converted aspect to be a new variable named
 120 “Southness” to quantify the degree of shade on slopes. We transformed the aspect values of
 121 181-360 to 179- 0, but kept the aspect values of 0-180. Therefore, the values of Southness range
 122 from north-facing (0 (N), sunny slopes) to south-facing (180 (S), shaded slopes).

123 The relationships between the five phenological indices (SOS, EOS, LOS, POS, P-NDVI) and
 124 the two topographic factors (Southness and Elevation) were analysed by linear regression
 125 models. In order to eliminate the effects of spatial autocorrelation which make the generalised
 126 least squares (OLS) models invalid, we chose the Simultaneous Auto-Regressive (SAR) model
 127 (Dormann et al. 2007) to calculate the relationships of the variables extracted from our spatial
 128 datasets. Three SAR model specifications (Formula 1-3) were used in this study, and which
 129 specification was suitable is determined by the Lagrange Multiplier (LM) test criteria
 130 (Elhorst 2014).

131 The SAR-lag model:

$$132 \quad Y = \rho WY + X\beta + u \quad (1)$$

133 The SAR-error model:

$$134 \quad Y = X\beta + \lambda W\varepsilon + u \quad (2)$$

135 The SAR-sac/sarar model:

$$136 \quad Y = \rho WY + X\beta + \lambda W\varepsilon + u \quad (3)$$

137 In above formulas, Y represents the vector of response variable (one of the five phenological
 138 indices); X is the matrix of predictor variables (Southness, Elevation, and both); β is the
 139 coefficient of a linear regression model; W is the spatial weights matrix; ρ is the autoregressive
 140 coefficient for the spatial lagged dependent variable Y; λ is the autoregressive coefficient for
 141 the spatial weighted error term; ε is a spatially dependent error term; u is a non-spatial error

142 term. In SAR-lag model, λ is assumed to be 0, while in SAR-error model ρ is assumed to be 0.
143 In SAR-sac/sarar model, ρ and λ are both non-zeros. In formula (3) the W s for spatial lagged
144 item (ρWY) and spatial error item ($\lambda W\varepsilon$) could be different, however, we used the same W for
145 both items in this study.

146 We used an inverse-distance-weight algorithm (“dnearneigh” function in the package “spdep”
147 on R platform (Bivand and Wong 2018)) to generate a specific spatial weights matrix (W) for
148 each SAR model. The main idea is that: In an image, one pixel’s spatial proximity to its
149 surrounding pixels can be described by certain weights which are negatively associated with
150 the distances of pixels (Fortin 2005). The hypothesis is that spatial autocorrelation affects
151 within a limited distance in a spatial dataset. Instead of calculating a W with all pixels, we used
152 a distance threshold to produce a specific W in each SAR model calculation. This threshold
153 was determined by a semi-variogram test (Hession and Moore 2011).

154 In order to evaluate the capacities of topographical factors in explaining the changes of
155 phenological indices, we used Akaike information criterion (AIC) to quantify the fitnesses of
156 different SAR models (Kissling and Carl 2008; Maggini et al. 2006). An intercept function (~ 1)
157 was used as null hypothesis. Four functions of southness and elevation against each
158 phenological index and for each year were assessed. The delta AIC averages were summarized
159 in Table 4.

160 All the calculations including SAR model regressions, Lagrange Multiplier tests, semi-
161 variogram tests, AIC assessments were accomplished with the package “spdep” on R platform
162 v3.5.3 (R Core Team 2020).

163 **3 Results**

164 **3.1 The correlation between aspect and growth phenology**

165 During our study period of time (2001-2016), the start of growing season (SOS) happened
166 earlier on southwest-facing slopes, but later on northeast- and east-facing slopes in all three
167 grassland types (Fig.3). The average dates of SOS were 149.9, 119.4 and 93.4 mDOY in Alpine
168 grassland, Tall Tussock grassland and Low Producing grassland, individually (Table 2). The
169 northeast-facing (sunny) habitats showed 3.3-7.1 days later of SOS than average in these
170 grassland types.

171 The end of growing season (EOS) was detected to be later on the north- and northeast-facing
172 slopes in Alpine grassland and Tall Tussock grassland, while EOS occurred later on the
173 southeast-facing areas in Low Producing grassland (Fig.3). The Alpine grassland on northeast
174 slopes (NE) exhibited approximately 29.9 days later of EOS than on southwest slopes (SW).
175 The trend of later EOS on more northeast-facing slopes was also observed in Tall Tussock
176 grassland, that the northeast-facing habitats showed roughly 22.0 days later of EOS than the
177 southwest-facing ones. In contrast, Low Producing grassland presented a different spatial trend

178 of EOS to the other two grassland types, that the southeast-facing Low Producing grassland
179 showed 16.4 days later of EOS than on the northwest-facing slopes (Table 2).

180 Alpine grassland and Tall Tussock grassland were detected to have a longer growing season
181 (LOS) on north-facing slopes, however shorter LOSs were found on the north-facing slopes in
182 Low Producing grassland. Specifically, there were about 20.2 days longer of LOS on north-
183 facing (sunny) Alpine grassland than on south-facing (shaded) side. For Tall Tussock grassland,
184 the difference was nearly 12.3 days longer of LOS on north-facing sites. Low Producing
185 grassland showed a reverse trend that the north-facing habitats had 11.6 days shorter of LOS.

186 The peak of growing season (POS) exhibited an identical pattern as EOS in the three grassland
187 types: the POS had been reached later on northeast-facing slopes in Alpine grassland and Tall
188 Tussock grassland, while on east-facing slopes in Low Producing grassland. The differences
189 of POS between the latest and the earliest are 22.3, 19.5 and 14.0 days in Alpine, Tall Tussock
190 and Low Producing grassland, individually.

191 Higher peak day NDVI (P-NDVI) was found on more east-facing slopes in all three grassland
192 types. Alpine grassland showed the highest mean P-NDVI of 0.357, with 0.052 higher than the
193 lowest mean value on west-facing slopes. Tall Tussock grassland's P-NDVI showed a
194 difference of 0.024 between the maximum (0.487 on east-facing slopes) and minimum (0.463
195 on west-facing slopes). The highest P-NDVI of Low Producing grassland appeared on
196 southeast-facing slopes with 0.029 higher than on the northwest-facing slopes.

197 **3.2 The correlation between elevation and growth phenology**

198 In all three grassland types, the start of growing season was observed to be later as elevation
199 rises. The SOS in Alpine grassland exhibited about 8.5 days later per 100m higher of elevation
200 (Fig.4). The delay of SOS in Tall Tussock grassland were 5.2 days per 100m higher on average,
201 and the degree of delaying became higher when elevation exceeds 1,200m. Similarly, the
202 delaying speed of SOS in Low Producing grassland was lower when under 1,200m of elevation
203 (3.7 days later per 100m upwards) but higher when above this elevation (8.2 days later per
204 100m upwards).

