

# Model-based reproduction and validation of the total spectrum of solar flare and their impact on the global environment at the X9.3 event of September 6, 2017

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## Full paper

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# Abstract

We attempted to reproduce the total electron content (TEC) variation in the Earth's atmosphere from the temporal variation of the solar flare spectrum of the X9.3 flare on September 6, 2017. The flare spectrum from the Flare Irradiance Spectral Model (FISM), and the flare spectrum from the 1D hydrodynamic model, which considers the physics of plasma in the flare loop, are used in the GAIA model, which is a simulation model of the Earth's whole atmosphere and ionosphere, to calculate the TEC difference. We then compared these results with the observed TEC. When we used the FISM flare spectrum, the difference in TEC from the background was in a good agreement with the observation. However, when the flare spectrum of the 1D-hydrodynamic model was used, the result varied depending on the presence or absence of the background. This difference depending on the models is considered to represent which extreme ultraviolet (EUV) radiation is primarily responsible for increasing TEC. From the flare spectrum obtained from these models and the calculation result of TEC fluctuation using GAIA, it is considered that the enhancement in EUV emission by approximately 15–35 nm mainly contributes in increasing TEC rather than that of X-ray emission, which is thought to be mainly responsible for sudden ionospheric disturbance. In addition, from the altitude/wavelength distribution of the ionization rate of Earth's atmosphere by GAIA (Ground-to-topside atmosphere and ionosphere model for aeronomy), it was found that EUV radiation of approximately 15–35 nm affects a wide altitude range of 120–300 km, and TEC enhancement is mainly caused by the ionization of nitrogen molecules. (265 words)

## 1. Introduction

[2] The radiation from the sun is the most important ionization and heating energy source for the Earth's upper atmosphere. When solar flares occur, the intensity of multi-wavelength electromagnetic emissions suddenly increases (e.g. Kane, 1974). Among these, the intensification of X-rays (<10 nm) to extreme ultraviolet (EUV; 10–121 nm) emissions accelerates the ionization and molecular dissociation of atmospheric components in the ionosphere and thermosphere, and it may cause a sharp increase in electron density (e.g. Solomon and Qian, 2005). It is generally considered that ultraviolet (UV) emissions with a long wavelength of 100 nm or more affect the lower atmosphere, whereas short-wavelength EUVs and X-rays may also affect the mesosphere (e.g. Woods et al., 2000). Moreover, these short-wavelength UV radiation reaches the ionosphere D region at 60–90 km, thereby causing an increase in electron density in this region. A rapid increase in the total electron content (TEC) may degrade the accuracy of single-frequency global navigation satellite system (GNSS) positioning, which has been widely used in recent years. In an air navigation system, GNSS positioning is critical for navigation and safety. As augmentation systems, the satellite-based augmentation system, which covers a wide area by geostationary satellites, and ground-based augmentation system have been developed and are in operation. Temporal and spatial variations of TEC can be a threat to these systems (International Civil Aviation Organization, 2010).

[3] The communication failure caused by the absorption of the short-wave due to the variations in electron density is known as the Dellinger phenomenon (Dellinger 1937). Generally, the Dellinger

phenomenon is considered to be caused by the occurrence of M-class ( $\sim 10^{-5}$  W/m<sup>2</sup> at 1–8 Å band) or higher solar flares and predicted by using the magnitude of the solar flares. However, reportedly, the Dellinger phenomenon sometimes occurred even in during C-class ( $\sim 10^{-6}$  W/m<sup>2</sup> at 1–8 Å band) flares; however, it did not occur even in during X-class ( $\sim 10^{-4}$  W/m<sup>2</sup> at 1–8 Å band) flares (e.g. Tao et al., 2020). From these results, it is considered that the flare emissions contributing to the occurrence of the Dellinger phenomenon may also be affected by emissions that are not proportional to the X-ray intensity. It has also been reported that fluctuations of the total electron content (TEC) depends not only on the flare class but also the occurrence position of solar flare on the solar surface (Qian et al. 2010; Le et al. 2012). The X-ray intensity from solar corona does not change depending on the location of flare. On the other hand, EUV emissions from the lower layer of the solar atmosphere, such as the chromosphere, attenuates as the flare location goes from the center of the solar surface to the limb because the thickness of the passing solar atmosphere changes. From this, it is thought that EUV emission is the main source on TEC fluctuations. In previous studies, it is suggested that 0-450 Å EUV band (Strickland 1995), and/or 0-170 Å bands (Woods et al. 2011) is an important indicator for the EUV irradiance input to the Earth's upper atmosphere.

[4] In recent years, the understanding of flare triggers and their physical processes has been advanced, and physics-based flare prediction research is rapidly progressing (e.g. Kusano et al., 2020). By applying these physics-based flare prediction methods to the prediction model of space weather, we would like to make one prediction model of the flare emission effects on the space weather in the future. The combination of physics-based predictive research for flare emissions and physics-based predictive research for Earth's atmosphere is considered to be compatible because the parameters to be used are common in the physical process between the two models. On the other hand, the mainstream of research for predicting the impact to the space weather, is basically non-physical-based research. For example, the Flare Irradiance Spectral Model (FISM; Chamberlin et al., 2006, 2007, 2008) is most famous prediction model of solar flare X-ray and EUV emissions, but FISM was constructed empirically using observational X-ray (1–8 Å) data of the *Geostationary Operational Environmental Satellites (GOES)* (Bornmann, 1996), EUV data of the Thermosphere Ionosphere Mesosphere Energetics and Dynamics Solar EUV Experiment (TIMED/SEE) (Woods et al., 2005) and so on. Usually, the empirical model well reproduces the observation results, but it is impossible to investigate the cause when a special phenomenon occurs. The advantage of physics-based model research is that we can understand where the problem lies in the model. In this study, we attempted to predict and evaluate the effects from the solar flare on the Earth's atmosphere using physics-based model studies.

[5] First, to reproduce the observed flare emissions, we constructed a new flare emission model based on physical processes (Imada et al., 2015; Kawai et al., 2020). In our proposed model, the physical process of the plasma in the flare loop is reproduced by combining the 1D hydrodynamic calculation using Coordinated Astronomical Numerical Software (CANS; <http://www-space.eps.s.u-tokyo.ac.jp/~yokoyama/etc/cans/>) 1D package with the CHIANTI atomic database (Dere et al., 1997, 2019). Using this model, we reproduced EUV dynamic spectra for some flare events and compared them

with the observed EUV spectra (Nishimoto et al., 2021) using the Extreme Ultraviolet Variability (EVE: Woods et al., 2012) onboard the *Solar Dynamics Observatory (SDO)*. We performed our calculations with the estimated loop length from the observed flare ribbon distance (Nishimoto et al., 2020). When we compared the observed *SDO/EVE* Multiple EUV Grating Spectrograph (MEGS)-A (60–370 Å) spectra with our calculation results, we found that our results clearly reproduced most of the EUV lines during flare.

[6] Then, to examine the effects of the constructed flare irradiance on the Earth's ionosphere, we use an Earth's global whole atmosphere and ionosphere model, GAIA (Ground-to-topside atmosphere and ionosphere model for aeronomy; Jin et al., 2011). The output of the solar irradiance spectrum from some models is applied to GAIA as an input source for ionospheric photoionization and thermospheric heating, using an ionization cross-section model by Solomon and Qian (2005). The vertical TEC was calculated using the results from GAIA runs and compared with GNSS-TEC observations.

[7] In this study, we attempted to evaluate the flare emissions following a flare event and its effect on the Earth's atmosphere by a physics-based model study. We focused on the X9.3 flare of September 6, 2017, which is the largest solar flare in the solar cycle 24. In section 2, we introduce observational data and the models for the flare spectrum and TEC used in this study. Section 3 presents the results of the GAIA runs, and Section 4 discusses and summarizes the results.

## 2. Observational Data And Models

[8] Two X-class flares occurred on AR102673 on September 6, 2017. Among them, we focused on the X9.3 flare that occurred at 11:53 UT. In this section, we present an overview of the observed solar flare emission spectra by *SDO/EVE* and the model calculation for its reproduction. Then we also present the observation and model calculation of the TEC fluctuation associated with the X9.3 flare.

### 2.1 Flare emission spectra and its prediction model

[9] The flare emissions from the X9.3 flare has been observed in X-rays (0.5–4 Å and 1–8 Å) by the *GOES*, and *SDO/EVE* MEGS-B observed the EUV emissions (360–1060 Å). However, emissions within the range of 60–370 Å, which have a large effect on the Earth's atmosphere, has not been observed due to a failure of the *SDO/EVE* MEGS-A. Therefore, we estimated the flare emission spectrum using the method of Kawai et al. (2020).

