

Contributions of winter and spring warming to the temporal shifts of leaf unfolding

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2 **Contributions of winter and spring warming to the temporal shifts of leaf**

3 **unfolding**

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15

16 **Abstract (150 words)** Changes in winter and spring temperatures have been widely
17 used to explain the diverse responses of spring phenology to climate change.
18 However, our understanding of their respective roles remain incomplete.
19 Using >300,000 *in situ* observations of leaf unfolding date (LUD) in Europe, we show
20 that the advancement of LUD since 1950 is due both to accelerated spring thermal
21 accumulation and changes in winter chilling which explain 61% and 39% of the LUD
22 shifts, respectively. Winter warming did not substantially retard the releasing of bud
23 dormancy, but increased the thermal requirement to reach leaf unfolding. The increase
24 of thermal requirement and decreased efficiency of spring warming on accelerating
25 thermal accumulation partly explained the temporally (1950s-2010s) decreasing
26 response of LUD to warming. Our study stresses the need to better assess the
27 antagonistic and heterogeneous effects of winter and spring warming on leaf
28 phenology, which is key to projection of future vegetation-climate feedbacks.

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31 Timing of leaf unfolding influences the duration of plant growing season, and
32 therefore their productivity, as well as land-surface physical properties like the surface
33 albedo and roughness¹⁻³. Shifts in leaf unfolding date (LUD) may also alter
34 competition between plants, community structure and ecosystem water, carbon and
35 nutrient cycles⁴⁻⁶. Many studies based on both *in situ* observations and satellite-
36 derived data have reported a general trend of advancing spring leaf phenology during
37 the past decades, in particular LUD, in response to climate warming in temperate and
38 boreal regions⁷⁻¹¹. However, although there is now mounting evidence that the spring
39 leaf phenology as a whole is shifting earlier in time, the magnitude (or even direction)
40 of this shift shows significant taxonomic, temporal and spatial variations¹²⁻¹⁶. In
41 particular, the sensitivity of LUD to climate warming (S_T , day °C⁻¹), defined as the
42 shift in LUD per °C warming, has significantly declined over the past 30 years in
43 many regions^{7,17}.

44 Temperature has been regarded as the most important environmental factor
45 controlling plant phenology¹⁸. Plants in temperate and boreal regions generally
46 require a certain number of days with cold temperatures to release bud dormancy, and
47 subsequently, a certain number of days with warmer conditions (called forcing
48 temperatures) to trigger cell growth and leaf development¹⁹⁻²². Several studies have
49 argued that changes in winter chilling and thermal accumulation caused by climate
50 change explain the divergent responses of LUD to rising temperature^{7,23-26}. Especially,
51 winter warming can affect LUD through two distinct effects: it may delay the timing
52 of bud dormancy release, and may increase the thermal requirement for bud break^{27,28}.
53 Indeed, experimental studies have shown that the thermal requirement for bud break
54 depends on the amount of winter chilling, that is, the more chilling received, the lower
55 the thermal requirement to reach bud break^{20,29}. These two opposing effects might

56 explain why the sensitivity of LUD to warming has significantly declined over the
57 past 30 years of warming^{7,24}. They might also explain why some species have shown
58 scant or no advances in LUD with climate warming^{8,23,25}. Even so, the impacts of
59 climate warming on winter chilling, spring thermal accumulation and the resulting
60 shifts in LUD have not been quantitatively estimated for multiple species at large
61 spatial scale, in particular when considering the asynchronicity of winter and spring
62 warming^{24,30,31}. Moreover, our mechanistic understanding of the shifts in LUD in
63 response to warming is still far from complete³², especially regarding the observed
64 decline in the sensitivity of LUD to continuously rising temperature⁷. This incomplete
65 understanding limits our ability to project regional and global changes in LUD under
66 future climate warming, as well as the resulting changes in ecosystem structure and
67 functions.

68 Here, we take a new step forward to advance our quantitative and mechanistic
69 understanding in ongoing changes in leaf phenology. Using long-term (1951-2019) *in*
70 *situ* observation data of LUD for six dominant tree species collected at 2945 sites in
71 central Europe³³ (see Methods and Fig. S1 in Supplementary Information, SI), this
72 study aims to 1) quantify the relative contribution of changes in winter and spring
73 warming to the temporal shifts in LUD in Europe; 2) elucidate the mechanisms that
74 can explain explain the declining sensitivity of LUD to rising temperature, as
75 observed over the past three decades⁷.

76 To achieve these aims, we applied the Unified phenology model^{19,28} at each of the
77 2945 observation sites for each tree species (see Methods). Plant phenological models
78 describe the causal relationships between winter and spring temperatures and bud
79 development^{34,35}. Existing phenological models (e.g. the Thermal Time, Sequential,
80 Parallel and Alternating models) generally rely on distinct assumptions regarding the

81 response of bud growth to spring thermal accumulation and/or winter chilling^{34,35}(see
82 Methods). The advantage of the Unified model used here is that it integrates these
83 different assumptions, and each of the earlier models can be regarded as particular
84 cases of the Unified model (see Methods and Fig. S2). A previous study also reported
85 that the Unified model provides a more accurate estimation of LUD in Europe than
86 other commonly used models⁷. For these reasons, we used the Unified model to
87 disentangle the various effects of climate warming on LUD.

88 **Contributions of winter and spring warming to the temporal shifts in LUD**

89 According to the Unified model¹⁹ (see Methods for details), bud dormancy is assumed
90 to be released when chilling accumulation meets plants' critical requirements (CHA_0
91 in Fig. S2). The forcing stage starts as soon as this chilling requirement is met and leaf
92 unfolding occurs when the thermal accumulation during the forcing stage exceeds a
93 given threshold. This threshold, denoted by TA_0 , declines exponentially with the total
94 amount of chilling received during the whole pre-growing season (CHA_{tot} , Fig. S2b),
95 defined as the period from the onset of chilling accumulation till the LUD. Therefore,
96 the temporal shifts in LUD are determined by the time when bud dormancy is released
97 (d_{f0}), which in turn depends on the chilling accumulation rate (CH_r , Eq. 1), by the
98 thermal accumulation rate (F_r , Eq. 2), and by the required amount of thermal
99 accumulation (TA_0 , Eq. 3) which in turn depends on the total amount of chilling
100 (CHA_{tot} , Eq. 4) received before the LUD (Fig. S2).