205 The season end was less sensitive to elevation changes, especially in Alpine grassland (Fig.4).
206 There was a weak trend of later EOS as elevation climbs in the Alpine grassland under 1,300
207 m. However, when above the elevation of 1,300m, EOS in Alpine grassland stayed the same
208 at about 312-315 mDOY (Table 3). For Tall Tussock grassland, the EOS happened 4.9 days
209 later per 100m upwards when below 1,300m elevation, and EOS also kept at 316-320 mDOY
210 when above 1,300m. In Low Producing grassland the EOS was delayed by 5.8 days per 100m
211 higher of elevation when below 1,300 m, but it fixed at 320 mDOY when above this elevation.

212 The length of growing season showed a strong negative correlation with elevation in the
213 grassland above 1,300 m (Fig.4). The Alpine grassland under 1,300m showed 190.4 days of

214 LOS on average, and it was not sensitive to elevation changes. Nevertheless, LOS in the Alpine
215 grassland located at 1,300-1,900m showed a negative relation with elevation, that LOS
216 decreased by 10.0 days per 100m upwards. It was also observed in Tall Tussock grassland that
217 LOS was shortened by 9.4 days per 100m uphill when above 1,300m. At the elevation of 900-
218 1,300m, Tall Tussock grassland exhibited a constant LOS of 200.0 days. The Low Producing
219 grassland above 1,400m showed a shorter LOS of 11.8 days per 100m upwards.

220 The peak of season was positively correlated with elevation (Fig.4). In the Alpine grassland
221 above 1,200m, POS was reached by 4.4 days later per 100m higher of elevation. The POS in
222 Tall Tussock grassland was delayed dramatically by elevation (13.3 days later per 100m
223 upwards) in low land (400 -700m) rather than in high land above 700m (3.5 days later per
224 100m upwards). The changing rate of POS in Low producing grassland was 5.3 days later per
225 100m higher on average.

226 P-NDVI was negatively related to elevation (Fig.4). The changing speed of P-NDVI were -
227 0.033, -0.023 and -0.012 per 100m upwards in Alpine grassland, Tall Tussock grassland and
228 Low Producing grassland. Oscillations in the changing rates of P-NDVI were observed in the
229 low land of Tall Tussock grassland and Low Producing grassland.

230 **3.3 Inter-annual relationships between topography and phenology**

231 Very weak relationships (less than ± 1.0 days per 10-degree-south) between SOS and southness
232 (aspect) were observed in all three grasslands through 16 years (Fig.5 (I) a). Instead, EOS was
233 highly correlated with southness in Alpine grassland and Tall Tussock grassland. The EOS in
234 Alpine grassland showed 1.0-2.0 days earlier per 10-degree-south in most of the 16 years,
235 except in year 2011 with 2.9 days earlier per 10-degree-south. In Tall Tussock grassland, the
236 sensitivities of EOS to aspect fluctuated among 0.5-1.5 days earlier per 10-degree-south
237 through the years. Weak positive correlations between EOS and southness had been observed
238 in Low Producing grassland during 2001-2016. LOS constantly showed negative relations in
239 Alpine grassland and Tall Tussock grassland in all the study period, with 1.4 and 0.8 days
240 shorter per 10-degree-south on average; In contrast, LOS in Low Producing grassland exhibited
241 positive correlations with southness with a mean of 0.7 days longer per 10-degree-south. The
242 Alpine grassland and Tall Tussock grassland showed the same trend of earlier POS on south-
243 facing slopes. On average, POS was reached in these two grassland types by 1.1 days earlier
244 per 10-degree-south in the study period. However, POS in Low Producing grassland showed
245 no response to aspect in all the years (Fig.5 (I) b). No more than ± 0.001 P-NDVI per 10-degree-
246 south of changes were detected during the 16 years (Fig.5 (I) c). The average changing rates
247 were +0.0004, -0.0005 and 0.0000 of P-NDVI per 10-degree-south in Alpine grassland, Tall
248 Tussock grassland and Low Producing grassland

249 All the five annual phenological indices were more sensitive to elevation changes during 2002-
250 2016 (Fig.5 (II)). In Alpine grassland, SOS exhibited strong correlations with elevation by 5.5-

251 10.3 days later per 100m upwards in these years, while the delaying trends of SOS were weaker
252 in Tall Tussock grassland (3.8-6.3 days later per 100m upwards) and Low Producing grassland
253 (1.3-4.7 days later per 100m upwards) (Fig.5 (II) a). On the other hand, the EOS in Alpine
254 grassland showed no relationship with elevation during the years, but weak later trends of EOS
255 along elevation had been detected in Tall Tussock grassland and Low Producing grassland. As
256 a result, the LOS in Alpine grassland was negatively related to elevation through years (7.4
257 days shorter per 100m upwards on average). Weaker negative correlations between LOS and
258 elevation were found in Tall Tussock grassland (3.0 days shorter per 100m upwards on average)
259 and in Low Producing grassland (0.9 days shorter per 100m upwards on average). POS in all
260 three grassland types were always delayed by elevation rise during the years (Fig.5 (II) b), with
261 average delays of 3.4 days in Alpine grassland, 3.6 days in Tall Tussock grassland and 5.1 days
262 in Low Producing grassland per 100m higher. P-NDVI constantly showed a negative relation
263 with elevation in the 16 years (Fig.5 (II) c), with means of -0.033, -0.023, -0.012 per 100m
264 upwards in three grassland types respectively.

265 **3.4 Interactive topographical effects on timing of growing seasons**

266 We displayed the distributions of five annual phenological on an aspect-elevation coordinate
267 to illustrate whether interactive topographical effects existed in our study (Fig.6). It showed
268 that SOS was mainly correlated to elevation only in grasslands (Fig.6a), except the SOS in the
269 grassland under 1,400m showing 1.0-5.0 days later on east-facing slopes than on west-facing
270 slopes at the same elevation.

271 Though elevation effect was still dominant, EOS exhibited a stronger correlation with aspect
272 than did SOS. For the alpine species of this study (including all of Alpine grassland and the
273 Tall Tussock grassland above the 900m treeline), there was a circle of EOS distribution in north
274 direction (Aspect 0 (N), Fig.6b) at 1,400-1,500m elevation, where EOS occurred later on north-
275 facing slopes than on other directions. Another circle can be seen on northwest slopes under
276 500m elevation in Low Producing grassland, indicating EOS on northwest slopes appeared
277 earlier than other regions.

278 The POS in Alpine grassland was correlated to both aspect and elevation (Fig.6c). POS
279 generally exhibited a delaying trend as elevation increases, meanwhile the grassland on east-
280 facing slopes reached POS later than on west-facing slopes at the same elevation. Differently,
281 POS in Tall Tussock grassland and Low Producing grassland was delayed by elevation in all
282 aspects.