[10] In our model (Kawai et al., 2020), we assumed a semicircular flare loop with a diameter that is the distance between two flare ribbons observed in this flare, for simulated the plasma motion in the flare loop with the 1D-hydrodynamic model. Main input parameters for CANS 1D are the half loop length and heating rate. For the X9.3 flare, the distance between the two ribbons can be estimated as 32.5 Mm from *SDO/AIA* 1600 Å data (<http://jsoc.stanford.edu/ajax/lookdata.html>), and then the loop length can be obtained as 51 Mm. The heating rate is  $120 \text{ erg cm}^{-3} \text{ s}^{-1}$ , which is estimated by comparing observed and estimated two channel of *GOES* X-ray peak intensities. The input heating duration and width along the loop of the flare are fixed at 60 s and 6000 km, respectively. These parameters are same as those used

for the X9.3 flare described in Kawai et al. (2020). Then, we obtained the EUV emissions strength and its time variation using CHIANTI. We calculated only for the X9.3 class flare in our model, and the result is shown in Figure 1 depicted by the blue line during 11:53-12:24 UT. Hereafter, this model will be called the “1D Hydro flare model” in this paper.

[11] Another prediction model for flare emissions is the FISM (Chamberlin et al., 2006, 2007, 2008). The FISM is an empirical model, that provides a spectrum of solar flare at a wavelength range of 1-1900 Å every minute with a wavelength resolution of 10 Å. The calculation results for the FISM are available on the Web site (<https://lasp.colorado.edu/lisird/>). Recently (October 2020), FISM2 with improved wavelength resolution (0.1–1900 Å at 1 Å spectral bins) was released (Chamberlin et al., 2020), however this research is verifying using FISM. FISM assumes that flare emission time evolution is proportional to *GOES* soft X-ray, and thus, the FISM time evolution X-ray profile of 1–10 Å X-ray, shown in Figure 1, is similar to that of the *GOES* X-ray profile of 1–8 Å. This is also the case for EUV line emissions, FISM has not been able to adequately reproduce the observed delay in EUV line emissions (Nishimoto et al., 2020, 2021).

[12] Figure 2 shows the flare spectrum at 12 UT on September 6, 2017, for each emission model. The red broken line in Figure 2 is the daily component of FISM, and this spectrum can be considered as background emissions. The red line is the FISM flare spectrum, and the green line is the 1D Hydro flare model spectrum. As the 1D Hydro flare model does not consider the continuum component, it can be observed that there is no continuum component at approximately 75-95 nm. The blue line is the sum of the 1D Hydro flare model spectrum and the FISM background spectrum.

[13] The flare spectra variations, which comprise of the Figure 2 spectra and Figure 1 light curves, were input to the GAIA model, which is described in Section 2.2. There are four types of input models: (1) FISM background (“bgd”), (2) FISM flare (“FISM”), (3) 1D Hydro flare or background (“1D Hydro/bgd”), and (4) 1D Hydro flare and background (“1D Hydro+bgd”).

## 2.2 TEC observation and GAIA

[14] The TEC in the Earth’s ionosphere can be obtained from the radio propagation delay between GNSS satellites and receivers (Otsuka et al., 2002; Tsugawa et al., 2018; Shinbori et al., 2020). Figure 3 shows the global distribution of verticalized TEC processed, and provided by National Institute of Information and Communications Technology (NICT) and Nagoya University. This is the TEC distribution observed at 12 UT on September 6, 2017, when the X9.3 flare occurred. As the flare occurred around 12 UT, we can observe the variations in TEC distribution due to the flare emissions using the TEC data around 0°E longitude. In the European region, sufficient data was obtained around 0°E longitude, and therefore, we used 0°E, 25–75°N data for comparative studies.

[15] Figure 4 shows the time variation of the TEC distribution at 0°E and 25–75°N. The upper left part of Figure 4 shows the average TEC during the time period from September 1 to 10, 2017, which can be used as the background TEC. The upper right of Figure 4 shows the TEC variation observed with the X9.3 flare,

and evidently, the TEC sharply increases at any observed latitude at approximately 12UT. The background data in the upper left panel in Figure 4 show that the amount of TEC variation increases in the afternoon, whereas the flare effect on TEC is significantly stronger than the diurnal change and is approximately 1.3 times stronger at the flare peak (Figure 4 bottom).

[16] Although the actual variations in TEC due to flare emissions are displayed in Figure 4, it is not evident which wavelength of the flare emissions causes the TEC variation. Therefore, to decipher, this we used the GAIA model which can calculate the ionospheric variation by inputting the flare emissions (discussed in Section 2.1).

[17] Flare emission ionization rates in the upper atmospheres depend on the ionization cross-sections and neutral species abundance. Figure 5 shows the cross-sections of photoionization and photo-dissociative ionization of oxygen atoms, oxygen molecules, and nitrogen molecules, which are the main constituents in the Earth's thermosphere. These cross-sections are based on the model of Solomon and Qian (2005). Notably, the X-ray cross-sections are two orders or more smaller than the cross sections of >15 nm EUVs. In addition, evidently, the cross-sections of >60 nm EUVs are considerably smaller depending on the ionization reaction.

[18] Figure 6 shows the change in ionization rate at each altitude as a result of inputting the flare spectrum of the FISM model (Figure 2) into GAIA. From this result, the wavelength ionized at each altitude in the upper atmosphere can be identified. X-rays mainly ionize oxygen and nitrogen atoms at 100–150 km; approximately 10 nm EUVs mainly ionize nitrogen atoms at similar altitude, and ~30 nm EUVs mainly ionize nitrogen molecules at a slightly higher altitude. Approximately 100 nm EUVs also ionizes oxygen molecules at an altitude of ~100 km.

## 3. Results

### 3.1 TEC obtained from GAIA simulations

[19] As shown in Figure 7, the four types of solar flare spectra presented in Section 2.1 were input to the GAIA model, and the TEC time variation was obtained. Run1 inputs FISM background spectrum, which is referred to as background; Run1, 2, 3 and 4 input the (1) "bgd", (2) "FISM flare", (3) "1D Hydro/bg d", and (4) "1D Hydro+bg d" spectral models, respectively. Evidently, TEC significantly increased in Run2, 3, and 4 during the X9.3 flare.

[20] Figure 8 shows the calculation results ratio between Run1 and Run2, 3, and 4. The increase in TEC due to flare emissions can be confirmed in all model calculations, and the increase rate was the largest in Run4 whereas the smallest in Run3. As the FISM flare model was input to Run2, a second TEC increase was observed after 12:30 UT.

### 3.2 Comparison with observed TEC data

[21] The time variation of TEC obtained from GAIA (Section 3.1) was compared with the observed TEC, which is shown in Figure 4. Figure 9 shows the TEC time variation at 0°E and 44°N. These variations do not change much even if the latitude and longitude are changed slightly, and the increase in TEC decreases as the latitude increases due to the difference in zenith angle. In the observed data, there is another prominent TEC increase from 12:30 UT after the peak at 12 UT. This is the effect of another flare event after 12:30 UT, which is seen in the X-ray light curve in Figure 1 and was captured by the Run2 GAIA calculation. In the case of the 1D Hydro flare model input (Run3 and 4), the background emission model was input after 12:30 UT, and therefore, the second increase is not observed. In the GAIA model calculation, a TEC increase was also observed at 15–20 UT; however, this was not noticed in the observation data and therefore an artifact of the GAIA model.

[22] As the absolute TEC value calculated by GAIA does not match the observed value, we decided to compare each increment from the TEC background. The difference in the observed TEC from the background, and the difference in calculated TEC by GAIA from the FISM background is shown in Figure 10. The error of the observed TEC value is about 0.01 TECU. The absolute value of the observation is about 1 to 3 TECU (Otsuka et al., 2002), but these effects can be eliminated by subtracting background value. The FISM model accurately reproduces the observed TEC, although it is slightly larger than the observed TEC due to the effects of previous X-class flare. The “1D Hydro/bgd” model and the “1D Hydro+bgd” model was 0.8 times and 2.4 times larger than the observed TEC, respectively.

## 4. Discussion And Summary

[23] We attempted to calculate TEC variations by inputting the temporal variation of the solar flare spectrum for the X9.3 flare into the GAIA model and compared it with the observed TEC. In particular, we analyzed the increments of TEC from background TEC at 0°E and 44°N, and the FISM model almost reproduced the observed values; the “1D Hydro/bgd” and “1D Hydro+bgd” models were 0.8 and 2.4 times larger than the observed value, respectively. Since the amount of TEC fluctuation depends on the zenith angle, it increases as the latitude decreases and decreases as the latitude increases, however, there was no latitude-dependent variation ratio of the difference between the flare spectrum models. Therefore, this discrepancy is considered to represent the intensity ratio of the EUV radiation that affects the TEC increment. By comparing this with the solar flare spectra in Figure 2, we can infer the primary wavelength causing TEC enhancement.

[24] From Figure 2, X-ray is the flare emission with the highest intensity, which has an intensity of one order or more compared to EUV emissions. However, the fluctuation in TEC is also caused by a combination of the emission spectrum, the photoionization cross-section (Figure 5), and the distribution of atmospheric densities. Furthermore, as shown in Figure 6(e), atmosphere ionization at approximately 100–150 km is primarily attributable to X-rays. However, from Figure 1, the background intensity of X-rays is two orders of magnitude smaller than the flare peak intensity of X-rays, indicating there is no significant difference between the “1D Hydro/bgd” and “1D Hydro+bgd” models. Figure 2 also shows that the background spectrum below 15 nm is smaller than the flare spectrum by one order or more, thereby

indicating that there is no difference in the X-ray intensity between the “1D Hydro/bgd” and “1D Hydro+bgd” models. In fact, the total emission energy of <10 nm was 1.3 times that of FISM in both models, with no difference between the presence and absence of background. This relationship implies that X-ray emissions are not the main wavelength causing TEC enhancement.

