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Comparison of Precipitation Projections of CMIP5 and CMIP6 Global Climate Models over Yulin, China

Mohammed Sanusi Shiru

Federal University Dutse

Eun-Sung Chung (Second eschung@second es

Seoul National University of Science and Technology https://orcid.org/0000-0002-4329-1800

Shamsuddin Shahid

Universiti Teknologi Malaysia

Xiao-Jun Wang

Nanjing Hydraulic Research Institute

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1 Corresponding Author

- 2 Eun-Sung Chung
- 3 Professor, Seoul National University of Science and Technology
- 4 Email: eschung@seoultech.ac.kr

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9	Mohammed Sanusi Shiru ^{1,2} , Eun-Sung Chung ^{1*} , Shamsuddin Shahid ³ , Xiao-Jun Wang ⁴
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11 12	¹ Department of Civil Engineering, Seoul National University of Science and Technology, Seoul 01811, South Korea.
13 14	² Department of Environmental Sciences, Faculty of Science, Federal University Dutse, P.M.B 7156, Dutse, Nigeria.
15 16	³ Department of Water and Environmental Engineering, School of Civil Engineering, Faculty of Engineering, Universiti Teknologi Malaysia (UTM), 81310 Johor Bahru, Malaysia
17 18	⁴ State Key Laboratory of Hydrology-Water Resources and Hydraulic Engineering, Nanjing Hydraulic Research Institute, Nanjing 210029
19	
20	*Correspondence: eschung@seoultech.ac.kr
21	
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Comparison of Precipitation Projections of CMIP5 and CMIP6 Global Climate Models over Yulin, China

25

26 Abstract

This study compared precipitation projections of Coupled Model Intercomparison Project 5 27 (CMIP5) and 6 (CMIP6) GCMs over Yulin City, China. The performance of CMIP5 and CMIP6 28 GCMs in replicating Global Precipitation Climatology Centre (GPCC) precipitation climatology 29 30 of the city was evaluated using different statistical metrics. The best performing GCMs common 31 to both CMIP5 and CMIP6 were selected and subsequently downscaled to GPCC resolution using linear scaling method to spatiotemporal changes in precipitation. The study revealed 32 BCC.CSM1.1(m), IPSL.CM5A.LR, MRI.CGCM3 and MIROC5 of CMIP5 and their equivalents 33 BCC-CSM2-MR, IPSL-CM6A-LR, MRI.ESM2.0 and MIRCO6 of CMIP6 as the most suitable 34 GCMs for the projection of rainfall in Yulin. Changes in precipitation were in the range of -14.0 35 -0.0% and -22.0 - 0.2% during 2021–2060 for RCP4.5 and SSP2-4.5 respectively. The highest 36 decrease of -29.7 - -22.0% was projected by MRI-ESM-2-0 for SSP2-4.5, while -28.0 - -20.0%37 by MIROC5 for RCP4.5. For RCP8.5 and SSP5-8.5, precipitation was projected to decrease in 38 the range of -17.0 - -2.0% and -32.0 - 0.0%, respectively during 2021 - 2060 by most of the 39 GCMs. An increase in precipitation up to 12.3% was projected only by IPSL-CM5A-LR for 40 RCP8.5 for this period. The highest decrease was projected by MIROC5 (-40.2 - -29.0%) for 41 RCP8.5 and IPSL-CM6A-LR (-40.2 - -26.0%) for SSP5-8.5. Overall, the results revealed a 42 43 higher decrease in precipitation in Yulin city by CMIP6 GCMs compared to those projected by their corresponding GCMs of CMIP5 for both scenarios. 44

45 Keywords: CMIP6, climate change, global circulation model, statistical indices, Yulin

47 1. Introduction

Climate change has been an issue of concern over several decades due to its devastating impacts. It has increased the risks of flooding (Manawi et al., 2020; Rahman et al., 2019), occurrences of droughts (Alamgir et al., 2019; Ayugi et al., 2020; Shiru et al., 2019a), heatwaves (Kang and Eltahir, 2018; Khan et al., 2019; Satyanarayana and Rao, 2020) and ecosystem damages (Kim et al., 2019; Pérez-Ruiz et al., 2018). Many sectors including water resources, agriculture, energy and health among others have also been affected by the changing climate.