283 All grassland types had shorter LOS as elevation rises, while aspect modified LOS's changing
284 rates (Fig.6d). Above 1,200m, the Alpine grassland on north-facing slopes had the longest LOS.
285 The Tall Tussock grassland above 1,200m showed decreasing LOS along elevation in all
286 directions, while under 1,200m LOS was shorter on north-facing habitats. There was a concave
287 of LOS in the contour of Low Producing grassland at 400m elevation, indicating that when

288 above 400m, longer LOS was detected on north-facing slopes and at higher elevations, whereas
289 when below 400m, the situation reversed.

290 Generally, P-NDVI in three grassland types decreased as elevation rises (Fig.6e). P-NDVI was
291 strongly correlated with elevation in both alpine and non-alpine ecosystems. Above 1,000m
292 elevation, P-NDVI declined at the same rate along elevation in all aspects. While in the Tall
293 Tussock grassland under 1,000m and in the Low Producing grassland under 800m, P-NDVI
294 was higher on south-facing slopes.

295 The delta AICs of SAR models showed that the additive and interactive formulas are more
296 effective than any single variable formulas when explaining the changes of the five annual
297 phenological indices (Table 4). The single variable function of elevation (“elev”) against SOS
298 contributed the most in decreases of AIC, showing that the changes of SOS can be well
299 estimated only by elevation. Aspect (“south”) was a more descriptive variable for EOS changes
300 in Alpine grassland, while in Low Producing grassland, elevation was a more effective factor
301 in explaining EOS’s variation. In Tall Tussock grassland, the two-variable formulas were more
302 suitable to describe EOS changes than single variable formulas. Aspect and elevation were
303 both substantial in explaining LOS changes in three grassland types. The two-variable formulas
304 can also effectively estimate POS changes in Alpine grassland and Tall Tussock grassland, but
305 elevation was the only descriptive variable for POS in Low Producing grassland. Elevation can
306 explain all the variations of P-NDVI in three grassland types.

307 Overall, the delta AICs of models (Table 4) demonstrates the fact that SOS was mainly
308 correlated to elevation, while EOS was likely controlled by aspect in alpine habitats but by
309 elevation in non-alpine habitats. Elevation showed a better capacity to describe LOS in alpine
310 habitats than in non-alpine ecosystems. The additive functions decreased the similar AIC as
311 the interactive functions did, which indicated that elevation and aspect had independent
312 correlations with the five phenological indices.

313 **4 Discussion**

314 **4.1 Alpine ecosystems’ phenological responses to topography**

315 This study showed that the timing of growing season in New Zealand’s alpine ecosystems was
316 sensitive to topographical factors. Aspect and elevation exhibited considerable effects on
317 phenology in different ways. Elevation mainly affects the start of season, while aspect mostly
318 modifies the end of season. Lower air temperature at high elevation regions might be the cause
319 for the delay of growing season start. However, the mechanism that how elevation “controls”
320 the growth season start was not deciphered in this research. For aspect’s modification on
321 growth phenology, that longer growing season was observed on north-facing slopes, a rational
322 explanation could be that the sunny habitats have more exposure to sunshine therefore have
323 more temperature accumulation (Zhang et al. 2004; Fontana et al. 2008). Researches also

324 showed that air temperature does not directly affect plants' phenology, and complex
325 interactions exist between environmental factors which limit the resources of plant growth
326 (Jolly et al. 2005). Temperature and radiation have been proven to be substantial controllers of
327 phenology, but at high elevation regions there are more uncertainties in their effects on
328 phenology due to other associated factors such as water stress (Hwang et al. 2011). Our result,
329 that aspect and elevation affected different phenological indices, indicates that alpine
330 ecosystems are subject to different environmental restrictions at different growing stages.

331 Because mountainous areas offer diverse environmental resources and montane plants well
332 adapt to the harsh but variant environmental conditions, alpine ecosystems may not have to
333 migrate upwards mountains to find suitable climates. Therefore, contemporary climate change
334 is not expected to exhaust alpine plants' ecological niches. The topographical heterogeneity in
335 mountains can provide a wide range of microclimatic conditions on the horizontal direction
336 which produces unique ecological niches (Bickford et al. 2011). Microclimate at a small scale
337 can provide "refugia" for montane plants to escape from the large-scale climate change by a
338 short migration (Opedal et al. 2015). The alpine grassland's plasticity could also lessen the
339 impacts of climate change by adjusting their phenology activities (Frei et al. 2014). Thus, alpine
340 plants have chance to survive under current global climate change due to the high heterogeneity
341 in their montane habitats.

342 **4.2 The difference of topographical effects in alpine and in non-alpine ecosystems**

343 The topographical factors exhibited different effects on the timing of growing season in alpine
344 and non-alpine ecosystems. The growing phenology of alpine plants showed a closer
345 association with topographical factors, especially when the habitats situated above 1,200m
346 elevation. In this study, the Alpine grassland type represented alpine ecosystems and the Low
347 producing grassland type represents non-alpine ecosystems. Tall Tussock grassland is a
348 transition type which includes both alpine and non-alpine components. The distribution of
349 phenological indices in the aspect & elevation coordinate (Fig.6) showed distinct patterns in
350 Alpine and Low producing grassland, while showed a combined pattern in Tall Tussock
351 grassland.

352 The growing phenology in alpine ecosystems was more sensitive to topographical effects than
353 in non-alpine ecosystems. The start of season shifted more days in alpine plants than in non-
354 alpine plants when elevation rises. The end of season in alpine ecosystems performed more
355 variations along aspect than did in non-alpine ecosystems. The similar situation was also
356 detected in the length of season and P-NDVI that they were more responsive to topographical
357 changes in alpine habitats. It would be due to the high diversity of topography in mountainous
358 land than in low land.

359 Furthermore, some of the phenological indices in non-alpine grassland showed opposite
360 responses to aspect changes as in alpine grassland. For example, later end of season and later

361 season peak in non-alpine grassland were observed more on shaded slopes, and the length of
362 season tended to be shorter on sunny slopes. Water deficiency on the sunny (north-facing)
363 slopes could be the reason for the shorter growing periods in non-alpine ecosystems at low
364 elevations.

365 **5 Conclusion**

366 The timing of growing season in New Zealand's natural grassland showed variant correlations
367 with topographical factors (aspect and elevation). The start of growing season was strongly
368 correlated with elevation, showing 7.5, 5.1 and 3.7 days delay per 100m higher the elevation
369 in Alpine grassland, Tall Tussock grassland and Low Producing grassland. The end of season
370 advanced by 1.7 and 1.3 days per 10-degree-south in Alpine grassland and Tall Tussock
371 grassland, while it was delayed by 0.3 days per 10-degree-south in Low Producing grassland.
372 In the alpine ecosystems (all Alpine grassland and the Tall Tussock grassland above treeline),
373 elevation strongly correlated to the start of season, while aspect highly affected the end of
374 season. So, the length of growing season had a strong negative relation with elevation and it
375 tended to be longer on north-facing (sunny) slopes in alpine ecosystems. The season peak
376 delayed as elevation rises and it appeared earlier on shaded slopes. The higher elevation showed
377 the lower peak NDVI in all grassland types. These topographical effects are more pronounced
378 at high elevations above 1,300m.