[25] However, large spectral differences depending on the type of flare emission model were observed for long wavelengths of >75 nm. In particular, because the 1D Hydro flare model does not consider the continuum component, the EUV flux of the “1D Hydro/bgd” model is two orders of magnitude smaller than that of the FISM model. In this wavelength band, the intensity of the “1D Hydro+bgd” model is approximately 1/3 times that of the FISM model. However, the TEC in the “1D Hydro+bgd” model at the flare peak at 0°E and 44°N is >2 times that of the FISM model, and therefore, the intensity ratio is not consistent with that of the EUV emissions. This indicates that long wavelength EUV radiation (>75 nm) does not affect the TEC.

[26] The wavelengths that become stronger in the order of “1D Hydro/bgd”, “FISM”, and “1D Hydro+bgd” are approximately 15–55 nm emission. Among this wavelength range, the EUV intensity at >40 nm is about an order of magnitude smaller than <40 nm (Figure 2), and at >35 nm, cross-section of photodissociation of nitrogen molecule is an order of magnitude smaller than that at <35 nm (Figure 5), there is not enough ion production rate to fluctuate the TEC (Figure 6(e)). Therefore, 15–35 nm is a candidate for radiation that is thought to be primarily effective in TEC fluctuation. Especially, at 15–20 nm wavelength band, the total emission energy of the “1D Hydro/bgd” and “1D Hydro+bgd” models were 0.8 and 2.1 times larger than the FISM, respectively, which is similar to the increase in each TEC. The EUV intensity in this wavelength band is  $\geq$  one order of magnitude smaller than that of X-rays, whereas approximately one order of magnitude larger than that of long-wavelength EUVs, which indicates that this EUV wavelength mainly affects the TEC of the Earth’s ionosphere. In addition, EUVs of ~30 nm wavelength band have a high ionization rate at a wide altitude of 120–300 km. Ionization is neutralized immediately at lower altitudes; however, 120–300 km is a relatively high altitude, and thus, it is likely that ionization is maintained for a considerable amount of time. Therefore, it is considered that EUVs of ~15–35 nm are the main cause of TEC increase.

[27] When reproducing the flare spectrum with our model’s 1 Å wavelength resolution, we can identify the dominant EUV line emission in the ~15–35 nm wavelength range. The strongest line emission in this wavelength range is the 304 Å He II line. In addition, lines such as Fe XII at 195 Å, He II at 256 Å, Fe XV at 284 Å, and Fe XVI at 335 Å are also stronger lines. Among these lines, the Fe lines come from several mega Kelvin (MK) plasmas whereas He II is mainly from the chromosphere. This indicates that the flare emission model can also reproduce the TEC variations correctly if it can reproduce the flare spectrum of several MK plasmas and emissions from the chromosphere.

[28] We discussed the TEC variation due to the X9.3 flare that occurred on September 6, 2017. As this analysis was limited to one case and quantified the specific feature of only one flare, we will conduct

statistical analysis of multiple flares to obtain global relationship between flare emission spectrum and TEC in future.

## List Of Abbreviations

CANS: Coordinated Astronomical Numerical Software

EUV: extreme ultraviolet

EVE: Extreme Ultraviolet Variability

FISM: Flare Irradiance Spectral Model

GAIA: Ground-to-topside atmosphere and ionosphere model for aeronomy

GBAS: Ground-based Augmentation System

GNSS: Global Navigation Satellite System

GOES: Geostationary Operational Environmental Satellite

MEGS: Multiple EUV Grating Spectrograph

NICT: National Institute of Information and Communications Technology

SBAS: Satellite-based Augmentation System

SDO: Solar Dynamics Observatory

SID: sudden ionospheric disturbance

TEC: total electron content

UT: universal time

UV: ultraviolet

## Declarations

### **Ethics approval and consent to participate**

No applicable

### **Consent for publication**

No applicable

## Availability of data and materials

We use the GOES X-ray data provided by NOAA National Centers for Environmental Information <https://www.ngdc.noaa.gov/stp/space-weather/solar-data/solar-features/solar-flares/x-rays/goes/xrs/>. The FISM results are available at <https://lasp.colorado.edu/lisird/> and observational EUV data of SDO/EVE also available at [http://lasp.colorado.edu/eve/data\\_access/eve-flare-catalog/index.html](http://lasp.colorado.edu/eve/data_access/eve-flare-catalog/index.html). The GNSS data collection and processing were performed using the NICT Science Cloud. The Receiver Independent Exchange (RINEX) format data used for GNSS-TEC processing were provided by the following organizations: UNAVCO (<ftp://data-out.unavco.org>), DGRSDUT (<http://gnss1.tudelft.nl>), Arecibo Observatory (<http://www.naic.edu>), CDDIS (<ftp://cddis.gsfc.nasa.gov>), CHAIN (<ftp://chain.physics.unb.ca>), CORS (<ftp://www.ngs.noaa.gov>), GDAF (<ftp://geodaf.mt.asi.it>), BKG (<ftp://igs.bkg.bund.de>), IGS (<ftp://rgpdata.ign.fr>), EUOLG (<ftp://olggps.oeaw.ac.at>), Geoscience Australia (<ftp://ftp.ga.gov.au>), IGSIGN (<ftp://igs.ensg.ign.fr>), KASI (<ftp://nfs.kasi.re.kr>), PNGA (<http://www.geodesy.cwu.edu>), IBGE (<ftp://geofp.ibge.gov.br>), RGCL (<ftp://ftp.itacyl.es>), TNG (<ftp://196.15.132.3>), SOPAC (<ftp://garner.ucsd.edu>), NRC (<ftp://wcda.pgc.nrcan.gc.ca>), GEONET (<ftp://163.42.5.1>), HRAO (<ftp://geoid.hartrao.ac.za>), GRN (<ftp://rinex.smartnetna.com>), GNNZ (<ftp://ftp.geonet.org.nz>), RENAG (<ftp://renag.unice.fr>), SONEI (<ftp://ftp.sonel.org>), FRDN (<ftp://www.crs.inogs.it>), LINZ (<ftp://apps.linz.govt.nz>), ROB (<ftp://gnss.oma.be>), GOP (<ftp://ftp.pecny.cz>), RGE (<ftp://62.99.86.141>), RGNA (<ftp://geodesia.inegi.org.mx>), CENAT (<ftp://ramsac.ign.gob.ar>), INGV (<ftp://bancadati2.gm.ingv.it>), REP (<ftp://158.49.61.10>), SWSBM (<ftp://ftp-out.sws.bom.gov.au>), CORS (<ftp://meristemum.carm.es>), AFREF (<ftp://ftp.afrefdata.org>), WHU (<ftp://igs.gnsswhu.cn>), TLALOCNET (<ftp://tla-locnet.udg.mx>), NCEDC (<ftp://www.ncedc.org>), ODT (<ftp://ftp.odot.state.or.us>), SWEPOS (<ftp://ftp-sweposdata.lm.se>), EUREF (<ftp://www.epncb.oma.be>), IGG (<ftp://ftp.glonass-iac.ru>), SUGAUR (<ftp://eos.ntu.edu.sg>), NMA (<ftp://ftp.statkart.no>), NERC (<ftp://128.243.138.204>), RAMSAC (<ftp://ramsac.ign.gob.ar>), and ERGNSS (<ftp://ftp.geodesia.ign.es>).

## Competing interests

The authors declare that they have no competing interests.

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## Authors' contributions

KW conducted the research and has responsibility for the results presented in this paper. HJ has responsibility for the GAIA calibration and supported this analysis. SN, SI, TK and TK contributed to providing the physics-based flare spectra, and discussion as experts of solar physics. YO, AS, TT and MN contributed to providing TEC data. All authors read and approved the final manuscript.

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## References

- Bornmann P. L., D. Speich, J. Hirman et al. (1996), GOES X-ray sensor and its use in predicting solar-terrestrial disturbances, *Proceedings of SPIE*, 2812, 291-298.
- Chamberlin P. C., T. N. Woods, and F. G. Eparvier (2006), Flare Irradiance Spectral Model (FISM) use for space weather applications, *Proceedings of the ILWS Workshop*, 153.
- Chamberlin P. C., T. N. Woods, and F. G. Eparvier (2007), Flare Irradiance Spectral Model (FISM): Daily component algorithms and results, *Space Weather*, 5, 7, S07005.
- Chamberlin P. C., T. N. Woods, and F. G. Eparvier (2008), Flare Irradiance Spectral Model (FISM): Flare component algorithms and results, *Space Weather*, 6, 5, S05001.
- Chamberlin P. C., T. N. Woods, L. Didkovsky et al. (2018), Solar Ultraviolet Irradiance Observations of the Solar Flares During the Intense September 2017 Storm Period, *Space Weather*, 16, 10, 1470-1487.
- Chamberlin, P. C., F. G. Eparvier, V. Knoer et al. (2020), The flare irradiance spectral model-version 2 (FISM2), *Space Weather*, 18, e2020SW002588.

