

A Superior Seismic Data Preprocessing Technique to Improve the Resolution of Surface wave Velocity Spectrums

Tarun Naskar (✉ tarunnaskar@iitm.ac.in)

Indian Institute of Technology Madras

Mrinal Bhaumik

Indian Institute of Technology Madras

Sayan Mukherjee

Indian Institute of Technology Madras

Research Article

Keywords: high-resolution surface wave velocity spectrum, dispersion image, seismic survey

Posted Date: July 15th, 2021

DOI: <https://doi.org/10.21203/rs.3.rs-712071/v1>

License:  This work is licensed under a Creative Commons Attribution 4.0 International License.

[Read Full License](#)

TITLE PAGE

A superior seismic data preprocessing technique to improve the resolution of surface wave velocity spectrums

Tarun Naskar^{1*}, Mrinal Bhaumik¹⁺ and Sayan Mukherjee¹⁺

¹Indian Institute of Technology, Madras-600036, India.

Email: tarunnaskar@iitm.ac.in *

⁺these authors contributed equally to this work

A superior seismic data preprocessing technique to improve the resolution of surface wave velocity spectrums

Abstract

A high-resolution surface wave velocity spectrum, also known as dispersion image, is of paramount importance for any seismic survey to accurately predict subsurface earth's properties. The presence of diversified noise in the field acquisition and dissimilar attenuation due to mechanical and radial damping makes it challenging for different wavefield transformation algorithm to produce a detailed and precise velocity spectrum. Standard seismic data preprocessing technique like trace normalisation or bandpass filter fails to address all issues appropriately. Here we have presented a new superior preprocessing technique that can eradicate most of the shortcomings adequately. Experimental field data and published results are used to demonstrate the accuracy of the proposed method. The proposed method also found to produce superior results when compared against the popular commercially available software package Surfseis 6. Overall, the proposed method improves the quality of the velocity spectrum significantly, and it produces a sharper dispersion image even for the extremely noisy data. The work presented here enhances our ability to interpret the surface wave data precisely and help explore accurate properties of the subsurface earth.

Introduction

Surface waves such as Rayleigh or Love waves get generated in the elastic half space from disturbances due to natural events like earthquake, tsunami, landslide etc., or from disruption due to man made events like blasting, construction activities and vehicular movements. The Rayleigh wave attenuates much slowly compared to body waves but carries almost 67% of the source's input energy¹. Therefore, surface waves are an excellent choice for exploring the earth's properties up to a shallow depth. The surface waves based tests are

25 relatively simple, non destructive in nature, cost-effective and considerably faster when
26 compared to conventional testing methodologies like SPT, seismic cross-hole, up-hole and
27 down-hole test etc. The surface wave testing technique can be divided into three major parts:
28 namely, i) the acquisition of experimental field record with the help of geophones, ii)
29 processing the field data to obtain the experimental dispersion image, and iii) performing
30 inverse analysis to predict the subterranean earth's properties²⁻⁴. Many wavefield transforms
31 available to generate high resolution multimodal dispersion image from the field seismogram
32 data with varying degree of accuracy. The two dimensional Fourier transform or frequency-
33 wave number ($f-k$) transform⁵, the slant stack transform or time intercept-phase slowness ($\tau-p$)
34 transform⁶, Park's wave field transform or phase shift method ($\omega-c$)³ and high-resolution linear
35 Radon transform (HRLRT)⁷, are few prominent ones. Park's $\omega-c$ method is currently the most
36 popular choice among researchers due to its high accuracy³ and ability to generate relatively
37 high-resolution dispersion plots especially when limited numbers of traces and offset range is
38 used^{3,8-10}. The $f-k$ transform and, to an extent, the $\tau-p$ transform has broader application and
39 are readily available as an inbuilt toolbox among all the coding platforms. Past researchers have
40 demonstrated the phase-shift method's superiority over $f-k$ and $\tau-p$ transform^{3,8,9}. Even after
41 applying different processing techniques, $f-k$ and $\tau-p$ transform failed to produce a velocity
42 spectrum on par with the phase-shift method^{3,8}. A relatively recent development is high-
43 resolution linear Radon transform (HRLRT) which reported producing a relatively sharper
44 velocity spectrum than the phase shift method^{7,11,12}. However, the HRLRT is a computationally
45 expensive method and it also struggles when poor quality data is used. To this end, we propose
46 a superior seismic data preprocessing technique that can be easily implemented for any
47 wavefield transformation method. The potential outcomes are, (i) the proposed technique
48 normalises each traces against dissimilar attenuation due to mechanical and radial damping,
49 (ii) superior handling of field data contaminated by noise, and (iii) it eliminates the primary

