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Research Article

Keywords: Extreme events, Global warming, Mutation analysis, detrended fluctuation analysis, climatic disaster

Posted Date: August 10th, 2021

DOI: <https://doi.org/10.21203/rs.3.rs-662878/v1>

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Version of Record: A version of this preprint was published at Theoretical and Applied Climatology on October 26th, 2021. See the published version at <https://doi.org/10.1007/s00704-021-03828-1>.

Spatiotemporal trends of temperature extremes in Bangladesh under changing climate using multi-statistical techniques

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Abstract

The rise in the frequency and magnitude of extreme temperature phenomena across the globe has led to the recurrent incidence of global climate hazards, which have had severe effects on socioeconomic development. The daily maximum and minimum temperature datasets of 27 sites in Bangladesh were used to detect spatiotemporal trends of temperature extreme over Bangladesh during 1980-2017 based on ten temperature extreme indices using multi-statistical modeling namely linear regression, Pearson correlation coefficient and factor analyses. Besides, mutation analyses based on the Mann-Kendall test, Sen's slope estimator

and Pettit test were employed to show the changing trend in extreme temperature. Results show that except for warmest days, the warm indices showed an increasing trend, mainly since the 2000s, while the growth rate was faster, and the response to global climate warming was sensitive. The cold indices demonstrated a reverse trend since the 2010s. Diurnal temperature range (DTR) and summer days (SU) increased faster, implying that the rising speed of daily max temperature was higher than of daily min-temperature in Bangladesh. The detrended fluctuation analysis (DFA) revealed a continuous increase in temperature extreme in the future except for cold days. The probability distribution functions (PDF) analysis revealed an evident variation of the curves in recent decades compared to the past three decades. Besides the warm night, DTR and SU primarily control the general warming trend of temperature extremes over Bangladesh during the study period. The mutation of the warm indices occurred before the cold index, indicating that the warm indices were more sensitive to global climate warming. The temperature extremes recognized in our research suggest that elevated warm temperature extremes due to global climate warming may have huge implications on the sustainable development of Bangladesh in the forthcoming period.

Keywords: *Extreme events, Global warming, Mutation analysis, detrended fluctuation analysis, climatic disaster.*

1. Introduction

Temperature extremes have critical effects on ecosystem imbalance, agricultural productivity, water resources and sustainable socioeconomic development (Easterling et al. 2000; Ciais et al. 2005; Schmidli and Frei 2005; Benestad and Haugen 2007; Allen et al. 2010; Rammig and Mahecha, 2015; Guo et al. 2019). Variations in temperature extremes in a warming climate have paid considerable attention in recent years due to the enormous impact of severe temperature occurrences on society and ecosystems (Bandyopadhyay et al., 2012; Sun et al., 2014). Furthermore, the scientific community has emphasized measure to tackle the climate change impact on ecosystems and human livelihoods for disaster prevention and mitigation (Smith 2011; Jiang et al. 2012; Endfield 2012; Garcia-Cueto et al. 2014; Sharma et al. 2018). As climate extreme events increased in many parts of the world, including Bangladesh (Hasan et al. 2013; Shahid et al. 2016; Mahmud et al. 2018; Khan et al. 2019). Thus, it is crucial to investigate spatiotemporal changes in Temperature Extremes for sustainable development over Bangladesh.

Extreme temperature phenomena are changing across the globe due to global warming (Alexander et al. 2006; Aguilar et al. 2005; Hidalgo-Muñoz et al. 2011; Coumou and Rahmstorf 2012; Abiodun et al. 2013; Coumou et al. 2013; Omondi et al. 2014). The fact is that global and local average temperatures are increasing, which have no bearing on the occurrence of extreme events (Finkel et al. 2018; Gleixner et al. 2020). However, severe temperatures may be more vital than mean temperatures for productivity and human survival, and local extreme temperature changes could significantly impact global mean temperature changes (Finkel et al., 2018; Gao et al., 2017; Katz et al., 1992). Several studies have reported that the frequency of warm temperature indices is increasing, while the frequency of cold temperature indices is decreasing (Alexander et al. 2006; Tank et al. 2006; Piccarreta et al. 2015; Sheikh et al. 2015; Guan et al. 2015; Sun et al. 2016; You et al. 2017; Ullah et al. 2019). All these studies looked at average temperature trends and temperature extremes, but the latter only looked at monthly

68 and yearly maximum and minimum values as extremes. The Expert Team on Climate Change Detection
69 and Indices (ETCCDI) has identified extreme temperature indices to illustrate better temperature
70 extremes, commonly used in climate change-related studies in various parts of the world (Zhang et al.,
71 2011). Therefore, a deep understanding of temperature extremes' trends and fluctuations is essential for
72 developing accurate estimates of future climate change projections and answering scientific researchers',
73 climate change analysts and decision-makers unique concerns.

74 Bangladesh is one of the world's top ten countries that could be severely affected by climatic extreme
75 (Eckstein et al., 2017). The most immediate impacts of climatic extreme in Bangladesh, as elsewhere
76 globally, are mainly due to increased daily temperatures and temperature-related extreme phenomena
77 (Shahid et al., 2016). Temperature extremes have been observed in the country's agriculture (Sikder and
78 Xiaoying, 2014) and other sectors (Shahid et al., 2016). Temperature variability is most likely to result
79 in significant yield reductions in the agricultural industry in the future (Ghose et al., 2021). Plant
80 development, pollination, and reproductive processes are all affected by higher temperatures (Tank et
81 al., 2006; Sacks and Kucharik, 2011). A short period of unusually high or low temperatures can
82 significantly negatively impact crop growth and yield (Mearns et al., 1984). Due to temperature
83 extremes, total rice production in Bangladesh is expected to decline by 7.4% per year from 2005 to 2050
84 (Sarker et al., 2012). Therefore, significant attempts should be made to estimate not only changes in
85 mean temperature sequence but also changes in the frequency, magnitude, and duration of extreme
86 temperature events (Easterling et al., 2000; Jones et al., 2001; Moberg and Jones, 2005; Alexander et al.,
87 2006). However, the characteristics of climate extremes are poorly understood at the regional level. So,
88 it is urgently needed to keep track of spatiotemporal variations in temperature extremes regularly at the
89 regional level, including Bangladesh.

90 In Bangladesh, many studies were performed to investigate the spatiotemporal variations of extreme
91 climate indices (Hasan et al., 2013; Shahid et al., 2016; Mahmud et al., 2018). These studies reported
92 that warm temperature events (cold temperature events) are rising (decreasing). However, the studies
93 mentioned above are either concentrated on the two regions, and a limited number of meteorological
94 stations (Mahmud et al. 2018), short time scales and a limited number of temperature indices (Khan et
95 al., 2019), different data sources (Hasan et al., 2013). Besides, it is unclear whether and how those
96 indices affect temperature extremes in Bangladesh both spatially and temporally. In addition to this,
97 previous studies have not shown the long-term connection among extreme temperature indices using
98 Detrended Fluctuation Analysis (DFA). To close the aforementioned research gaps, this study has four
99 objectives:

- 100 1. To examine the spatiotemporal trends of temperature extremes from 1980 to 2017 in Bangladesh.
- 101 2. To identify the association among extreme temperature indices.
- 102 3. To predict long-term connection among ten temperature extremes indices.
- 103 4. To analyze the factor affecting the extreme temperature variation over Bangladesh.

104 The novel aspect of this research is that the trends and the associated connection of extreme temperature
105 were analyzed in Bangladesh to understand their spatial and temporal variability. The findings will serve
106 as a scientific foundation for future severe event prediction and hazard mitigation and prevention.