- Dellinger J. H. (1937), Sudden disturbances of the ionosphere, *Journal of Applied Physics*, 8, 11, 732-751.
- Dere K. P., E. Landi, H. E. Mason, B. C. Monsignori, and P. R. Young (1997), CHIANTI - an atomic database for emission lines, *Astronomy & Astrophysics Supplement*, 125, 149-173.
- Dere K. P., G. Del Zanna, P. R. Young, E. Landi, and R. S. Sutherland (2019), CHIANTI - An Atomic Database for Emission Lines. XV. Version 9, Improvements for the X-Ray Satellite Lines, *Astrophysical Journal Supplement*, 241, 22.
- Imada S., I. Murakami, and T. Watanabe (2015), Observation and numerical modeling of chromospheric evaporation during the impulsive phase of a solar flare, *Physics of Plasmas*, 22, 101206.
- International Civil Aviation Organization (2010), *International Standard and Recommended Practices, Annex 10 to the Convention on International Civil Aviation, Volume I, Radio Navigation Aids*.
- Jin H., Y. Miyoshi, H. Fujiwara et al. (2011), Vertical connection from the tropospheric activities to the ionospheric longitudinal structure simulated by a new Earth's whole atmosphere-ionosphere coupled model, *J. Geophys. Res.*, 116, A01316.
- Kane S. R. (1974), Impulsive (flash) Phase of Solar Flares: Hard X-Ray, Microwave, EUV and Optical Observations, *Coronal Disturbances: Proc. 57th IAU Symp.*, 105.
- Kawai T., S. Imada, S. Nishimoto, K. Watanabe, and T. Kawate (2020), Nowcast of an EUV dynamic spectrum during solar flares, *Journal of Atmospheric and Solar-Terrestrial Physics*, 205:105302.
- Kusano K., Iju, Y. Bamba, S. Inoue (2020), A physics-based method that can predict imminent large solar flares, *Science*, 369, 6503, 587-591.
- Le H., L. Liu, W. Wan (2012), An analysis of thermospheric density response to solar flares during 2001–2006, *J. Geophys. Res.*, 117, A03307.
- Nishimoto S., K. Watanabe, S. Imada et al. (2020a), Statistical and observational research on solar flare EUV spectra and geometrical features, *Astrophysical Journal*, 904, 1, 31.
- Nishimoto S., K. Watanabe, S. Imada et al. (2020b), Validation of EUV Emission Spectra During Solar Flares, *Earth, Planets and Space*, submitted to this issue.
- Otsuka, Y., T. Ogawa, A. Saito, T. Tsugawa, S. Fukao, S. Miyazaki (2002), A new technique for mapping of total electron content using GPS network in Japan, *Earth, Planets and Space*, 54(1), 63–70.
- Qian, L., A. G. Burns, P. C. Chamberlin, and S. C. Solomon (2010), Flare location on the solar disk: Modeling the thermosphere and ionosphere response, *J. Geophys. Res.*, 115, A09311.
- Qian L., A. G. Burns., P. C. Chamberlin., S. C. Solomon (2011), Variability of thermosphere and ionosphere responses to solar flares, *Journal of Geophysical Research: Space Physics*, 116, A10, A10309.

Shinbori, A., Y. Otsuka, T. Sori, T. Tsugawa, M. Nishioka (2020), Temporal and spatial variations of total electron content enhancements during a geomagnetic storm on 27 and 28 September 2017, *Journal of Geophysical Research: Space Physics*, 125, e2019JA026873.

Solomon, S. C., and L. Qian (2005), Solar extreme-ultraviolet irradiance for general circulation models, *Journal of Geophysical Research*, 110, A10306.

Strickland, D. J., Evans, F. S., Paxton, L. J. (1995), Satellite remote sensing of thermospheric O/N<sub>2</sub> and solar EUV: 1. Theory, *Journal of Geophysical Research*, 100:12217-12226.

Tao C., M. Nishioka, S. Saito et al. (2020), Statistical analysis of short-wave fadeout for extreme space weather event estimation, *Earth, Planets and Space*, 72, 1, 173.

Tsugawa, T., M. Nishioka, M. Ishii, K. Hozumi, S. Saito, A. Shinbori, et al. (2018), Total electron content observations by dense regional and worldwide international networks of GNSS, *Journal of Disaster Research*, 13(3), 535–545.

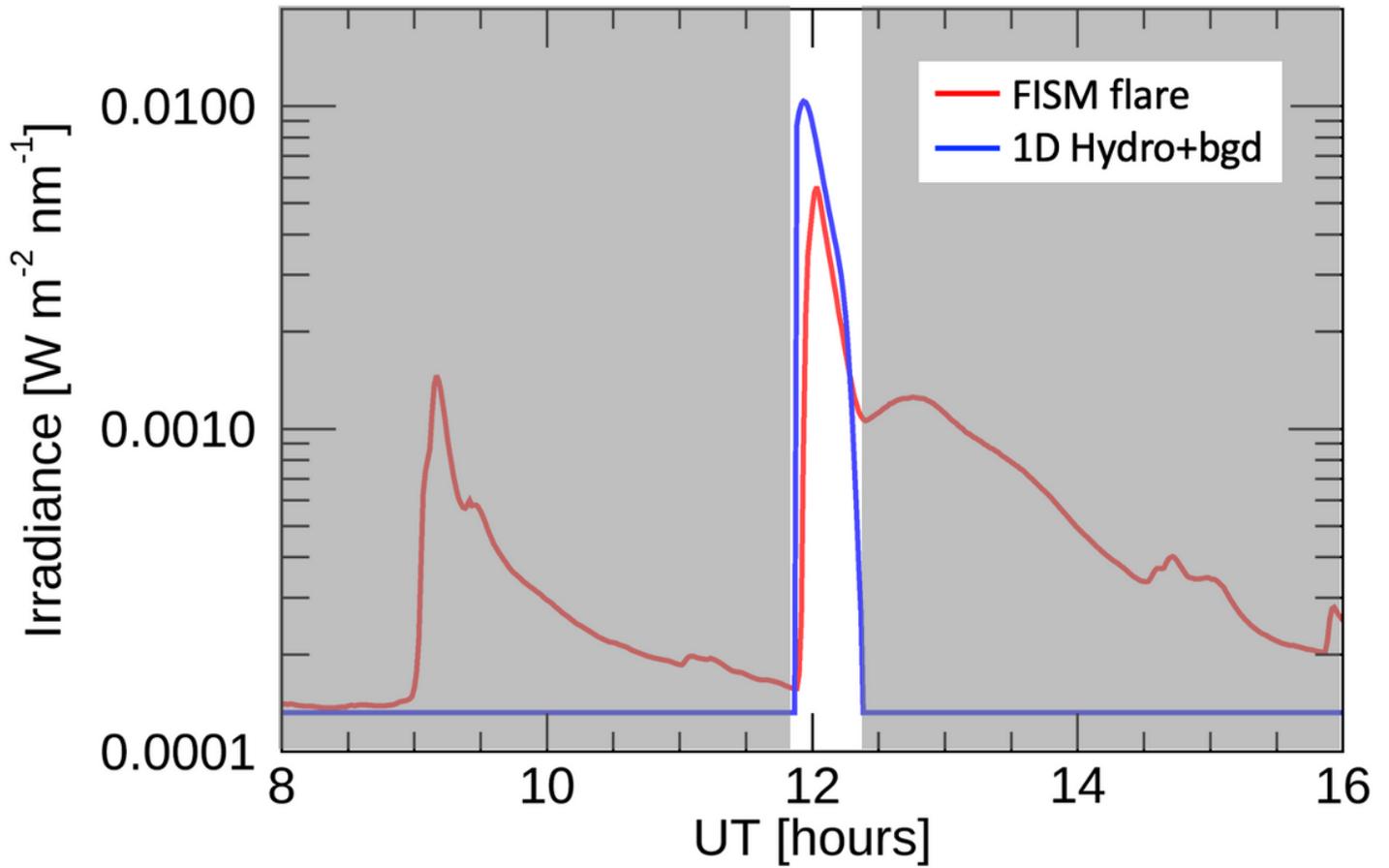
Woods T. N., G. Rottman, J. Harder et al. (2000), Overview of the EOS SORCE Mission, *Proceedings of SPIE - The International Society for Optical Engineering*, 4135-21.

Woods T. N., F. G. Eparvier, S. M. Bailey et al. (2005), Solar EUV Experiment (SEE): Mission overview and first results, *Journal of Geophysical Research: Space Physics*, 110, A1, A01312.

Woods T. N., R. Hock, F. Eparvier et al. (2011), New Solar Extreme-ultraviolet Irradiance Observations during Flares, *Astrophysical Journal*, 739, 2, 59.