50 shortcomings associated with different wavefield transformation algorithm, thus produces
51 superior velocity spectrum. The proposed technique is simple, and it can be easily
52 implemented. To illustrate the efficacy of the proposed technique, experimental and published
53 results were compared. The proposed method produced a significantly better result for the $f-k$,
54 τ - p method even for the extremely noisy data. Furthermore, using HRLRT, the proposed
55 method produced a superior result against the latest version of commercially available software
56 Surfseis 6, which uses its own processing tool and traditional trace normalisation. The proposed
57 scheme unable to improve the quality of the dispersion image for the phase shift method, which
58 is known to be insensitive towards the data preprocessing technique⁸.

59 **Improvement for $f-k$ and τ - p method**

60 The present consensus among geophysicists is that the ω - c method generates far
61 superior dispersion image compared to $f-k$ transform and τ - p transform. Against this long-
62 cherished belief, the proposed preprocessing technique enables the $f-k$, τ - p method to produce
63 a dispersion image on par with the ω - c method. To demonstrate the effectiveness of the
64 proposed preprocessing technique, velocity spectrums published by Dal Moro *et al.*⁸ are
65 compared with velocity spectrums obtained using the proposed approach (Fig. 1). The original
66 seismogram was digitised to reproduce the published results, and the velocity spectrums were
67 reconstructed using our own Matlab code. Dal Moro *et al.*⁸ compared the efficacy of the
68 different data preprocessing techniques such as trace normalisation, high pass filter, lowpass
69 filter and bandpass filter using the three most popular wavefield transformation algorithms,
70 namely, $f-k$ transform, τ - p transform and Park's ω - c method. They had used 24 channel
71 common receiver shot data with a source to first sensor distance of 5m and sensor spacing of
72 2m. Dal Moro *et al.*⁸ observed that the ω - c method outperformed the other two wavefield
73 transform algorithm and generated the best quality velocity spectrums in terms of accuracy. In

74 Fig. 1 we have compared the result published by Dal Moro *et al.*⁸ with the result obtained using
75 the proposed preprocessing technique. For $f-k$ and $\tau-p$ transform, the proposed method
76 significantly improves the quality of the velocity spectrum. By implementing preprocessing
77 technique proposed in this paper, both $f-k$ and $\tau-p$ transform produced velocity spectrums on
78 par with Park's $\omega-c$ method. Furthermore, Dal Moro *et al.*⁸ reported that compare to $\omega-c$
79 method the $f-k$ and $\tau-p$ method suffers from spatial aliasing and noises, which leads to
80 ambiguous interpretation. Using the proposed approach, we have observed no sign of such
81 spatial aliasing and noises are relative low. Dal Moro *et al.*⁸ also concluded that Park's phase
82 shift transform is poorly sensitive to any data preprocessing technique. For Park's phase shift
83 algorithm, trace normalisation and the proposed approach produced identical velocity spectrum
84 (Fig 1a,b). This outcome is in accordance with the observation made by Dal Moro *et al.*⁸

85 **Superior preprocessing of noisy data**

86 In the seismic field data acquisitions, noise contamination is an unavoidable issue that
87 can occur due to low energy sources, non-source generated waves, defect in sensors, faulty
88 connections etc. Each of these issues will affect the individual traces and corresponding seismic
89 data uniquely. The proposed method can effectively preprocess the noisy data better than any
90 methods currently in use. To demonstrate the efficacy of the proposed approach, 48 channel
91 common receiver shot gathered with the help of 4.5 Hz vertical geophones was employed. The
92 source was kept at a distance of 2.5m from the first sensor and 0.5 m sensor spacing was used.
93 Different types of field noises are added on 12 randomly selected traces to represent assorted
94 noisy field acquisitions on the ground. The velocity spectrum for this noisy acquisition is
95 generated using the $\tau-p$ method with the help of the proposed preprocessing technique and trace
96 normalisation technique. Finally, velocity spectrums for original and noisy data are compared
97 (Fig. 2). As expected, the proposed method produced a superior result than the trace
98 normalisation method for both initial acquisitions and noisy data. However, it is noisy data