107 **2. Data and Method**

108 ***2.1 Study Area Description***

109 Bangladesh is a sub-tropical country of South Asia situated in between latitude 20° 34' N to 26° 38' N
110 and longitude 88° 01' E to 92° 41' E (Fig. 1) with complex hydro-geologic settings (Jerin et al., 2021;
111 Ghose et al., 2021). Excluding some hilly area of Bangladesh, maximum portions of the land area in the
112 floodplain (80%) region (Rahman and Islam, 2019; Islam et al., 2021). In Bangladesh, the monthly

average temperature ranges from 5.8°C (January) and 35.7°C (August), whereas the monthly average precipitation varies between 1mm (January) and 350mm (July). The annual average temperature is 26°C, and the average yearly rainfall is about 2400 mm (Islam et al. 2020; Jerin et al., 2021). Seasonal variations in rainfall are indistinguishable characteristics of its climate. The country's climate is characterized by a hot and humid summer with heavy rain and a dry and mildly cold winter. Winter (December to February), pre-monsoon (March to May), monsoon (June to September), and post-monsoon (October to November) are the four predominant seasons of the country (Islam et al. 2021; Salam et al. 2020). Bangladesh has an average daily mean relative humidity of 80% and a 3.72 mm/day evapotranspiration rate, respectively (Salam and Islam, 2020). The coldest month is January and the hottest months in Bangladesh between April and October.

2.2 Data source and quality control of the dataset

Daily minimum and maximum temperature data are collected from Bangladesh Meteorological Department (BMD). Though BMD has 43 weather stations across the country, 27 sites were selected for this study because of the lack of availability of long-term temperature data. Thus, the daily Tmin and Tmax data of 38 years from 1980 to 2017 were used in this study with missing values of less than 4%. The selected stations are uniformly scattered all over the country, which is assumed to be a perfect representation for the whole country. Our study was analyzed variations in extreme temperature indices based on daily minimum and maximum temperatures for 1980–2017. Many temperature series were omitted from our analysis due to inhomogeneity. Missing data at each site were filling-up from the records of nearby locations. In addition, the sites discarded due to the unavailability of data for a more extended period was also used to fill up the missing values (Rahman et al. 2019). The BMD followed the World Meteorological Organization (WMO) guidelines for weather data record and collection.

135 Nevertheless, the quality control of the dataset is still imperative before investigating climatic extremes
 136 because incorrect outliers influence the extremes significantly (Gao et al., 2015).
 137 Quality control of site observation was initially done through systematic checking of data, namely,
 138 positive records of climate variables; Tmin is lower than Tmax, and temperature less than 45°C. The
 139 time-series data were identified to be homogeneous and consistent at all locations (Hans, 1986). The
 140 BMD staff also approved all data records through a data quality check. Serial autocorrelation is one of
 141 the critical problems in trend analysis (Praveen et al., 2020). Assessment of serial autocorrelation in time
 142 series for different lags showed correlation at $p < 0.05$, except for a few cases. Data were quality-
 143 controlled by RClimDex1.1.
 144 In this study, 10 extreme temperature indices were chosen from the indicators recommended by the
 145 Expert Team for Climate Change Detection Monitoring and Indices (ETCCDMI). Recently, these 10
 146 indices have been extensively adopted in extreme temperature study (Zhou et al., 2020; Islam et al.,
 147 2021). The indices were computed by using the RClimDex software package developed by ETCCDI
 148 (<http://cccma.seos.uvic.ca/ETCCDI>). Table 1 presents a detailed description of these ten indices.

149 ***2.3 Pettitt's Test***

150 Pettitt's test is a non-parametric test introduced by Pettitt (1979) applied to detect a change point in any
 151 time series data with its significance test (Islam et al., 2020). In this study, Pettitt's test is used for
 152 detecting change points among different extreme temperature indices. This test employed Mann-
 153 Whitney statistic U_t that examines if the two sets of sample $x_1, x_2, x_3, \dots, x_t$ and $x_{t+1}, x_{t+2}, x_{t+3} \dots x_n$ are
 154 from the similar population or not (Mu et al. 2007). The U_t can be expressed as:

$$U_t = \sum_{i=1}^t \sum_{j=t+1}^n \text{sign}(x_i - x_j) \quad (1)$$

$$\text{sign}(x_t - x_j) = \begin{cases} 1, & \text{if}(x_i - x_j) > 0 \\ 0, & \text{if}(x_i - x_j) = 0 \\ -1, & \text{if}(x_i - x_j) < 0 \end{cases} \quad (2)$$

155 The K (test statistic) and ρ (confidence level) for the n (sample length) are defined by Eq. (3-4):

$$K = \text{Max } |U_t| \quad (3)$$

$$\rho = \exp\left(\frac{-K}{n^2 + n^3}\right) \quad (4)$$

156 The null hypothesis is rejected if ρ is lower than the specified significance level. The p (significance
157 probability) can be expressed as:

$$p = 1 - \rho \quad (5)$$

158 **2.4 Detrended Fluctuation Analysis (DFA)**

159 Detrended fluctuation analysis (DFA) is a novel method for assessing the long-term relationship in the
160 non-stationary time series data analysis, which is firstly presented by Peng et al. (1994) for the
161 investigation of DNA. The DFA method can prevent the wrong identification of artificial relationships
162 by eliminating local trends of the different time series (Rahman and Islam, 2019; Islam et al., 2021).
163 Nowadays, this method has been employed mainly to identify long-range associations among natural
164 systems following continuous improvement (Li and Zhang, 2007). In this paper, DFA is used for
165 predicting upcoming trends in extreme temperature indices. The following procedure can calculate it:

166 For extreme climate series $\{x_k, k = 1, 2, \dots, N\}$, In which N represents the series length and x represents
167 mean. The cumulative deviation of the original series can be computed as:

$$y(i) = \sum_{k=1}^n (x_k - \bar{x}) \quad (i = 1, 2, \dots, N) \quad (6)$$

168 Then the latest series y_i is ordered into N_s different sub-intervals through s length by Eq. (7):

$$N_s = \text{int}\left(\frac{N}{s}\right) \quad (7)$$

169 For confirming the information as the series is not divisible, it is categorized again into inverse way. As
 170 a result, a total of $2N_s$ subintervals is created. Then Polynomial fitting is executed for each sub-interval
 171 v ($v = 1, 2, \dots, 2N_s$) data to achieve the trend function $y_{v(i)}$ in that series. The original series trend in sub-
 172 function is filtered out by executing Eq. (8):

$$y_{s(i)} = y(i) - y_{v(i)} \quad i = (1, 2, \dots, N) \quad (8)$$

173 wherein, $y_{v(i)}$ is second order polynomial though it can be first or higher order polynomial. To eliminate
 174 the trend and computing of the each-interval variance is calculated by Eq. (9-10):

$$F^2(v, s) = \frac{1}{s} \sum_{i=1}^s \{y[(v-1)s + i] - y_v(i)\}^2 \quad (i = 1, 2, \dots, N_s) \quad (9)$$

$$F^2(v, s) = \frac{1}{s} \sum_{i=1}^s \{y[N - (v - N_s)s + i] - y_v(i)\}^2 \quad (i = N_s + 1, N_s + 2, \dots, 2N_s) \quad (10)$$

175 The estimation of the second order wave function of the entire series by following Eq. (11):

$$F(s) = \sqrt{\frac{1}{2N_s} \sum_{v=1}^{2N_s} F^2(v, s)} \quad (11)$$

176 Power law relationship series of $F(s)$ and s variations are computed by using Eq. (12):

$$F(s) \sim s^a \text{ or } \ln F(s) = a \ln s + b \quad (12)$$

177 The datasets are close fitted using least square method in dual logarithmic where 'a' (slope) of the trend
 178 is scaled DFA index. The whole procedure is randomly divided and independent if there is $a=0$ and $0 < a$
 179 < 0.5 denotes the short-term relation with dependent process demonstrate that the time series data are
 180 inversed to the prevailing trend. On the other hand, $0.5 < a < 1$ indicates the continuous series and
 181 approaching trend is same to the previous. When 'a' is nearby to 1, the greater the change of similarity
 182 and $a=1$ denotes the procedure is $1/f$ sequence like a non-stationary casual cycle along with $1/f$ spectrum

classified by scale invariance and a long-term relationship. Moreover, $a \geq 1.5$ denotes the procedure is brown-noise sequence.