Woods T. N., F. G. Eparvier, R. Hock et al. (2012), Extreme Ultraviolet Variability Experiment (EVE) on the Solar Dynamics Observatory (SDO): Overview of Science Objectives, Instrument Design, Data Products, and Model Developments, *Solar Physics*, 275, 115-143.

## Figures



**Figure 1**

Light curve of soft X-ray (1–10 Å) emission “FISM flare” (red line) is proportional to the light curve of GOES soft X-ray (1–8 Å). “bgd” is the daily component of FISM which is a constant value. “1D Hydro flare” reproduces only the light curve of the X9.3 flare during 11:53–12:24 UT. “1D Hydro+bgd” (blue line) is the sum of the light curves of “bgd” and “1D Hydro flare”. Gray regions have no 1D Hydro calculation data.

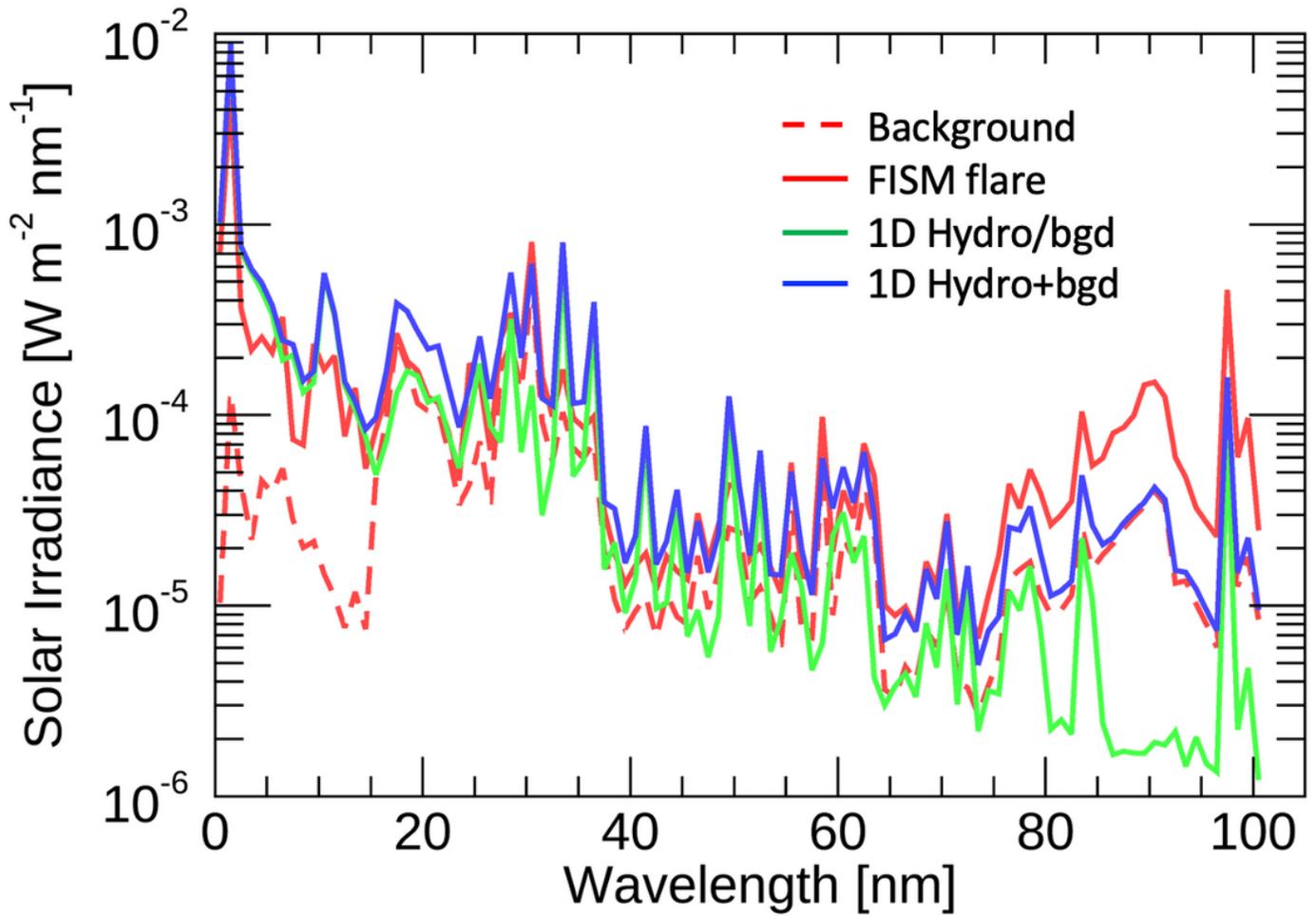
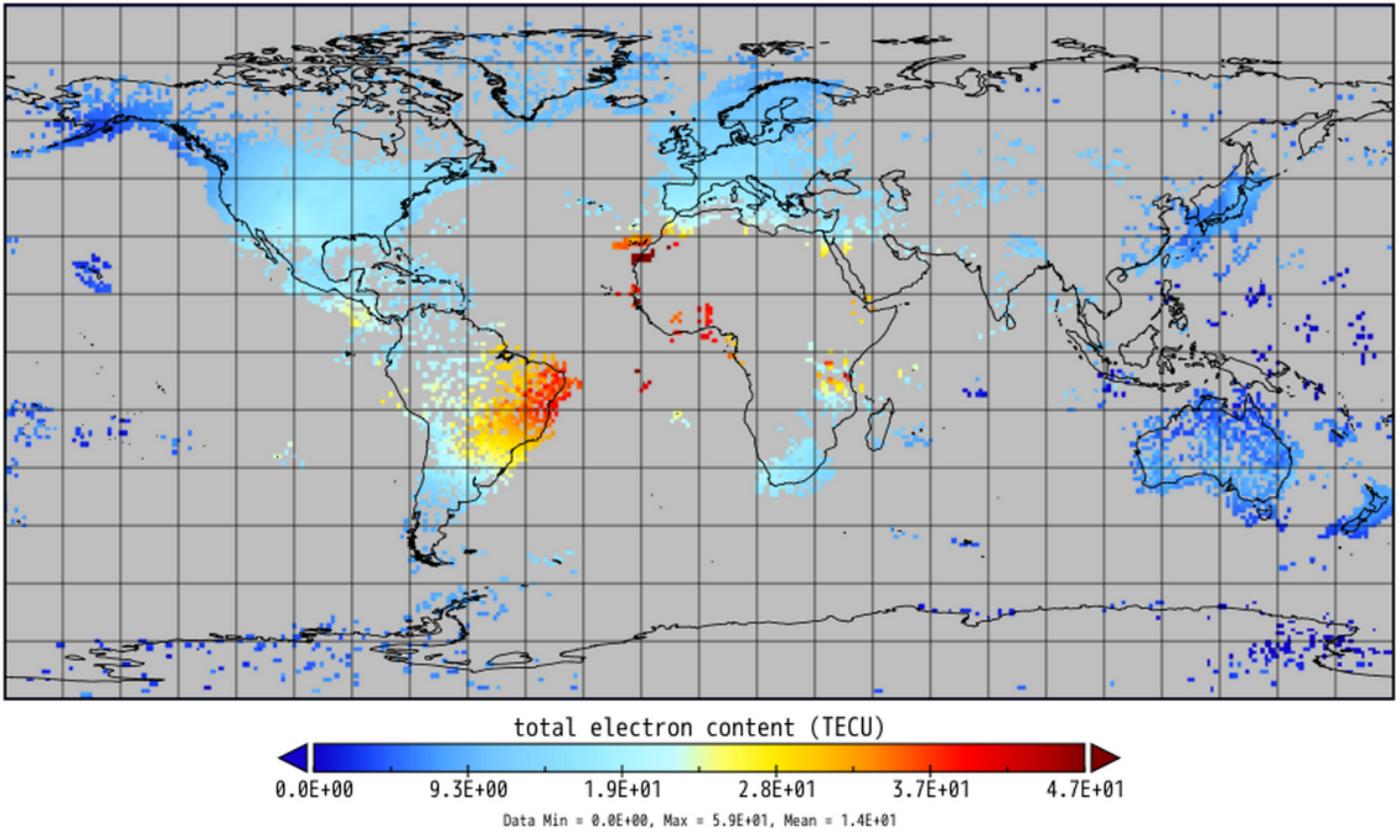