379 In contrast, the non-alpine ecosystems' (all Low Producing grassland and the Tall Tussock
380 grassland below treeline) end of season was less responsive to aspect changes, but it was
381 delayed as elevation climbs. The start of season in non-alpine ecosystems showed a weak
382 relationship with elevation. Topography had weak effects on the length of season in non-alpine
383 habitats. Weak responses to topographical changes were observed in the peak of season and
384 the peak day NDVI in all non-alpine grassland.

385 Overall, aspect and elevation are important factors in explaining the changes of growth
386 phenology in alpine ecosystems, while elevation can explain most of the changes in the timing
387 of growing season in non-alpine ecosystems. Aspect and elevation independently affect the
388 timing of growing season in New Zealand's natural grassland.

389

390 **Declarations**

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393 **Conflicts of interest/Competing interests:** None.

394 **Ethics approval:** Not applicable.

395 **Consent to participate:** Not applicable.

396 **Consent for publication:** Not applicable.

397 **Availability of data and material:** The MODIS NDVI data that support the findings of this
398 study are available from the co-author Pascal Sirguy (pascal.sirguy@otago.ac.nz) upon
399 reasonable request. The climate data (VCSN) that support the findings of this study are openly
400 available in NIWA at <https://data.niwa.co.nz>.

401 **Code availability:** The code that support the findings of this study is available from the
402 corresponding author Xiaobin Hua (Xiaobin.hua@lincoln.ac.nz) upon reasonable request.

403 **Authors' contributions:** RO and XH conceived the study. XH developed and conducted all
404 analyses and wrote a first draft of the manuscript. PS collated and prepared all MODIS data
405 and provided guidance on the time series analyses. All authors contributed to writing the final
406 manuscript.

407

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517

518 **Tables**

519 **Table 1** The five annual growth phenology indices investigated in this study (see Methods
 520 and Fig.2 for details of how these indices are calculated).

No	Index	Abbreviation	Definition
1	Start of season	SOS	The day of the season on which NDVI exceeds 50% of the season peak NDVI for the first time.
2	End of season	EOS	The day of the season on which NDVI drops to 50% of the season peak NDVI after the peak day of season.
3	Length of season	LOS	The number of days between start and end of the season.
4	Peak Day of season	POS	The day of the season on which NDVI reaches its peak of this growing season.
5	Peak NDVI	P-NDVI	The fitted NDVI maximum of a growing season

521

522 **Table 2** The 16-year averages of five phenological indices on aspects. Aspect values (0-360)
 523 are divided into 8 bands to summarize phenological indices' averages. Each aspect band has a
 524 45-degree range with the central aspect value as the band's name: For instance, North (N, 0
 525 (360)) covers the range of 337.5-360 degree plus 0-22.5 degree.

Index	Grassland type	Aspect (Deg.) and Direction								
		225 SW	270 W	315 NW	0(360) N	45 NE	90 E	135 SE	180 S	Mean
Start of season	Alpine	144.3	144.3	148.3	153.7	153.2	153.2	153.3	148.9	149.9
	Tall Tussock	113.6	113.8	119.3	124.4	126.5	124.0	118.8	116.4	119.4
	Low Producing	87.9	88.1	89.9	93.9	99.2	100.1	95.6	90.6	93.4
End of season	Alpine	297.5	306.1	319.6	324.6	327.4	323.7	309.5	299.6	313.6
	Tall Tussock	301.6	310.4	318.4	323.3	324.0	320.5	312.5	303.0	314.1
	Low Producing	274.5	270.9	265.4	269.3	279.9	282.8	281.8	277.7	275.3
Length of season	Alpine	153.2	161.8	171.3	170.9	174.2	170.5	156.1	150.7	163.7
	Tall Tussock	188.0	196.6	199.1	198.9	197.5	196.5	193.7	186.6	194.7
	Low Producing	186.5	182.8	175.4	175.4	180.8	182.7	186.2	187.0	181.9
Peak of season	Alpine	221.4	226.5	236.6	241.8	243.7	241.5	232.9	224.9	233.8
	Tall Tussock	209.7	215.0	222.1	227.6	229.2	225.6	218.4	211.9	219.7
	Low Producing	179.7	177.8	175.7	179.8	188.0	189.7	187.2	182.9	182.7
Peak day NDVI	Alpine	0.305	0.305	0.329	0.326	0.346	0.357	0.327	0.333	0.331
	Tall Tussock	0.464	0.463	0.463	0.472	0.483	0.487	0.485	0.473	0.474
	Low Producing	0.581	0.578	0.575	0.583	0.596	0.600	0.604	0.599	0.589

526

Topographical effects on phenology

527 **Table 3** The 16-year averages of five phenological indices on elevation bands. Units: Start of
 528 season (mDOY), End of season (mDOY), Length of season (days), Peak of season (mDOY),
 529 Peak day NDVI (unitless). The ‘n/a’ means there is no data (distribution) at this elevation range.

Index	Grassland type	Elevation (m)								Mean
		0-300	301-600	601-900	901-1200	1201-1500	1501-1800	1801-2100	2101-2400	
Start of season	Alpine	n/a	n/a	n/a	118.6	135.4	156.7	175.7	182.8	149.9
	Tussock	82.3	89.5	105.5	108.6	121.9	146.6	169.2	191.5	119.4
	Low Producing	74.3	84.3	98.3	106.9	113.4	138.2	n/a	n/a	93.4
End of season	Alpine	n/a	n/a	n/a	308.9	315.3	313.3	311.8	319.8	313.6
	Tussock	265.6	273.8	303.7	310.6	321.2	320.6	316.2	316.4	314.1
	Low Producing	252.5	260.4	282.4	296.2	321.2	317.7	n/a	n/a	275.3
Length of season	Alpine	n/a	n/a	n/a	190.4	179.9	156.6	136.1	137.0	163.7
	Tussock	183.3	184.3	198.2	202.1	199.3	174.0	147.0	124.8	194.7
	Low Producing	178.1	176.2	184.1	189.2	207.7	179.4	n/a	n/a	181.9
Peak of season	Alpine	n/a	n/a	n/a	218.1	228.3	236.6	243.8	251.1	233.8
	Tussock	170.0	179.8	207.5	212.9	224.9	236.3	244.9	254.5	219.7
	Low Producing	159.8	169.8	189.0	201.8	221.8	230.3	n/a	n/a	182.7
Peak day NDVI	Alpine	n/a	n/a	n/a	0.446	0.384	0.309	0.219	0.173	0.331
	Tussock	0.474	0.514	0.533	0.497	0.458	0.412	0.333	0.215	0.474
	Low Producing	0.590	0.619	0.586	0.530	0.444	0.349	n/a	n/a	0.589