2.5 Trend analysis using multi-statistical techniques

The nonparametric Mann-Kendall (MK) test is employed to detect a trend in extreme temperature indices (Islam et al., 2019; Islam et al., 2021). It is renowned for its robustness in analysing non-normally sequenced datasets and lower sensitivity to missing value (Islam et al., 2021). A trend free pre-whitening (TFPW) method was employed to remove serial auto-correlation (Praveen et al. 2020; Jerin et al., 2021). Correspondingly, Sen's slope (SS) estimator (Sen, 1968) was employed to determine the frequency of change in extreme temperature indices. The elaborate process of the MK test and SS are found in Islam et al. (2019) and Praveen et al. (2020).

Pearson's correlation coefficient was applied to reveal the relationship among extreme temperature indices. The factor analysis was employed to detect extreme temperature factors across Bangladesh (Rahman and Islam, 2019). The univariate linear regression analysis was used to identify the trend rate of extreme temperatures indices in Bangladesh and each site (Donat et al., 2014). We Used the R studio software to perform the M-K mutation test (Gallant and Karoly, 2010; Panda et al., 2014) and use the student t-test to confirm the mutation change point and, increasing the credibility of the mutation outcome.

3. Result

3.1 Linear Regression Trends Analysis

The rate of the inter-decadal tendency of every single extreme temperatures index was quantified by employing the linear regression technique, and its significance was examined (Figure 2). Based on the time scale, the warm indices of extreme temperature demonstrated an increasing trend except for TXX showed a decreasing trend during 1980-2015. The rate of the magnitude of DTR, SU, TXX, TNX,

TX90p, and TN90p were 8.7e days/decades, 0.21 days/decades, 0.037 oC/decades, 0.01 oC/decades, 0.11 days/decades, and 0.12 days/decades, respectively (Figure 2), denoting that the warm indices of extreme temperature are rising significantly in Bangladesh. Among them, the rising rate of DTR and TNX were relatively shorter during the study period 1980 to 2015. At the same time, the increasing rate of other indices was comparatively larger, mainly the rising rate of SU, which was the fastest, whose tendency of inter-decadal magnitude rate attained 0.21 days/decades (Figure 2b). As shown in Figure 2, except for TXX, the rest of the warm indices demonstrated an increasing trend from the 1990s to 2000, representing that there can be a warming occurrence in this period. DTR, SU and TX90 were changed smoothly before 2010, while abrupt upward change started since 2010 (Figure 2a, b, f). TN90 showed a comparatively short variation before the 1990s; after that, it had large fluctuations and showed an upward trend (Figure 2e). TNX exhibited changes that started after 1995; before 1995, the fluctuation was short (Figure 2h). The overall decreasing trend showed in TXX during the total study period, and two significant changes were in the 1990s and 2010s (Figure 2j). Based on the discussion mentioned above, the warm indices of extreme temperature exhibited an overall upward trend in the study period except for TXX in this research. Since 1995, the warm indices showed a rapid upward trend in line with the global warming trend.

The cold indices of extreme temperatures TNN, TXN, TN10 and TX10 all exhibited decreasing trend, and the rate of magnitude were -0.019 oC/decades, -0.049 oC/decades, -0.15 days/decades and -0.078 days/decades, respectively. This demonstrates that these cold indices had a significant decreasing trend, of which the TN10 had the biggest decline and TNN had the most minor decrease that was also a certain tendency of global warming. Figure 2c shows that the TN10 significantly fluctuated before 2010, and from 2011 to 2015, the fluctuation was extremely short. TX10 demonstrated a decreasing trend where significant fluctuation had started since 1985 (Figure 2d). In Figure 2g, the TNN had the most noticeable

change and fluctuation from the 2000s, while the most significant fluctuation was after 2010. The temporary fluctuation showed from 1980 to 1995 while the most considerable fluctuation and variation exhibited after 1995 with decreasing trend in TXN (Figure 2f). By observing the above assessment, TX10, TN10, TXN and TNN all exhibited a downward trend; notably, TX10, TN10 showed a rapid decline speed.

3.2 Long-Term Connection analysis Using DFA

To explore the trend behaviours of fluctuation in forthcoming temperature extremes, a long-term association assessment was conducted employing DFA, and the results are demonstrated in Figure 3a – j. The Figure 3 illustrates that the DFA ranging exponent in DTR, SU, TN10, TX10, TN90, TX90, TNN, TNX, TXN, and TXX was 0.96, 0.65, 0.86, 0.38, 0.76, 0.85, 0.63, 0.71, 0.75 and 0.86 respectively. In the study period 1980 to 2015, they all had a strong association, showing that the forthcoming behaviours of trend in each single warm and cold temperature extremes are similar to the fluctuation trend. Indeed, the warm temperature extremes will rise, and cold temperature extremes will decline in the forthcoming period. The rate of change in the warm temperature indices (SU, TNX, and TN90) will increase continuously than the other warm indices. The DFA values of these warm indices are not closer to 1. In cold temperature indices, TN10 will continue to decrease than the other cold temperature indices except for TX10 due to its DFA value being larger than the other cold indices. We explored that the sequence values of the DFA exponent of warm indices (SU, TNX, TN90) are less than the cold temperature extreme (TN10). It demonstrates that the forthcoming trend pattern in temperature extreme has a strong long-term association with the current state and forthcoming trend pattern. By contrast, the DFA exponent value of the warm indices is higher than DTR, TXX, TX90, while in cold indices, the exponent value of DFA is lower than TNN, TXN and TX10.

251 In Table 1, the correlation matrix of 10 temperature indices has been exhibited. We assessed the
252 significant associations between the Warm and cold indices ($p < 0.01$) in this current study. There is a
253 strong significant negative relationship between warm indices and cold indices. In contrast, the warm
254 indices SU, DTR and TXX have a significant positive relationship with cold indices TNN, TXN and
255 TN10, respectively, where their coefficient was 0.395, 0.389 and 0.348, respectively ($p < 0.01$). The most
256 significant positive relationship is between TX90 and SU, where the coefficient was 0.756 ($p < 0.01$). The
257 warm temperature indices were a significantly positive association with one another. On the other hand,
258 the significant negative relationship has shown in cold temperature indices, while the relationship
259 between TXN and TNN has a significant positive association ($p < 0.05$). The index selected in this
260 present work can have good indicators of climate warming over Bangladesh.

261 *3.3 Spatial patterns of extreme temperature indices*

262 The spatial changes in temperature extremes are exhibited in Figure 4. Nearly -0.81 to 1.12 days/decade
263 was the change rate of DTR where southern and western parts had an increasing trend, and at the same
264 time, another part faced the decreasing trend while 58% and 29% of areas under decreasing and
265 increasing trends, respectively (Figure 4a). In Figure 4b, 65% of sites remain under the increasing trend
266 of SU, and the change tendency rate was -0.93 to 1.44 days/decade; the southern part demonstrated
267 decreasing tendency. TN10 had a decreasing trend in the south and south eastern regions, where 84% of
268 areas occupied the decreasing trend (Fig. c). The trend change rate of TN10 was -0.5 to 1.22
269 days/decade in Figure 4c. The trend rate of TX10 was -1.37 to 0.3 days/decade, decreasing trend found
270 in northern and southern parts, and 23% of areas remain under increasing direction (Figure 4d). The
271 change rate of TN90 was -0.25 to 1.63 days/decade, and the northern and southern parts met the
272 increasing trend where 50% of areas were under increasing trend (Figure 4e). Figure 4f exhibits that
273 TX90 had -1.6 to 1.45 days/decade change rate and 67% area under increasing and 25% area under

decreasing rate where western and southern part faced increasing trend and the other part met decreasing trend. The change rate of TNN was -1.50 to 1.51 days/decade, with 69% of station stayed under increasing tendency in the northern, western and southern regions (Figure 4g). TNX had a -1.14 to 1.50 rate of change; 80% of areas demonstrated a declining trend in the western and southern region (Figure 4h). In Figure 4i, the change rate was within -1.44 to 1.15 days/decade for TXN, and the higher decreasing rate was in the southern region; 73% station was under decreasing trend. The variation rate in TXX varied from -1.41 to 1.03 days/decade, where 56% of areas show a declining pattern in the southern part of the study area (Figure 4j). The trending behaviour of extreme temperature indices varied spatially, further assured by temporal distribution in the earlier part.