Like many other parts of the globe, China is also experiencing the impacts of climate change. 54 Flooding is a common occurrence in different parts of China including cities like Yulin (Huang 55 et al., 2015; Yang et al., 2017) and it often causes damages to sectors such as power, agriculture, 56 57 health and sometimes leads to loss of lives (He et al., 2018). In 2003, heavy precipitation caused river flooding leading to the destruction of 17,438 acres of farmlands, destruction of roads, and 58 damages to about 1,458 houses (He et al., 2018). Heavy precipitation affected more than 150,000 59 people in 2012 with 15 people missing and11 deaths, destruction of roads and disruption of 60 aviation due to the heavy rain (China-Daily 2012). The disaster caused a direct economic loss in 61 the Yulin and Yan'an region up to 134 million yuan (\$21 million). Similarly, the flood in 2016 62 affected more than 21,000 people and evacuation of about 1200 persons (He et al., 2018). The 63 flood led to the damage of more than 850 houses, destruction of more than 1100 kilometers of 64 roads, damages to 30 bridges and culverts, and caused 11 landslides and other related geological 65 disasters. The economic implication of the disaster was about 150 million yuan. The region is 66 67 also known to be affected by droughts (Wang et al., 2020; Yin et al., 2020).

⁶⁸ Understanding the expected changes in climate is crucial for the areas susceptible to ⁶⁹ disasters in order to develop adaptation and mitigation plans against climate change. This is ⁷⁰ particularly important using the recently released global climate models (GCM) of the Coupled ⁷¹ Model Inter-comparison Phase 6 (CMIP6). It is also crucial to assess how the CMIP6 differ in ⁷² projection from the previous Coupled Model Inter-comparison Phase 5 (CMIP5) in order to ⁷³ streamline the existing adaptation measures based on CMIP5 projections (Jiang et al., 2020; ⁷⁴ Song et al., 2021).

GCMs are developed by different institutions and have different performances in different parts of the globe (Chen et al., 2017). Hence, the selection of the most realistic ones for a reliable projection of climate for a region is crucial. The assessment of the performances of GCMs has been conducted using different statistical indices (Rivera and Arnould, 2020; Sreelatha and
Anand Raj, 2019). However, due to contradictory outputs from different statistical measures,
supporting such outputs with other measures can be beneficial in reaching a compromise.

As GCMs are characterized by coarser spatial resolutions, their applications in climate 81 projections and climate change impact studies at local and regional scales can be unreliable 82 (Onyutha et al., 2016; Pour et al., 2018; Salman et al., 2018). Therefore, they are required to be 83 downscaled (Almazroui et al., 2020a; Sa'adi et al., 2019; Shiru et al., 2020) using either the 84 85 dynamical downscaling or the statistical downscaling (SD) methods. The SD method is known to have the advantages of computational efficiency and cost-effectiveness, the possibility of 86 incorporation of observations directly into methods, and provision of point-scale climatic 87 projections from GCM-scale (Fowler et al., 2007). 88

89 An array of statistical metrics was used to select the best performing GCMs for the projection of precipitation in the study area. Equivalent GCMs of the CMIP5 and CMIP6 for 90 91 selected for the comparison of projections. The selected GCMs were downscaled using linear scaling method. This study employs only the GCMs which can reliably simulate the exiting 92 93 climate of the study area and thus, capable to provide a trustworthy comparison of CMIP5 and CMIP6 projects. It is expected that the comparison of precipitation projections of CMIP5 and 94 95 CMIP6 would help in streamlining the existing adaptation measures formulated based on CMIP5 projections or deriving new measures based on new scenarios of CMIP6. 96