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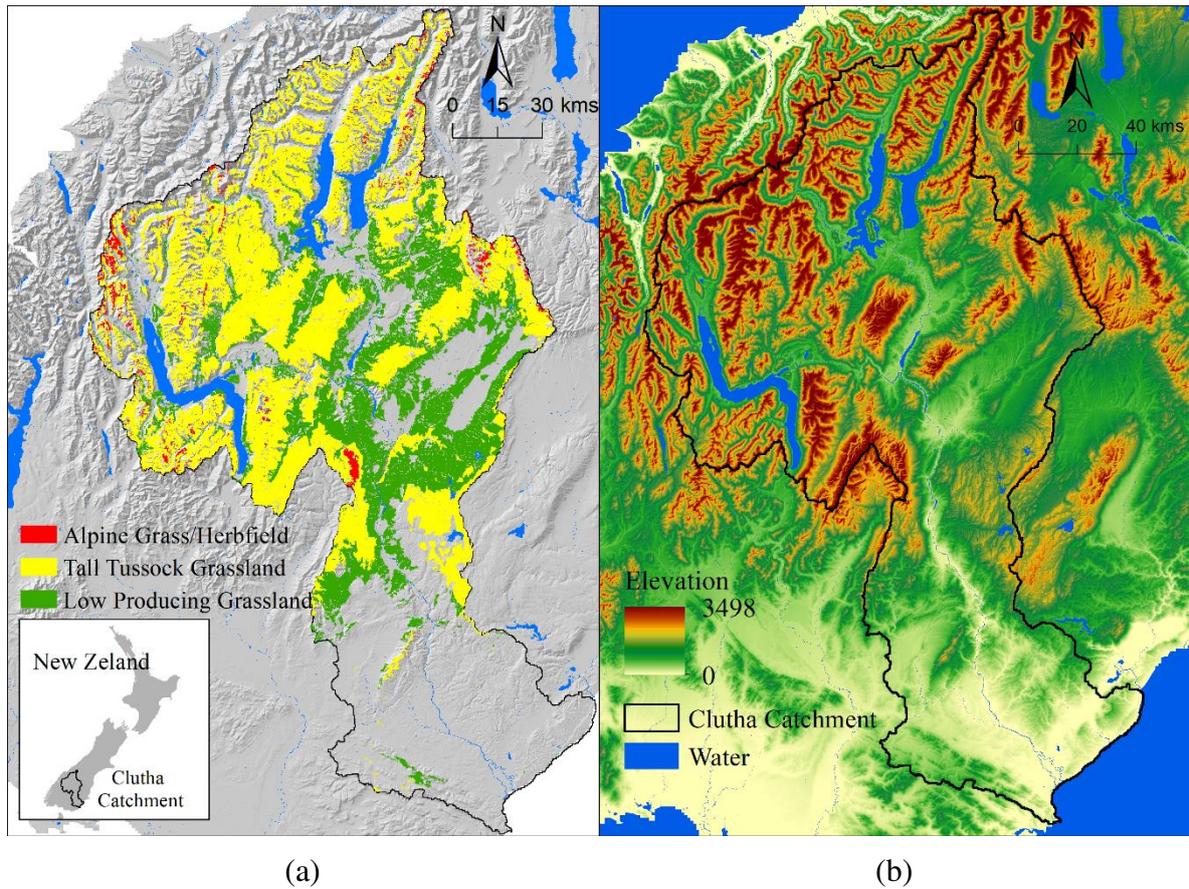
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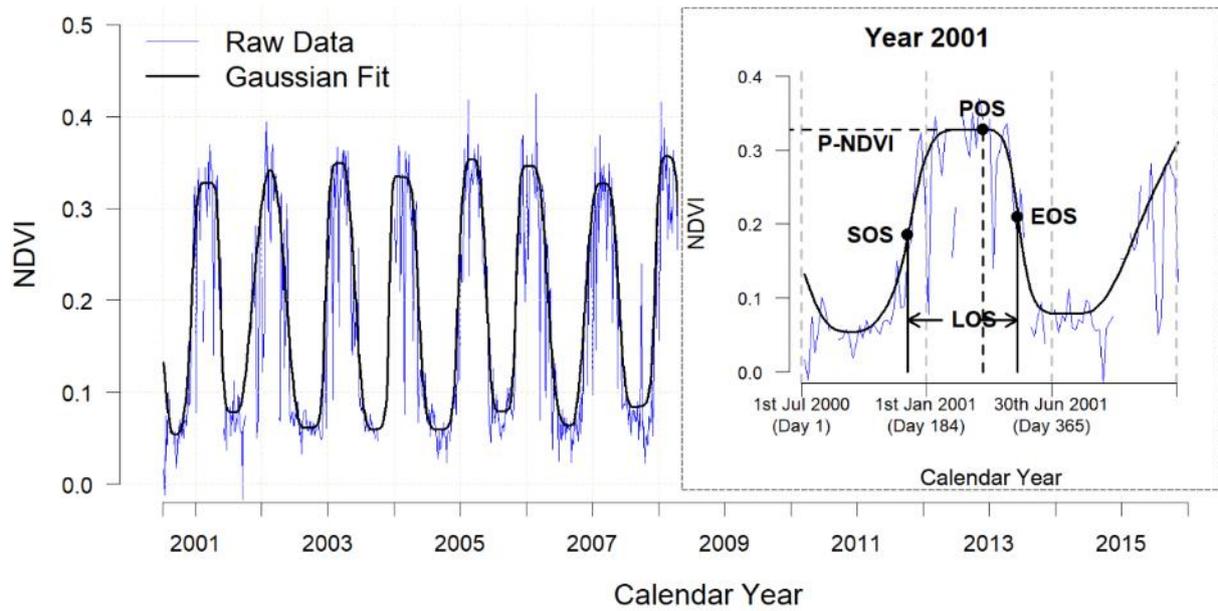
532 **Table 4** Delta AICs (Δ AIC) of the spatial autoregressive (SAR) models simulating the five
 533 phenological indices against aspect and elevation. Four functions including single variable,
 534 additive variables and multiplied variables are tested under null hypothesis (see Methods).
 535 Lower Δ AIC indicates better capability of a function to explain the changes of a phenological
 536 index. Abbreviations of variable names: SOS—Start of season, EOS—End of season, LOS—
 537 Length of season, POS—Peak of season, P-NDVI—Peak day NDVI, south—southness, elev—
 538 elevation.