3.4 Factor analysis of temperature extreme indices

Factor analysis was carried out to detect the most influential factor affecting extreme temperature indices. F1, comprising TN90, SU and TNX indices of temperature, the overall variance of temperature data is 40.76%, which denotes the warm night primarily controls the general warming trend of temperature extremes over Bangladesh from 1980 to 2015 (Table 3). Daily temperature range (DTR), SU, and TX90 control F2 calculated 22.40% of the whole variance, ensuring the rising of DTR is the principal factor that affects changes in yearly temperature. The DTR reflects the association between the minimum and maximum temperatures. About 11.83% of the total variance accounted for F3, where Min Tmax (TXN) dominates the third factor. For F4, the coldest days (TNN) and warmest days (TXX) are the major dominant factors that represent the overall variance of 9.49% and 4.88%, indicating the warming climate has occurred daily minimum and maximum temperatures to rise significantly. In most cases, warm and cold extremes have a significant statistical association with the average annual temperature (You et al., 2011).

3.5 Variation of probability distribution functions of temperature extreme

Due to the variations in temperature extremes, the relative indices were chosen to explore the probability distribution function (PDF) of extreme temperature phenomena (Figure 5) in the previous four decades. The frequency has declined in cold extremes while the frequency has increased in warm indices in a recent decade over Bangladesh, which also is consistent with temporal trend analysis. The cold extremes were moved negatively (Figure 4c, 4e, 4g and 4i) while warm extremes (Figure 4a, 4b, 4d, 4f, 4h and 4j) positively shifted to their upper trails. The movement of negative and positive of the curve of PDF denotes that frequency is declined of cold extremes and increased in frequency of warm extremes. It is worth mentioning that the variations of the PDF curve are more evident in recent decades than in the past three decades. In warm indices, the highest peak was found in 1990 to 1999 and 2000 to 2009 for DTR, SU and TXX. The recent two decades were for TNX, and 1980 to 1989 was for TN90, and 1990 to 1999 was for TX90. In cold indices, the maximum peak was found in the decades of 2000 to 2009 for TN10, TX10 and TXN and 1980 to 1989 was for TNN.

3.6 Rate of change analysis using Sen's Slope Estimation

The ordinary non-parametric system improved by (Zaiontz 2020) was employed to account for the slopes present in the rate of trend using the Sen's slope estimators (Fig. 6). The positive mark represents the increasing slope, and the declining slope denotes the negative impact. Figure 6a shows that the lowest negative Sen's slope value was -0.764, primarily found in the northeastern part, and the highest value was found in the southeastern region. The highest positive Sens value was 8.821 for SU (Figure 6b), which generally exhibited in the northeastern part, while -3.019 was the lowest negative value for TN10 (Figure 6c) which maximum distributed in the northern region of Bangladesh. In Figure 6d, the lowest negative value was -2.279 for TX10, primarily found in Bangladesh's northeastern and southern parts. The north and northeastern parts of Bangladesh have distributed the highest Sen's slope value 3.649 for TN90 (Figure 6e). The highest positive value was 5.312 for the TX90 (Figure 6f), which

320 mainly occupied southern and northern parts. The lowest value was -0.906 for the TNN (Figure 6g),
321 which was commonly found in the country's southern region. Figure 6h exhibits that the northeastern
322 part met the lowest negative value of TNX, which was -0.089. For TXN (Figure 6i), the lowest negative
323 value was -1.4, mostly occupying the northern part of Bangladesh. The highest positive value was 0.406,
324 which was distributed in the southern part of Bangladesh.

325 ***3.7 Mutation Analysis of Extreme Temperatures Indices***

326 The M-K change point detection analysis was performed for each extreme temperature index, and the
327 change year was selected by combining the sliding t-test technique (Table 4). Table 4 shows that overall
328 Bangladesh, the warm indices (DTR, SU, and TX90) and the cold indices (TNN and TX10) did not
329 change, but the remaining warm (TN90, TNX, TXX) and cold indices (TN10, and TXN) were changed
330 significantly ($p < 0.05$). The station-wise warm and cold indices were muted. For warm index, in different
331 stations, the points of change happened in the last 20th and the beginning of the 21st century where
332 stations like Patuakhali, Rangpur for SU, Mymensingh, Patuakhali and Sitakundu for TXX, and Bhola,
333 Jashore, Patuakhali and Mymensingh for TX90 met the change point in the 1980s and later in 2008. In
334 cold indices, their mutation point was also the beginning of the 21st and last of the 20th, where
335 Rangamati and Sitakundo for TNN, Chattogram and Teknaf for TN10 faced the mutation point after
336 1987 and 2005. The points of mutation of all indexes of temperature extremes passed 0.05 level of test
337 of significance. It can be observed that the mutation point is varied from region to region. In those
338 stations, warm indices such as TXX and TX90 were muted in 1986, 1987 while cold indices including
339 TNN and TN10 were muted in 1988 and 1989. Overall, the mutation points of warm indices were before
340 cold indices and other indices. Similarly, it implied that the warm indices were more sensitive to global
341 climate warming than cold indices.

342 **4. Discussion**

343 Extreme weather occurrences have drawn attention to the massive dominance of extreme climate
344 changes on life, nature and human production in recent years (Weaver et al. 2014; Li et al., 2013). Due
345 to getting how extreme climate influences the natural environment and society. It is significantly
346 essential to assess climate extremes' spatial and temporal change trend. In this present study, we
347 evaluated the temporal and spatial changes in temperature extremes in Bangladesh in the previous 36
348 years (1980-2015). The results exhibited that the trend of warm indices increased while TXX showed
349 decreasing trend and the cold indices are meeting decreasing trend. This result is consistent with Li et al.
350 2020, where they explored that the warm indices including WSDI, TR20, SU25, Tx90p and Tn90p are
351 increasing, and the cold indices such as CSDI, ID0, FD0, Tx10p and Tn10p are decreased. Zhou et al.
352 (2020) explored those warm indices in China showed an increasing trend while cold indices exhibit a
353 declining tendency except TXN and TNN, which agrees with the present study. Sun et al. (2015)
354 investigated that cold extremes are significantly declining in China, similar to the current research. The
355 warm occurrences are increased and decreased considerably the cold events stated by Ren et al. (2010),
356 which is in line with the present study. The warm days are increased by 0- 0.3 days/ decade in china (Shi
357 et al. 2018). Jiang et al. (2016) assessed that the cold indices TX10 and TN10 are declined whether the
358 warm indices TX90 and TN90 increase in the Tibetan plateau. Cold extremes and warm extremes
359 exhibited downward and upward trend, respectively, in the assessment of Yu and Li (2016), which is
360 similar to the current research findings. We have explored the correlation between warm indices and
361 cold indices. Warm and cold indices with each other were investigated. Those warm temperature indices
362 (SU, TNX, and TN90) will be increasing in future than the other warm indices, while cold indices TN10
363 will decline in coming days than other cold indices by using the DFA exponent.