97

98 2. Study Area and Datasets

99 2.1. Study Area

The study area, Yulin (Figure 1) is located in the northern Shaanxi province of China (Longitude: 107°15′47″–111°14′44″E; Lattitude: 36°49′07″–39°34′47″N). Yulin covers a total area of 385 km by 263 km (Zha, et al. 2008). The terrain of the area descends from 1,907 m in the east to 585 m in the west above the mean sea level. Yulin has a semi-arid temperate continental monsoon type climate which is characterized by dry and little precipitation in spring and winter and high precipitation during summer and autumn. The annual average precipitation in Yulin is around 450 mm. The annual average temperature in the area is 9.6°C (Wang, 2016).





108

Figure 1 Map of Yulin showing the grid points considered for projection in the study.

109 2.2. Gauged based gridded precipitation

The GPCC full data reanalysis product of the Deutscher Wetterdienst (Becker et al., 2013; 110 111 Schneider et al., 2014) was used in this study as the reference data. The GPCC precipitation amongst most other precipitation products has the advantages of (1) good data quality for 112 hydrological studies, (2) availability for a longer period, (3) better performance as being 113 114 developed using the highest number of collected precipitation records, and (4) completeness of time series for the recent decades (Ahmed, et al. 2017; Spinoni, et al. 2014). This study used the 115 monthly precipitation data for the period 1961 - 2005. Data were collected from a total of 100 116 grid points to cover the whole Yulin. The location of the grid points is shown in Figure 1. 117

118

119 2.3 Global Climate Models

120 This study uses the historical and future simulations of GCMs of CMIP5 and CMIP6. The 121 CMIPs are sets of globally coordinated GCM simulations which comprises of historical and 122 future climate simulations assembled from different climate modeling groups. The CMIP5 offers

significant improvements compared to the CMIP3 (Taylor et al., 2012). The CMIP5 comprises of 123 four scenarios called representative concentration pathways (RCPs). In the CMIP6, the four RCP 124 scenarios of CMIP5, RCP2.6, RCP4.5, RCP6.0 and RCP8.5 have been updated as Shared 125 Socioeconomic Pathways (SSPs) scenarios, SSP1-2.6, SSP2-4.5, SSP4-6.0, and SSP5-8.5 126 respectively. Each GCM also considers 2100 radiative forcing levels. The GCMs of CMIP6 are 127 developed through improved emission scenarios, land use data, physical processes and model 128 parameterization to provide more realistic projections of future climate (Eyring, et al. 2016; 129 O'Neill, et al. 2016). In this study, historical and future simulation of 10 GCMs of CMIP5 and 130 CMIP6 were considered. The GCMs were chosen based on their availability from the same 131 institution. The first ensemble members for both CMIP5 and CMIP6 were considered. The 132 GCMs chosen for this study are provided in Table 1. 133