Models	SOS		EOS		LOS		POS		P-NDVI	
	Δ AIC	(%)	Δ AIC	(%)						
Alpine										
~1	0	0.00	0	0.00	0	0.00	0	0.00	0	0.00
~south	9	0.03	697	2.01	247	0.65	435	1.35	2	0.02
~elev	1984	5.63	91	0.26	863	2.26	608	1.89	2284	27.65
~south+elev	2016	5.72	742	2.14	1325	3.47	1034	3.22	2379	28.81
~south*elev	2066	5.86	773	2.23	1338	3.50	1109	3.45	2393	28.97
Tall Tussock										
~1	0	0.00	0	0.00	0	0.00	0	0.00	0	0.00
~south	78	0.18	508	1.26	169	0.38	426	1.08	7	0.07
~elev	1775	4.18	655	1.62	410	0.92	2080	5.29	2643	24.25
~south+elev	1849	4.36	1254	3.10	605	1.36	2625	6.68	2676	24.55
~south*elev	1943	4.58	1373	3.40	773	1.74	2653	6.75	2722	24.98
Low Producing										
~1	0	0.00	0	0.00	0	0.00	0	0.00	0	0.00
~south	46	0.13	18	0.05	77	0.20	-28	-0.08	10	0.10
~elev	652	1.81	601	1.55	77	0.20	955	2.64	623	5.88
~south+elev	705	1.95	633	1.63	159	0.41	975	2.70	629	5.94
~souht*elev	719	1.99	656	1.69	169	0.43	993	2.75	650	6.14

539

540 **Figures**



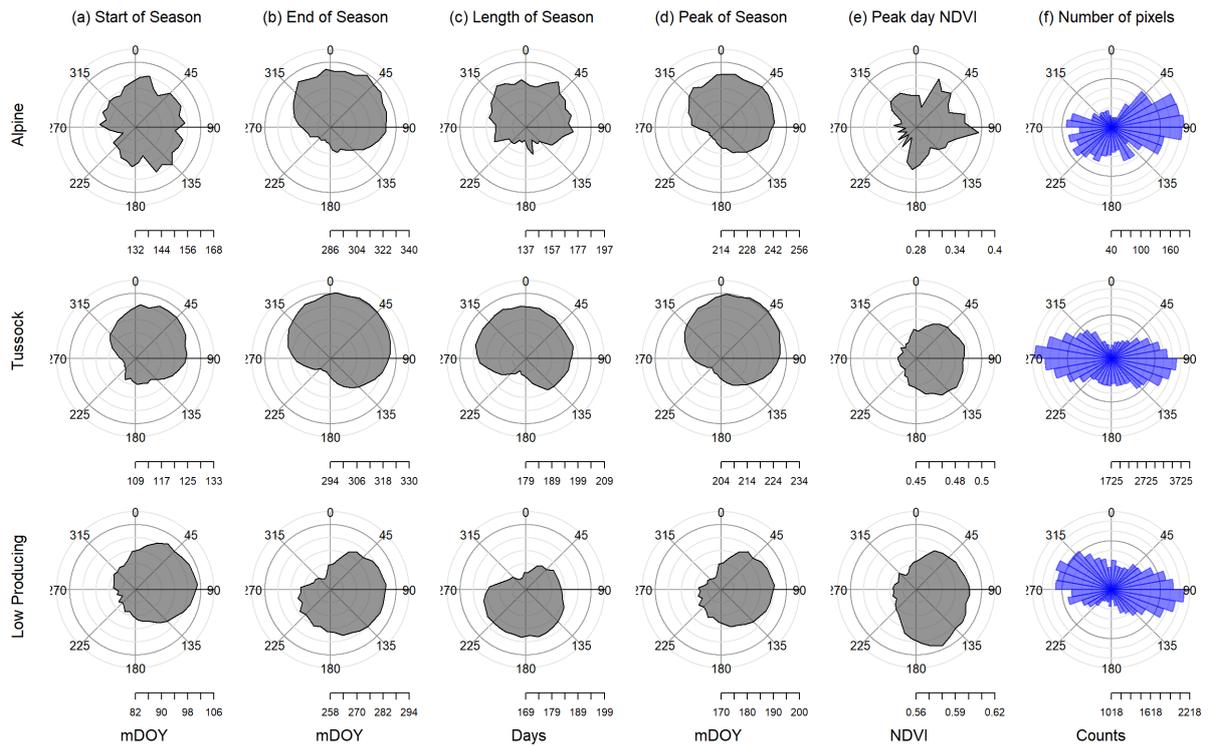


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551 **Fig.2** The MODIS NDVI 2001-2016 time series for an example pixel and derivation of the five
 552 growth phenology indices (SOS = start of season, EOS = end of season, LOS = length of season,
 553 POS = peak of season, P-NDVI = peak day NDVI). In the time series, the blue colour indicates
 554 raw data and the black line shows the fitted curve derived by TIMESAT (see Methods for
 555 details). We used a modified day-of-year (mDOY) to describe growth phenology in the South
 556 Hemisphere, so that a growing year starts on 1st July one year and ends on 30th June the year
 557 after (mDOY 1= 1st July, mDOY 365 = 30st June).