101 The Unified model with nine optimized model parameters (see Methods and Table
102 S2) captured well the observed LUD for each species at each site (Fig. S3). For the six
103 species considered here (*Aesculus hippocastanum*, *Alnus glutinosa*, *Betula pendula*,
104 *Fagus sylvatica*, *Fraxinus excelsior* and *Quercus robur*, Table S1), the model

105 explained 60-80% of the spatiotemporal variations in LUDs (Fig. S3). For
106 approximately 90% of the entire phenological records (1951-2019), the absolute error
107 in simulated LUDs were shorter than one week (Fig. S3) which is the observation
108 interval of the phenological data used in this study³³.

109 The average LUDs advanced by 1.9 (± 2.3 , standard deviation of both inter-site and
110 inter-species variations) and 5.8 (± 2.9) days during the periods 1980-1999 and 2000-
111 2019, respectively, compared to the reference period 1951-1979 (Δ LUD, Fig. 1a and
112 Fig. S4). Changes in LUD over time for the different species were overall similar, the
113 inter-species differences in average LUD advancement rarely exceed one day for both
114 two periods. Yet, differences in LUD temporal shifts were larger across sites whatever
115 the species, ranging from an advance of more than 10 days to a delay of more than 2
116 days (Figs. S4 & S6). Generally, advances in LUD occurred more often at sites with
117 stronger warming (e.g. northern Germany), while delays in LUD mostly occurred at
118 sites with temperature declines (e.g. southeastern Germany and Austria) during the
119 period 1951-2019 (Figs. S6 & S7). 1 °C rise in mean annual temperature, mean spring
120 temperature and mean winter temperature resulted in an average advance in LUD by
121 4.4 (± 1.8), 3.2 (± 1.2) and 1.4 (± 0.7) days in central Europe, respectively (Fig. S8).

122 Advances in LUD were mainly caused by the acceleration in thermal accumulation
123 induced by the warming of spring (Δ D_{FD}, Fig. 1). For the periods 1980-1999 and
124 2000-2019, the durations of the forcing stage were 4.0 (± 3.8) and 8.7 (± 7.0) days
125 shorter than during the period 1951-1979, respectively. In contrast, the increase in
126 thermal requirements (TA₀) induced by the decline in total pre-season chilling
127 accumulation (CHA_{tot}) delayed LUDs by 1.2 (± 1.6) and 2.3 (± 2.0) days in 1981-2000
128 and 2000-2019 respectively, compared to the earlier period (Δ D_{TA0}, Fig. 1). Changes
129 in dormancy release day (Δ D_{dft0}, which is also the onset of the forcing stage)

130 contributed less to the shifts in LUDs, generally by less than 2 days. From the period
131 1951-1979 to the periods 1980-1999 and 2000-2019, the dormancy release day
132 advanced by 0.3 (± 2.1) and 1.7 (± 3.9) days on average, respectively (Fig. 1a).
133 Temporal shifts in dormancy release day differed significantly across species and
134 were overall larger for *Fraxinus excelsior* and *Quercus robur*, which have
135 comparatively later LUDs (Fig. S9) than other species (Figs. 1 and S4). Overall, the
136 changes in duration of forcing accumulation (ΔD_{FD}), plants' thermal requirement
137 (ΔD_{TA0}) and dormancy release day (ΔD_{df0}) between the reference period 1951-1979
138 and the most recent period 2000-2019 explained 61%, 22% and 17% of the total LUD
139 advance (Fig. 1b). In other words, the acceleration in thermal accumulation caused by
140 climate warming (ΔD_{FD}) explained 61% of the temporal shift in LUD, while changes
141 in chilling accumulation ($\Delta D_{df0} + \Delta D_{TA0}$) explained the remaining 39% through their
142 influence on plants' thermal requirement and dormancy release day (Fig. 1b).

143 **Revealing the complex and antagonistic effects of temperature on LUD**

144 The warming-induced change in the dormancy release day has been widely invoked to
145 explain why the LUDs of some plants show a weak or no delayed shifts with climate
146 warming^{7,8,16}. A common hypothesis relies on the premises that warm temperatures in
147 winter delay the timing at which chilling requirements for dormancy release are met,
148 therefore postponing the start of the forcing stage and mitigating the advancement of
149 the LUD caused by the spring warming. Our results however suggest that the winter
150 warming only marginally changed the timing of dormancy release, especially from
151 1951-1979 to 1980-1999 (Fig. 1). At more than 98% of the sites, the optimized
152 functions for calculating daily chilling rates showed an optimum (see the pattern
153 plotted in Fig. S10b), with the optimal chilling temperature (T_{op}) mostly varying
154 between 0 °C and 12 °C with a mean value of 6.5 (± 2.1) °C (Fig. S11c). Therefore,

155 the widely used index-chilling days^{20,35} (Fig. S10a), which assumes identical effect of
156 chilling temperature below a threshold, was less efficient in explaining the observed
157 LUD and, thus, was not retained by the optimization algorithm. Winter temperatures
158 recorded at the observation sites were mostly 3 to 12 °C lower than the local T_{op}
159 during the period 1951-1979, and then gradually increased towards T_{op} with time (Fig.
160 2a). This decreasing difference between winter and optimal chilling temperatures
161 translated into an important increase in the chilling rates during winter days when
162 temperatures are lower than T_{op} (e.g. days between d_{c1} and d_{c2} in Fig. 3). As a result
163 of these increased chilling rates, the warming-induced loss in chilling rates in late
164 autumn and early spring (e.g. days before d_{c1} and days after d_{c2} in Fig. 3) were
165 actually partly or sometimes even fully offset. This explains why the dormancy
166 release days (ΔD_{df0}) for all species were not substantially postponed and were in fact
167 mostly occurring slightly earlier in time (Figs. 1 and S4).