134

	135	Table 1.	Descripti	ions of th	e GCMs	of CMIP5	and CMIP	'6 used in	this st	tudy.
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Institution - Country	CMIP5 Model Name	Resolution (lon/lat in °)	CMIP6 Model Name	Resolution (lon/lat in °)	
Australian Community Climate and Earth System	ACESS1 2	10×12	ACCESS ESMI 5	1 0~1 2	
Simulator - Australia	ACESSI.5	1.9 × 1.2	ACCESS.ESM1.5	1.9×1.2	
Beijing Climate Center - China	BCC.CSM1.1(m)	1.125×1.125	BCC.CSM2.MR	1.1×1.1	
Canadian Centre for Climate Modelling and		28428		2.8×2.8	
Analysis - Canada	CanESM2	2.8 × 2.8	CanESM5		
NASA Goddard Institute for Space Studies -		2520		2.5×2.0	
United States	GISS.E2.R	2.5 × 2.0	GISS.E2.1.G		
Marchuk Institute of Numerical Mathematics -		2 ~ 1 5		2×1.5	
Russia	INM.CM4	2 × 1.5	INM.CM4.8	2×1.5	
Institut Pierre-Simon Laplace - France	IPSL.CM5A.LR	3.75×1.875	IPSL.CM6A.LR	2.5×1.3	
The University of Tokyo, National Institute for					
Environmental Studies, and Japan Agency for	MIROC5	1.4×1.4	MIROC6	1.4× 1.4	
Marine-Earth Science and Technology - Japan					
Max-Planck-Institut für Meteorologie - Germany	MPI.ESM.LR	1.9×1.9	MPI.ESM1.2.LR	1.9×1.9	
Meteorological Research Institute - Japan	MRI.CGCM3	1.25×1.25	MRI.ESM2.0	1.1×1.1	
Norwegian Meteorological Institute - Norway	NorESM1.M	2.5×1.875	NORESM2.MM	2.9 x 1.9	

136 137

3. Methodology

138 The methods used in this study are described in this section. For an unbiased comparison of 139 model performance, the GPCC and the precipitation simulations of all GCMs were re-gridded to a uniform resolution of $1^{\circ}\times1^{\circ}$ (latitude×longitude) using bilinear interpolation to have a uniform resolution. Bilinear interpolation is often conducted for transforming spatially coarse GCM data into finer data through GCM data interpolation from the four nearest neighboring grid points (Ahmed, et al. 2019; Almazroui, et al. 2020a). After selection of GCMs, the selected GCMs and the GPCC precipitation data were re-gridded to 0.25° and used for spatiotemporal projection of precipitation.

146 *3.1 Statistical Indices*

The ability of the different GCMs in reproducing the properties of GPCC precipitation at all the 147 148 100 grid points of the study area was assessed using four statistical indices: normalized root mean square error (NRMSE), Percentage of Bias (Pbias), Volumetric Efficiency (VE), and 149 150 Coefficient of Determination (R²). Besides, probability density function (pdf) and spatial relationship of the mean monthly precipitation of the different GCMs were compared with 151 152 GPCC precipitation to assess their abilities in replicating the precipitation climatology of the study area. Details about the statistical metrics used in this study are as follows. The expressions 153 154 used to describe statistical metrics used here: $x_{pred,i}$ and $x_{obs,i}$ are the *i*-th GCM and GPCC data.

The magnitude of the errors in estimation can be summarized by NRMSE (Willmott 1982). It is a normalized statistic that provides a relative magnitude of the residual variance to the variance of the measured data. Smaller NRMSE values (preferably zero) indicate better performance of the model. NRMSE is defined as follows

159
$$NRMSE = \frac{\left[\frac{1}{n}\sum_{i=1}^{n} (x_{pred,i} - x_{obs,i})^{2}\right]^{1/2}}{\frac{1}{n}\sum_{i=1}^{n} (x_{pred,i})}$$
(1)

160 The tendency of GCM to underestimate or overestimate the GPCC precipitation is measured 161 using Pbias. Model performance is better when the Pbias is closer to zero. The PBIAS is a 162 statistical metric that gives the estimate of over estimation or under prediction of a model 163 (Wagena, et al. 2018). The evaluation of Pbias is conducted as follows.

164
$$Pbias = 100 \times \left[\frac{\sum_{i=1}^{n} (x_{pred,i} - x_{obs,i})}{\sum_{i=1}^{n} x_{pred,i}}\right]$$
 (2)

165 The VE measures the ratio between GCM and GPCC precipitation volumes over a period, where 166 a VE value of 1 indicates a perfect estimation. It can be calculated using the following equation.