Figure 2

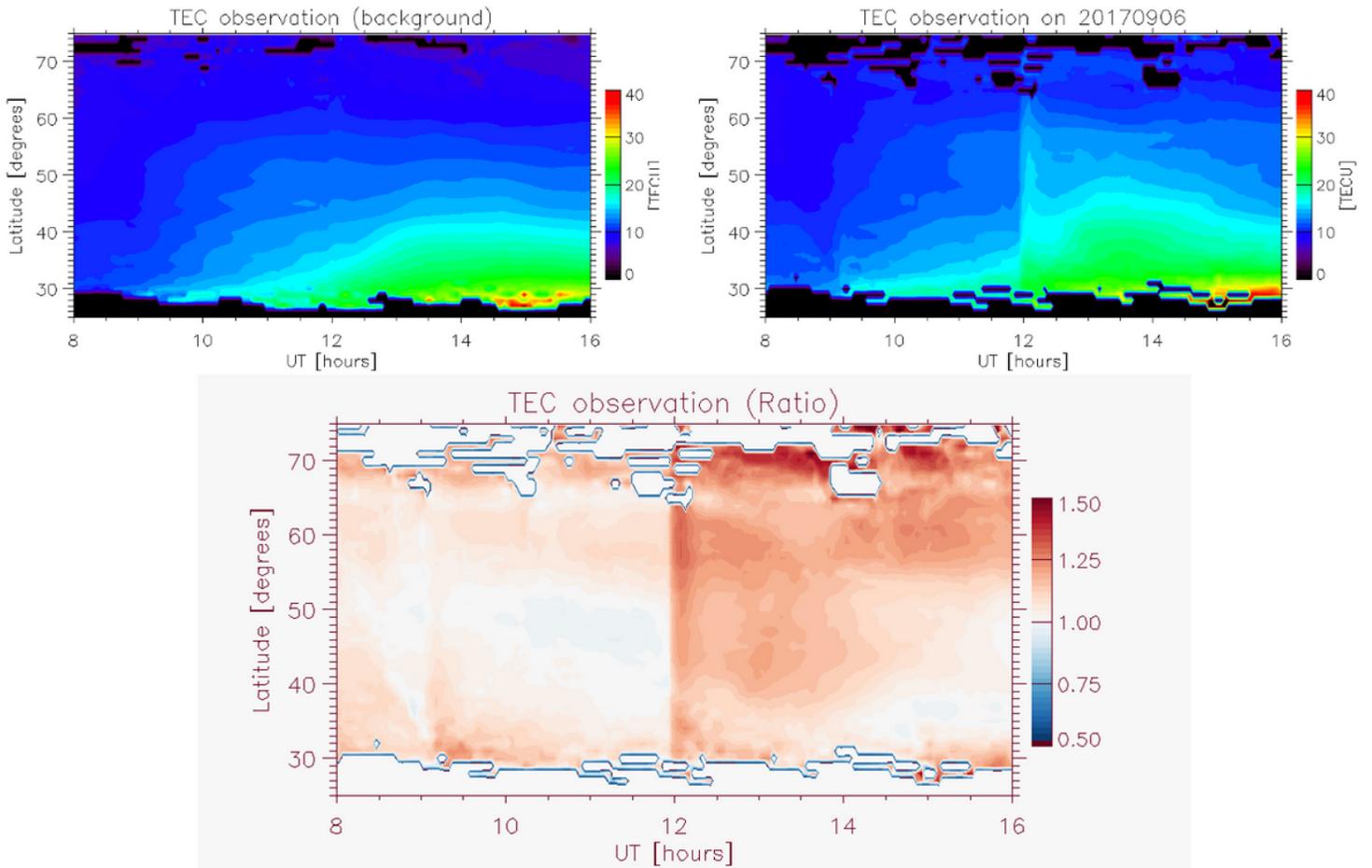
Emission spectra of X9.3 flare at 12 UT. The red solid line shows the FISM spectrum and the green solid line shows the 1D Hydro flare model spectrum. The blue solid line is the sum of the 1D Hydro flare model spectrum (green solid line) and the background (red dashed line). Because the 1D Hydro flare model does not consider the continuum component, there is no continuum component around 70-90 nm.

total electron content



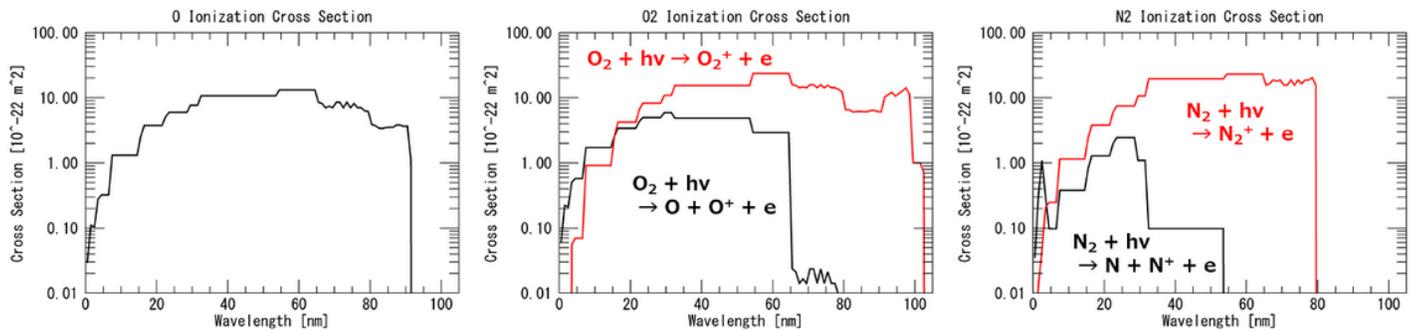
**Figure 3**

Distribution of TEC observed by global TEC (Drawing-TEC) at X9.3 class flare (12 UT) Distribution of TEC observed by global TEC (Drawing-TEC) at X9.3 class flare (12 UT). The European region was in the daytime at the time of the flare, and has sufficient data. Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its authorities, or concerning the delimitation of its frontiers or boundaries. This map has been provided by the authors.



**Figure 4**

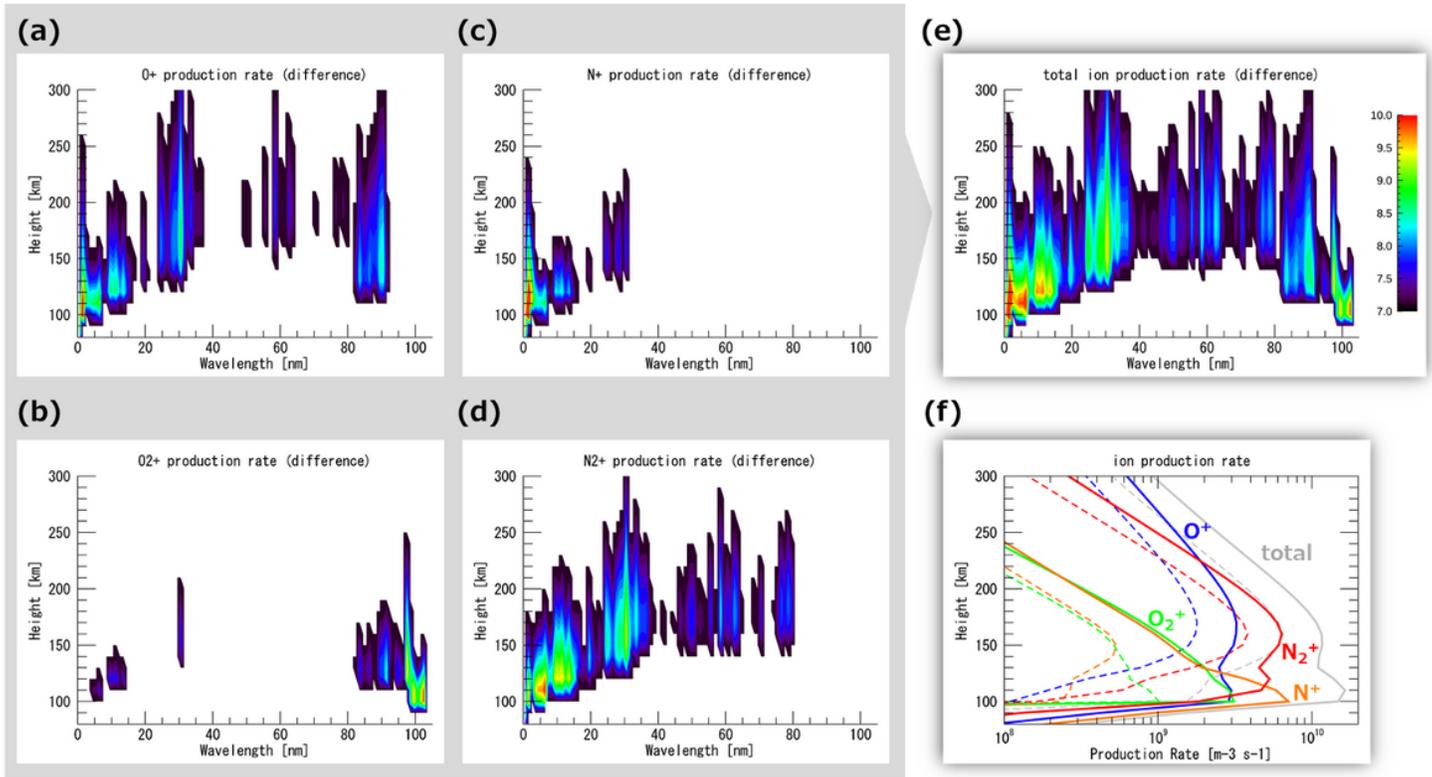
TEC observations at 0°E 25–75°N (Top left) Averaged TEC from September 1 to 10, 2017. This is the background. (Top right) TEC on September 6, 2017. (Bottom) Ratio of TEC from flare (top-right figure) and the background (top-left figure). TEC significantly increases at any latitude when flare occurs.