168 The decrease in chilling accumulation rates after dormancy release from the period
169 1951-1979 to 2000-2019 because of the moving away of temperature from T_{op} due to
170 the warming of spring (Figs. 2b and 3), additionally resulted in an overall decrease of
171 the total amount of chilling accumulated before bud break (Fig. 2c). As a result,
172 while dormancy could be released earlier due to more efficient chilling in cool winter
173 during the last decades, the total chilling accumulation declined by $0.8(\pm 0.9)\%$ on
174 average from 1951-1979 to 1980-1999, and by $2.0(\pm 1.2)\%$ from 1951-1979 to 2000-
175 2019 (Fig. 2c). This decline translated into a slight but significant increase in the
176 critical forcing accumulation required for leaf unfolding (TA0), by $0.5(\pm 0.5)\%$ and
177 $1.2(\pm 0.8)\%$ between the reference period and the periods 1981-2010 and 2000-2019,
178 respectively (Fig. 2d). This slight increase in TA0 was however largely compensated
179 by an important increase of the forcing rate (F_r , Fig. 2e) during spring of $8.9(\pm 10.4)\%$

180 and 22.9(\pm 13.9)% between the reference period and the 1981-2010 and 2000-2019
181 periods respectively. As a result, the plants' thermal requirements were reached
182 within a shorter time interval in the recent decades compared to the period 1951-1979,
183 resulting in a significant advancement of the LUD (Fig. 1).

184 **Why does the temperature sensitivity of LUD decline with climate warming?**

185 Consistent with previous studies conducted in Europe and China^{7,17}, we also find a
186 general decline in temperature sensitivity of LUD (S_T , i.e. change in LUD per $^{\circ}\text{C}$
187 warming), although this decline seems to have stopped in the last 20 years in central
188 Europe (Fig. S12). Linear regression between S_T and different metrics of winter
189 chilling and forcing temperature reveal that the average daily chilling rate during the
190 dormancy stage (CH_{rave} from d_{c0} to d_{f0}), the total chilling accumulation during the
191 whole pre-growing season (CHA_{tot} from d_{c0} to LUD) and the average daily forcing
192 rate during the forcing stage (F_{rave} from d_{f0} to LUD) across all sites explained
193 17.6(\pm 18.3)%, 18.4(\pm 18.9)% and 21.0(\pm 20.5)% of the temporal variation of S_T on
194 average, respectively (Fig. 4). CH_{rave} and CHA_{tot} , which together represents the
195 overall impacts of winter chilling on LUD, explained 33.4(21.5)% of the variation in
196 S_T . By further including F_{rave} , the three metrics together accounted for 44.5(\pm 21.0)%
197 of the variation in S_T . The forcing duration (FD) alone, which is determined by
198 CHA_{tot} and F_{rave} , explained 21.3(20.6)% of the variation in S_T . Irrespective of the
199 metrics, their explanatory power varied drastically among observation sites, as
200 demonstrated by the large interquartile (thick) and 95% confidence intervals of the R^2
201 in Fig. 4. Note that the pre-growing season for calculating S_T in Fig. 4 is defined as
202 the period from d_{c0} to LUD and thus differs from the one applied in several previous
203 studies^{7,17,36}. In earlier work, the pre-growing season was usually defined as the period
204 for which the absolute value of the correlation coefficient between LUD and air

205 temperature was largest (i.e. the period which is most relevant to leaf unfolding, see
206 Methods). Nevertheless, our analysis suggests that the abilities of different chilling
207 and forcing metrics to explain the S_T variation calculated based on these two distinct
208 definitions of the pre-growing season were overall similar (Figs. 4 and Fig. S13).

209 Changes in winter chilling due to warming have been proposed as the main
210 explanation of the temporal variation in S_T ^{7,17}. Our results show that the overall effect
211 of winter warming on S_T is indeed higher than that of spring warming. They further
212 support the hypothesis that winter warming results in a decline of S_T by inducing a
213 decrease in the total amount of chilling received and a subsequent increase in the
214 critical forcing requirements for leaf unfolding (Figs. 2, 4 and Eq. 3). However, they
215 do not support the hypothesis that winter warming leads to a decline of S_T by delaying
216 the timing of dormancy release (d_{f0}), at least up to 2019. Indeed, while changes in
217 CH_{rave} due to winter warming can induce changes in d_{f0} (Fig. 2) and explain a part of
218 the temporal variation in S_T (Fig. 4), we found very limited changes in CH_{rave} and d_{f0} .
219 In fact, at many observation sites, d_{f0} actually occurred slightly earlier and not later in
220 time in response to the winter warming (Figs. 1 and S4).

221 For all species, the increase of forcing rate consistently explained more variation in S_T
222 than did chilling rate and chilling accumulation taken alone (Fig. 4). Since dormancy
223 break date did not change significantly (Fig. 1) and the forcing rate increased
224 substantially, the duration of the forcing stage decreased which explained most of the
225 advancement in LUD. Therefore, the variation in S_T was, to date, also strongly
226 dependent on the duration of the forcing stage and the average temperature during this
227 stage. For example, in Fig. 3, when the LUD advances from d_0 to d_{0+6} , the forcing
228 temperatures on days between d_{0+6} and d_0 become useless to leaf unfolding (but still
229 useful to late leaf growing). The daily average forcing rate (F_{rave}) from d_{f0} to d_{0+3} in

230 the year with +3 °C warming (Y_{0+3}) has to be larger than the F_{rave} from d_{f0} to d_0 in the
231 reference year Y_0 , and the increment of F_{rave} from Y_{0+3} to the year with +6 °C warming
232 (Y_{0+6}) should be even larger than the increment from year Y_0 to Y_{0+3} , due to the
233 further decreased DF. This explains why the temperature sensitivity of LUD is
234 significantly related to the DF (Fig. 4).