364 The warm and cold extremes were significantly negatively correlated except for the relationship
365 between SU and TNN, DTR and TXN, TXX and TN10. This outcome is consistent with Yu et al.

(2015). A similar study has also done by You et al. (2013); You et al. (2011), Zhai and Pan, (2003). DTR is increasing in the southern and western part of Bangladesh for the SU, and southern part met the decreasing tendency, south and southeastern part demonstrating the decreasing trend for TN10, TX10 met the decreasing trend in the south and northern part. TN90 show the increasing trend in the north and southern part of Bangladesh, south and western region met increasing trend for TX90. Approximately 69% of the northern, western and southern parts of Bangladesh met increasing tendency of TNN, -1.14 to 1.50 was the change rate of TNX and different parts of the country were faced decreasing. A rising trend, southern part experiences the decreasing trend for TXN and TXX analyzed by spatial distribution in this current research work. Similar research work is also done by Zhou et al. (2020) where they found that the warm and cold indices are facing increasing and decreasing pattern in different parts of the study area. Li et al. (2020) found that TNN and TXX spatially varied, which agrees with the present study's outcomes. Some studies performed in North-Eastern regions of India are in good agreement with our finding (Jhajharia et al. 2014). Dabral et al. (2016) indicated that minimum temperature is rising in the north-eastern part of India which is quite dissimilar to this finding. Results indicate that the frequency and intensity of warm night is enhancing in the North-Eastern region of India (Jhajharia and Singh 2011; Dabral et al. 2016), similar to our study.

Yu and Li (2014) investigated the spatial distribution of temperature indices in line with the current study's findings. The almost same study conducted by Zhao and Chen (2021), You et al. (2011), Nie et al. (2012). The present study assessed the principal factor, which was DTR, and it also influences the variation of annual temperature, and its total variance was 22.40%. This outcome is consistent with Yu and Li 2014 where they found that SU25 is the principal factor of yearly temperature change. You et al. 2011 also explored similar findings. By using the probability distribution function, it discovered that warm indices are shifted positively. Cold indices are moved negatively, which further indicated that the

389 warm indices are following an increasing trend while cold indices are decreasing trend and the density
390 peak was different for each temperature extremes. The variations were most remarkable in the recent
391 decades in this research study. This result is almost similar to Yu and Li (2014). Liu et al. (2021) found
392 that the temperature extremes varied differently in different regions while the significant variations were
393 in recent decades, consistent with the present study. The outcome of Fu and Ding (2021) is in line with
394 the current research.

395 In this study, we have analyzed the change point of temperature extremes where we found that the
396 change point of warm and cold indices was last of the twentieth and beginning of the twenty-one
397 century. Zhou et al. (2020) investigated that the mutation point for warm indices was beginning in the
398 twenty-first century, which is in good agreement with the present research work. Liu et al. (2021)
399 explored the change point of temperature extremes. By contrast, the daily maximum temperature altered
400 quicker than the daily minimum temperature. The main reason is the probable urbanization impact on
401 the extreme temperature indices (Zhou and Ren, 2011). The urbanization impact may be aggravated by
402 extreme temperature trends associated with daily maximum temperature in Bangladesh. However, the
403 rapid urbanization impacts of extreme temperature indices related to daily minimum temperatures were
404 generally trivial (Duan et al., 2012).

405 Therefore, the increasing pattern of warm indices and decreasing cold indices have already influenced
406 vegetation succession, soil and farm production over Bangladesh. There will be more intense
407 consequences and uprising extreme weather occurrences and temperature forecasted in the coming days.
408 The achieved outcomes have scientific and practical implications, which assist policy-makers in
409 developing suitable measures to safeguard vegetation change and lessen the detrimental impacts
410 triggered by extreme climatic phenomena. Our study aims to understand extreme climatic phenomena in
411 Bangladesh. Environmental factors such as soil moisture content, potential evapotranspiration, pan

412 evaporation and land use/land cover changes should be considered in future studies. The extreme
413 temperature influences vegetation dynamics in future based on CMIP6 datasets in Bangladesh deserves
414 further investigation.

415 **5. Conclusion**

416 This paper aims to detect temporal and spatial trends of extreme temperatures in Bangladesh in the past
417 37 years (1980–2017) using multi-statistical modeling approaches. The results show that warm
418 temperature indices in Bangladesh were increasing in the last four decades except for the warmest days
419 (TXX), especially before the 2000s, while cold temperature indices showed a decreasing trend after
420 2000. DTR and SU showed an upward trend in the country, indicating that the rising rate of daily
421 maximum temperature in the past four decades was more than that of daily minimum temperature. This
422 confirms that Bangladesh is uninterruptedly developing towards a warmer trend, which is a negative
423 response to global climate warming. Spatially, the change rate of warm indices was the largest in the
424 northwest region and the smallest in the eastern part. Besides, the change speed of DTR at each station
425 was less than that of other indices, whose change tendency rate was the smallest in the central region
426 and the decrease speed was more prominent in the northern part. The outcomes of DFA showed a long-
427 range association among extreme temperature indices, implying that warm and cold indices will
428 continue their present trend in the upcoming years, except for cold days will not sustain their current
429 trend in the future. By using the probability distribution functions, the variations of the curves are more
430 evident in recent decades than in the past three decades. Based on factor analysis, the warm night
431 primarily controls the general warming trend of temperature extremes over Bangladesh from 1980 to
432 2017. Mutation analysis revealed that the mutation points of the warm index were before cold index and
433 other indices, indicating that warm indices were more sensitive to global climate warming than cold
434 indices. The increase in the warm indices and the decrease in the cold indices have eventually impacted

435 agricultural crop production, soil fertility and vegetation dynamics in northwest Bangladesh. Rising
436 temperature and the increase in extreme climatic phenomena forecasted in the future will have a more
437 intense impact on sustainable development. Thus, to confirm the country's sustainable development,
438 governments at all levels should systematically take adequate countermeasures based on climate change
439 characteristics and constantly develop their capacity to cope with extreme climatic phenomena.

440 **Acknowledgement**

441 The authors extend their appreciation to the Deanship of Scientific Research at King Khalid University
442 for funding this work through Research Group under grant number (R.G.P.2 /194/42). We are grateful to
443 the Department of Disaster Management, Begum Rokeya University, Rangpur for all sort of assistant
444 provided during this study. Furthermore, we would like to thank the Bangladesh Meteorological
445 Department (BMD) for providing required data for this research.

446 **Ethical approval**

447 Not applicable

448 **Consent to Participate**

449 Not applicable

450 **Consent to Publish**

451 Not applicable

452 **Data availability**

453 Data are available upon request on the corresponding author

454 **Author contributions**

455 A.R.M.T.I., J.M., and H.M.T.I., B.G., designed, planned, conceptualized, drafted the original manuscript, and
456 H.M.T.I, and Y.R., were involved in statistical analysis, interpretation; H.M.T.I., Y.R., and J.M., contributed
457 instrumental setup, data analysis, validation; Z.H., and S.A.B., contributed to editing the manuscript, literature
458 review, proofreading; B.G., J.M., Z.H., and A.R.M. T.I., were involved in software, mapping, and proofreading
459 during the manuscript drafting stage.

460 **Conflict of interest**

461 There is no conflict of interest to publish this work.

462 **Funding statement**

463 The authors extend their appreciation to the Deanship of Scientific Research at King Khalid University for funding this
464 work through Research Group under grant number (R.G.P.2 /194/42).