99 where the proposed method notably outperforms the trace normalisation method (Fig. 2c,d).
100 For noisy data with trace normalisation, only a few modes are visible and the modes are not
101 clearly distinguishable. On the other hand, the proposed method produced a velocity spectrum
102 more or less similar to velocity spectrums without noise. Almost all the seven modes are visible
103 and they are clearly distinguishable. The dispersion image created using the proposed method
104 for noisy data is marginally superior to the dispersion image produced using trace normalisation
105 for original data without noise. From the dispersion image produced using the proposed
106 method, the fundamental mode and first mode can be separately identified, while it is not
107 possible in the case of trace normalisation. The results indicate that the proposed technique is
108 promising. It is worth mentioning that no frequency filtering technique such as low pass, high
109 pass or bandpass filter is used to obtain the present results.

110 Similar results can be obtained for the f - k and HRLRT method, but we have not included
111 them in the present manuscript due to the limitation on the number of display items.

112 **Improvement for HRLRT method**

113 The previous sections' results are a clear indicator of the significant improvements over
114 the conventional data preprocessing technique. In addition to these popular algorithms, we
115 tested our method against the latest commercially available software package, Surfseis 6. The
116 Surfseis is one of the most popular and frequently used software and 6 indicates the latest
117 release which incorporates the HRLRT algorithm and trace normalisation. The HRLRT
118 method is arguably the most advanced wavefield transform and it generates the highest quality
119 velocity spectrum^{7,11-13}. Using our experimental data, we have compared the proposed data
120 processing technique's efficacy for the HRLRT method. Fig. 3a represents the velocity
121 spectrum produced using Surfseis 6 with HRLRT and trace normalisation, Fig. 3b illustrates
122 the velocity spectrum produced by HRLRT and trace normalisation using our Matlab code and

123 finally, Fig. 3c represent the velocity spectrum produced by HRLRT with the proposed method
124 using our own Matlab code. The proposed preprocessing technique produced the best quality
125 velocity spectrum, while Surfseis produced relatively an inferior quality velocity spectrum. The
126 velocity spectrum produced by Surfseis has lots of noise, especially at higher frequencies and
127 between 0 to 500m velocity ranges. Moreover, the different propagation modes are not clearly
128 identifiable and separable. Our HRLRT code using trace normalisation produced a relatively
129 better quality velocity spectrum compared to Surfseis; with no visible noises present, different
130 modes are identifiable and separable except a few. Our HRLRT code with the proposed
131 preprocessing technique produced the best quality velocity spectrum. There are no visible
132 noises present; all the modes are clearly identifiable and separable. It should be noted that
133 Surfseis software does not display different preprocessing technique they have used while
134 implementing the HRLRT algorithm. For Surfseis, we have used the default integrity factor
135 value of 2 as it produced the best possible results.

136 **Conclusions**

137 A new seismic data preprocessing technique is proposed to process noisy field data and
138 compensate for the losses due to mechanical and radial damping. The accuracy of the proposed
139 approach has been demonstrated using published results as well as using experimental data.
140 The present method works well with all the available wavefield transform algorithms except
141 ω - c method. In agreement with the literature, the present study found that the ω - c method is
142 poorly sensitive to the data preprocessing technique. It is noted that the proposed approach
143 improves the quality of the dispersion image significantly when compared to a traditional
144 preprocessing technique like trace normalisation. The present method produced a sharper
145 velocity spectrum containing multiple easily separable modes with little to no noise. The
146 proposed method performs better than the latest commercially available software package,
147 Surfseis 6. Overall, the seismic data preprocessing technique proposed in this paper will help

148 in producing sharper dispersion image and predicting accurate layer properties of subsurface
149 earth.

150 **Methods**

151 Although the traditional trace normalisation improves the dispersion image quality by
152 compensating losses due to radial damping, it can not compensate for the mechanical damping.
153 The mechanical damping tends to dissipate the energy through molecular friction inside the
154 material. Unlike radial damping, which affects the wave as per distance travelled, mechanical
155 damping effects the wave as per the number of the cycle completed while travelling unit
156 distances. As the higher modes or high frequency wave has to complete more number cycle for
157 travelling the same lengths, mechanical damping affects them severely. The traditional data
158 preprocessing techniques are ineffective against these type of energy losses. Furthermore, the
159 traditional techniques perform poorly for high noise contaminated data. In this paper, we have
160 proposed a new seismic data preprocessing technique that normalises the traces in such a way
161 that the energy losses due to radial damping and mechanical damping can be compensated. It
162 also suppresses the noise present in the data more effectively. Furthermore, the proposed
163 method preserves the phase information of each frequency; thus, it can be applied to all the
164 available transformation techniques. Let $u(x, t)$ is the seismogram data in the distance (x) -
165 time (t) domain . The proposed normalisation technique can be achieved by reconstructing the
166 data in the following way :