235 The weakening effect of spring warming on S_T can be explained by the non-linear
236 shape of the relationship between daily forcing rate (F_r) and daily mean temperature
237 (Eq. 2). At most sites, this relationships indeed followed a sigmoid pattern (e.g.
238 forcing rate presented in Fig. S2a), rather than a nearly linear change pattern as the
239 widely used forcing metric-growing degree-days (defined as the difference between
240 daily mean temperature and a base temperature, e.g. Fig. S10d). Therefore, the effects
241 of warming on accelerating forcing accumulation increases as temperature moves
242 towards T_{50} , the temperature inducing 50% of optimal growth in forcing rate and the
243 inflection point of the function, but decreases as temperature moves away from T_{50} .
244 For example, in Fig. S2a, the forcing rate increases substantially when daily
245 temperature rises from 5 °C to 10 °C, but increases limitedly when the temperature
246 further rises to 15 °C. The T_{50} estimates in Europe mostly varied between 8 to 11°C
247 on average (Fig. S11). In 1951-1979, the temperature during the forcing stage was
248 $0.35(\pm 2.36)$ °C lower than T_{50} on average (Fig. S14). However, in 1980-1999 and
249 2000-2019, the temperatures during the forcing stage were $0.08(\pm 2.28)$ and
250 $1.55(\pm 2.51)$ °C higher than T_{50} on average, respectively. Therefore, spring mean
251 temperature moved towards T_{50} from 1951-1979 to 1980-1999 while it moved away
252 from it afterwards, which explains why the effect of warming on accelerating forcing
253 accumulation decreased the last 20 years while it had increased in the earlier periods.

254 Although temperature is the dominant factor of spring phenology¹⁸, other factors,
255 especially photoperiod, may also influence the response of LUD to climate warming
256 in some species^{5,37}. Temperature can vary strongly within a short time (e.g. a few
257 days) and may show large interannual variations. Sensitivity to photoperiod protects
258 plants from the potentially fatal consequences (e.g. frost damage) short warm spells
259 during autumn and winter³⁸. Photoperiod has also been shown to impact bud growth
260 after dormancy break for a few species, among which *Fagus sylvatica* which shows
261 the greatest sensitivity to this factor^{39,40}. Our results indeed reveal that the three
262 temperature metrics (CH_{rave} , CH_{tot} and F_{rave}) together explain less than half of the
263 temporal changes in S_T at many sites (Fig. 4). Therefore, part of the remaining
264 variation might be explained by an effect of photoperiod. In addition, we found that
265 the relationships between temperature sensitivity of LUD and the duration of the
266 forcing stage and the forcing rate are likely nonlinear at many sites (e.g. Fig. S15).
267 Using linear regression functions may therefore underestimate the impacts of these
268 variables on the temporal variation of S_T .

269 Understanding the controlling mechanism of climate warming on LUD is important
270 for accurately projecting future vegetation dynamics and ecosystem carbon cycle^{5,37}.
271 This study quantified the potential contributions of warming-induced changes in
272 dormancy release day, forcing rate and thermal requirement for leaf unfolding to the
273 advances of LUD of six dormant tree species in central Europe. It showed that the
274 overall effect of the winter warming explained most of the decline in the sensitivity of
275 LUD to temperature in Europe from 1951-2019 although not through a delayed
276 dormancy break date. It showed additionally that the spring warming by increasing
277 the forcing rate and thereby decreasing the duration of the forcing stage also
278 explained a large part of the decline in the temperature sensitivity of LUD because of

279 a lower increase of forcing rate per degree warming. Our results revealed the
280 importance of representing the antagonistic effects and their inhomogeneous
281 effectiveness of the chilling and forcing temperatures on bud development, especially
282 when considering the seasonally asynchronous climate change^{30,31}. These results are
283 helpful to improve the simulation of plant phenology in earth system models,
284 therefore decreasing the uncertainties in projections of future vegetation dynamics and
285 carbon cycle-climate feedbacks. Note that our findings are based on phenological
286 observations from central Europe, they may not hold true for other regions. Therefore,
287 we encourage more studies based on global field observations or control laboratory
288 experiments to further explore the mechanisms driving plant phenology in a larger
289 range of climatic conditions.

290 **Methods**

291 **Phenological and climatic data** Phenological dataset of leaf unfolding day (LUD)
292 during the period 1951- 2019, as well as the geographical location and elevation for
293 each observation site, were obtained from the Pan European Phenology (PEP)
294 network (<http://www.pep725.eu>), an open access repository of *in situ* phenological
295 records for multiple plant species across Europe³³. The LUDs were defined by the
296 BBCH (Biologische Bundesanstalt, Bundessortenamt und Chemische Industrie) code
297 as stage 11 (first leaf unfold). Specifically, we selected the LUD records of the six
298 mostly observed tree species at 2,944 sites in central Europe (Table S1 and Fig. S1 in
299 Supplementary Information (SI)). For each species, only the sites with phenological
300 observations of at least 40 years were included in our analysis. The median absolute
301 deviation (MAD) method^{41,42} was then used to identify and exclude potentially
302 erroneous records of LUD. For each species at each site, the MAD was calculated as:
303 $MAD = \text{median} (|LUD_i - \text{median} (LUD_1, LUD_2, \dots, LUD_n)|)$, where $LUD_{1, i, n}$ are the

304 LUD in the 1th, ith and nth observation year, respectively. Any record deviating by
305 more than 3 times the MAD was considered as an outlier and removed from the
306 original dataset used in this study. After MAD pre-treatment, > 300,000 LUD records
307 were left for further analysis.

308 Daily mean elevation and air temperature at each site were obtained from the gridded
309 database E-OBS (<http://ensembles-eu.metoffice.com>) at a spatial resolution of 0.1°
310 (approximately 10 km)⁴³. Due to the relatively coarse spatial resolution of the E-OBS
311 database, the elevation of some phenological sites may be vastly different from the
312 mean elevation of the grid cells where they locate, especially in mountainous areas.
313 To reach a more accurate quantification of temperature at each site, the temperature
314 data from E-OBS were adjusted using the difference between the actual elevation at
315 each site (as specified in the PEP database) and the mean elevation of the
316 corresponding grid cell (as specified in E-OBS). The temperature lapse rate is set to
317 0.64 °C per 100 m change in elevation⁴⁴.

318 **The Unified phenology model** Phenological models are generally developed based
319 on distinct assumptions regarding the response of bud growth to spring thermal
320 accumulation and/or winter chilling^{34,35}. For example, the Thermal Time model
321 considers that only forcing temperatures explain the LUD²². More sophisticated
322 models also consider the effect of chilling temperatures on bud dormancy breaking.
323 For example, the Sequential model assumes that the effect of forcing temperatures
324 cannot be effective unless chilling requirements have already been fulfilled⁴⁵. In
325 contrast, the Parallel model assumes that forcing temperatures can be active
326 concomitantly with chilling accumulation⁴⁵. Furthermore, in other models such as the
327 Alternating model, the thermal accumulation required for leaf unfolding is assumed to
328 decline exponentially with increasing chilling accumulation²⁰. These models are non-

329 nested within each other and , therefore, cannot be used to test the effects of the above
330 assumptions on the response of LUD to chilling and forcing temperatures
331 simultaneously¹⁹. To circumvent these limitations, we applied the Unified phenology
332 model¹⁹ in our study.