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Figures

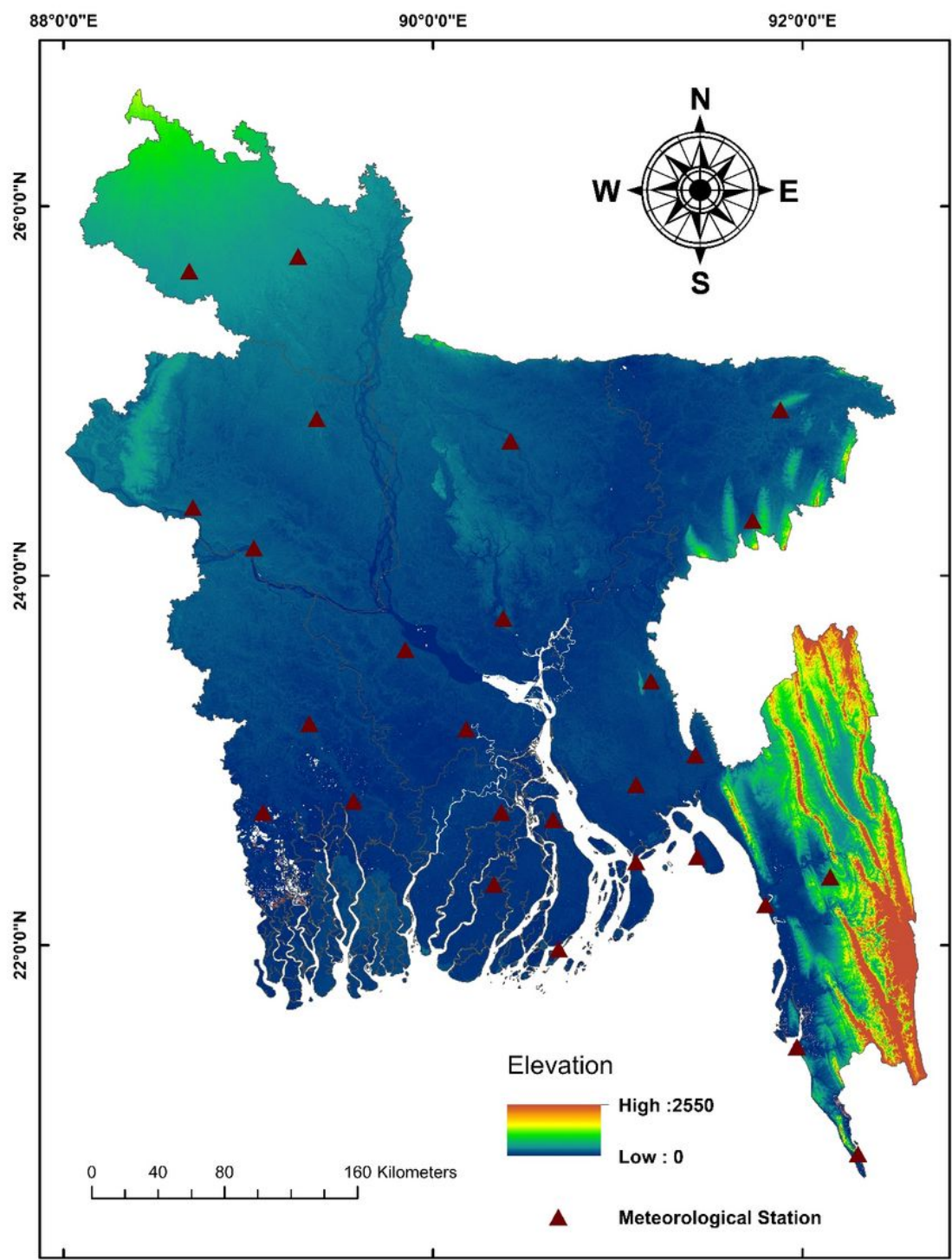


Figure 1

Geographical location of the study area showing meteorological sites

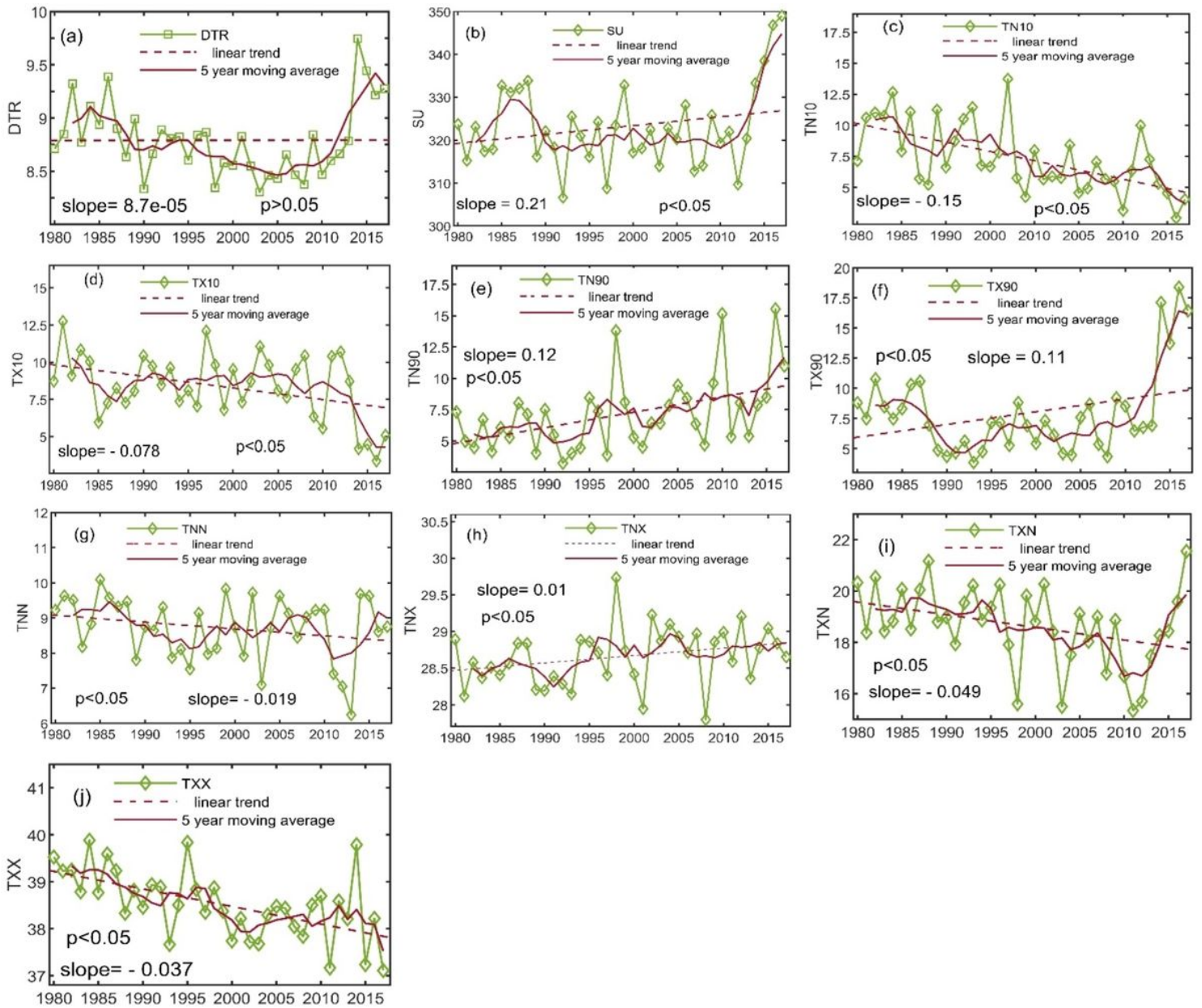


Figure 2

Linear regression trends for 10 extreme temperature indices during 1980-2017 over Bangladesh