$$167 \quad U(x, t) = \frac{1}{2\pi} \int_{-\infty}^{+\infty} \left[\int_{-\infty}^{+\infty} \frac{u(x, t)}{|\int_{-\infty}^{+\infty} u(x, t) e^{-i\omega t} dt|} e^{-i\omega t} dt \right] e^{i\omega t} d\omega \quad (1)$$

168 where the circular frequency $\omega = 2\pi f$ and $i = \sqrt{-1}$. In the above expression, $U(x, t)$
169 is the normalised form of the original data $u(x, t)$. To generate the dispersion image for
170 different wavefield transform, one need to use $U(x, t)$ instead of $u(x, t)$. Fig. 4 represents the

171 seismogram corresponds to original data, trace normalised data and data processed using the
172 proposed method. From these images (Fig. 4), we can observe that, compared to the original
173 seismogram, trace normalisation amplifies the signal associated with the furthest sensor while
174 the proposed method does the best job. The trace normalisation can only raise the amplitude of
175 the dominating waves travelling close to median velocities, while the proposed method can
176 augment the amplitude of waves travelling with the entire velocity spectrum. Therefore, the
177 proposed method can obtain an enhanced dispersion image with low noise.

178 **Data availability**

179 All the data used for this research findings are available and can provided based on the
180 request to the corresponding authors.

181 **Acknowledgements**

182 The authors gratefully acknowledge the financial support provided by the Ministry of
183 Human Resource and Development, Govt. of India. Grant No. SB20210856CEMHRD008957.

184

185 **Reference**

- 186 1. Miller, G. . & Pursey, H. On the partition of energy between elastic waves in a semi-
187 infinite solid. *Proc. R. Soc. London. Ser. A. Math. Phys. Sci.* **233**, 55–69 (1955).
- 188 2. Tokimatsu, K., Tamura, S. & Kojima, H. Effects of Multiple Modes on Rayleigh Wave
189 Dispersion Characteristics. *J. Geotech. Eng.* **118**, 1529–1543 (1992).
- 190 3. Park, C. B., Miller, R. D. & Xia, J. Imaging dispersion curves of surface waves on
191 multi-channel record. in *SEG Technical Program Expanded Abstracts 1998* 1377–
192 1380 (Society of Exploration Geophysicists, 1998). doi:10.1190/1.1820161.

- 193 4. Kumar, J. & Naskar, T. A fast and accurate method to compute dispersion spectra for
194 layered media using a modified Kausel-Roësset stiffness matrix approach. *Soil Dyn.*
195 *Earthq. Eng.* **92**, 176–182 (2017).
- 196 5. Yilmaz, Ö. *Seismic Data Analysis. Seismic Data Analysis* (Society of Exploration
197 Geophysicists, 2001). doi:10.1190/1.9781560801580.
- 198 6. McMechan, G. A. & Yedlin, M. J. Analysis of dispersive waves by wave field
199 transformation. *GEOPHYSICS* **46**, 869–874 (1981).
- 200 7. Luo, Y., Xu, Y., Liu, Q. & Xia, J. Rayleigh-wave dispersive energy imaging and mode
201 separating by high-resolution linear Radon transform. *Lead. Edge* **27**, 1536–1542
202 (2008).
- 203 8. Dal Moro, G., Pipan, M., Forte, E. & Finetti, I. Determination of Rayleigh wave
204 dispersion curves for near surface applications in unconsolidated sediments. in *SEG*
205 *Technical Program Expanded Abstracts 2003* vol. 22 1247–1250 (Society of
206 Exploration Geophysicists, 2003).
- 207 9. Park, C. B. Imaging dispersion of MASW data - Full vs. Selective offset scheme. *J.*
208 *Environ. Eng. Geophys.* **16**, 13–23 (2011).
- 209 10. Naskar, T. & Kumar, J. A faster scheme to generate multimodal dispersion plots for
210 Rayleigh wave propagation. *Soil Dyn. Earthq. Eng.* **117**, 280–287 (2019).
- 211 11. Luo, Y. *et al.* Rayleigh-wave mode separation by high-resolution linear Radon
212 transform. *Geophys. J. Int.* **179**, 254–264 (2009).
- 213 12. Shen, C., Wang, A., Wang, L., Xu, Z. & Cheng, F. Resolution equivalence of
214 dispersion-imaging methods for noise-free high-frequency surface-wave data. *J. Appl.*
215 *Geophys.* **122**, 167–171 (2015).