167
$$VE = 1 - \frac{\sum_{i=1}^{n} (x_{pred,i} - x_{obs,i})}{\sum_{i=1}^{n} x_{obs,i}}$$
(3)

The R^2 can be defined as the square of the Pearson's product-moment correlation coefficient (i.e. R² = r²) describing the proportion of the total variance in the GPCC precipitation which is explainable by GCM (Legates and McCabe Jr 1999). R² values can range between -1.0 and 1.0, in which the higher absolute value indicates a better agreement. Computation of R² is as follows.

172
$$R^{2} = \frac{\sum_{i=1}^{N} \left(x_{obs,i} - \overline{x_{obs}} \right) \left(x_{\Pr ed,i} - \overline{x_{\Pr ed}} \right)}{\sqrt{\sum_{i=1}^{N} \left(x_{\Pr ed,i} - \overline{x_{\Pr ed}} \right)^{2} \sum_{i=1}^{N} \left(x_{obs,i} - \overline{x_{obs}} \right)^{2}}}$$
(4)

173

174 *3.2 Downscaling of precipitation of selected GCMs of CMIP5 and CMIP6*

The linear scaling (LS) method was applied for the downscaling of the selected GCMs. LS (Lenderink, et al. (2007) uses the monthly correction values obtained from the difference in GPCC and GCMs simulated monthly precipitation for the reference period. The monthly scaling factor is then applied to raw GCM data. The monthly precipitation, P is corrected using the following equation:

$$\begin{array}{l} 180\\ 181 \qquad P^* = \alpha P \end{array} \tag{5}$$

183 where
$$\alpha = P_o/P_s$$
 (6)

184

 P_o is the monthly mean of GPCC precipitation whereas P_s is the monthly mean of the GCM simulated precipitation. LS method requires less information such as only monthly data to calculate the scaling factor (Lafon, et al. 2013) and thus, widely used for precipitation downscaling.

189

190 **4. Results**

191 *4.1 Performance assessment of GCMs using statistical metrics*

192 The results of the statistics used for the performance evaluation of different historical GCMs are

193 presented in Table 2. It shows a variation in the performances of different GCMs. For CMIP5,

- the GCMs with the best performance metrics are ACESS1.3, BCC.CSM1.1(m), IPSL.CM5A.LR,
- and MIROC5. The GCMs of the same institutions also showed good performances for CMIP6
- 196 except for ACCESS.ESM1.5 which showed a relatively poor performance.
- 197

198 Table 2. Statistical metrics showing the performances of GCMs of CMIP5 and CMIP6 in replicating GPCC

199 precipitation over Yulin (Best metrics are presented in bold).

	CMIP	CMIP6							
GCM	NRMSE	Pbias	VE	R ²	GCM	NRMSE	Pbias	VE	R ²
ACESS1.3	76.3	10.9	0.68	0.73	ACCESS.ESM1.5	154.4	23.8	0.67	0.77
BCC.CSM1.1(m)	144.8	21.2	0.79	0.79	BCC.CSM2.MR	115	19.8	0.8	0.63
CanESM2	292.9	49.8	0.5	0.16	CanESM5	136.1	23.7	0.73	0.46
GISS.E2.R	135.8	29.1	0.68	0.22	GISS.E2.1.G	175.8	53.3	0.47	0.04
INM.CM4	274.7	38.8	0.61	0.46	INM.CM4.8	279.9	41.5	0.59	0.45
IPSL.CM5A.LR	81.0	-4.5	0.77	0.36	IPSL.CM6A.LR	70.2	6.1	0.8	0.59
MIROC5	177.2	27.5	0.72	0.76	MIROC6	181.6	23.3	0.72	0.65
MPI.ESM.LR	174.3	35.5	0.65	0.41	MPI.ESM1.2.LR	123.3	23.6	0.76	0.54
MRI.CGCM3	173.8	-44.6	0.70	0.85	MRI.ESM2.0	85.8	-21.3	0.78	0.78
NorESM1.M	397.3	55.1	0.45	0.21	NORESM2.MM	182.6	29.8	0.7	0.62