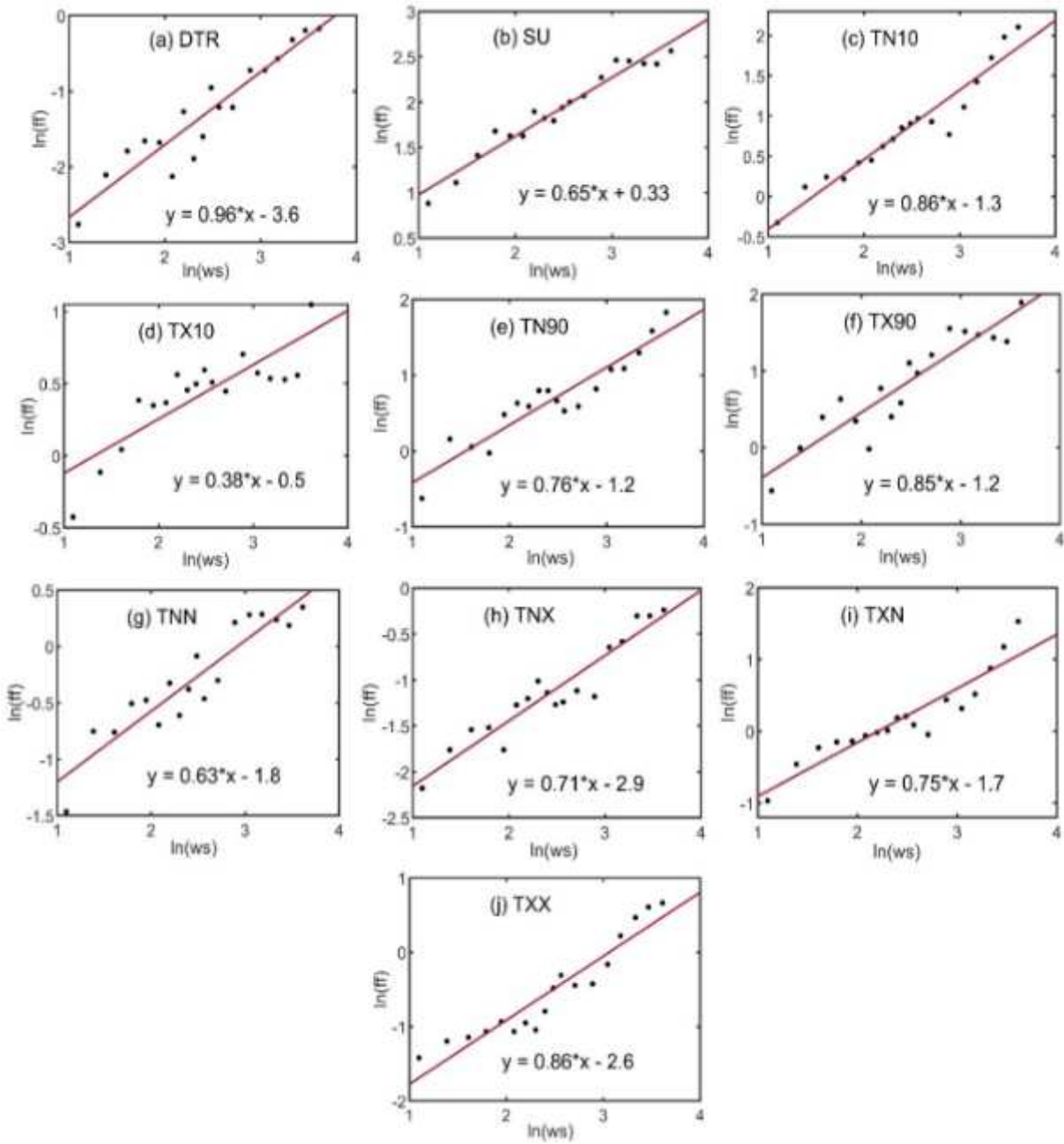


Figure 3

DFA long-term relationships of extreme temperature indices in the forthcoming period across Bangladesh

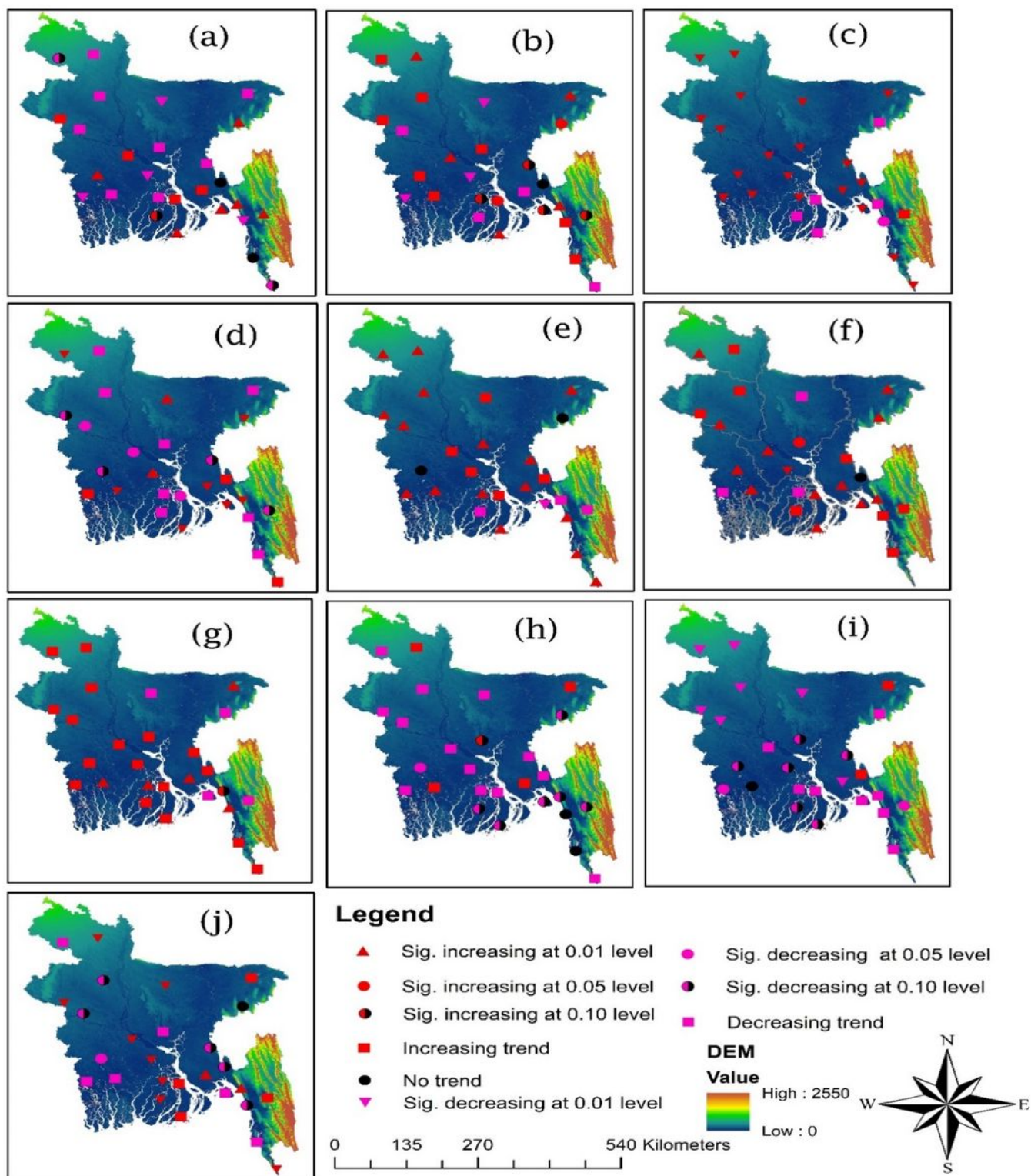


Figure 4

Spatial distribution patterns of decadal trends in 10 temperature indices over Bangladesh during 1980 to 2017.