216 13. Ivanov, J., Miller, R., Feigenbaum, D. & Schwenk, J. Benefits of using the high-
217 resolution linear Radon transform with the multichannel analysis of surface waves
218 method. in *SEG Technical Program Expanded Abstracts 2017* 2647–2653 (Society of
219 Exploration Geophysicists, 2017). doi:10.1190/segam2017-17793766.1.

220 **FIGURE CAPTIONS**

221 Fig. 1| Comparison of the velocity spectrum obtained by, a, Phase shift transform with trace
222 normalization (Dal Moro et al.). b, Phase shift transform with proposed normalization. c, f-k
223 transform with trace normalization (Dal Moro et al.). d, f-k transform with proposed
224 normalization. e, τ -p transform with trace normalization (Dal Moro et al.). f, τ -p transform
225 with proposed normalization.

226 Fig. 2| Dispersion image obtained by τ -p transform for, a, Original data with proposed
227 normalization. b, Original data with trace normalised. c, Noisy data with proposed
228 normalization. d, Noisy data with trace normalization.

229 Fig. 3| Dispersion image by HRLRT method using, a, Surfseis software with trace
230 normalization. b, Our own HRLRT code with trace normalization. c, Our own HRLRT code
231 with proposed normalization.

232 Fig. 4| 48 channel common shot gather. a, Recorded field seismogram. b, Trace normalised
233 seismogram. c, Seismogram obtained after employing the proposed normalization.