**Figure 5**

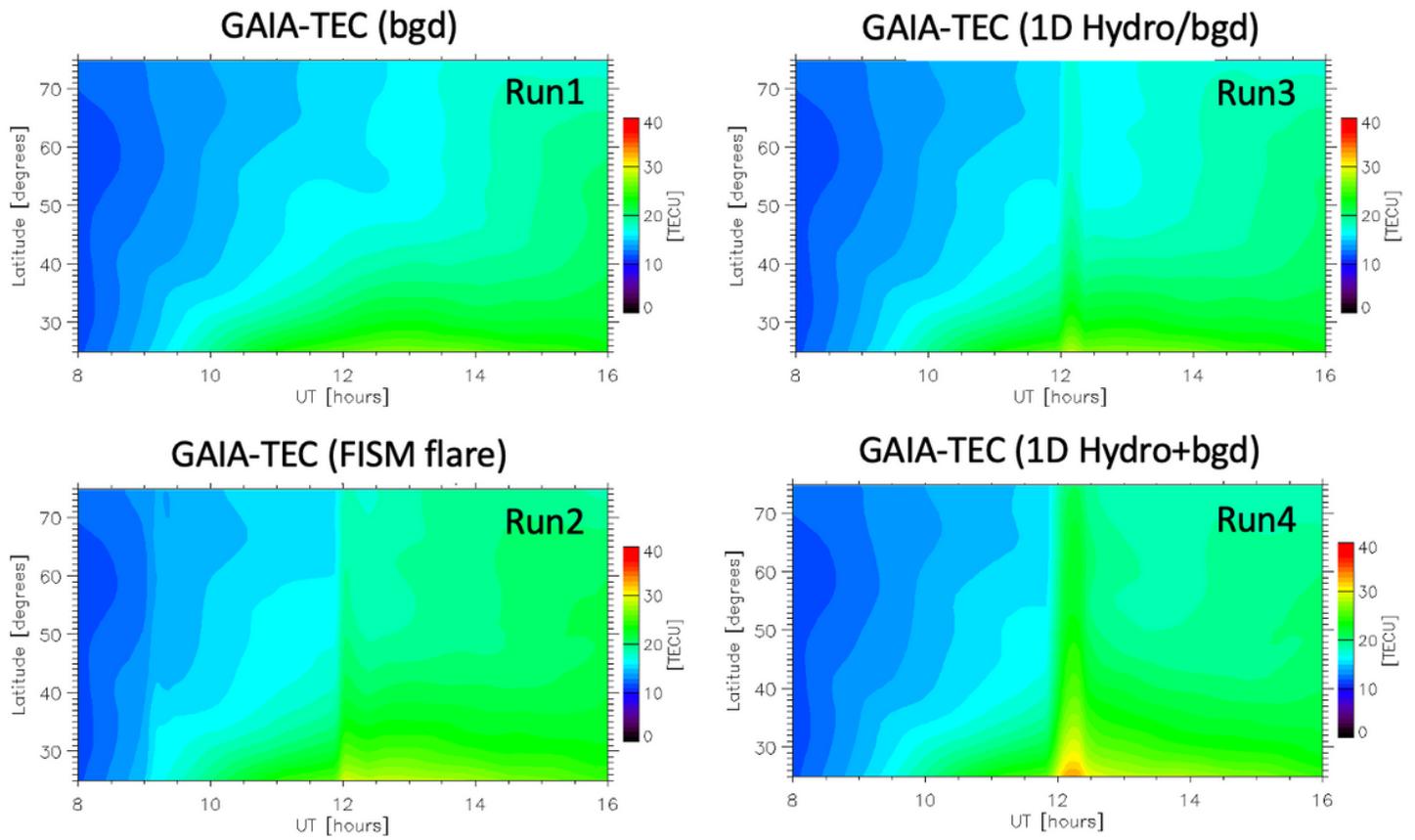
Photoionization and photo-dissociative ionization cross sections of major solar atmosphere atoms and molecules Photoionization and photo-dissociative ionization cross-sections for each wavelength. (Left) Photoionization cross-section of oxygen atom, (middle) cross-section of photoionization and photo-

dissociative ionization of oxygen molecule, and (right) photoionization and photodissociation of nitrogen molecule. Cross-section data are based on Solomon and Qian (2005).



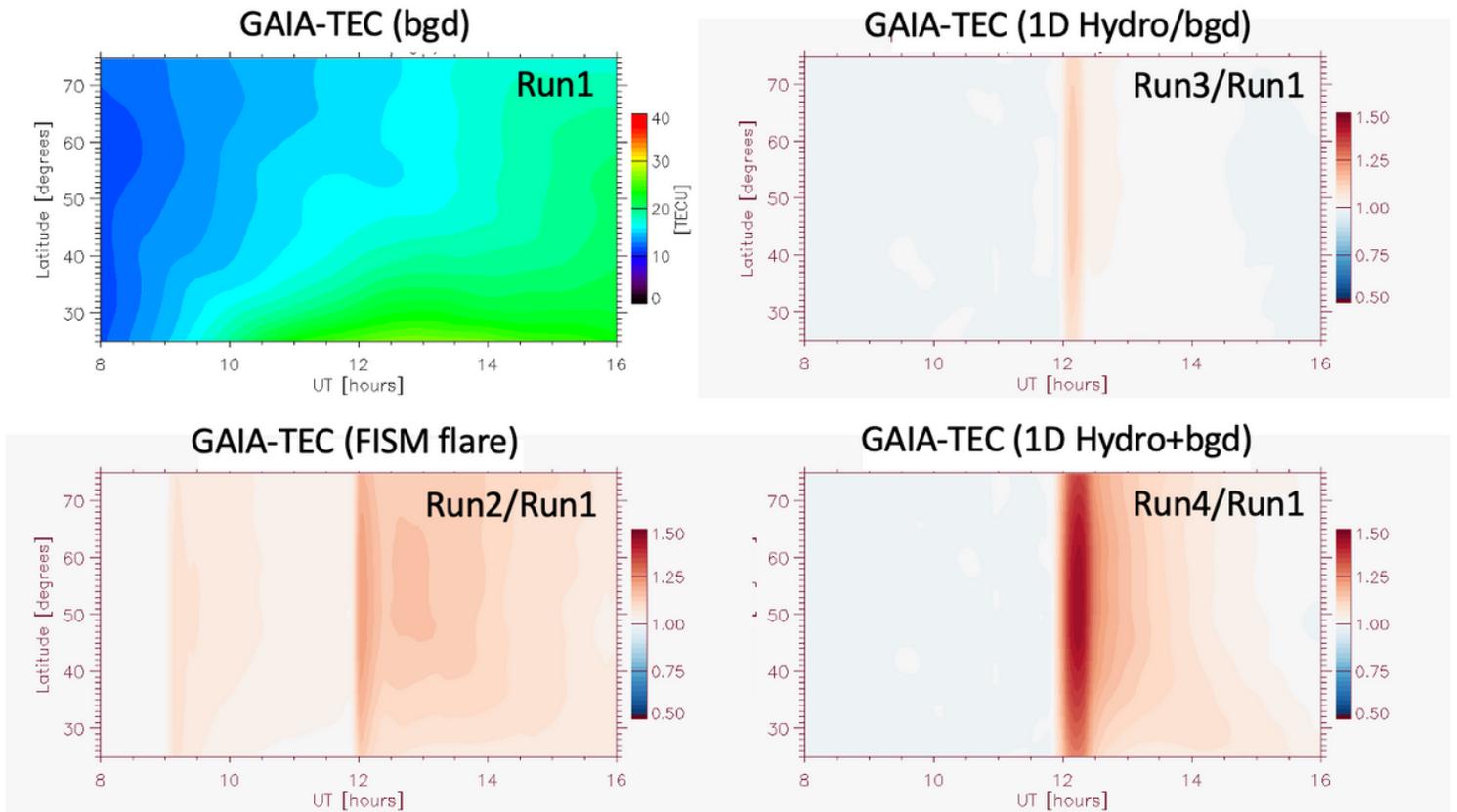
**Figure 6**

Difference in ionization rate due to X9.3 flare that occurred in September 6, 2017. The difference in ionization rate of each ion at 0°E and 30°N due to X9.3 flare on September 6, 2017, including ionization by solar irradiance and secondary photoelectrons. The FISM flare spectrum is used for the input spectrum to GAIA. Differences in ionization rate before and after flare for (a) O<sup>+</sup>, (b) O<sub>2</sub><sup>+</sup>, (c) N<sup>+</sup>, (d) N<sub>2</sub><sup>+</sup>, and (e) total ion as a function of the wavelength of solar irradiance (unit: Log<sub>10</sub>(W m<sup>-2</sup> nm<sup>-1</sup>)). (f) Differences in the altitude distribution of the ionization rate of each ion. Each solid line starts immediately after the flare (12:00 UT), and each dotted line starts before the flare (11:30 UT).