333 The Unified phenology model¹⁹ allows for a direct estimation of the response of
334 spring phenology to both chilling and forcing temperatures, and of the periods when
335 these temperatures affect the plant phenology. In addition, it integrates the main
336 assumptions of simpler phenology models, in particular the sequential⁴⁵, parallel³⁴ and
337 alternating²⁰ models, which can in fact be regarded as particular cases of the Unified
338 model. Moreover, previous studies have indicated that the Unified model generally
339 provides more accurate estimations of the LUD of European trees than other
340 commonly used models⁷.

341 In the Unified model, the daily chilling rate (CH_r , unitless) during cold days is
342 calculated using a unimodal function (Fig. S2a) of daily mean air temperature (T , °C):

$$343 \quad CH_r = \frac{1}{e^{c_1(T-T_{op})^2} + c_2(T-T_{op})} \quad (1)$$

344 where T_{op} is the optimal chilling temperature, and c_1 and c_2 are two calibration
345 coefficients (Table S2). With specific values of T_{op} , c_1 and c_2 , Eq.1 can capture other
346 widely used chilling metrics¹⁹ such as, the chilling days ($CH_r = 1$ if $T \leq 5$ °C; $CH_r = 0$
347 if $T > 5$ °C) (Figs. S10a,b). Bud dormancy is released (d_{f0}) when the accumulation of
348 daily chilling rate since a specific day (d_{c0} , the start day of chilling accumulation)
349 exceeds the plants' critical chilling requirement (CHA_0 , unitless), i.e. when
350 $\sum_{d_{c0}}^{d_{f0}} CH_r > CHA_0$. Note that d_{f0} also corresponds to the start day of forcing (thermal)

351 accumulation (Fig. S2). The daily forcing rate (F_r , unitless) is calculated using a
 352 sigmoid function (Fig. S2a) of daily mean air temperature (T , °C):

$$353 \quad F_r = \frac{1}{1.0 + e^{c_3(T - T_{50})}} \quad (2)$$

354 where T_{50} is the mid-response temperature, which induces 50% of optimal growth in
 355 forcing rate and is the inflection point of the function. c_3 is a calibration coefficient.
 356 Leaf unfolding occurs when the forcing accumulation ($\sum_{d_{f0}}^{LUD} F_r$) exceeds a certain
 357 thermal requirement (TA_0) which declines exponentially with the total chilling
 358 accumulation (CHA_{tot}) during the whole pre-growing season (i.e. period from the start
 359 day of chilling accumulation to the LUD) (Fig. S2b in SI):

$$360 \quad TA_0 = c_4 e^{c_5 CHA_{tot}} \quad (3)$$

$$361 \quad CHA_{tot} = \sum_{d_{c0}}^{LUD} CH_r \quad (4)$$

362

363 where c_4 and c_5 are two calibration coefficients. The important point to understand in
 364 this model is that after dormancy break, chilling temperature continue to have an
 365 effect on CHA_{tot} and forcing requirement decreases as CHA_{tot} increases.

366 **Parameter estimation and model application** Nine parameters of the Unified
 367 model, including the start day of chilling accumulation (d_{c0}), the critical chilling
 368 requirement for releasing bud dormancy (CHA_0), the optimal chilling temperature
 369 (T_{op}), the mid-response temperature (T_{50} in Eq. 2) and the five coefficients c_1 to c_5 in
 370 Eqs. 1-3, were optimized for each species at each site using an effective global
 371 optimization algorithm-the shuffled complex evolution algorithm (SEC-UA)⁴⁶. Prior
 372 value and the range of each parameter to be optimized are listed in Table S2. Root
 373 mean square error (RMSE, Eq. 5) between simulated (LUD_{sim_i}) and observed

374 (LUD_{obs_i}) LUDs was used as the objective function, and parameter values that
375 minimized the RMSE were regarded as optimal.

$$376 \quad RMSE = \sqrt{\left(\frac{\sum_{i=1}^n (LUD_{sim_i} - LUD_{obs_i})^2}{n}\right)} \quad (5)$$

377 where n is the number of LUD records (years) for each species at each site.

378 Using the optimized parameters, the Unified model was then applied to estimate the
379 day of bud dormancy release, total winter chilling accumulation, daily forcing rate
380 and the LUD for each year, each species and each site. We divided the whole 1951-
381 2019 period into three shorten timespans: 1951-1979, 1980-1999 and 2000-2019.
382 Analysis of the time series of LUD, mean winter temperature and mean spring
383 temperature showed that the LUD and temperature were overall stable in Europe
384 during 1951-1979 (Figs. S15 and S16). During the 1980-1999 period, the spring
385 temperature in Europe increased significantly and the LUD advanced quickly, by up
386 to 0.5 day yr⁻¹. During 2000-2019, the spring temperature in Europe was again overall
387 stable, while winter temperature increased significantly (Figs. S15 and S16). For each
388 of these three timespan, in addition to the average value of LUD, spring and winter
389 temperatures, we also calculated the metrics needed to constrain winter chilling and
390 spring forcing. The changes in each variable was then computed and their significance
391 was evaluated using an independent t-test.