200

201 4.2 Spatial relationship between GCMs and GPCC precipitations

The ability of different GCMs in replicating the spatial distribution of GPCC precipitation for the study area are presented in Figure 2. The performances of the GCMs were found to vary widely in reproducing the GPCC precipitation. Among CMIP5 GCMs, the highest overestimations were by CanESM2 and NorESM1-M, while an underestimation was by MRI-CGCM3 followed by IPSL-CM5A-LR in some parts. For CMIP6, GISS-E2-1G showed the highest overestimation of precipitation. Generally, the GCMs with better performance metrics (Table 2) showed better skills in replicating the precipitation climatology of GPCC for the study area.



Figure 2 Spatial distribution of average annual precipitation of GCMs of CMIP5 and CMIP6 compared to

that of GPCC during 1961–2005.

212

213 *4.3 Comparison using probability density function*

The PDFs of monthly GCM precipitation was compared with the PDF of GPCC precipitation for the study area. The results for CMIP5 and CMIP6 are presented in Figure 3(a) and 3(b) respectively. Figures show that most GCMs were able to replicate the precipitation properties of the GPCC, especially the skewness. However, the distribution of precipitation was found better for the GCMs which showed a better performance in terms of statistical metrics presented in Table 2.



Figure 3 Probability density function of monthly precipitations of GPCC and the GCMs of (a) CMIP5 and(b) CMIP6.

223

4.4 Comparison of GCMs in reproducing monthly mean precipitation

The mean monthly GCM precipitation was compared with mean monthly GPCC precipitation for the period 1961–2005. Obtained results for the GCMs of CMIP5 and CMIP6 are presented in Figure 4(a) and (b) respectively. There was a variation in the estimation of GPCC precipitation by the GCMs, especially during the wet season. CanESM2 and NorESM1-M of CMIP5 overestimated the precipitation for all the months, while an underestimation was by MRI-CGCM3. For the CMIP6, overestimation was by GISS-E2-1-2G and underestimation by MRI-ESM2-0 and IPSL-CM5A-LR, especially during the wet period. Though most of the CMIP6 GCMs were found to overestimate the GPCC precipitation, overall they were found morecapable in replicating the mean monthly precipitation of GPCC compared to CMIP5 GCMs.



Figure 4 Comparison of mean monthly precipitation of GCMs of (a) CMIP5 and (b) CMIP6 with that of GPCC
 precipitation.

239 4.5 Selection of the best performing GCMs

The performance of the GCMs based on statistical indices and replication of PDF, and spatial precipitation distribution patterns were considered. Based on the statistics, BCC.CSM1.1(m), IPSL.CM5A.LR, and MIROC5 were the better performing GCMs for both CMIP5 and CMIP6. Besides, ACESS1.3 showed overall a better performance compared to the rest of the CMIP5 GCMs, while its equivalent under CMIP6 showed a poor performance than MRI.ESM2.0. As CMIP6 is a more recent simulation, the MRI.CGCM3 for CMIP5 and its equivalent GCM for MICP6, MRI.ESM2.0 were prioritized as the fourth model for projection of precipitation.

247

248 4.7 Projection of precipitation from the selected GCMs of CMIP5 and CMIP6

The spatial projections of precipitation for the study area by CMIP5 GCMs for RCP 4.5 and 249 CMIP6 GCMs for SSP2-4.5 for two future periods, 2021–2060 and 2061 – 2100 are presented in 250 Figure 5. Large heterogeneity in precipitation changes was projected by different GCMs for 251 RCPs and SSPs and the two projection periods. During 2021 - 2060, the highest decrease in 252 253 precipitation was projected by MRI-ESM-2-0 for SSP2-4.5 while the highest decrease for RCP 4.5 was projected by MIROC5. For the same period, BCC-CSM2-MR projected an increase in 254 255 precipitation by 1.2% at the extreme west of the study area. Percentage change in precipitation was in the range of -14.0 - 0.0% for RCP4.5 while it was -22.0 - 1.2% for SSP2-4.5. During 256 257 2061 – 2100, all the GCMs projected decreases in precipitation, with the highest decrease of -29.7 - -22.0% by MRI-ESM-2-0 for SSP2-4.5 and -28.0 - -20.0% by MIROC5 for RCP4.5. The 258 259 lowest decrease during this period was projected by MRI-CGCM3 and BCC-CSM2-MR for RCP4.5 and SSP2-4.5 respectively. 260