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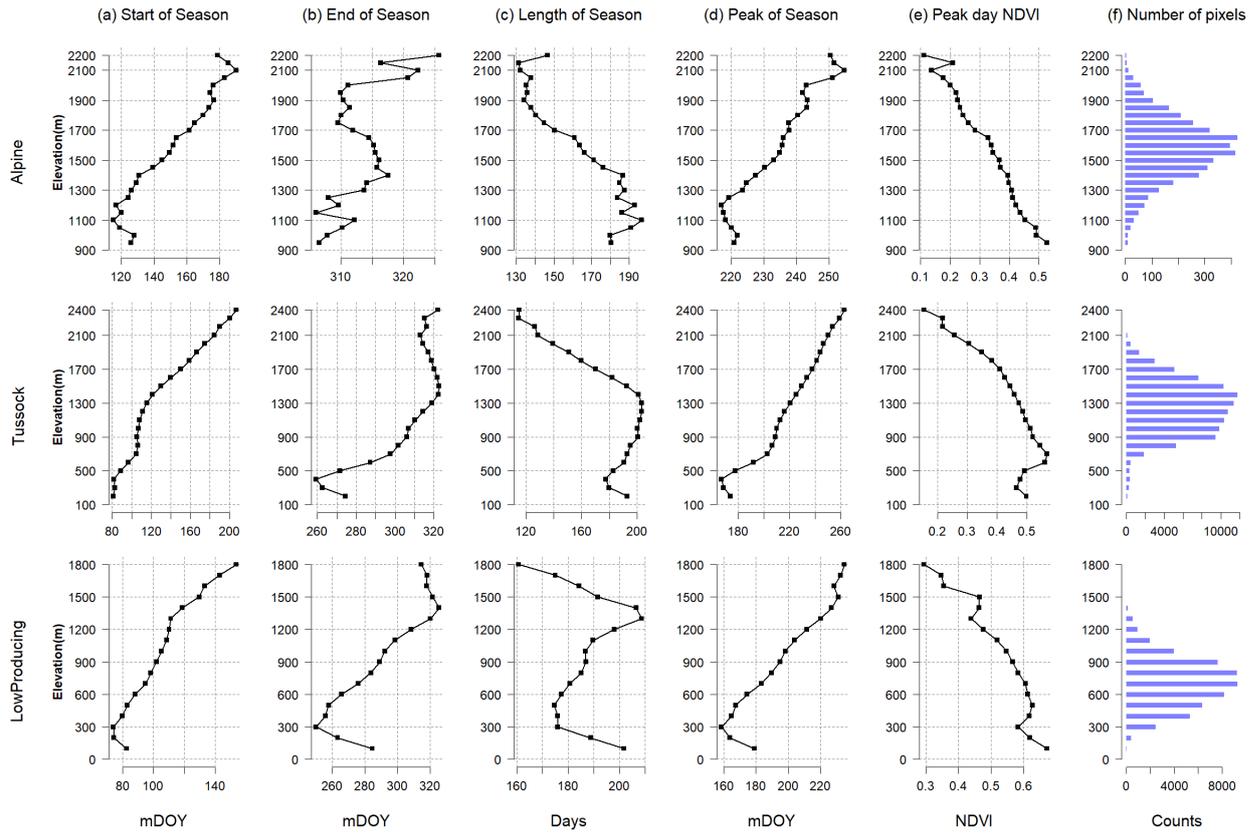


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561 **Fig.3** The 16-year averages of five growth phenological indices (a-e) in relation to aspect in
 562 the study area. For each aspect circle, 0 represents direct north, 90 represents direct east, 180
 563 represents direct south and 270 represents direct west. The shade in each panel illustrates the
 564 16-year average value of each phenological index at each 10-deg of aspect in each grassland
 565 type. The total number of pixels at each 10-deg aspect is shown in panels (f).

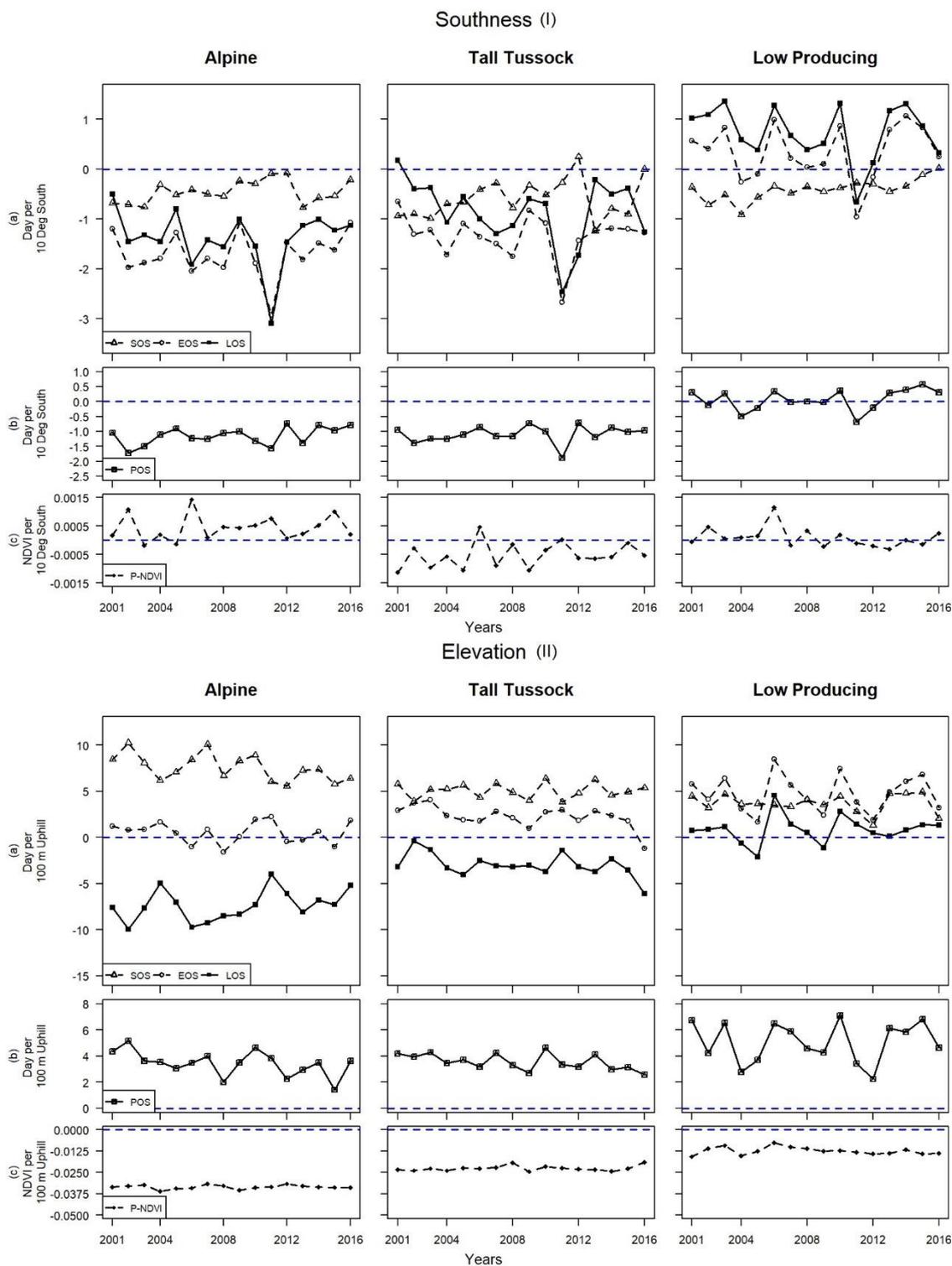
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Topographical effects on phenology



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Fig. 4 The 16-year average of five growth phenological indices (a-e) along the elevation gradient in the study area..mDOY = modified day of year (1 = 1st July; see Methods). The total number of pixels in each 100m elevation band is shown in panels (f).