392 In the Unified model, the LUD is determined by the time when bud dormancy is
393 released (which itself depends on the daily rate of chilling accumulation), the rate of
394 forcing accumulation and the required amount of forcing (TA_0) required for leaf
395 unfolding (Fig. S2). To assess the relative contributions of these three factors to the
396 temporal shifts in LUD, we expressed their effects in number of days. We first
397 calculated the changes in the date of bud dormancy break (d_{f0} in Fig. S2c) from the

398 period 1951-1979 (d_{f0_1970s}) to periods 1980-1999 (d_{f0_1990s}) and 2000-2019 (d_{f0_2010s})
 399 (Eqs. 6 & 7) for each tree species at each site (ΔD_{df0} , Figs. 2 & S4).

$$400 \quad \Delta D_{df0_1990s} = d_{f0_1990s} - d_{f0_1970s} \quad (6)$$

$$401 \quad \Delta D_{df0_2010s} = d_{f0_2010s} - d_{f0_1970s} \quad (7)$$

402 Second, to estimate the shifts in LUD caused by the increased thermal accumulation
 403 (ΔD_{TA0} in Fig. 1) required for leaf unfolding due to warming-induced loss of winter
 404 chilling, we simulated the LUD for each species at each site for the two more recent
 405 timespans (1981-2019) by fixing the TA_0 in the Unified model to the average TA_0
 406 corresponding to the 1951-1979 period. From the period 1951-1979 to periods 1980-
 407 1999 and 2000-2019, the ΔD_{TA0} was then calculated as the difference between the
 408 simulated LUDs using the 1951-1979 TA_0 (LUD_{TA0_1970s}) and the LUDs calculated on
 409 the basis of the actual total winter chilling accumulation for each year using Eq. 3
 410 (LUD_{TA0_1990s} and LUD_{TA0_2010s}) (Eqs. 8 & 9).

$$411 \quad \Delta D_{TA0_1990s} = LUD_{TA0_1990s} - LUD_{TA0_1970s} \quad (8)$$

$$412 \quad \Delta D_{TA0_2010s} = LUD_{TA0_2010s} - LUD_{TA0_1970s} \quad (9)$$

413 Third, the contribution of warming-induced changes in daily forcing rates to shifts in
 414 LUDs (ΔD_{FD}) was calculated as the changes in the duration of the forcing stage (FD,
 415 i.e. the number of days from d_{f0} to LUD) from the period 1951-1979 (FD_{1970s}) to the
 416 periods 1980-1999 (FD_{1990s}) and 2000-2019 (FD_{2010s}) respectively, but excluding the
 417 influence of TA_0 changes on this temporal shift:

$$418 \quad \Delta D_{FD_1990s} = FD_{1990s} - FD_{1970s} - \Delta D_{TA0_1990s} \quad (10)$$

$$419 \quad \Delta D_{FD_2010s} = FD_{2010s} - FD_{1970s} - \Delta D_{TA0_2010s} \quad (11)$$

420 Finally, the relative contributions of ΔD_{df0} ($p\Delta D_{df0}$), ΔD_{FD} ($p\Delta D_{FD}$) and ΔD_{TA0}
 421 ($p\Delta D_{TA0}$) to the shifts in LUD were calculated as follows:

$$422 \quad p\Delta D_{df0} = \left(\frac{|\Delta D_{df0}|}{|\Delta D_{df0}| + |\Delta D_{FD}| + |\Delta D_{TA0}|} \right) \times 100\% \quad (12)$$

$$423 \quad p\Delta D_{FD} = \left(\frac{|\Delta D_{FD}|}{|\Delta D_{df0}| + |\Delta D_{FD}| + |\Delta D_{TA0}|} \right) \times 100\% \quad (13)$$

$$424 \quad p\Delta D_{TA0} = \left(\frac{|\Delta D_{TA0}|}{|\Delta D_{df0}| + |\Delta D_{FD}| + |\Delta D_{TA0}|} \right) \times 100\% \quad (14)$$

425 **Temperature sensitivity of LUD** The temperature sensitivity of LUD (S_T , day °C⁻¹)

426 was first calculated for each species at each site based on a linear least square
 427 regression analysis of LUD and mean pre-season temperature, defined as the period
 428 from the start date of chilling accumulation (d_{co} in Fig. 3) to the mean LUD. The
 429 slope of the linear regression line was then used to quantify S_T . In addition, we also
 430 calculated the S_T when the pre-season for each species at each site is defined as the
 431 period (with 5-day steps) for which the absolute value of the correlation coefficient
 432 between LUD and air temperature was highest during 1951-2019, i.e. the period
 433 which is most relevant to leaf unfolding⁷. In this case, we first calculated the mean
 434 temperature during each of the 27 periods ranging from 20 to 150 days (i.e. 20,
 435 25, ..., 145, 150, each at 5 day intervals) preceding the mean LUD. Pearson's
 436 correlation coefficients were calculated between the LUD and the mean temperature
 437 during each of these 27 periods. The period during which the absolute value of
 438 correlation coefficient between LUD and mean air temperature was highest was
 439 regarded as the optimal pre-season.

440 To estimate the trend in the temporal change of the temperature sensitivity of LUD,
 441 we also conducted a reduced major axis regression for each species at each site with a

442 15-year moving window from 1951 to 2019, that is, we calculated the temperature
443 sensitivity for each continuous 15 years over the entire record period.

444

445 **Data and code availability**

446 Phenology data are available from the Pan European Phenology (PEP) network
447 (<http://www.pep725.eu>). Climate data can be downloaded from E-OBS site:
448 <http://ensembles-eu.metoffice.com>. The codes of the Unified model and the program
449 (SCE-UA algorithm) used for parameterization can be found in the supplementary
450 materials. All programs for data analysis used in this study can be obtained by
451 contacting the corresponding author.

452

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458 collecting and providing the phenological data.