Figure 5. Spatial distribution of projected precipitation by the GCMs of CMIP5 and CMIP6 during 2021–
2060 and 2061–2100 for RCP4.5 and SSP2-4.5.

The spatial distribution of projected precipitation in the study area by CMIP5 GCMs for RCP 8.5 and CMIP6 GCMs for SSP5-8.5 for the periods 2021-2060 and 2061-2100 are presented in Figure 6. Compared to RCP 4.5 and SSP2-4.5, RCP8.5 and SSP5-8.5 showed higher decreases in precipitation. During 2021 - 2060, the projected decrease in precipitation was in the range of -17.0 - 2.0% by the CMIP5 GCMs for RCP8.5 while the decrease was projected between -32.0 and 0.0% by the CMIP6 GCMs for SSP5-8.5. Increases in precipitation of 0.0 - 12.3% were noticed only for IPSL-CM5A-LR for RCP8.5. During 2061 – 2100, the decrease in precipitation was projected the highest by MIROC5 (-40.2 – -29.0%) for RCP8.5 and IPSL-CM6A-LR (-40.2 – -26.0%) for SSP5-8.5.



273

Figure 6. Spatial distribution of projected precipitation by the GCMs of CMIP5 and CMIP6 during 2021–

275 2060 and 2061–2100 for RCP8.5 and SSP5-8.5.

277 **5. Discussion**

Climate change remains a major challenge in many parts of the globe as they have 278 devastating impacts on several sectors. Many studies have shown that the expected changes in 279 climate will result in increased temperatures and more erratic precipitations in many parts of the 280 281 world. Projection of precipitation under CMIP5 in Nigeria showed that while precipitation will increase in some parts of the country, particularly the semi-arid and arid regions where they are 282 283 usually low, the other parts where they used to be higher will experience decreases (Shiru, et al. 284 2019b). (Homsi, et al. 2020) used CMIP5 GCMs in Syria and showed that precipitation would increase by up to 87% in some parts and decrease by up to -85% in the coastal areas. Many other 285 studies have also shown both increase and decrease in precipitation in many parts of the world 286 287 using CMIP5 GCMs (Iqbal, et al. 2020; Narsey, et al. 2020; Shiru and Park 2020; Ullah, et al. 288 2020).

The studies for the recently released CMIP6 GCMs have also shown such changes in 289 precipitation. The study conducted over South Asian countries Almazroui, et al. (2020c) showed 290 that the annual mean precipitation will increase by 27.3% in India, 18.9% in Bhutan, 26.4% in 291 292 Pakistan, 19.5% in Nepal, 25.1% in Sri Lanka and 17.1% in Bangladesh in the last part of the century under SSP5-8.5 scenario. Over Africa, (Almazroui, et al. 2020b) projected precipitation 293 294 under CMIP6 and showed that while the northern and the southern parts of Africa are analyzed to witness a reduction in precipitation, the central parts are expected to have increases of 6.2%, 295 296 6.8%, and 9.5% for SSP1-2.6, SSP2-4.5, and SSP5-8.5, respectively.