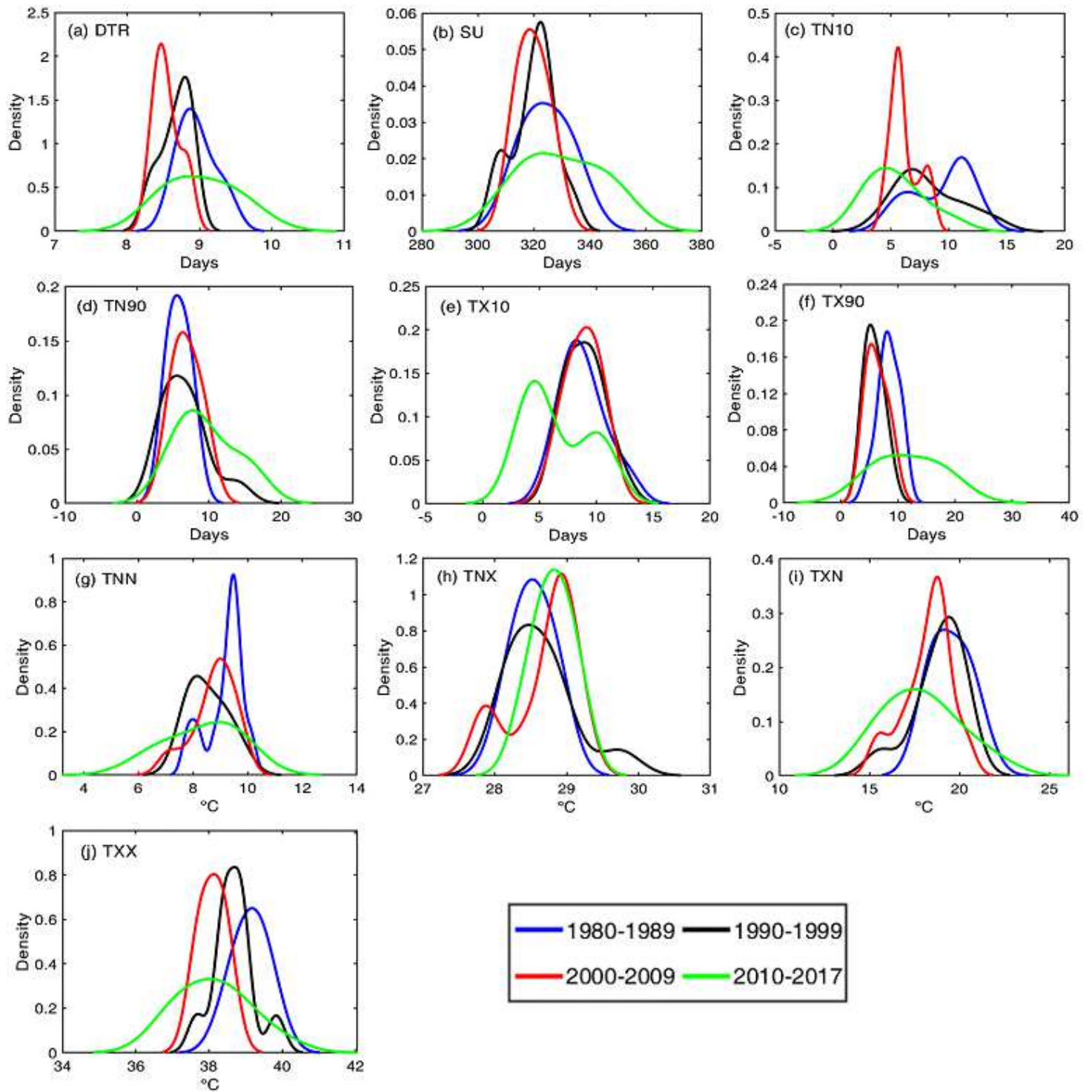


Figure 5

The probability distribution function (PDF) of extreme temperature indices in the last four decades in Bangladesh

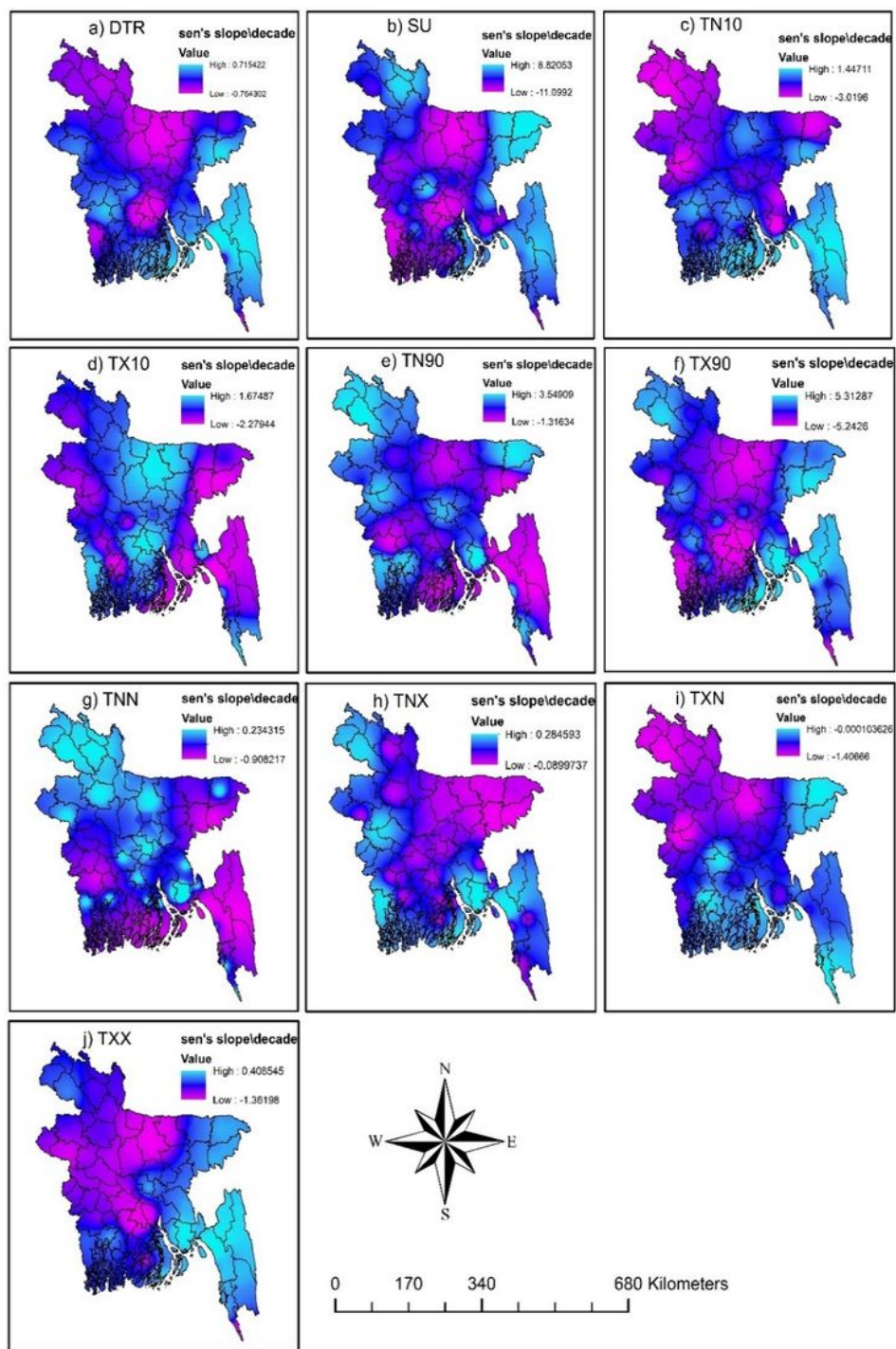


Figure 6

The rate of trend pattern of extreme temperature indices from 1980-2017 in Bangladesh