Figures

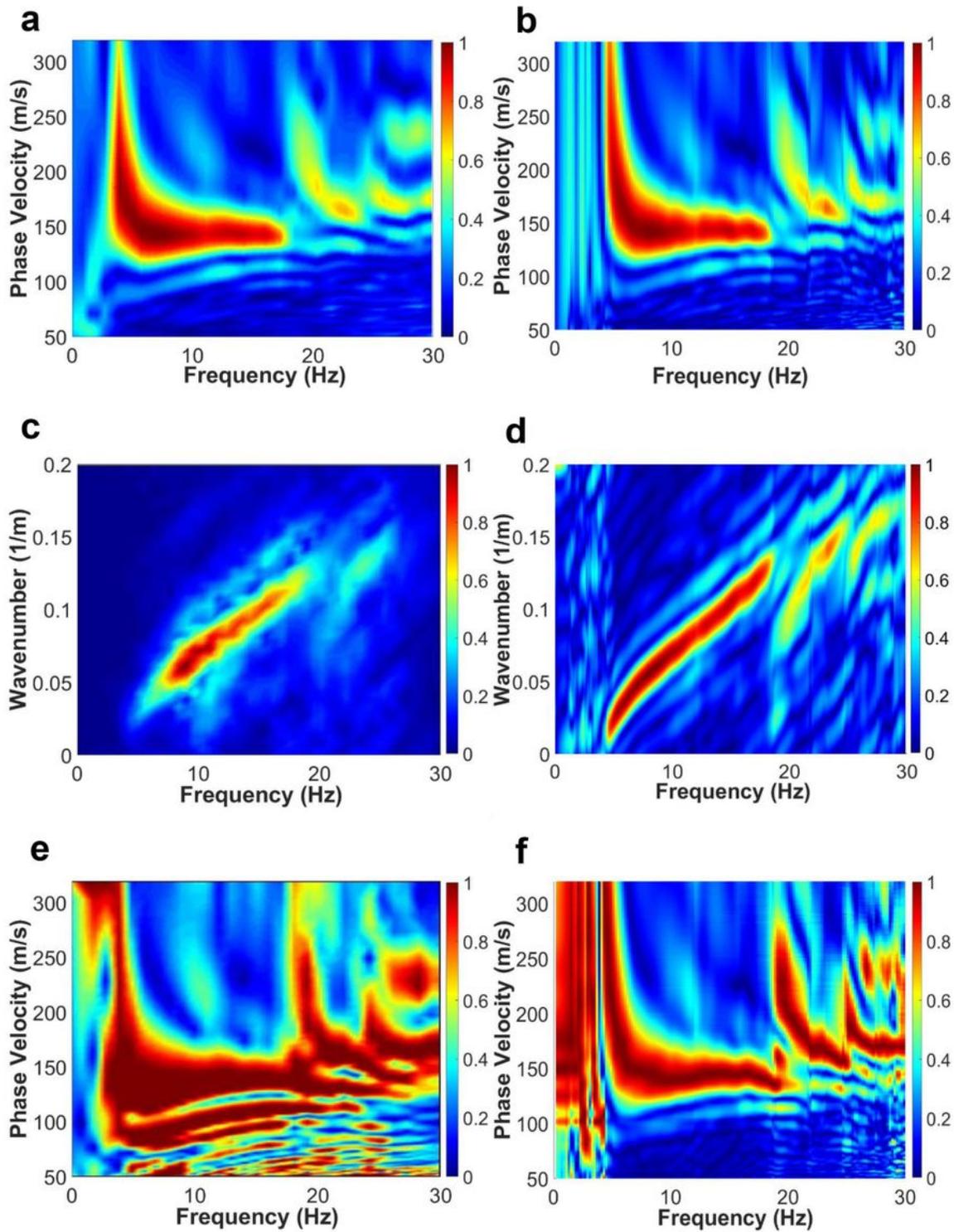


Figure 1

Comparison of the velocity spectrum obtained by, a, Phase shift transform with trace normalization (Dal Moro et al.). b, Phase shift transform with proposed normalization. c, f-k transform with trace

normalization (Dal Moro et al.). d, f-k transform with proposed normalization. e, τ -p transform with trace normalization (Dal Moro et al.). f, τ -p transform with proposed normalization.