In China, studies showed variations in the projected precipitations for both CMIP5 and 297 CMIP6 with some studies showing mostly increases while others showed a simultaneous 298 increase and decrease depending on the region. The annual precipitation was projected to 299 increase significantly relative to the present day for CMIP5 (Chen 2013). This study revealed 300 that the increase of precipitation in the Northwest of China is primarily due to the increase in 301 light showers while the increases in precipitation in the north and northeastern parts are due to an 302 increase in medium precipitation. The increases in precipitation are expected in the southern 303 304 parts of China due to an increment in heavy precipitation. Zhou, et al. (2019) showed both increases and decreases in daily precipitation under different warming conditions over China 305 using CMIP5. The increase in warming of 1.5°C was projected to cause an increase in the 306

frequency and intensity of precipitation in northeast China, north China, and the Qinghai–Tibet Plateau while there would be a decrease in total precipitation in south China and southwest China. The number of wet days were projected to increase in the north while the decreases in the south under 2°C.

The projection of the changes in precipitation over northwestern China for CMIP6 (Su-311 Yuan, et al. 2020) showed that there would be lesser warming compared to that previously 312 expected, which would affect the patterns in precipitation changes. Unlike in the historical period 313 (1850 - 2014) when the rate of warming was 0.05° C per decade, the annual mean temperature is 314 projected to increase up to 0.06°C, 0.26°C and 0.59°C per decade for SSP1-2.6, SSP2-4.5 and 315 SSP5-8.5, respectively for the period 2015–2099. The total annual precipitation for the area is 316 projected to increase by 5.6, 6.4 and 8.0 mm/decade for SSP1-2.6, SSP2-4.5 and SSP5-8.5, 317 318 respectively.

In this study, both increases and decreases in precipitation are projected for the study area. Zhou, et al. (2019) projected the increases in precipitation over China, except some northwestern parts where Yulin belongs, which supports the findings of this study. The projected the increases in some parts of the study area are also corroborated by other studies.

323

324 6. Conclusions

325 This study compares the projections of precipitation by CMIP5 and CMIP6 GCMs over Yulin city of China. Different statistical and graphical metrics were used in assessing the ability 326 of 10 GCMs in replicating the precipitation properties of the study area. Finally, four GCMs 327 having the highest ability in replicating the properties of GPCC precipitation were selected for 328 329 the projection of precipitation over the study area. This study revealed that ACESS1.3, BCC.CSM1.1(m), IPSL.CM5A.LR and MIROC5 of CMIP5 and their equivalents in CMIP6, 330 BCC-CSM2-MR, IPSL-CM6A-LR, MRI.ESM2.0 and MIROC6 have better abilities in 331 replicating historical precipitation properties over Yulin. Projection of precipitation showed an 332 overall decrease in precipitation over Yulin by the GCMs of both CMIPs. The decrease in 333 334 precipitation would be more in the far future compared to the near future. The expected decreases in precipitation can increase the frequency and intensity of droughts in the area. The 335

findings of the study can be used as a guide in the development of adaptation and mitigation

measures against climate change in the area. In the future, more GCMs common to both CMIP5

and CMIP6 can be employed when they will be available for CMIP6 for the selection of best

performing GCMs. Besides, a more reliable approach can be utilized for the selection of GCMs

- to avoid the dispute in selection using conventional statistical metrics.
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- 342

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Sanusi Shiru, Eun-Sung Chung, Shamsuddin Shahid, and Xiao-Jun Wang; Methodology,
Mohammed Sanusi Shiru, Eun-Sung Chung, Shamsuddin Shahid, and Xiao-Jun Wang, Writing –
original draft, Mohammed Sanusi Shiru, Eun-Sung Chung; Writing – review & editing,
Mohammed Sanusi Shiru, Eun-Sung Chung, Shamsuddin Shahid, and Xiao-Jun Wang.

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This article does not contain any studies with human or animal participants performed by any ofthe authors.

489

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491 The authors declare that they have consent to participate in this paper.

492

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- 497 The authors declare that they have no known competing interests that authors must disclose all
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502 The authors can support all relevant data if requested.