459

460 **Author contributions**

461

462 **Competing interests**

463 The authors declare no competing interests.

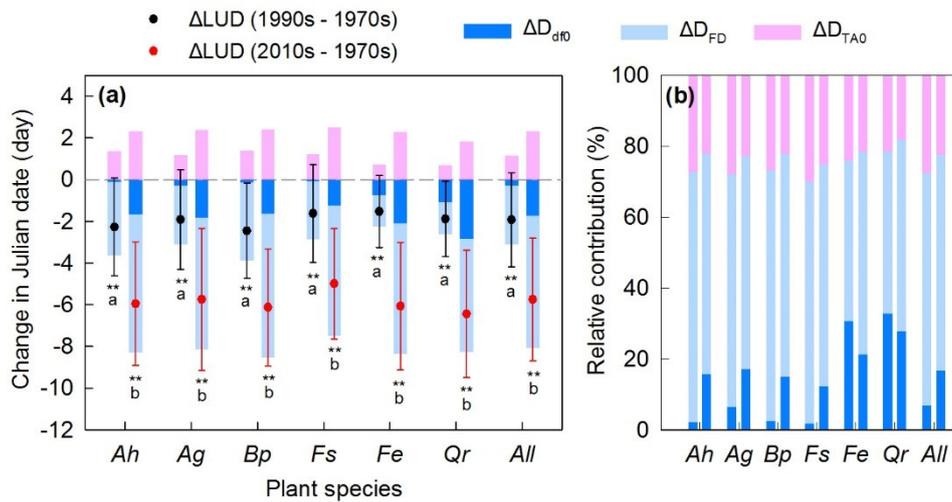
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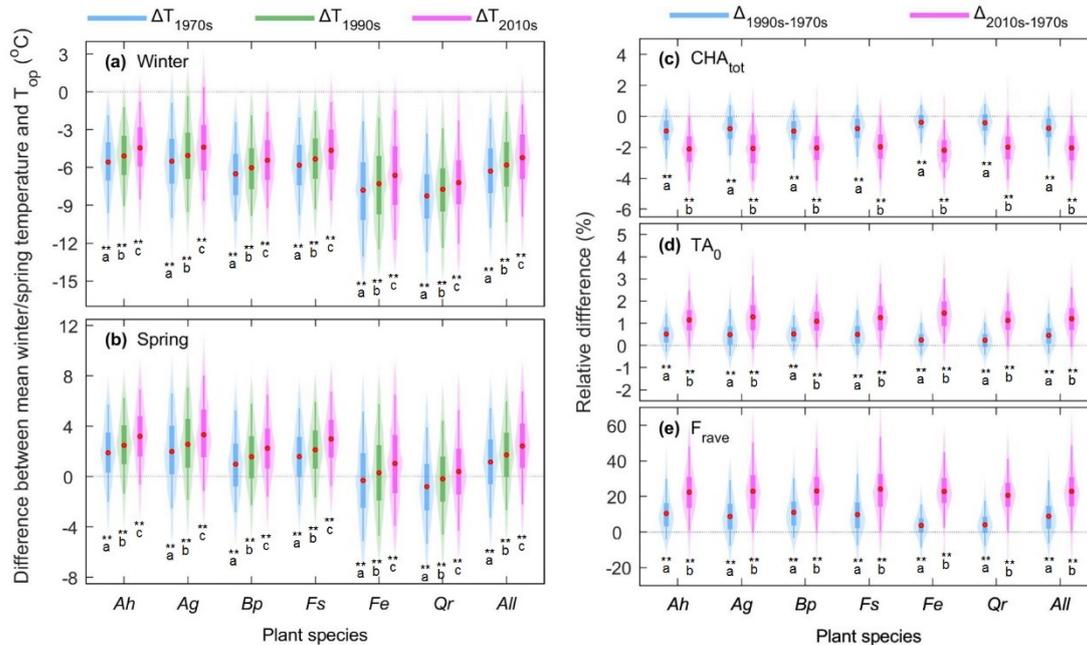
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598 **Figure 1 | Changes in the leaf unfolding days (Δ LUD) of six plant species in**
 599 **Europe from 1951 to 2019 and the potential absolute (a) and relative (b)**
 600 **contributions of changes in winter chilling and spring forcing to the Δ LUD.** The
 601 negative value denotes an advance of LUD; otherwise, the negative value denotes a
 602 delayed LUD. The black and red dots show the average Δ LUD between 1980-1999
 603 (black), 2000-2019 (red) and 1951-1979, respectively. Error bars denote the standard
 604 deviation of Δ LUD at different observation sites. The asterisks (**) indicate the
 605 changes in LUD are significantly different from zero ($p < 0.01$), and the different
 606 letters (a, b) below asterisks means the average LUDs in 1980-1999 are significantly
 607 different from that in 2000-2019 ($p < 0.01$). ΔD_{dfo} is the changes in date when
 608 dormancy is released. ΔD_{FD} is the potential changes in the duration (day) of forcing
 609 stage caused by changes in spring forcing temperatures. ΔD_{TA0} is the potential shifts
 610 in LUD caused by changes in plants' critical requirement of thermal accumulation.
 611 *AH* denotes *Aesculus hippocastanum*; *AG* denotes *Alnus glutinosa*; *BP* is *Betula*
 612 *pendula*; *FS* is *Fagus sylvatica*; *FE* is *Fraxinus excelsior*; *QR* is *Quercus robur*; *All*
 613 represents the average values for all of the six species.