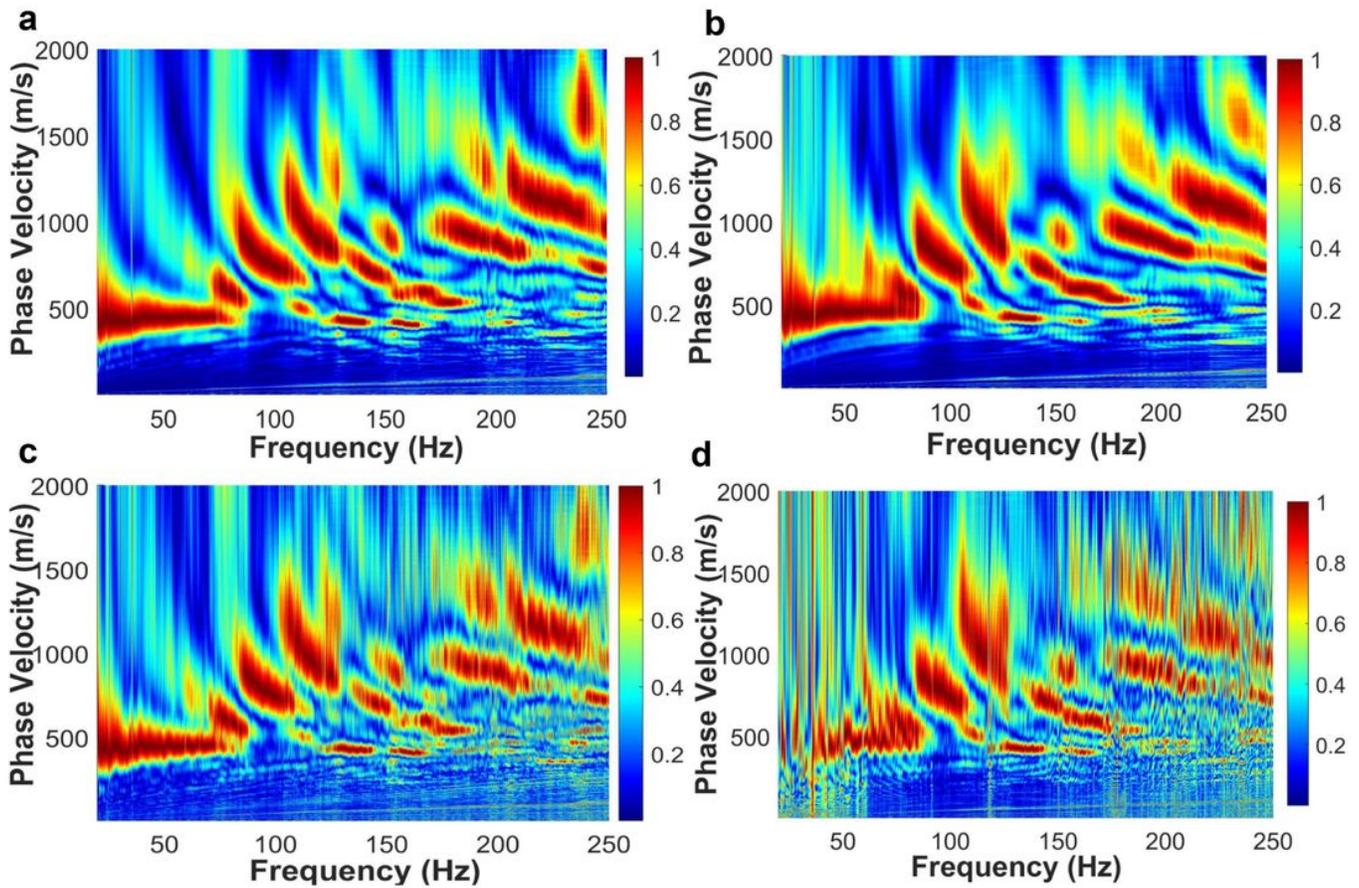


Figure 2

Dispersion image obtained by τ - p transform for, a, Original data with proposed normalization. b, Original data with trace normalised. c, Noisy data with proposed normalization. d, Noisy data with trace normalization.