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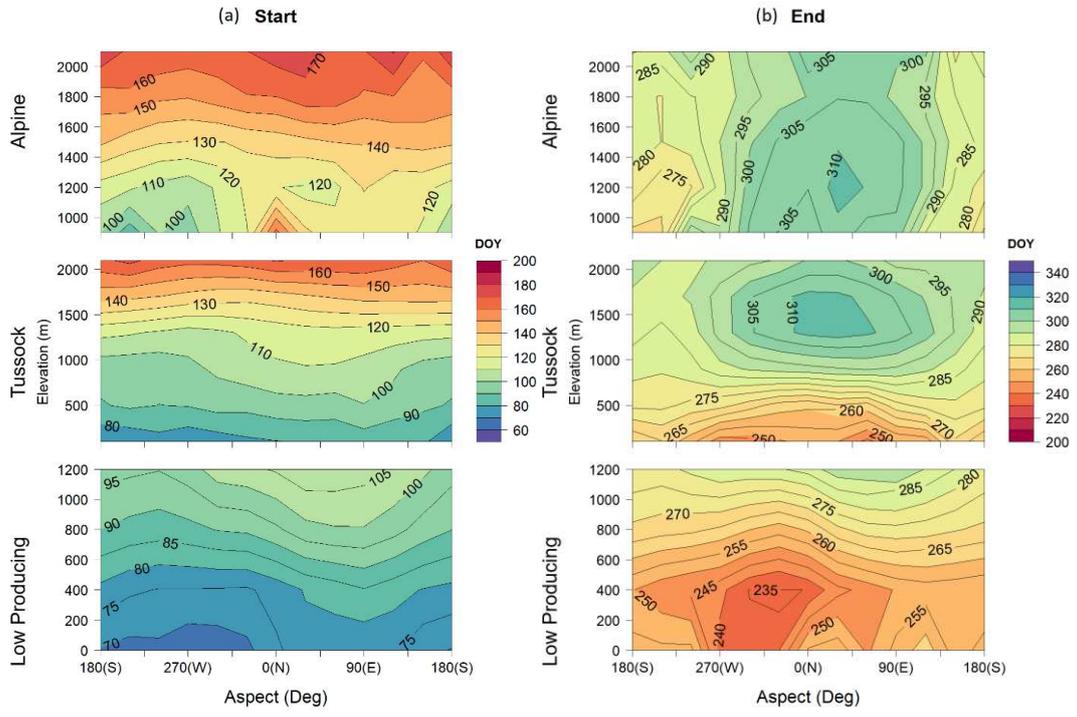
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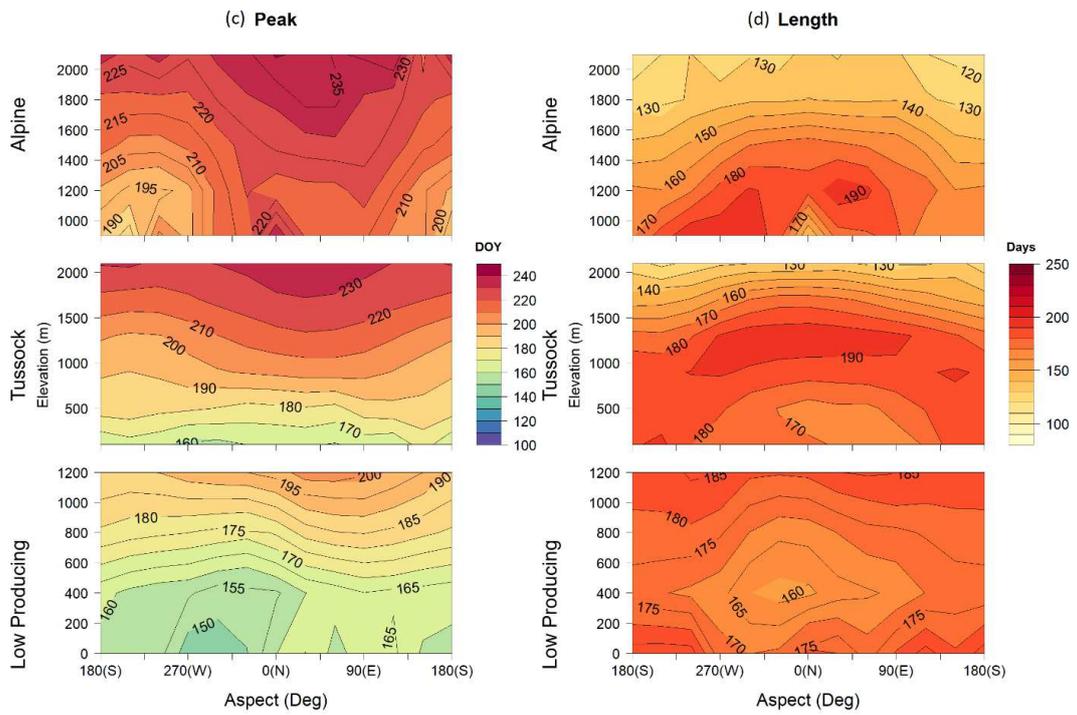
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Fig. 5 The strength of the relationship between phenology and (I) aspect and (II) elevation for each year (2001-2016) and grassland type. The strength of the relationship between the five phenological indices (a-c) and aspect/elevation is expressed as the slope of linear SAR model index~aspect/slope. Positive values (above blue line) indicate later start, end, peak of season; a longer season and higher NDVI on the peak day (see Methods for details). SOS = start of season, EOS = end of season, LOS = length of season, POS = peak of season, P-NDVI = peak day NDVI.

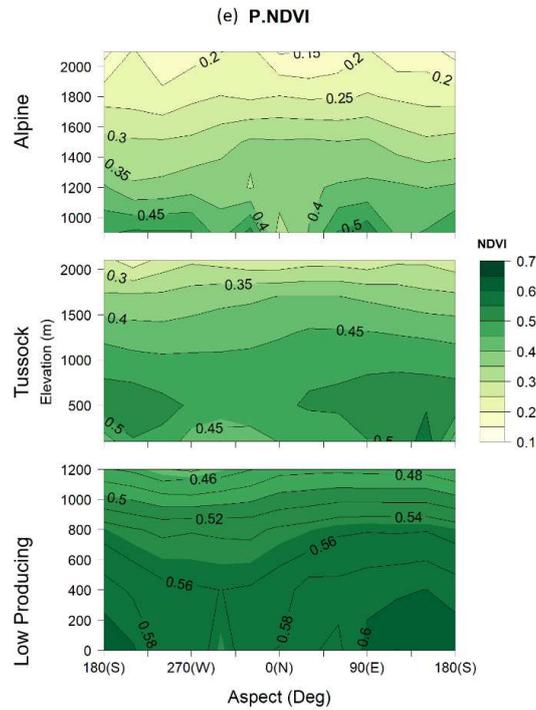
Topographical effects on phenology



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 585 **Fig.6** The interaction effects of aspect and elevation on the five phenological indices (a-e)
 586 Pixels were assigned to bins in aspect-elevation space and 16-yr averages for each index were
 587 calculated for each bin. Bins of similar values were connected with contour lines.
 588 Abbreviations: (a) Start = start of season, (b) End = end of season, (c) Peak = peak of season,
 589 (d) Length = length of season, (f) P-NDVI = peak day NDVI.