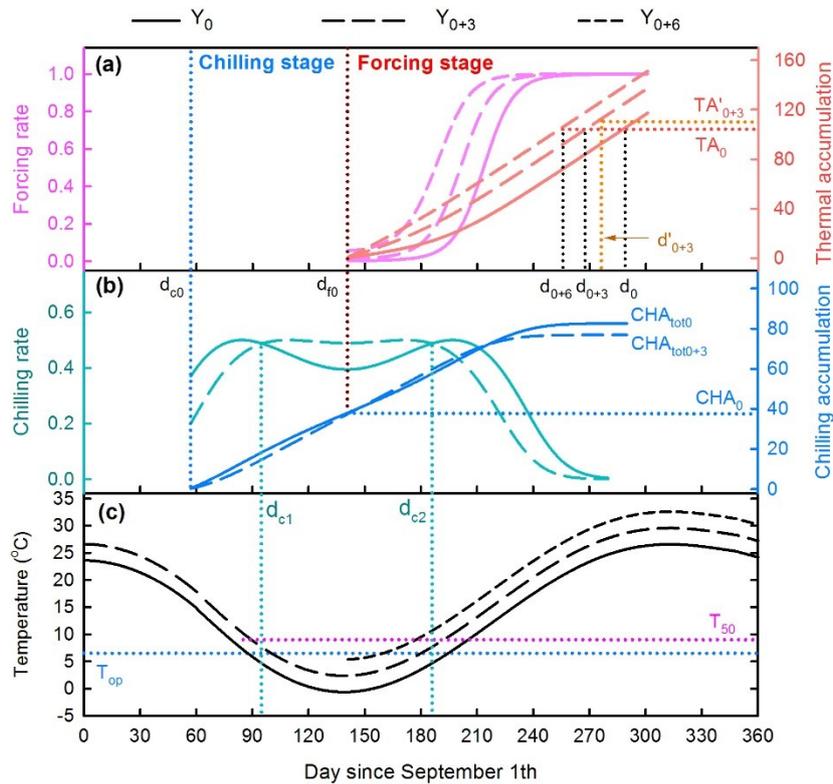
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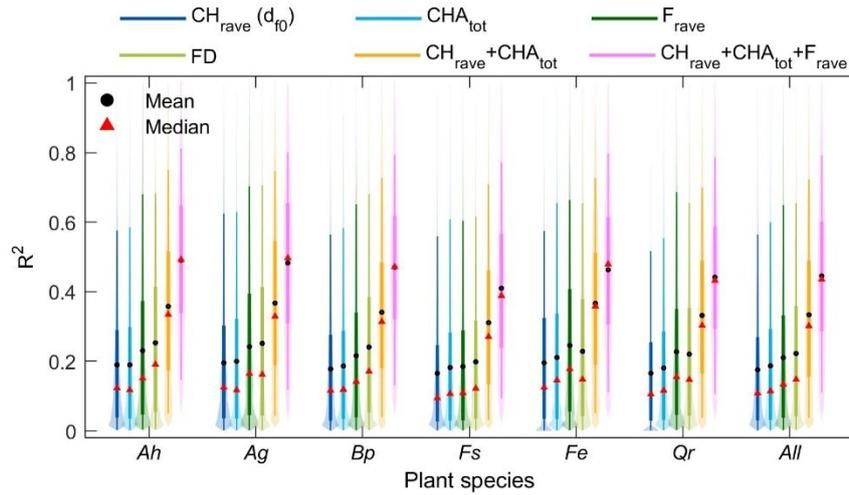
616 **Figure 2 | Differences between mean winter (December-February) (a), mean**
 617 **spring (March-May) (b) temperatures and the optimal chilling temperature**
 618 **(T_{op} , °C) in 1951-1979 (ΔT_{1970s}), 1980-1999 (ΔT_{1990s}) and 2000-2019 (ΔT_{2010s}), and**
 619 **the relative changes in total chilling accumulation (CHA_{tot}) (c), the critical**
 620 **forcing accumulation required for leaf unfolding (TA_0) (d) and the average daily**
 621 **forcing rate during forcing stage (F_{rave}) (e) from 1951-1979 to 1980-1999 ($\Delta_{1990s-1970s}$ -**
 622 **1970s) and to 2000-2019 ($\Delta_{2010s-1970s}$), respectively. In each violin plot, red dot is the**
 623 **mean value and balloon representing the probability density distribution of each**
 624 **value. Whiskers indicate the interquartile (thick) and 95 % confidence intervals (thin).**
 625 **The asterisks (**) indicate that the differences between winter/spring temperatures**
 626 **and T_{op} (a,b) or the relative changes in CHA_{tot} (c) or TA_0 (d) are significantly**
 627 **different from 0 ($p < 0.01$). For each plant species, the violins marked with different**
 628 **letters (i.e. a, b and c below the asterisks) are significantly different ($p < 0.01$). AH**
 629 **denotes *Aesculus hippocastanum*; AG denotes *Alnus glutinosa*; BP is *Betula pendula*;**
 630 **FS is *Fagus sylvatica*; FE is *Fraxinus excelsior*; QR is *Quercus robur*; All denotes the**
 631 **average values for all of the six species.**

632



633

634 **Figure 3 | Schematic plot showing the impacts of rising temperatures (+3 and**
 635 **+6 °C) on winter chilling, spring forcing and leaf unfolding day.** T_{op} is the optimal
 636 chilling temperature (°C). CHA_0 is the critical chilling requirement for releasing
 637 dormancy. CHA_{tot0} and CHA_{tot+3} are the total chilling accumulation in the reference
 638 year (Y_0) and the year with 3 °C warming (Y_{0+3}), and TA_0 and TA_{+3} are the
 639 corresponding critical forcing requirements for leaf. d_{c0} and d_{f0} are the start dates of
 640 chilling and forcing accumulation, respectively in the reference year (Y_0). d_0 , d_{0+3} and
 641 d_{0+6} are the leaf unfolding dates in the reference year Y_0 and the years with +3°C and
 642 +6°C warming if the chilling accumulation would have been identical to the reference
 643 year (CHA_{tot0}). d'_{0+3} is the leaf unfolding date in Y_{0+3} when considering the lower
 644 chilling accumulation (CHA_{tot+3}) induced by warming and responsible for a higher
 645 forcing requirements. d_{c1} and d_{c2} represents the start and end dates of the period when
 646 temperature in Y_{0+3} is more efficient for chilling accumulation (because closer to T_{op})
 647 than the temperature in Y_0 .
 648



649

650 **Figure 4 | Determining coefficients (R^2) of the linear regression function between**
 651 **the temperature sensitivity of LUD and different metrics of winter chilling and**
 652 **spring forcing.** CH_{rave} is the average daily chilling rate from the start date of chilling
 653 accumulation (i.e. d_{c0} in Fig. 3) to the start date of forcing accumulation (d_{f0}). CHA_{tot}
 654 is the total chilling accumulation in the whole pre-growing season. F_{rave} is the average
 655 daily forcing rate during the forcing stage. DF is the duration of forcing stage (day).
 656 $CH_{rave} + CHA_{tot}$ refers to the regression using both CH_{rave} and CHA_{tot} as independent
 657 variables, and $CH_{rave} + CHA_{tot} + F_{rave}$ refers to the regression using CH_{rave} , CHA_{tot} and
 658 F_{rave} together as independent variables. In each violin plot, balloon represents the
 659 probability density distribution of each gradient of R^2 . Whiskers indicate the
 660 interquartile (thick) and 95 % confidence intervals (thin). *AH* denotes *Aesculus*
 661 *hippocastanum*; *AG* denotes *Alnus glutinosa*; *BP* is *Betula pendula*; *FS* is *Fagus*
 662 *sylvatica*; *FE* is *Fraxinus excelsior*; *QR* is *Quercus robur*; *All* denotes the average
 663 values for all of the six species.

664

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