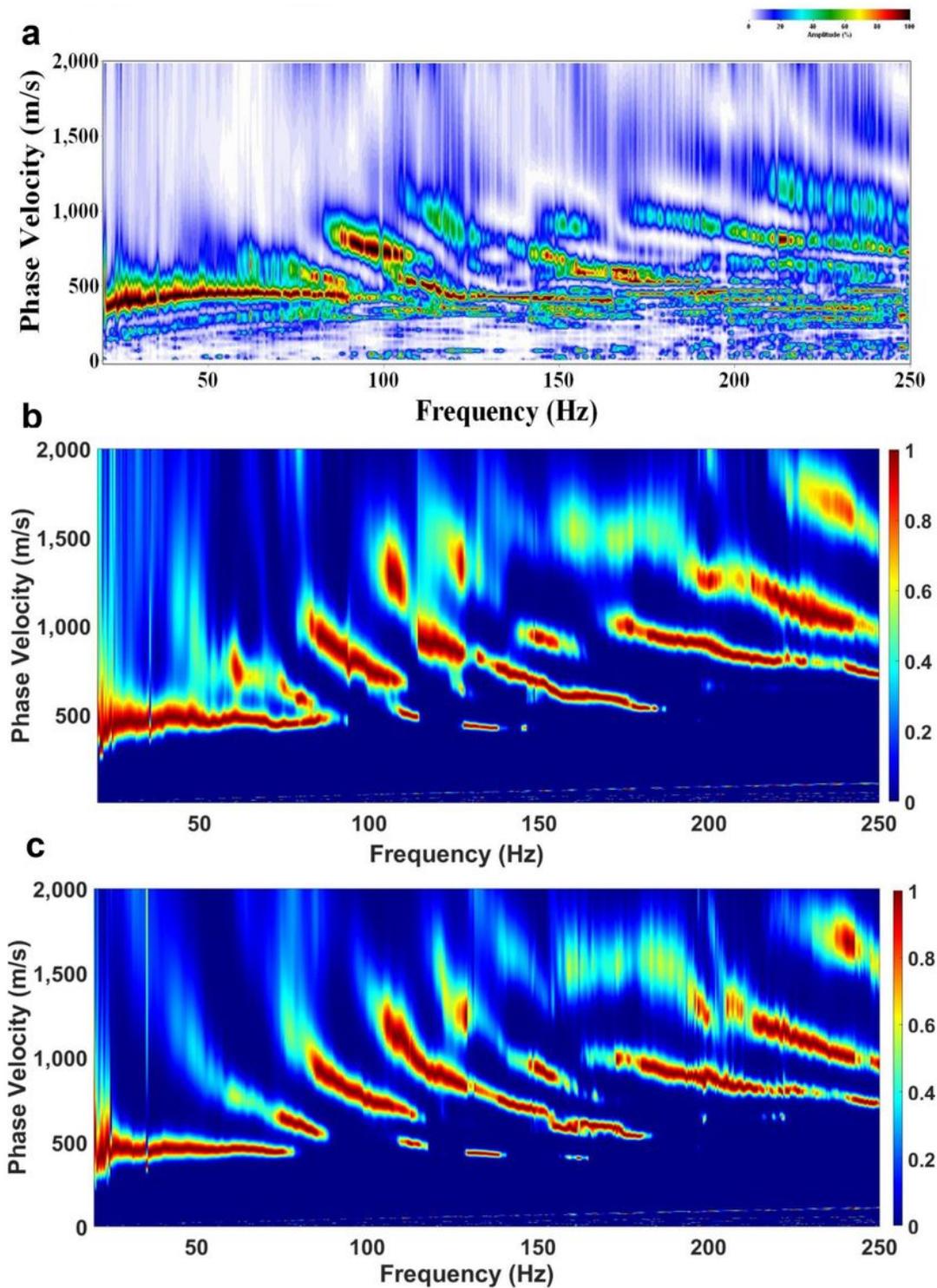


Figure 3

Dispersion image by HRLRT method using, a, Surfseis software with trace normalization. b, Our own HRLRT code with trace normalization. c, Our own HRLRT code with proposed normalization.

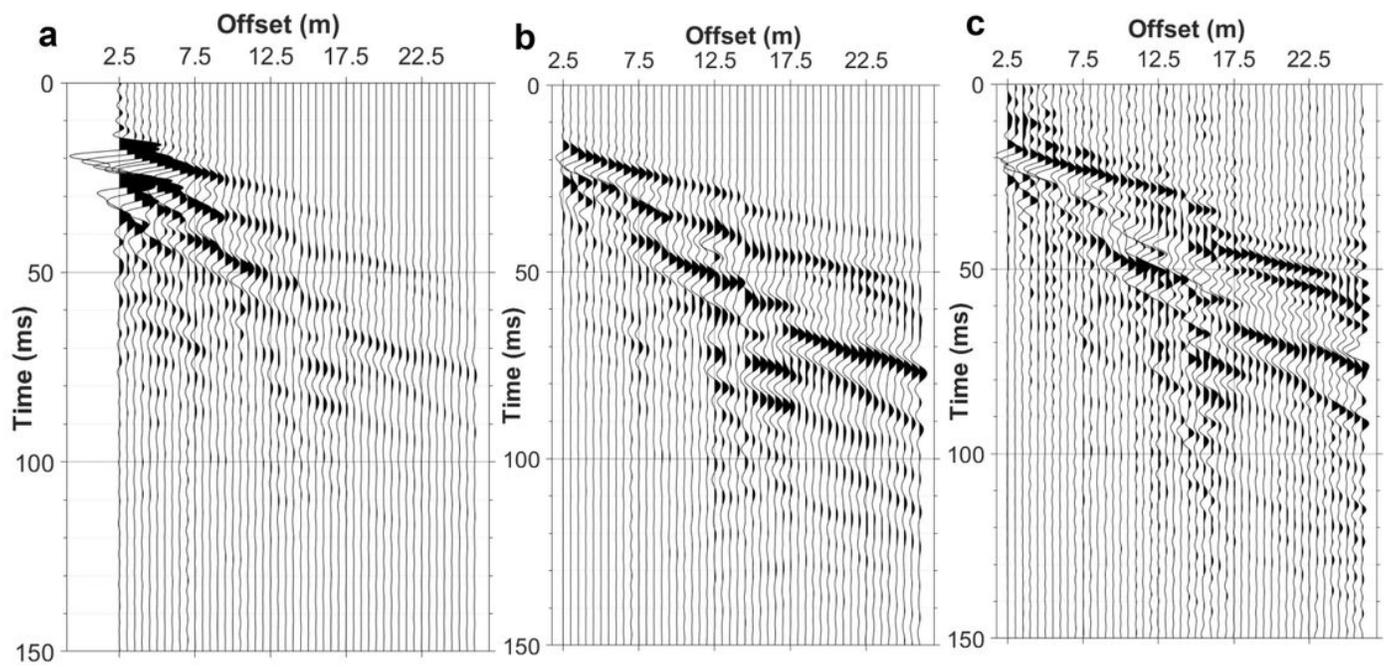


Figure 4

48 channel common shot gather. a, Recorded field seismogram. b, Trace normalised seismogram. c, Seismogram obtained after employing the proposed normalization.