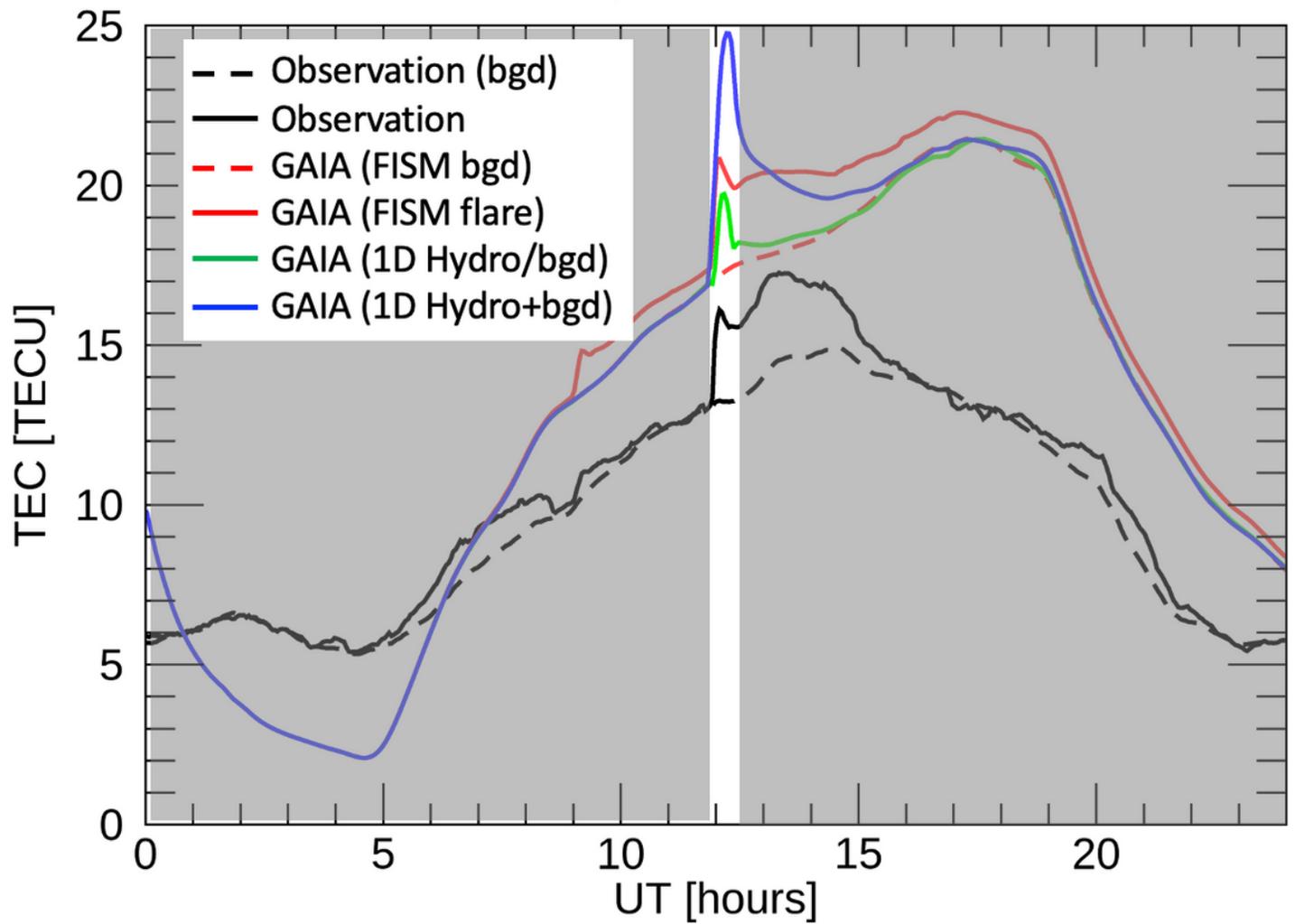
**Figure 7**

TEC distribution at 0°E 25–75°N calculated by GAIA TEC distribution derived by inputting (upper left: Run1) background, (lower left: Run2) “FISM”, (upper right: Run3) “1D Hydro/bgd”, and (lower right: Run4) “1D Hydro+bgd” spectral changes.



**Figure 8**

Ratio of TEC at 0°E 25–75°N calculated by GAIA and the background Upper left: background TEC distribution. Same as the top-left figure in Figure 7. Lower left: ratio of TEC for the “FISM” input (Run2/Run1). Upper right: ratio of TEC for the “1D Hydro/bgd” input (Run3/Run1). Lower right: ratio of TEC for the “1D Hydro+bgd” input (Run4/Run1).



**Figure 9**

Variation in TEC at 0°E and 44°N The solid black line is the observed value (top-right panel in Figure 4), and the dashed black line is the background of the observation (top-left panel in Figure 4). The red dashed line, red solid line, green line, and blue line represent the calculation results of GAIA with background input (Figure 7: Run1), “FISM” input (Figure 7: Run2), “1D Hydro/bgd” input (Figure 7: Run3), and “1D Hydro+bgd” input (Figure 7: Run4), respectively. Gray regions have no 1D Hydro model data.

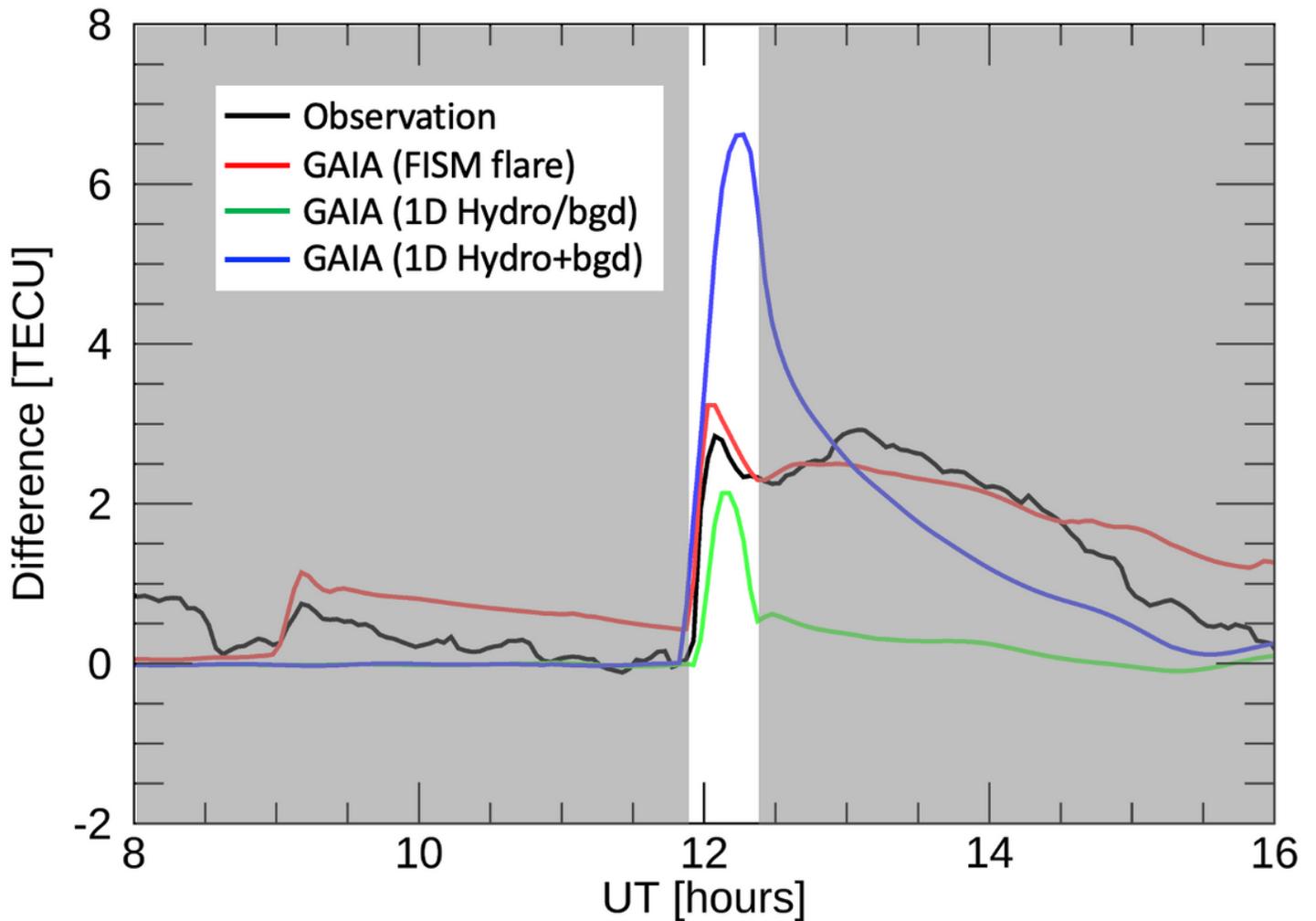


Figure 10

Difference in TEC from background at 0°E 44°N The black line is the difference between the observed TEC and the background. The red, green, and blue lines represent the difference in the calculation results of GAIA with “FISM” input, “1D Hydro/bgd” input, and “1D Hydro+bgd” input from the background, respectively. Gray regions have no 1D Hydro model data.

## Supplementary Files

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